

# Brief communication: Comparison of in-situ ephemeral snow depth measurements over a mixed-use temperate forest landscape

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**Abstract.** The accuracy and precision of snow depth measurements depend on the measuring device and the conditions of the site and snowpack in which it is being used. This study compares collocated snow depth measurements from a magnaprobe snow depth probe and a Federal snow tube in an ephemeral snow environment. We conducted three snow depth sampling campaigns from December 2020 to February 2021 that included 39 open field, coniferous, mixed, and deciduous forest sampling sites in Durham, New Hampshire, United States. For all sampling campaigns and land cover types with a total of 936 paired observations, the magnaprobe snow depth measurements were consistently deeper than the snow tube. There was a 12% average difference between the magnaprobe (14.9 cm) and snow tube (13.2 cm) average snow depths with a greater difference in the forest (1.9 cm) than the field (1.3 cm). This study suggests that snow depth measurements using a Federal snow tube can avoid overprobing with an ephemeral snowpack in forested environment.

**Short Summary.** This study compares snow depth measurements from two manual instruments in a field and forest. Snow depths measured using a magnaprobe were typically 1 to 3 cm deeper than those measured using a snow tube. These differences were greater in the forest than the field.

## 1 Introduction

Snow depth is one of the easier snowpack properties to measure in the field and is an observation that can be measured relatively precisely without considerable expertise or expense. Hundreds of snow depth measurements can readily be taken in a single day and automated samplers can substantially increase that number (Sturm and Holmgren 2018). In-situ snow depth observations can be measured manually or automatically. While automated measurements are increasing in use (Bongio et al. 2021; Kinar and Pomeroy 2015; Kopp et al. 2019), in-situ measurements remain the mainstay of data collection research and operations (Kinar and Pomeroy 2015; Pirazzini et al. 2018). Manual in-situ snow depth measurements are typically made using snow stakes, rulers, or narrow diameter snow probes (Kinar and Pomeroy 2015; Pirazzini et al. 2018). Snow tube samplers, which have been in use since the 1930s, also measure snow depth. The magnaprobe, an automatic snow depth probe that records snow

38 depth and GPS measurements, has considerably increased the number of georeferenced snow depth observations  
39 that can be made in a single day and is used extensively for snow depth research campaigns (Sturm and Holmgren  
40 2018; Walker et al. 2020). Measurement variability and errors are sometimes reduced by repeating the  
41 measurement, typically three times (Leppänen et al. 2016). Because snow depth is assumed to have greater spatial  
42 variability than snow density (Elder et al. 1998), a snow survey often makes numerous snow depth measurements  
43 per snow density measurement then combines to obtain snow water equivalent (SWE) (López-Moreno et al. 2013).  
44 If depth can be well constrained, then density becomes the source of uncertainty (Raleigh and Small 2017).

45

46 SWE measurement errors associated with snow tube samplers are relatively well understood and characterized.  
47 Known issues include biases as compared to snow pit measurements (Dixon and Boon, 2012; Farnes et al., 1983;  
48 Goodison, 1978; Sturm et al., 2010), accuracies around +/- 5% to 10% for an individual instrument, and  
49 differences among SWE from different snow tube models (e.g., the Meteorological Service of Canada, the Federal  
50 or Mt. Rose, the Adirondack, and the Snow-Hydro) that can exceed 10% (Farnes et al. 1983). Less is understood  
51 about the errors in snow depth measurements. Lopez-Moreno et al.'s (2020) comparison of nine snow core  
52 samplers found that snow depths were relatively consistent when taken over a paved surface. However, over  
53 uneven ground, the snow depth differences among samplers was much greater and replicate snow depth  
54 measurements had larger variability as compared to the snow density. The magnaprobe, which measures snow  
55 depth with a precision of less than 0.1 mm, has the potential for low biases if its basket settles into soft surface  
56 snow, but those biases are typically less than 1 cm (Sturm and Holmgren 2018). When the rod penetrates the  
57 substrate (over-probing), the error depends on the ground surface and the operation. Solid or frozen ground  
58 surfaces have negligible over-probing, but unfrozen natural surfaces may have considerable penetration (Derry et  
59 al. 2009) with biases on the order of 5 to 10 cm (Berezovskaya and Kane 2007; Sturm and Holmgren 2018). These  
60 errors can have profound effects on SWE estimates in shallow snow environments and represent a challenge for  
61 error accounting in hydrological modelling.

