Microstructure-based modelling of snow mechanics: experimental evaluation on the cone penetration test

3 Clémence Herny^{1,2}, Pascal Hagenmuller¹, Guillaume Chambon², Isabel Peinke¹, Jacques Roulle¹

⁴ ¹Univ. Grenoble Alpes, Université de Toulouse, Météo-France, CNRS, CNRM, Centre d'Etude de la Neige, Grenoble, France

⁵ ²Univ. Grenoble Alpes, CNRS, INRAE, IRD, Grenoble INP, IGE, Grenoble, France

6 *Correspondence to*: Clémence Herny (clemence.herny@gmail.com)

7 Abstract. Snow is a complex porous material presenting a variety of microstructural patterns. This microstructure largely 8 controls the mechanical properties of snow, although the relation between the micro and macro properties remains to be better 9 understood. Recent developments based on the discrete element method (DEM) and three-dimensional microtomographic data 10 make it possible to reproduce numerically the brittle mechanical behaviour of snow. However, these developments lack 11 experimental evaluation so far. In this study, we evaluate a DEM numerical model by reproducing cone penetration tests on 12 centimetric snow samples. The microstructures of different natural snow samples were captured with X-ray microstructures of different natural snow samples. 13 before and after the cone penetration test, from which the grain displacements induced by the cone could be inferred. The tests 14 were conducted with a modified Snow MicroPenetrometer (5 mm cone diameter), which recorded the force profile at a high 15 resolution. In the numerical model, an elastic brittle cohesive contact law between snow grains was used to represent the 16 cohesive bonds. The initial positions of the grains and their contacts were directly derived from the tomographic images. The numerical model was evaluated by comparing the measured force profiles and the grain displacement fields. Overall, the model 17 18 satisfactorily reproduced the force profiles in terms of mean macroscopic force (mean relative error of about 20%) and the 19 amplitude of force fluctuations (mean relative error of about 55%), while the correlation length of force fluctuations was more 20 difficult to reproduce (mean relative error of about 40% for two samples out of four and by a factor ≥ 8 for the other two). 21 These characteristics were, as expected, highly dependent on the tested sample microstructure, but they were also sensitive to 22 the choice of the micro-mechanical parameters describing the contact law. A scaling law was proposed between the mechanical 23 parameters, the initial microstructure characteristics and the mean macroscopic force obtained with the DEM numerical model. 24 The model could also reproduce the measured deformation around the cone tip (mean grain displacement relative error of 57% 25 along the horizontal axis), with a smaller sensitivity to the contact law parametrisation in this case. These detailed comparisons 26 between numerical and experimental results give confidence in the reliability of the numerical modelling strategy and opens 27 promising prospects to improve the understanding of snow mechanical behaviour.

28 1 Introduction

29 Snow is a brittle and porous material existing on Earth close to its melting point. The thermodynamical conditions in the clouds 30 govern the snowflake morphology and, once deposited on the ground, snow continues to evolve via metamorphism. The snow 31 material is thus characterised by a large variety of microstructural patterns (grain size, grain shape, density) classified into 32 different snow types (Fierz et al., 2009). It has been established that the snow microstructure controls the properties of snow 33 (Shapiro et al., 1997; Johnson and Schneebeli, 1999; Schneebeli, 2004). For instance, weak layers involved in avalanche 34 triggering (Schweizer et al., 2003) are usually constituted of specific snow types (depth hoar, surface hoar, precipitation 35 particle, faceted crystals) characterised by low cohesion and low strength (Jamieson and Johnston, 1992). The link between 36 the snow microstructure and its properties, especially its mechanical properties, is still not well understood, even if it is crucial 37 for many applications, such as avalanche forecasting (Schweizer et al., 2003, Jamieson and Johnston, 1992), snowpack 38 modelling (Calonne et al. 2014), ice core interpretation (Montagnat et al. 2020) or geotechnics (Shapiro et al., 1997). In 39 particular, the brittle failure occurring at high shear rates (> 10^{-4} s⁻¹) during the release of an avalanche remains represented by 40 very coarse empirical laws (Brun et al., 1992; Bartelt, et al. 2002; Vionnet et al. 2012). In this elastic-brittle regime (rapid and 41 large deformations), the mechanical behaviour of snow is thought to be mainly controlled by bond failures and grain 42 rearrangements (Narita, 1983).

