

1 **Brief communication: Comparison of in-situ ephemeral snow**
2 **depth measurements over a mixed-use temperate forest**
3 **landscape**

4
5 Holly Proulx¹, Jennifer M. Jacobs^{1,2}, Elizabeth A. Burakowski², Eunsang Cho^{3,4}, Adam G.
6 Hunsaker^{1,2}, Franklin B. Sullivan², Michael Palace^{2,5}, Cameron Wagner¹

7 ¹Department of Civil and Environmental Engineering, University of New Hampshire, Durham, NH, 03824, USA

8 ²Earth Systems Research Center, Institute for the Study of Earth, Oceans, and Space, University of New
9 Hampshire, Durham, NH, 03824, USA

10 ³Hydrological Sciences Laboratory, NASA Goddard Space Flight Center, Greenbelt, MD, USA

11 ⁴Earth System Science Interdisciplinary Center, University of Maryland, College Park, MD, USA

12 ⁵Department of Earth Sciences, University of New Hampshire, Durham, NH, 03824, USA

13 *Correspondence to:* Jennifer M. Jacobs (Jennifer.jacobs@unh.edu)

14 **Abstract.** The accuracy and precision of snow depth measurements depend on the measuring device and the
15 conditions of the site and snowpack in which it is being used. This study compares collocated snow depth
16 measurements from a magnaprobe snow depth probe and a Federal snow tube in an ephemeral snow environment.
17 We conducted three snow depth sampling campaigns from December 2020 to February 2021 that included 39
18 open field, coniferous, mixed, and deciduous forest sampling sites in Durham, New Hampshire, United States.
19 For all sampling campaigns and land cover types with a total of 936 paired observations, the magnaprobe snow
20 depth measurements were consistently deeper than the snow tube. There was a 12% average difference between
21 the magnaprobe (14.9 cm) and snow tube (13.2 cm) average snow depths with a greater difference in the forest
22 (1.9 cm) than the field (1.3 cm). This study suggests that snow depth measurements using a Federal snow tube
23 can avoid overprobing with an ephemeral snowpack in forested environment.

24
25 **Short Summary.** This study compares snow depth measurements from two manual instruments in a field and
26 forest. Snow depths measured using a magnaprobe were typically 1 to 3 cm deeper than those measured using a
27 snow tube. These differences were greater in the forest than the field.

28 **1 Introduction**

29 Snow depth is one of the easier snowpack properties to measure in the field and is an observation that can be
30 measured relatively precisely without considerable expertise or expense. Hundreds of snow depth measurements
31 can readily be taken in a single day and automated samplers can substantially increase that number (Sturm and
32 Holmgren 2018). In-situ snow depth observations can be measured manually or automatically. While automated
33 measurements are increasing in use (Bongio et al. 2021; Kinar and Pomeroy 2015; Kopp et al. 2019), in-situ
34 measurements remain the mainstay of data collection research and operations (Kinar and Pomeroy 2015; Pirazzini
35 et al. 2018). Manual in-situ snow depth measurements are typically made using snow stakes, rulers, or narrow
36 diameter snow probes (Kinar and Pomeroy 2015; Pirazzini et al. 2018). Snow tube samplers, which have been in
37 use since the 1930s, also measure snow depth. The magnaprobe, an automatic snow depth probe that records snow

38 depth and GPS measurements, has considerably increased the number of georeferenced snow depth observations
39 that can be made in a single day and is used extensively for snow depth research campaigns (Sturm and Holmgren
40 2018; Walker et al. 2020). Measurement variability and errors are sometimes reduced by repeating the
41 measurement, typically three times (Leppänen et al. 2016). Because snow depth is assumed to have greater spatial
42 variability than snow density (Elder et al. 1998), a snow survey often makes numerous snow depth measurements
43 per snow density measurement then combines to obtain snow water equivalent (SWE) (López-Moreno et al. 2013).