62

63 The goal of this brief study is to determine 1) if the magnitude of the snow depth measurements using a  
64 magnaprobe and a Federal tube are significantly different in an ephemeral snow environment with shallow snow  
65 and 2) if the differences vary by land cover type. We hypothesize that the snow depth measurements from the  
66 magnaprobe will be deeper than those from the snow tube. This hypothesis is based on the understood errors and  
67 biases associated with each the magnaprobe and the Federal tube, including the smaller surface area of the probe  
68 which allows for greater penetration through snowpacks and leaf litter. Three snow depth sampling campaigns  
69 were conducted from December 2020 to March 2021 over field and forest plots at Thompson Farm in Durham,  
70 New Hampshire, USA.

## 71 **2 Site, Methods, and Data**

### 72 **2.1 Study Site**

73 This study was conducted at the University of New Hampshire's Thompson Farm Research Observatory in  
74 southeast New Hampshire, United States (N 43.11°, W 70.95°, 35 m above sea level, ASL). The site has mixed  
75 hardwood forest and open field land covers (Perron et al. 2004) that are characteristic of the region (**Fig. 1**). The

76 agricultural fields are managed pasture grass with unmown grass in local areas. The deciduous, mixed, and  
77 coniferous forest is composed primarily of white pine (*Pinus strobus*), northern red oak (*Quercus rubra*), red  
78 maple (*Acer rubrum*), shagbark hickory (*Carya ovata*), and white oak (*Quercus alba*) (Perron et al. 2004). The  
79 forest soils are classified as Hollis/Charlton very stony fine sandy loam and well-drained; field soils are  
80 characterized as Scantic silt-loam and poorly drained.

81

82 In-situ sampling was conducted at 39 sites located along three parallel transects (**Fig. 1**). The approximately 145  
83 m long transects were laid out from east to west. The transects were separated by approximately 10 m, north to  
84 south. From east to west, each transect started in the open field area, then transitioned to the coniferous, then  
85 mixed, and finally, deciduous forested areas. Each of the three transects had 13 sampling sites; four sites were in  
86 the open field area, three in the coniferous forest, three in the mixed forest, and three in the deciduous forest,  
87 which were each marked with a stake. The stake locations were geolocated using a Trimble<sup>®</sup> Geo7X GNSS  
88 Positioning Unit and Zephyr<sup>™</sup> antenna with an estimated horizontal uncertainty of 2.51 cm (standard deviation  
89 0.95 cm) and 4.17 cm (standard deviation 4.60 cm) for the field and forest, respectively, after differential  
90 correction. Three soil frost tubes were located in the field approximately 25 m south of the field transect and  
91 another three in the forest about 100 m southwest of the study area.

## 92 **2.2 In-Situ Sampling Methods**

93 Snow depth was measured using a magnaprobe and a Federal snow sampler, also known as a snow tube. The  
94 Federal snow tube with its long operational history (Clyde 1932) served as a historical reference against the  
95 magnaprobe. A magnaprobe consists of an avalanche probe-like rod of about 1.5 m in length that contains a  
96 magnetostrictive device and a sliding magnetic disk-shaped basket with a 25 cm diameter. The rod has a 1.27 cm  
97 diameter with an affixed tip that tapers to a point to help penetrate ice layers. The magnaprobe was operated by  
98 inserting the pole into a snowpack until the tip of the pole reached the ground surface, allowing the basket to slide  
99 down to float on top of the snow. A handheld portable keypad connected to a datalogger recorded the snow depth  
100 between the tip of the pole and the bottom of the basket.

101

102 A Federal snow sampler is an aluminium tube, about 76 cm in length with a 4.13 cm inner diameter, that is used  
103 to measure snow depth and SWE (Clyde 1932). To measure snow depth, the snow tube was inserted vertically  
104 into the snowpack until it reached the ground, and a depth was read at eye level. Snow depth was recorded to the  
105 nearest 0.5 cm. To measure snow density, the snow tube was then lifted out of the snowpack, using a spatula as  
106 needed to ensure that snow did not fall out of the tube. The snow and snow tube were weighed using a digital  
107 hanging scale (CCi HS-6 Electronic Scale, 2-gram resolution).