43 The snow microstructure and its evolution can be captured at high resolution (typically 10-50 µm) with X-ray micro 44 tomography imaging (µCT) (Coléou et al., 2001; Freitag et al., 2004; Schneebeli, 2004; Heggli et al., 2011). This non-45 destructive method preserves the snow microstructure and resolves the shape of snow grains, grain bonds and porosity which 46 is of primary importance for mechanical studies. In particular structural properties of snow, such as density, specific surface 47 area (SSA), correlation length, bond characteristics, can be evaluated from tomographic data (e.g. Schneebeli, 2004; 48 Schneebeli et al., 2004; Hagenmuller et al., 2014a; Calonne et al., 2014; Proksch et al., 2015). The tomographic data are also 49 used as a basis for numerical modelling (Schneebeli, 2004; Schneebeli et al., 2004; Hagenmuller et al., 2015) or 50 calibration/validation data of statistical empirical models retrieving grain-scale physical and mechanical properties from other 51 measurements (e.g. Proksch et al., 2015; Reuter et al., 2019). However, tomographic imaging is time-expensive and not adapted 52 to routine measurements in the field.

The mechanical properties of snow are commonly derived from Cone Penetration Test (CPT) measurements, which is an objective and relatively easy-to-set-up method (Schneebeli and Johnson, 1998). This method has been widely used to characterise soil stratigraphy (Lunne et al., 1997) and adapted to snowpack stratigraphy (Gubler, 1975; Schaap and Fohn, 1987; Dowd and Brown, 1986; Schneebeli and Johnson, 1998; Mackenzie and Payten, 2002; McCallum, 2014). The CPT provides a force profile by measuring the resisting force exerted on a conic tip penetrating, at a constant rate, into a material. The development of high-resolution digital penetrometers dedicated to snow studies (Schneebeli and Johnson, 1998; Mackenzie and Payten, 2002; McCallum, 2014) has provided the possibility to resolve the force profile at a microscopic scale

- and capture the high-frequency fluctuations of the force signal up to a metre depth. Such force penetration profiles contain
- 61 valuable information on the snow structural parameters at macro- and micro-scale (Löwe and van Herwijnen, 2012).

Interpretation of the CPT requires a good understanding of the interactions between the cone tip and the snow grains. Several studies aimed to investigate the grain displacement field around the tip. Particle Image Velocimetry (PIV) imaging was performed to quantify the 2D displacement field of snow grains while the tip penetrates into the material (Floyer and Jamieson, 2010; Herwijnen, 2013; LeBaron et al., 2014). Peinke et al. (2020) developed a grain tracking algorithm to reconstruct from µCT the 3D displacement field of snow grains induced by a CPT. All these studies revealed the development of a compaction zone (CZ) in front of the tip.

68 Various mechanical or statistical models have been developed to interpret the CPT penetration signal in terms of mechanical 69 properties. The cavity expansion model (CEM) (Bishop et al., 1945; Yu and Carter, 2002) has been applied to snow by Ruiz 70 et al. (2016) and Peinke et al. (2020). This model considers snow as a continuum and describes the elastic-plastic deformation 71 of the material around the tip in order to retrieve macroscopic material properties (cohesion, friction, etc.). The continuum 72 assumption becomes invalid for a ratio between cone diameter and mean grain diameter lower than 20 typically (Bolton et al. 73 1993), leading to potentially erroneous interpretations of the CPT results. Alternatively, the shot noise model interprets the 74 force signal and its fluctuations as a superposition of independent elastic-brittle ruptures occurring next to the tip (Schneebeli 75 and Johnson, 1999; Marshall and Johnson, 2009; Löwe and van Herwijnen, 2012) and retrieves microstructural properties 76 (bond rupture force, etc.) The penetration process is generally modelled as a Homogeneous Poisson Process (HPP) with a 77 constant intensity (Löwe and van Herwijnen, 2012). Peinke et al. (2019) have generalised the HPP method to account for the 78 transient phase of the penetration process, attributed to the development of the CZ (Peinke et al., 2019). These authors used a 79 Non-Homogeneous Poisson Process (NHPP) considering a depth dependency of the intensity (number of bond failures per 80 penetration increment). Yet, the assumption of independent elastic-brittle rupture events essentially neglects the development 81 of a CZ (Johnson and Schneebeli, 1999; Schneebeli, 2001; Herwijnen, 2013; LeBaron et al., 2014; Ruiz et al. 2017). Therefore, 82 none of these two models appear to fully account for the specificity of snow deformation induced by CPT. Additional 83 investigations are required to better understand the tip interaction with snow and better interpret the force measurements.