44 If depth can be well constrained, then density becomes the source of uncertainty (Raleigh and Small 2017).

45

46 SWE measurement errors associated with snow tube samplers are relatively well understood and characterized.
47 Known issues include biases as compared to snow pit measurements (Dixon and Boon, 2012; Farnes et al., 1983;
48 Goodison, 1978; Sturm et al., 2010), accuracies around +/- 5% to 10% for an individual instrument, and
49 differences among SWE from different snow tube models (e.g., the Meteorological Service of Canada, the Federal
50 or Mt. Rose, the Adirondack, and the Snow-Hydro) that can exceed 10% (Farnes et al. 1983). Less is understood
51 about the errors in snow depth measurements. Lopez-Moreno et al.'s (2020) comparison of nine snow core
52 samplers found that snow depths were relatively consistent when taken over a paved surface. However, over
53 uneven ground, the snow depth differences among samplers was much greater and replicate snow depth
54 measurements had larger variability as compared to the snow density. The magnaprobe, which measures snow
55 depth with a precision of less than 0.1 mm, has the potential for low biases if its basket settles into soft surface
56 snow, but those biases are typically less than 1 cm (Sturm and Holmgren 2018). When the rod penetrates the
57 substrate (over-probing), the error depends on the ground surface and the operation. Solid or frozen ground
58 surfaces have negligible over-probing, but unfrozen natural surfaces may have considerable penetration (Derry et
59 al. 2009) with biases on the order of 5 to 10 cm (Berezovskaya and Kane 2007; Sturm and Holmgren 2018). These
60 errors can have profound effects on SWE estimates in shallow snow environments and represent a challenge for
61 error accounting in hydrological modelling.

62

63 The goal of this brief study is to determine 1) if the magnitude of the snow depth measurements using a
64 magnaprobe and a Federal tube are significantly different in an ephemeral snow environment with shallow snow
65 and 2) if the differences vary by land cover type. We hypothesize that the snow depth measurements from the
66 magnaprobe will be deeper than those from the snow tube. This hypothesis is based on the understood errors and
67 biases associated with each the magnaprobe and the Federal tube, including the smaller surface area of the probe
68 which allows for greater penetration through snowpacks and leaf litter. Three snow depth sampling campaigns
69 were conducted from December 2020 to March 2021 over field and forest plots at Thompson Farm in Durham,
70 New Hampshire, USA.

71 **2 Site, Methods, and Data**

72 **2.1 Study Site**

73 This study was conducted at the University of New Hampshire's Thompson Farm Research Observatory in
74 southeast New Hampshire, United States (N 43.11°, W 70.95°, 35 m above sea level, ASL). The site has mixed
75 hardwood forest and open field land covers (Perron et al. 2004) that are characteristic of the region (Fig. 1). The

76 agricultural fields are managed pasture grass with unmown grass in local areas. The deciduous, mixed, and
77 coniferous forest is composed primarily of white pine (*Pinus strobus*), northern red oak (*Quercus rubra*), red
78 maple (*Acer rubrum*), shagbark hickory (*Carya ovata*), and white oak (*Quercus alba*) (Perron et al. 2004). The
79 forest soils are classified as Hollis/Charlton very stony fine sandy loam and well-drained; field soils are
80 characterized as Scantic silt-loam and poorly drained.

81

82 In-situ sampling was conducted at 39 sites located along three parallel transects (**Fig. 1**). The approximately 145
83 m long transects were laid out from east to west. The transects were separated by approximately 10 m, north to
84 south. From east to west, each transect started in the open field area, then transitioned to the coniferous, then
85 mixed, and finally, deciduous forested areas. Each of the three transects had 13 sampling sites; four sites were in
86 the open field area, three in the coniferous forest, three in the mixed forest, and three in the deciduous forest,
87 which were each marked with a stake. The stake locations were geolocated using a Trimble® Geo7X GNSS
88 Positioning Unit and Zephyr™ antenna with an estimated horizontal uncertainty of 2.51 cm (standard deviation
89 0.95 cm) and 4.17 cm (standard deviation 4.60 cm) for the field and forest, respectively, after differential
90 correction. Three soil frost tubes were located in the field approximately 25 m south of the field transect and
91 another three in the forest about 100 m southwest of the study area.

92 **2.2 In-Situ Sampling Methods**

93 Snow depth was measured using a magnaprobe and a Federal snow sampler, also known as a snow tube. The
94 Federal snow tube with its long operational history (Clyde 1932) served as a historical reference against the
95 magnaprobe. A magnaprobe consists of an avalanche probe-like rod of about 1.5 m in length that contains a
96 magnetostrictive device and a sliding magnetic disk-shaped basket with a 25 cm diameter. The rod has a 1.27 cm
97 diameter with an affixed tip that tapers to a point to help penetrate ice layers. The magnaprobe was operated by
98 inserting the pole into a snowpack until the tip of the pole reached the ground surface, allowing the basket to slide
99 down to float on top of the snow. A handheld portable keypad connected to a datalogger recorded the snow depth
100 between the tip of the pole and the bottom of the basket.

101

102 A Federal snow sampler is an aluminium tube, about 76 cm in length with a 4.13 cm inner diameter, that is used
103 to measure snow depth and SWE (Clyde 1932). To measure snow depth, the snow tube was inserted vertically
104 into the snowpack until it reached the ground, and a depth was read at eye level. Snow depth was recorded to the
105 nearest 0.5 cm. To measure snow density, the snow tube was then lifted out of the snowpack, using a spatula as
106 needed to ensure that snow did not fall out of the tube. The snow and snow tube were weighed using a digital
107 hanging scale (CCi HS-6 Electronic Scale, 2-gram resolution).