108

109 Sampling campaigns were conducted on 18 December 2020, 4 February 2021, and 24 February 2021. A total of  
110 936 paired magnaprobe and Federal snow tube snow depth observations were collected during the three sampling  
111 campaigns. At each of the 39 sampling locations, nine measurements were made in a 1x1 m area. At each location,  
112 a 1x1 m square polyvinyl chloride (PVC) grid was placed on the snow surface with one vertex located coincident  
113 with a stake. The orientation of two adjacent sides of the grid was recorded using a compass. Nine magnaprobe  
114 depth measurements were made at an approximately even spacing within the grid. Immediately after the

115 magnaprobe measurements, snow tube snow depth measurements were made at the same nine locations by  
116 positioning the snow tube directly over each magnaprobe sampling location. At a 10<sup>th</sup> location within each 1x1 m  
117 grid, the snow tube was used to make a snow density measurement. For the 24 February 2021 campaign, after the  
118 magnaprobe measurements were completed for the two northern transects, the instrument was transferred to a  
119 new operator who made measurements on the southernmost transect (Transect 1). Transect 1 data for that date  
120 were removed from the analysis because the QA/QC process identified notable errors for observations from that  
121 transect.

122

123 Moultrie Wingscapes Birdcam Pro Field Cameras were used to capture images of the snowpack every 15 minutes  
124 relative to a 1.5 meter marked PVC pole following the method used in NASA's 2020 SnowEx field camera  
125 campaign in Grand Mesa, CO (personal communication, 16 November 2020). Three cameras were used; one was  
126 in the open field, one was in the coniferous forest, and one was in the deciduous forest (**Fig. 1**). Snow depth was  
127 derived by manual inspection of the photos and recorded to the nearest cm.

### 128 **2.3 Ancillary Soils and Vegetation Cover Data**

129 Daily soil frost depth data were collected at field and forest locations at the Thompson Farm Research Observatory  
130 using Cold Regions Research and Engineering Laboratory style frost tubes (Gandahl 1957). The frost tubes have  
131 flexible, polyethylene inner tubing filled with methylene blue dye whose color change is easy to differentiate  
132 when extruded from ice. The outer tubing consists of PVC pipe installed between 0.4 to 0.5 m below the soil  
133 surface. The field and forest sites each had three soil frost tubes.

134

135 Leaf litter depth was measured on 2 April 2021 after the spring snowmelt. The leaf litter depth was measured at  
136 each snow depth sample location. Sampling was conducted using a PVC collar or round ring that is 8 cm in depth  
137 and 10 cm in diameter (Kaspari and Yanoviak 2008). The collar was placed in the leaf litter and was pushed down  
138 until it was through the leaf litter layer. If sticks or larger stones were in the way, they were either carefully  
139 removed or the collar was moved slightly to an adjacent location. Measurements were taken using a wooden ruler  
140 at four cardinal points in the collar. The four measurements were recorded and their average to the nearest cm  
141 was used as the final litter depth. The range of leaf litter depths measured in the forest using the collar was typically  
142 3 to 7 cm with an average leaf litter depth of 3.9 cm. Magnaprobe leaf litter penetration depth measurements, also  
143 made on 2 April 2021 in the forest, had an average value of 5.8 cm.

## 144 **3 Results**

145 The three sampling campaigns, 18 December 2020, 4 February 2021, and 24 February 2021, all had shallow  
146 snowpacks. The field camera observations indicate that the snowpacks had similar depths, between 10 and 15 cm,  
147 on the three sampling dates with modestly deeper snow in the field than the forest. The deepest snow was on 4  
148 February 2021 with 15 cm in the field and 9.3 cm in the forest. Between the 18 December and the 4 February  
149 sampling campaigns, there was a melt event in which the entire 10 cm snowpack on 18 December ablated. The  
150 next significant snowfall event (15 cm) occurred on 1 February 2021. The snowpack experienced little additional  
151 accumulation or ablation between 4 February and 24 February. The 4 February (0.15 g/cm<sup>3</sup>) and 24 February

152 (0.20– 0.24 g/cm<sup>3</sup>) snowpack density values were higher than those in December (~ 0.10 g/cm<sup>3</sup>). There were  
153 shallow soil frost depths (< 4 cm) during the early winter 18 December campaign in the forest and the field.  
154 Deeper soil frost depths of 15.1 cm in the field and 5.9 cm in the forest occurred on 4 February 2021, with similar  
155 soil frost conditions on 24 February 2021.