84 Recently, numerical approaches have been developed to study the mechanical response of snow by explicitly accounting for 85 the microstructure (Johnson and Hopkins, 2005; Gaume et al., 2015, 2017; Hagenmuller et al., 2015; Wautier et al., 2015; 86 Mede et al. 2018b, 2020; Bobillier et al., 2020, 2021). Snow is described as a granular material and modelled by the discrete 87 element method (DEM) in a high shear rate regime. The complexity of the snow microstructure can be taken into account by 88 feeding the DEM simulations with high-resolution 3D reconstructions obtained with µCT. These simulations have provided 89 new insights into the snow mechanical behaviour, such as the dependence of snow strength to microstructure properties 90 (Hagenmuller et al., 2015) or the identification of different failure modes in shear loading (Mede et al., 2018b, 2020). The 91 downside of this method is that it is time-consuming, and simulations can only be performed on small samples (up to a few 92 centimetres). Furthermore, these numerical models still lack direct experimental evaluation.

93 In this context, the aim of this study was to evaluate a microstructure-based DEM model using recent CPT experimental data 94 performed in a controlled environment (Peinke et al., 2020). The dataset includes uCT images of the samples acquired before and after the tests. The deformation induced by the CPT (strain rate of about 10² s⁻¹, Reuter et al., 2019) belongs to the elastic-95 96 brittle regime (Narita, 1983; Floyer and Jamieson, 2010) and is therefore suitable for DEM simulation. The results of the 97 numerical model are directly compared to experimental data in terms of (1) macroscopic force profile and associated statistical 98 indicators and (2) grain displacements induced by the cone penetration. A systematic sensitivity analysis to DEM mechanical 99 parameters, including Young's modulus, cohesion and friction coefficient, was performed to find the combinations of 100 parameters that best reproduce experimental results. Finally, the role of the microstructure was also investigated by performing 101 DEM simulations for different snow types. The evaluation of the numerical model provides the opportunity to better understand 102 the mechanisms at play during snow deformation in an elastic-brittle regime and better interpret CPT profiles.

We first present the experimental dataset and the numerical methods. The data processing used to compare experimental and numerical results is also explained. The results of the DEM, the sensitivity analysis to mechanical parameters and the comparison to experimental results are then presented. The relevance of the DEM model and the limits of our approach are eventually discussed before concluding.

107 **2 Methods**

108 2.1 Experimental measurements

109 The experimental dataset used in this study has been acquired by Peinke et al. (2020) and is only briefly presented in this paper.

110 The methodology comprises collection and preparation of snow samples, acquisition of high-resolution micro-tomographic

111 images and cone penetration tests (CPT).

112 **2.1.1 Snow sample preparation**

Blocks of natural snow were sampled in the French Alps near Grenoble and stored at -20°C in a cold room. The materials collected were representative of the variety of seasonal snow types (Table 1), namely rounded grains (RG), large rounded grains (RGlr), depth hoar (DH) and precipitation particles (PP), with distinct bulk densities and specific surface areas (SSA).

The samples were then prepared in a cold room at -10° C by sieving the different snow types into aluminium cylinders of 20

117 mm height and 20 mm diameter. All samples were prepared at least 24 hours before the measurements in order for the bonds

118 between grains to rebuild after sieving.