108

109 Sampling campaigns were conducted on 18 December 2020, 4 February 2021, and 24 February 2021. A total of
110 936 paired magnaprobe and Federal snow tube snow depth observations were collected during the three sampling
111 campaigns. At each of the 39 sampling locations, nine measurements were made in a 1x1 m area. At each location,
112 a 1x1 m square polyvinyl chloride (PVC) grid was placed on the snow surface with one vertex located coincident
113 with a stake. The orientation of two adjacent sides of the grid was recorded using a compass. Nine magnaprobe
114 depth measurements were made at an approximately even spacing within the grid. Immediately after the

115 magnaprobe measurements, snow tube snow depth measurements were made at the same nine locations by
116 positioning the snow tube directly over each magnaprobe sampling location. At a 10th location within each 1x1 m
117 grid, the snow tube was used to make a snow density measurement. For the 24 February 2021 campaign, after the
118 magnaprobe measurements were completed for the two northern transects, the instrument was transferred to a
119 new operator who made measurements on the southernmost transect (Transect 1). Transect 1 data for that date
120 were removed from the analysis because the QA/QC process identified notable errors for observations from that
121 transect.

122

123 Moultrie Wingscapes Birdcam Pro Field Cameras were used to capture images of the snowpack every 15 minutes
124 relative to a 1.5 meter marked PVC pole following the method used in NASA's 2020 SnowEx field camera
125 campaign in Grand Mesa, CO (personal communication, 16 November 2020). Three cameras were used; one was
126 in the open field, one was in the coniferous forest, and one was in the deciduous forest (Fig. 1). Snow depth was
127 derived by manual inspection of the photos and recorded to the nearest cm.

128 2.3 Ancillary Soils and Vegetation Cover Data

129 Daily soil frost depth data were collected at field and forest locations at the Thompson Farm Research Observatory
130 using Cold Regions Research and Engineering Laboratory style frost tubes (Gandahl 1957). The frost tubes have
131 flexible, polyethylene inner tubing filled with methylene blue dye whose color change is easy to differentiate
132 when extruded from ice. The outer tubing consists of PVC pipe installed between 0.4 to 0.5 m below the soil
133 surface. The field and forest sites each had three soil frost tubes.

134

135 Leaf litter depth was measured on 2 April 2021 after the spring snowmelt. The leaf litter depth was measured at
136 each snow depth sample location. Sampling was conducted using a PVC collar or round ring that is 8 cm in depth
137 and 10 cm in diameter (Kaspri and Yanoviak 2008). The collar was placed in the leaf litter and was pushed down
138 until it was through the leaf litter layer. If sticks or larger stones were in the way, they were either carefully
139 removed or the collar was moved slightly to an adjacent location. Measurements were taken using a wooden ruler
140 at four cardinal points in the collar. The four measurements were recorded and their average to the nearest cm
141 was used as the final litter depth. The range of leaf litter depths measured in the forest using the collar was typically
142 3 to 7 cm with an average leaf litter depth of 3.9 cm. Magnaprobe leaf litter penetration depth measurements, also
143 made on 2 April 2021 in the forest, had an average value of 5.8 cm.

144 3 Results

145 The three sampling campaigns, 18 December 2020, 4 February 2021, and 24 February 2021, all had shallow
146 snowpacks. The field camera observations indicate that the snowpacks had similar depths, between 10 and 15 cm,
147 on the three sampling dates with modestly deeper snow in the field than the forest. The deepest snow was on 4
148 February 2021 with 15 cm in the field and 9.3 cm in the forest. Between the 18 December and the 4 February
149 sampling campaigns, there was a melt event in which the entire 10 cm snowpack on 18 December ablated. The
150 next significant snowfall event (15 cm) occurred on 1 February 2021. The snowpack experienced little additional
151 accumulation or ablation between 4 February and 24 February. The 4 February (0.15 g/cm³) and 24 February

Deleted: The

153 (0.20– 0.24 g/cm³) snowpack density values were higher than those in December (~ 0.10 g/cm³). There were
154 shallow soil frost depths (< 4 cm) during the early winter 18 December campaign in the forest and the field.
155 Deeper soil frost depths of 15.1 cm in the field and 5.9 cm in the forest occurred on 4 February 2021, with similar
156 soil frost conditions on 24 February 2021.