### 156 **3.1 Magnaprobe vs. Snow Tube**

157 The full experiment yielded individual 936 pairs of snow depth measurements from the snow tube and the  
158 magnaprobe (**Fig. 2a**). For the comparison between measurement techniques, the orthogonal Deming regression  
159 method was applied to consider measurement errors in both variables. Overall, there was moderate agreement ( $R$   
160 = 0.74) between the two datasets for all three sampling campaigns (**Table S1**). The snow depths measured by the  
161 magnaprobe (14.9 cm average snow depth) were deeper than the snow tube (13.2 cm average snow depth) with  
162 an overall bias of 1.7 cm. The magnaprobe snow depth was at least 0.5 cm deeper than the snow tube in 74% of  
163 the 936 measurement pairs. Only 6.3% of the pairs had snow tube snow depths exceeding magnaprobe snow  
164 depths by 0.5 cm or more. Conversely, 7.4% of the pairs' magnaprobe snow depths were over 5.0 cm deeper than  
165 the snow tube. In eight pairs of measurements, when the magnaprobe measured snow depth greater than 15 cm,  
166 the magnaprobe snow depths were more than double the snow tube snow depth.

167

168 The majority of the nine sampling locations in each grid had magnaprobe snow depth values that were deeper than  
169 those measured using the snow tube. For all the grids, an average of 8.7, 7.7, and 7.0 out of the nine sampling  
170 locations had deeper magnaprobe snow depths on 18 December 2020, 4 and 24 February 2021, respectively. As  
171 hypothesized, t-test results showed that the magnaprobe snow depth values were significantly greater than those  
172 measured using the snow tube for 39 and 31 of the 39 sampling locations on 18 December 2020 and 4 February  
173 2021, respectively, but only 11 out of the 26 sampling locations on 24 February 2021. The mean differences were  
174 2.3, 1.4, and 1.6 cm, with root mean square difference (RMSD) values of 3.0, 2.3, and 3.3 cm, on 18 December  
175 2020, 4 and 24 February 2021, respectively, which is on the order of 15 to 25% of the overall depth observed  
176 during these campaigns. Despite the biases, the average within cell snow depth variability was nearly identical for  
177 the magnaprobe and the snow tube in the field (1.3 cm standard deviation for the magnaprobe). In the forest, the  
178 Magaprobe's 2.0 cm within-cell standard deviation modestly exceeded the snow tube's 1.5 cm standard deviation.  
179 A slightly reduced agreement was found on 24 February when there was a 1 to 4 cm thick ice layer at the bottom  
180 of the snowpack in local depressions.

181

182 The overall agreement between the snow tube and magnaprobe was better when the nine measurements within a  
183 single 1x1 m grid cell were averaged at each of the sampling locations (**Fig. 2b** and **Table S1**). There is a notable  
184 improvement in grid cell statistics, and the correlation is stronger (overall  $R = 0.87$ ), with slopes closer to one,  
185 intercepts closer to zero, and the RMSD values reduced to 2.5 cm or less. Although averaging has no impact on  
186 the overall bias, the range of differences among pairs narrowed. The difference between the magnaprobe and the  
187 snow tube is typically constrained to less than 3 cm with a limited number of outliers. The magnaprobe snow  
188 depth was at least 0.5 cm deeper than the snow tube in almost all grid cells (86.7%), but only three grid cells had  
189 differences greater than 5 cm. Among the grid averaged magnaprobe snow depths, there were no instances in  
190 which there was a doubling of snow depth when compared to the snow tube measurements.

### 191 3.2 Magnaprobe vs. Snow Tube by Land type

192 The magnaprobe and snow tube snow depths differ by land type, with the field having deeper snow and more  
193 spatial variability than the forest land types (Fig. 3). Among the three forest types, the deepest snow was in the  
194 deciduous-dominated forest, with mixed and coniferous forest having similar snow depths. The mean difference  
195 between the magnaprobe and snow tube snow depths is a modest 1.3 cm in the field and 1.9 cm in the forest, with  
196 differences of 1.9, 2.0, and 1.9 cm in the deciduous, mixed, and coniferous land types, respectively. However, the  
197 differences between the magnaprobe and snow tube snow depths in the forest were higher on 18 December (2.5  
198 cm), than on 4 February, and 24 February, 1.7, and 1.4 cm, respectively. Based on t-test results, the magnaprobe  
199 measured significantly deeper snow depth compared to the snow tube in both the field and the forest regardless  
200 of whether individual locations ( $p$ -value  $< 0.001$ ) or grid cell average snow depths ( $p$ -value = 0.02) were used.  
201 Based on Welch's adjusted ANOVA test, there are no significant differences in over-probing among forest land  
202 types ( $p$ -value = 0.24). The RMSD values between the magnaprobe and snow tube snow depths are 3.0 cm (2.3  
203 cm) and 2.5 cm (2.0 cm) for the forest and field sampling sites (grid average values), respectively. Thus, the  
204 sampling method has a different impact in the field than the forest and the RMSD and bias values provide an  
205 indicator of the different errors associated with in-situ measurements based on land type when used for model or  
206 remote sensing validation. While these differences are significant, the average litter depths exceeded the  
207 differences between the magnaprobe and snow tube snow depths in the forest, which were 2.5, 1.7, and 1.4 cm on  
208 18 December, 4 February, and 24 February, respectively.