119 2.1.2 Micro-Tomography (μCT)

120 Tomographic scans of each sample were acquired before and after performing the CPT to capture the initial and final 121 microstructure of the snow, respectively. An X-ray tomograph (DeskTom130, RX Solutions) operating at a pixel size of 15 122 μ m pix⁻¹, a voltage of 80 kV and a current of 100 μ A was used. During tomographic scanning, the samples were maintained 123 at a constant and uniform temperature of -10°C in a cryogenic cell (CellDyM, Calonne et al. (2015)). Each scan, consisting of

124 1440 2D radiographs, was reconstructed to obtain 3D grayscale images representing the attenuation coefficients of the different

125 materials composing the samples. The grayscale images were then transformed into binary (ice matrix – pore space) segmented

126 images using an energy-based segmentation algorithm (Hagenmuller et al., 2013).

127 **2.1.3 Cone Penetration Test (CPT)**

128 After the initial micro-tomography scan, a CPT was performed on the snow samples using a modified SnowMicroPenetrometer 129 (SMP version 4, Schneebeli and Johnson, 1998). The specific rod used by Peinke et al. (2020) displays a conic tip with an 130 apex angle a of 60° and a maximum cone radius equal to the rod radius R of 2.5 mm. The rod was inserted vertically into the 131 snow sample at a constant penetration speed v of 20 mm s⁻¹. The resisting force applied on the penetrometer (cone and rod) 132 was recorded at every 4 µm of penetration increment (i.e., 5 kHz frequency). The SMP sensor (Kistler sensor type 9207) can 133 measure forces up to 40 N with a resolution of 0.01 N. The tip was stopped at depths between 7 and 15 mm, i.e., 5 to 13 mm 134 above the sample bottom, to avoid boundary effects (Peinke et al., 2020). The experimental force profiles are presented in 135 Figure S26.

136 2.2 Numerical modelling

- Snow is here considered as a granular cohesive material. The high strain rate $(> 10^{-4} \text{ s}^{-1})$ induced by the tip penetration in the snow sample is considered to lead to brittle deformations, with inter-granular damage and grain rearrangements (Narita, 1983; Johnson and Hopkins 2005; Hagenmuller et al., 2015). We adopted an approach based on DEM to simulate the cone penetration tests in the measured snow samples. The mechanical model, based on YADE software (Šmilauer et al., 2015), is adapted from the work of Hagenmuller et al. (2015) and Mede et al. (2018a, b and 2019).
- The setting-up of the simulations involves different steps, namely the generation of the initial conditions based on measured snow microstructures, the definition of the contact laws between the snow grains, and the setting of the boundary conditions to reproduce the CPT configuration.

145 **2.2.1 Grain segmentation and grain shape representation**

- The DEM model was fed by the 3D ice-air images derived from μ CT. The continuous ice matrix was first segmented into individual grains based on geometrical criteria, as described by Hagenmuller et al. (2013). The main idea of the approach is to detect potential mechanical weakness zones (i.e., the bonds) based on the principal minimal curvature κ_T and a contiguity parameter c_T . The threshold on curvature κ_T was set to 1.0 for RG, RGIr and DH samples and to 0.7 for PP sample; the contiguity parameter was set to 0.1 for all the samples (see Hagenmuller et al., 2013 for details).
- To construct the DEM sample, the irregular shape of the grains was approximated by filling the grain volume with a population of overlapping spheres (Fig. 1). The position of these spheres was derived from the medial axis of the structure (Coeurjolly et
- al., 2007; Mede et al., 2018a) and redundant spheres were discarded based on a power diagram filter (Coeurjolly et al., 2007).

154 This grain shape representation by a multitude of spheres preserves the capability of YADE to handle sphere-sphere contact 155 detection. However, a high number of spheres slows down the simulations. We thus further decimated the number of spheres 156 by approximating the grain shape. We only selected the spheres with a radius larger than a threshold L (voxel) and with a 157 relative coverage larger than S (i.e., the ice volume associated with the sphere according to the power diagram should be larger 158 than S times the sphere volume (Coeurjolly et al., 2007). A trade-off must be found between this grain shape approximation, 159 influencing the simulation accuracy, and the number of spheres influencing the numerical cost. Eventually, the spheres 160 belonging to the same grain were clumped together in rigid aggregates constituting single discrete elements (DE). A detailed 161 sensitivity analysis was conducted (see supplementary material, Table S1 and Fig. S1) to determine the optimal values of L162 and S parameters. Note that this grain shape approximation might also lead to delete the smallest grains in the numerical 163 samples, as they cannot be covered with the chosen parameters L and S. The grain number difference and shape approximation 164 of the numerical sample compared to initial the segmented µCT image can be quantified by computing the volumetric error 165 E_V . The final chosen L and S values for each snow type, with the associated volumetric E_V and mechanical E_M errors (defined 166 in Sect. S1.1), can be found in Table 1.