157 3.1 Magnaprobe vs. Snow Tube

158 The full experiment yielded individual 936 pairs of snow depth measurements from the snow tube and the
159 magnaprobe (**Fig. 2a**). For the comparison between measurement techniques, the orthogonal Deming regression
160 method was applied to consider measurement errors in both variables. Overall, there was moderate agreement (R
161 = 0.74) between the two datasets for all three sampling campaigns (**Table S1**). The snow depths measured by the
162 magnaprobe (14.9 cm average snow depth) were deeper than the snow tube (13.2 cm average snow depth) with
163 an overall bias of 1.7 cm. The magnaprobe snow depth was at least 0.5 cm deeper than the snow tube in 74% of
164 the 936 measurement pairs. Only 6.3% of the pairs had snow tube snow depths exceeding magnaprobe snow
165 depths by 0.5 cm or more. Conversely, 7.4% of the pairs' magnaprobe snow depths were over 5.0 cm deeper than
166 the snow tube. In eight pairs of measurements, when the magnaprobe measured snow depth greater than 15 cm,
167 the magnaprobe snow depths were more than double the snow tube snow depth.

168
169 The majority of the nine sampling locations in each grid had magnaprobe snow depth values that were deeper than
170 those measured using the snow tube. For all the grids, an average of 8.7, 7.7, and 7.0 out of the nine sampling
171 locations had deeper magnaprobe snow depths on 18 December 2020, 4 and 24 February 2021, respectively. As
172 hypothesized, **t-test results showed that** the magnaprobe snow depth values were significantly greater than those
173 measured using the snow tube for 39 and 31 of the 39 sampling locations on 18 December 2020 and 4 February
174 2021, respectively, but only 11 out of the 26 sampling locations on 24 February 2021. The mean differences were
175 2.3, 1.4, and 1.6 cm, with root mean square difference (RMSD) values of 3.0, 2.3, and 3.3 cm, on 18 December
176 2020, 4 and 24 February 2021, respectively, which is on the order of 15 to 25% of the overall depth observed
177 during these campaigns. Despite the biases, the average within cell snow depth variability was nearly identical for
178 the magnaprobe and the snow tube in the field (1.3 cm standard deviation for the magnaprobe). In the forest, the
179 Magaprobe's 2.0 cm within-cell standard deviation modestly exceeded the snow tube's 1.5 cm standard deviation.
180 A slightly reduced agreement was found on 24 February when there was a 1 to 4 cm thick ice layer at the bottom
181 of the snowpack in local depressions.

182
183 The overall agreement between the snow tube and magnaprobe was better when the nine measurements within a
184 single 1x1 m grid cell were averaged at each of the sampling locations (**Fig. 2b** and **Table S1**). There is a notable
185 improvement in grid cell statistics, and the correlation is stronger (overall $R = 0.87$), with slopes closer to one,
186 intercepts closer to zero, and the RMSD values reduced to 2.5 cm or less. Although averaging has no impact on
187 the overall bias, the range of differences among pairs narrowed. The difference between the magnaprobe and the
188 snow tube is typically constrained to less than 3 cm with a limited number of outliers. The magnaprobe snow
189 depth was at least 0.5 cm deeper than the snow tube in almost all grid cells (86.7%), but only three grid cells had
190 differences greater than 5 cm. Among the grid averaged magnaprobe snow depths, there were no instances in
191 which there was a doubling of snow depth when compared to the snow tube measurements.

192 **3.2 Magnaprobe vs. Snow Tube by Land type**

193 The magnaprobe and snow tube snow depths differ by land type, with the field having deeper snow and more
194 spatial variability than the forest land types (Fig. 3). Among the three forest types, the deepest snow was in the
195 deciduous-dominated forest, with mixed and coniferous forest having similar snow depths. The mean difference
196 between the magnaprobe and snow tube snow depths is a modest 1.3 cm in the field and 1.9 cm in the forest, with
197 differences of 1.9, 2.0, and 1.9 cm in the deciduous, mixed, and coniferous land types, respectively. However, the
198 differences between the magnaprobe and snow tube snow depths in the forest were higher on 18 December (2.5
199 cm), than on 4 February, and 24 February, 1.7, and 1.4 cm, respectively. Based on t-test results, the magnaprobe
200 measured significantly deeper snow depth compared to the snow tube in both the field and the forest regardless
201 of whether individual locations (p-value < 0.001) or grid cell average snow depths (p-value = 0.02) were used.
202 Based on Welch's adjusted ANOVA test, there are no significant differences in over-probing among forest land
203 types (p-value = 0.24). The RMSD values between the magnaprobe and snow tube snow depths are 3.0 cm (2.3
204 cm) and 2.5 cm (2.0 cm) for the forest and field sampling sites (grid average values), respectively. Thus, the
205 sampling method has a different impact in the field than the forest and the RMSD and bias values provide an
206 indicator of the different errors associated with in-situ measurements based on land type when used for model or
207 remote sensing validation. While these differences are significant, the average litter depths exceeded the
208 differences between the magnaprobe and snow tube snow depths in the forest, which were 2.5, 1.7, and 1.4 cm on
209 18 December, 4 February, and 24 February, respectively.