### 209 4 Discussion and Suggestions

210 This study quantifies the differences between snow depth measurements made with a magnaprobe and with a  
211 snow tube. The differences seem to be primarily associated with greater over-probing by the magnaprobe into  
212 vegetation/organic layers and thawed soils. Though in some cases the large differences could instead be due to  
213 the larger diameter snow tube hitting a branch from a down tree or debris that the magnaprobe bypassed. The  
214 result was that the magnaprobe snow depth measurements were higher than snow tube measurements, with a  
215 greater difference in the forest than in the field. An average of 5 cm bias occurred in the tundra mat during the  
216 Cold Land Processes Experiment (CLPX) Alaska campaign (Sturm and Holmgren 2018). Also in the open tundra  
217 environment found a 7.6 cm average over-probe penetration for approximately 40 cm deep snow (Canada 2018).  
218 Berezovskaya and Kane (2007) also noted over-probing of 5 to 9 cm with a magnaprobe as compared to a snow  
219 tube found a bias in northern Alaska for snow depths between 29 and 48 cm. In this study, the over-probing, 1.3  
220 cm in the field and a 1.9 cm in the forest, was less than previous studies probably due to the lower range of snow  
221 depth and different surface conditions as compared to previous studies.

222

223 We also agree with López-Moreno et al. (2020) that it is important to understand the snowpack and land conditions  
224 for which an individual sampler was designed to select the most appropriate sampler. Understanding leaf litter or  
225 vegetation depths and underlying soils may potentially reduce and help to account for the over-probing errors of  
226 magnaprobe snow depth measurements. Sturm and Holmgren (2018) suggested that operators need to learn to  
227 push a magnaprobe through snow, yet not penetrate it too deeply into underlying vegetation/organic layers by  
228 developing a sense for the base of the snowpack. This recommendation may be difficult to implement (e.g., over

229 soft vegetation) where the probe easily penetrates the vegetation and problematic if multiple operators apply a  
230 different amounts of force (Berezovskaya and Kane 2007). If operators over-probe into the base of the (frozen)  
231 soils, one option is to consistently measure the depths in the same way (which would be snow depth *plus*  
232 vegetation) and then subtract typical vegetation depths in the study area from the depths. When leaf litter is  
233 evident, penetration into the organic layer should be quantified by using independent leaf litter measurements,  
234 preferably using the snow depth sampling instrument, and use to bias correct snow depths.

235

236 As observed in this study, leaf litter and soil frost may differentially impact in-situ snow depth sampling methods.  
237 The earliest sampling campaign had limited soil frost and likely reduced litter compaction. Distinct contributions  
238 of forest leaf litter depth to magnaprobe and snow tube snow depths may occur because the narrow magnaprobe  
239 fully penetrates the leaf litter and the larger diameter snow tube only partially penetrates the litter, or the  
240 magnaprobe may only partially penetrate the leaf litter but the snow tube does not break through the leaf litter.  
241 Partial penetration of the magnaprobe into the leaf litter layer (i.e., over-probing) may vary by the freeze-thaw  
242 state of the duff layer and/or mineral soil layers beneath the leaf litter layer. The horizontally aligned, matted leaf  
243 litter could also limit snow tube penetration. High spatial variability of leaf litter depth could also be a factor,  
244 though this was not quantified here. Thus, the increased differences among in-situ methods in forested areas  
245 observed in this study point to the particular importance of in-situ validation in forested areas and, more generally,  
246 sampling with multiple methods in an area with a nonuniform underlying substrate.

247

248 In summary, there are three major suggestions from this work below.

249 1) With an ephemeral snowpack in forested environment, snow depth measurements using a Federal  
250 snow tube likely avoid over-probing that can frequently occur when a magnaprobe is used.

251 2) The use of the average of multiple point samples within a grid is recommended instead of single  
252 measurements, because the average of multiple point samples can reduce the point-to-point variability  
253 and spatial representativeness errors.

254 3) Measurements of vegetation, leaf litter, and soil frost can help to account for the errors of in-situ snow  
255 depth observations, particularly when using a magnaprobe.