210 **4 Discussion and Suggestions**

211 This study quantifies the differences between snow depth measurements made with a magnaprobe and with a
212 snow tube. The differences seem to be primarily associated with greater over-probing by the magnaprobe into
213 vegetation/organic layers and thawed soils. Though in some cases the large differences could instead be due to
214 the larger diameter snow tube hitting a branch from a down tree or debris that the magnaprobe bypassed. The
215 result was that the magnaprobe snow depth measurements were higher than snow tube measurements, with a
216 greater difference in the forest than in the field. An average of 5 cm bias occurred in the tundra mat during the
217 Cold Land Processes Experiment (CLPX) Alaska campaign (Sturm and Holmgren 2018). Also in the open tundra
218 environment found a 7.6 cm average over-probe penetration for approximately 40 cm deep snow (Canada 2018).
219 Berezovskaya and Kane (2007) also noted over-probing of 5 to 9 cm with a magnaprobe as compared to a snow
220 tube found a bias in northern Alaska for snow depths between 29 and 48 cm. In this study, the over-probing, 1.3
221 cm in the field and a 1.9 cm in the forest, was less than previous studies probably due to the lower range of snow
222 depth and different surface conditions as compared to previous studies.

223
224 We also agree with López-Moreno et al. (2020) ~~that it is important to understand the snowpack and land conditions~~
225 for which an individual sampler was designed to select the most appropriate sampler. Understanding leaf litter or
226 vegetation depths and underlying soils may potentially reduce and help to account for the over-probing errors of
227 magnaprobe snow depth measurements. Sturm and Holmgren (2018) suggested that operators need to learn to
228 push a magnaprobe through snow, yet not penetrate it too deeply into underlying vegetation/organic layers by
229 developing a sense for the base of the snowpack. This recommendation may be difficult to implement (e.g., over

Deleted: Lopez-Moreno et al.'s (2020)

Deleted: finding

232 soft vegetation) where the probe easily penetrates the vegetation and problematic if multiple operators apply a
233 different amounts of force (Berezovskaya and Kane 2007). If operators over-probe into the base of the (frozen)
234 soils, one option is to consistently measure the depths in the same way (which would be snow depth *plus*
235 vegetation) and then subtract typical vegetation depths in the study area from the depths. When leaf litter is
236 evident, penetration into the organic layer should be ~~quantified by using independent leaf litter measurements,
237 preferably using the snow depth sampling instrument, and use to bias correct snow depths.~~

Deleted: also

Deleted: considered

Deleted: . In this study, we found that the 2.0 cm snow depth differences were approximately half of the end of winter forest leaf litter depth (3.9 cm).

238
239 As observed in this study, leaf litter and soil frost may differentially impact in-situ snow depth sampling methods.
240 The earliest sampling campaign had limited soil frost and likely reduced litter compaction. Distinct contributions
241 of forest leaf litter depth to magnaprobe and snow tube snow depths may occur because the narrow magnaprobe
242 fully penetrates the leaf litter and the larger diameter snow tube only partially penetrates the litter, or the
243 magnaprobe may only partially penetrate the leaf litter but the snow tube does not break through the leaf litter.
244 Partial penetration of the magnaprobe into the leaf litter layer (i.e., over-probing) may vary by the freeze-thaw
245 state of the duff layer and/or mineral soil layers beneath the leaf litter layer. The horizontally aligned, matted leaf
246 litter could also limit snow tube penetration. High spatial variability of leaf litter depth could also be a factor,
247 though this was not quantified here. Thus, the increased differences among in-situ methods in forested areas
248 observed in this study point to the particular importance of in-situ validation in forested areas and, more generally,
249 sampling with multiple methods in an area with a nonuniform underlying substrate.

250
251 In summary, there are three major suggestions from this work below.

- 252 1) With an ephemeral snowpack in forested environment, snow depth measurements using a Federal
253 snow tube likely avoid over-probing that can frequently occur when a magnaprobe is used.
- 254 2) The use of the average of multiple point samples within a grid is recommended instead of single
255 measurements, because the average of multiple point samples can reduce the point-to-point variability
256 and spatial representativeness errors.
- 257 3) Measurements of vegetation, leaf litter, and soil frost can help to account for the errors of in-situ snow
258 depth observations, particularly when using a magnaprobe.