256

## 257 **5 Conclusion**

258 Manual in-situ sampling snow depth measurements can be made quickly and easily, but making consistent,  
259 representative, and unbiased measurements can be challenging when the surface is irregular, vegetation/organic  
260 layers and unfrozen soils result in over-probing, and the leaf litter compacts during the winter. This study  
261 quantified the differences between snow depth measurements made with a magnaprobe and a Federal snow tube  
262 in a mixed-use temperate forest landscape with ephemeral snowpack. For all sampling campaigns and land cover  
263 types, the magnaprobe snow depth measurements (mean 14.9 cm) were usually, but not always deeper than the  
264 snow tube measurements (13.2 cm) and had a 1.7 cm average difference. For these shallow snowpacks, this  
265 amounts to a 12% difference, but in a deeper snow pack the relative impact of this difference would be much  
266 smaller. Biases were significantly higher in the forest (1.9 cm) than the field (1.3 cm). The difference between the  
267 two instruments was 50% higher in early winter campaign than the later campaigns. The differences among  
268 measurement techniques in this present study reflect the current study area, surface conditions for a single season,

269 and the operation of the instruments by this project team. Further studies to understand the errors from in-situ  
270 sampling using snow probe are warranted in various snow environments with different vegetation and soil  
271 conditions to provide guidance on best practices for using in-situ snow probe datasets under conditions when over-  
272 probing is likely.

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### 279 **Data Availability**

280 The in-situ snow observations are available in supporting information.

### 281 **Author Contributions**

282 HP, JJ, EB, AH, FS, MP, and EC designed the research. HP, CW, JJ, AH, FS, MP, EB, and EC conducted field  
283 work to obtain lidar and/or in-situ snow observations. HP, CW, JJ, EB, AH, and MP and performed the analysis.  
284 HP, EC, and AH produced the figures. HP, JJ, EB, and EC wrote the initial draft. All authors contributed to  
285 manuscript review and editing.

### 286 **Competing Interests**

287 The authors declare that they have no conflict of interest.

### 288 **References**

- 289 Berezovskaya, S., and D. L. Kane, 2007: Measuring snow water equivalent for hydrological applications: part 1,  
290 accuracy of observations. *Proceedings of the 16 th International Northern Research Basins Symposium and*  
291 *Workshop, Petrozavodsk, Russia*, 29-37.
- 292 Bongio, M., A. N. Arslan, C. M. Tanis, and C. De Michele, 2021: Snow depth time series retrieval by time-lapse  
293 photography: Finnish and Italian case studies. *The Cryosphere*, **15**, 369-387.
- 294 Canada, E. a. C. C., 2018: Snow depth data from trail valley creek. .
- 295 Clyde, G. D., 1932: Utah snow sampler and scales for measuring water content of snow.
- 296 Derry, J., D. Kane, M. Lilly, and H. Toniolo, 2009: Snow-course measurement methods, North Slope, Alaska.  
297 *University of Alaska Fairbanks, Water and Environmental Research Center, Report INE/WERC*, **15**.
- 298 Elder, K., W. Rosenthal, and R. E. Davis, 1998: Estimating the spatial distribution of snow water equivalence in  
299 a montane watershed. *Hydrological Processes*, **12**, 1793-1808.
- 300 Farnes, P. E., B. E. Goodison, N. R. Peterson, and R. P. Richards, 1983: Metrication of manual snow sampling  
301 equipment. *Final report Western Snow Conference*, 19-21.
- 302 Gandahl, R., 1957: Determination of the depth of soil freezing with a new frost meter. *Text in Swedish) Rapport*,  
303 **20**, 3-15.
- 304 Kaspari, M., and S. P. Yanoviak, 2008: Biogeography of litter depth in tropical forests: evaluating the phosphorus  
305 growth rate hypothesis. *Functional Ecology*, **22**, 919-923.



306 Kinar, N., and J. Pomeroy, 2015: Measurement of the physical properties of the snowpack. *Rev. Geophys.*, **53**,  
307 481-544.

308 Kopp, M., Y. Tuo, and M. Disse, 2019: Fully automated snow depth measurements from time-lapse images  
309 applying a convolutional neural network. *Science of The Total Environment*, **697**, 134213.

310 Leppänen, L., A. Kontu, H.-R. Hannula, H. Sjöblom, and J. Pulliainen, 2016: Sodankylä manual snow survey  
311 program. *Geoscientific Instrumentation, Methods and Data Systems*, **5**, 163-179.

312 López-Moreno, J. I., and Coauthors, 2013: Small scale spatial variability of snow density and depth over complex  
313 alpine terrain: Implications for estimating snow water equivalent. *Advances in water resources*, **55**, 40-52.

314 López-Moreno, J. I., and Coauthors, 2020: Intercomparison of measurements of bulk snow density and water  
315 equivalent of snow cover with snow core samplers: Instrumental bias and variability induced by observers.  
316 *Hydrological Processes*, **34**, 3120-3133.

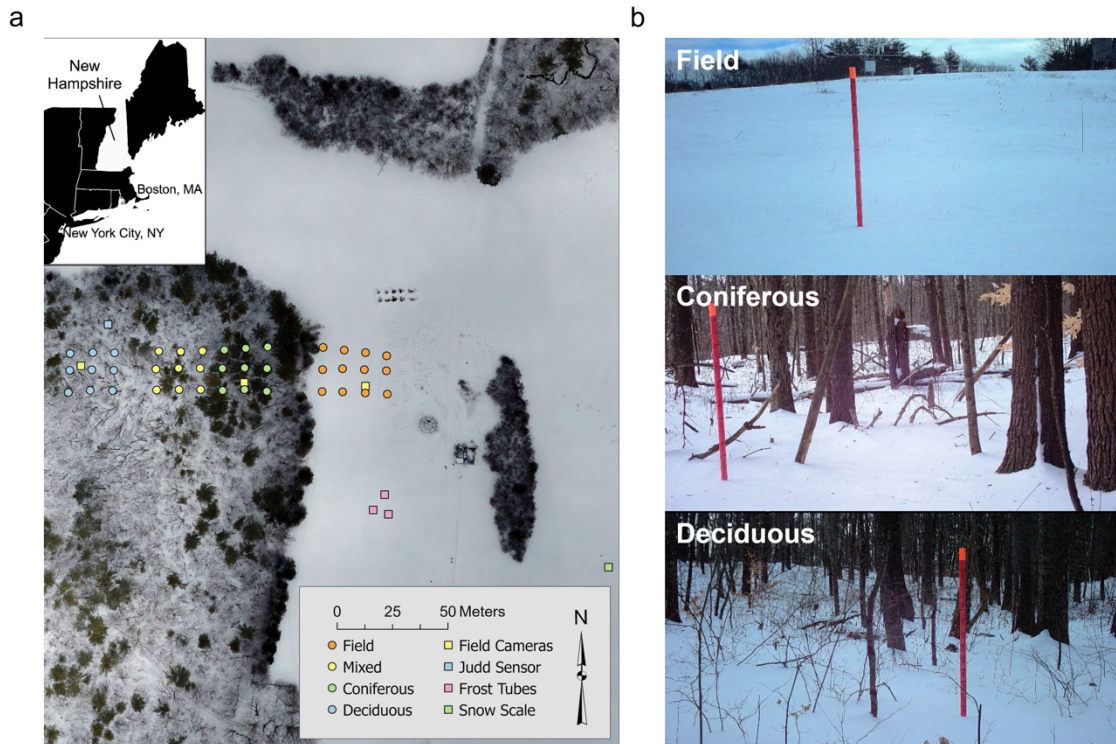
317 Perron, C. J., K. Bennett, and T. D. Lee, 2004: Forest stewardship plan: Thompson farm. NH:  
318 University of New Hampshire. Ossipee Mountain Land Company, West Ossipee.  
319 <https://colsa.unh.edu/sites/default/files/thompson-farm-plan.pdf>.

320 Pirazzini, R., and Coauthors, 2018: European in-situ snow measurements: Practices and purposes. *Sensors*, **18**,  
321 2016.

322 Raleigh, M. S., and E. E. Small, 2017: Snowpack density modeling is the primary source of uncertainty when  
323 mapping basin-wide SWE with lidar. *Geophysical Research Letters*, **44**, 3700-3709.

324 Sturm, M., and J. Holmgren, 2018: An automatic snow depth probe for field validation campaigns. *Water  
325 Resources Research*, **54**, 9695-9701.