259 5 Conclusion

260
261 Manual in-situ sampling snow depth measurements can be made quickly and easily, but making consistent,
262 representative, and unbiased measurements can be challenging when the surface is irregular, vegetation/organic
263 layers and unfrozen soils result in over-probing, and the leaf litter compacts during the winter. This study
264 quantified the differences between snow depth measurements made with a magnaprobe and a Federal snow tube
265 in a mixed-use temperate forest landscape with ephemeral snowpack. For all sampling campaigns and land cover
266 types, the magnaprobe snow depth measurements (mean 14.9 cm) were usually, but not always deeper than the
267 snow tube measurements (13.2 cm) and had a 1.7 cm ~~average difference. For these shallow snowpacks, this
268 amounts to a 12% difference, but in a deeper snow pack the relative impact of this difference would be much
269 smaller.~~ Biases were significantly higher in the forest (1.9 cm) than the field (1.3 cm). The difference between the
270 two instruments was 50% higher in early winter campaign than the later campaigns. The differences among
271 measurement techniques in this present study reflect the current study area, surface conditions for a single season,

Deleted: or 12% average difference

278 and the operation of the instruments by this project team. Further studies to understand the errors from in-situ
279 sampling using snow probe are warranted in various snow environments with different vegetation and soil
280 conditions to provide guidance on best practices for using in-situ snow probe datasets under conditions when over-
281 probing is likely.

282 **Acknowledgements**

283 This material is based upon work supported by the Broad Agency Announcement Program and the Cold Regions
284 Research and Engineering Laboratory (ERDC-CRREL) under Contract No. W913E518C0005 and
285 W913E521C0006. The authors are grateful to Lee Friess for providing a technical review of the draft manuscript,
286 Mahsa Moradi Khaneghahi for supporting manuscript preparation, and Brigid Ferris for training the team on litter
287 depth sampling. Christina Herrick for post-processing GPS data.

288 **Data Availability**

289 The in-situ snow observations are available in supporting information.

290 **Author Contributions**

291 HP, JJ, EB, AH, FS, MP, and EC designed the research. HP, CW, JJ, AH, FS, MP, EB, and EC conducted field
292 work to obtain lidar and/or in-situ snow observations. HP, CW, JJ, EB, AH, and MP performed the analysis.
293 HP, EC, and AH produced the figures. HP, JJ, EB, and EC wrote the initial draft. All authors contributed to
294 manuscript review and editing.

295 **Competing Interests**

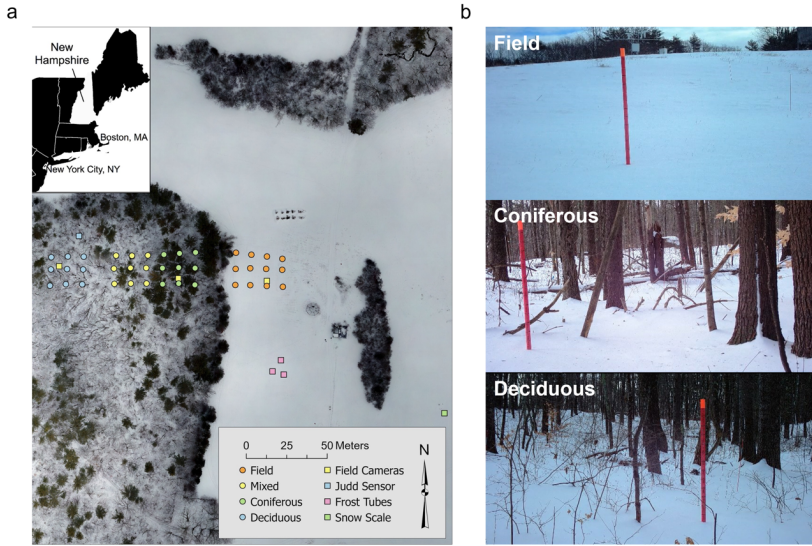
296 The authors declare that they have no conflict of interest.