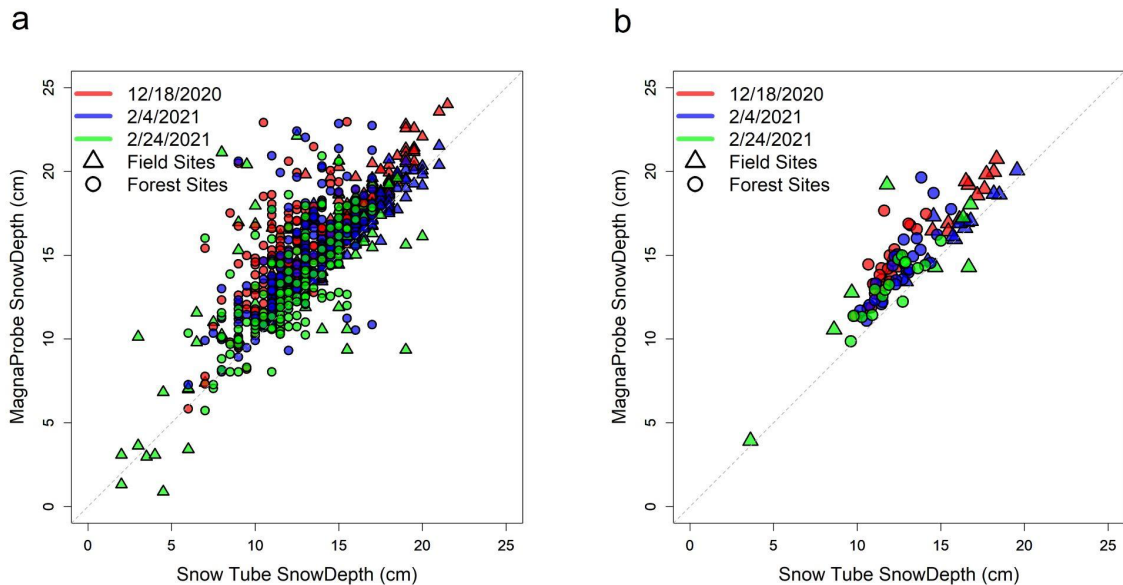
326 Walker, B., E. J. Wilcox, and P. Marsh, 2020: Accuracy assessment of late winter snow depth mapping for tundra  
327 environments using Structure-from-Motion photogrammetry. *Arctic Science*, 1-17.  
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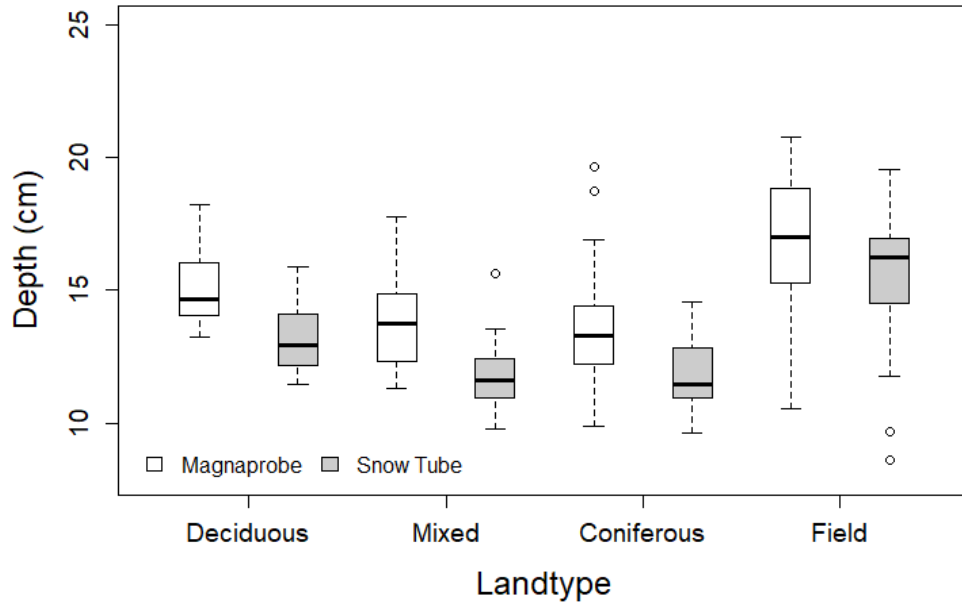
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332 **Figure 1:** (a) The 4 February 2021 aerial optical image of Thompson Farm, Durham NH, USA showing both  
 333 forest and field regions with snow sampling sites in the field, coniferous, mixed, and deciduous forested areas as  
 334 well as the locations of the CRREL soil frost tubes; and field cameras. (b) Field camera images in the field,  
 335 coniferous, and deciduous forested areas taken on 4 February 2021 by the field cameras.



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337 **Figure 2:** Comparison of snow depths measured by magnaprobe and snow tube for the three sampling campaigns  
 338 using (a) the sampling individual points ( $n = 936$ ) and (b) using grid cell average values ( $n=104$ ).



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340 **Figure 3:** Boxplots of snow depths by land type measured by the magnaprobe and the snow tube for the three  
 341 sampling campaigns using the grid cell average values.

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