297 **References**

- 298 Berezovskaya, S., and D. L. Kane, 2007: Measuring snow water equivalent for hydrological applications: part 1,
299 accuracy of observations. *Proceedings of the 16 th International Northern Research Basins Symposium and*
300 *Workshop, Petrozavodsk, Russia*, 29-37.
- 301 Bongio, M., A. N. Arslan, C. M. Tanis, and C. De Michele, 2021: Snow depth time series retrieval by time-lapse
302 photography: Finnish and Italian case studies. *The Cryosphere*, **15**, 369-387.
- 303 Canada, E. a. C. C., 2018: Snow depth data from trail valley creek .
- 304 Clyde, G. D., 1932: Utah snow sampler and scales for measuring water content of snow.
- 305 Derry, J., D. Kane, M. Lilly, and H. Toniolo, 2009: Snow-course measurement methods, North Slope, Alaska.
306 *University of Alaska Fairbanks, Water and Environmental Research Center, Report INE/WERC*, **15**.
- 307 Elder, K., W. Rosenthal, and R. E. Davis, 1998: Estimating the spatial distribution of snow water equivalence in
308 a montane watershed. *Hydrological Processes*, **12**, 1793-1808.
- 309 Farnes, P. E., B. E. Goodison, N. R. Peterson, and R. P. Richards, 1983: Metrication of manual snow sampling
310 equipment. *Final report Western Snow Conference*, 19-21.
- 311 Gandahl, R., 1957: Determination of the depth of soil freezing with a new frost meter. *Text in Swedish) Rapport*,
312 **20**, 3-15.
- 313 Kaspari, M., and S. P. Yanoviak, 2008: Biogeography of litter depth in tropical forests: evaluating the phosphorus
314 growth rate hypothesis. *Functional Ecology*, **22**, 919-923.

315 Kinar, N., and J. Pomeroy, 2015: Measurement of the physical properties of the snowpack. *Rev. Geophys.*, **53**,
316 481-544.
317 Kopp, M., Y. Tuo, and M. Disse, 2019: Fully automated snow depth measurements from time-lapse images
318 applying a convolutional neural network. *Science of The Total Environment*, **697**, 134213.
319 Leppänen, L., A. Kontu, H.-R. Hannula, H. Sjöblom, and J. Pulliainen, 2016: Sodankylä manual snow survey
320 program. *Geoscientific Instrumentation, Methods and Data Systems*, **5**, 163-179.
321 López-Moreno, J. I., and Coauthors, 2013: Small scale spatial variability of snow density and depth over complex
322 alpine terrain: Implications for estimating snow water equivalent. *Advances in water resources*, **55**, 40-52.
323 López-Moreno, J. I., and Coauthors, 2020: Intercomparison of measurements of bulk snow density and water
324 equivalent of snow cover with snow core samplers: Instrumental bias and variability induced by observers.
325 *Hydrological Processes*, **34**, 3120-3133.
326 Perron, C. J., K. Bennett, and T. D. Lee, 2004: Forest stewardship plan: Thompson farm. NH:
327 University of New Hampshire. Ossipee Mountain Land Company, West Ossipee.
328 <https://colsa.unh.edu/sites/default/files/thompson-farm-plan.pdf>.
329 Pirazzini, R., and Coauthors, 2018: European in-situ snow measurements: Practices and purposes. *Sensors*, **18**,
330 2016.
331 Raleigh, M. S., and E. E. Small, 2017: Snowpack density modeling is the primary source of uncertainty when
332 mapping basin-wide SWE with lidar. *Geophysical Research Letters*, **44**, 3700-3709.
333 Sturm, M., and J. Holmgren, 2018: An automatic snow depth probe for field validation campaigns. *Water*
334 *Resources Research*, **54**, 9695-9701.
335 Walker, B., E. J. Wilcox, and P. Marsh, 2020: Accuracy assessment of late winter snow depth mapping for tundra
336 environments using Structure-from-Motion photogrammetry. *Arctic Science*, 1-17.
337

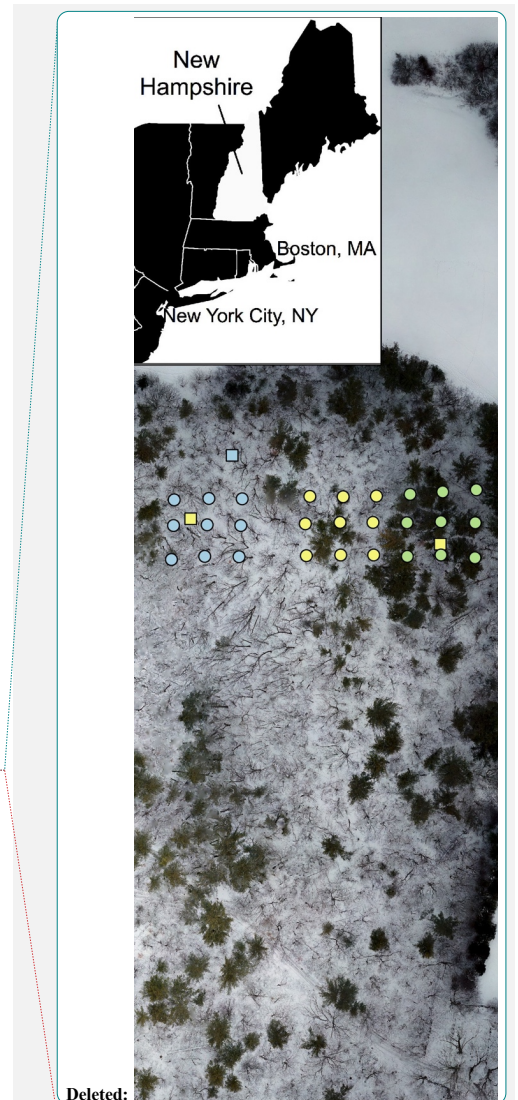
338



339

340

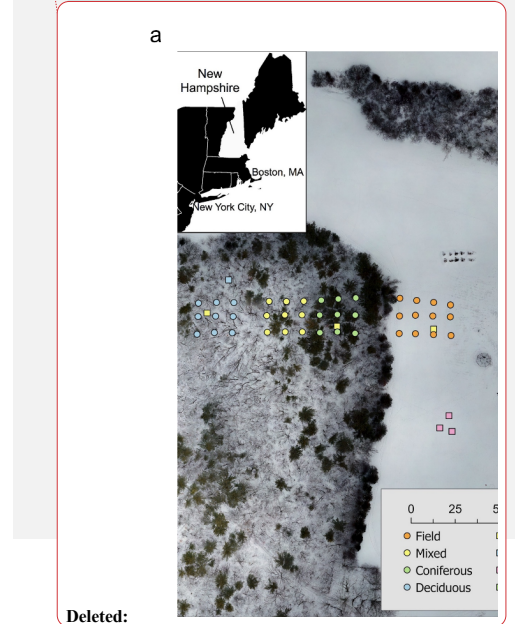
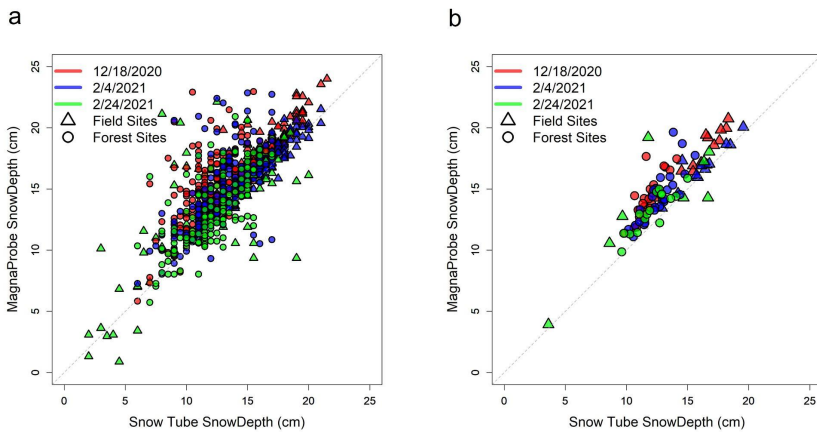
341 **Figure 1:** (a) The 4 February 2021 aerial optical image of Thompson Farm, Durham NH, USA showing both
 342 forest and field regions with snow sampling sites in the field, coniferous, mixed, and deciduous forested areas as
 343 well as the locations of the CRREL soil frost tubes; and field cameras. (b) Field camera images in the field,
 344 coniferous, and deciduous forested areas taken on 4 February 2021 by the field cameras.



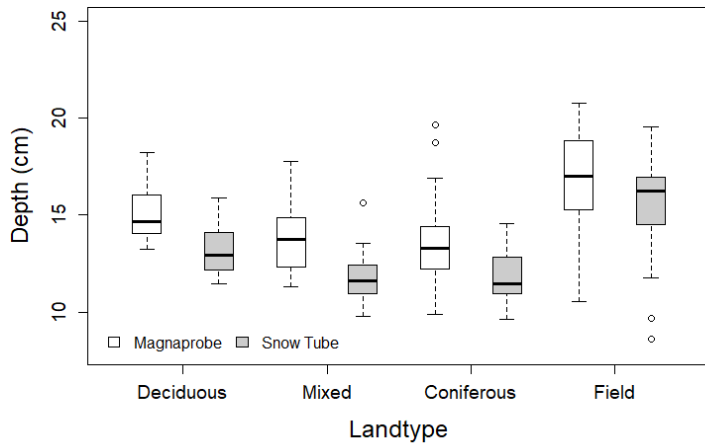
Deleted:

345

346 **Figure 2:** Comparison of snow depths measured by magnaprobe and snow tube for the three sampling campaigns
 347 using (a) the sampling individual points (n = 936) and (b) using grid cell average values (n=104).



Deleted:



352

353 **Figure 3:** Boxplots of snow depths by land type measured by the magnaprobe and the snow tube for the three
 354 sampling campaigns using the grid cell average values.

355