167

Sample name	Snow type	Sieve size (mm)	Bulk density (kg m ⁻³)	SSA (m² kg ⁻¹)	L (vx)	S	Number of spheres	Number of grains	Number of initial cohesive interactions between grains	Initial contact density V	Ev (%)	Ем (%)
RG	Rounded Grains	1.6	289	23.0	5	0.3	514917	27560	47736	0.55	42.3	5.3
RGlr	Large Rounded Grains	1	530	10.1	5	0.3	270143	8488	24005	1.63	14.6	4.2
DH	Depth Hoar	1.6	364	15.9	5	0.2	743546	11211	24258	0.86	24.7	14.3
PP	Precipitat ion Particle	1.6	91.3	53.5	2	0.5	1797567	95022	125805	0.13	32.2	10.3

168Table 1: Overview of the snow samples analysed in this study and parameters of DEM grain shape representation. Sample names169were given according to the snow type classification (Fierz et al., 2009). The sample density and specific surface area (SSA) were170derived from the micro-tomographic images (Peinke et al., 2020). The initial contact density was computed according to Eq. 10. The171minimum radius of the sphere L and the minimum sphere coverage S were determined through a sensitivity analysis presented in172Sect. S1.1. The resulting number of spheres, grains and cohesive grain-grain interactions are indicated, as well as the volumetric173error E_V and the mechanical error E_M associated with each grain shape representation.

174

175 **2.2.2 Interactions and contact law**

176 The contacts between adjacent grains were identified during the grain segmentation phase. In the DEM simulations, each grain

177 contact is represented by several sphere-sphere interactions. The interactions between spheres are described by an elastic brittle

178 cohesive contact law characterised by four parameters, namely the normal and the shear contact stiffness K_N and K_S , the 179 adhesion *A*, and the friction angle φ . The normal force F_N between two spheres is computed as proportional to the distance 180 between the two sphere surfaces x_N , and limited by the adhesion value in the tensile regime ($x_N > 0$):

$$181 F_N = K_N x_N \le A. (1)$$

182 The shear force F_s is proportional to the shear displacement between the spheres $x_{s,}$, with a maximal value given by the sum 183 of adhesion and friction:

$$184 F_S = K_S x_S \le A + F_N \tan(\varphi) . (2)$$

185 If the force exceeds the threshold, either in tension or in shear, the cohesive bond is broken. As long as the spheres remain in 186 contact after the bond is broken, friction remains active in shear. In the initial state, all interactions in the numerical sample 187 are considered cohesive. While the sample deforms, grain displacements lead to progressive breakage of the initial cohesive 188 interactions and the potential creation of new contacts. These new interactions are frictional only (no cohesion), meaning that 189 sintering mechanisms are not considered in this study.

The force of a given intergranular cohesive contact corresponds to the sum of all the associated sphere-sphere interactions. Based on the total contact surface between two grains (obtained from the μ CT image) and the number of associated spheresphere interactions, each sphere-sphere interaction *i* can be associated with a representative contact surface D_i . In order to recover the correct cohesion strength between two grains, the adhesion parameter *A* was defined for each sphere-sphere interaction as:

$$195 \quad A_i = D_i C, \tag{3}$$

with C (Pa) the cohesion of ice. In YADE, by default, the contact stiffnesses are computed based on the radii of the spheres in interaction and two elastic material parameters, namely the Young's modulus E and the Poisson ratio v. For our computations, to ensure that all cohesive sphere-sphere interactions between two grains break at the same separation distance, the computation of the normal stiffness was redefined as:

$$200 K_{N,i} = \frac{D_i E}{r_{mean}}, (4)$$

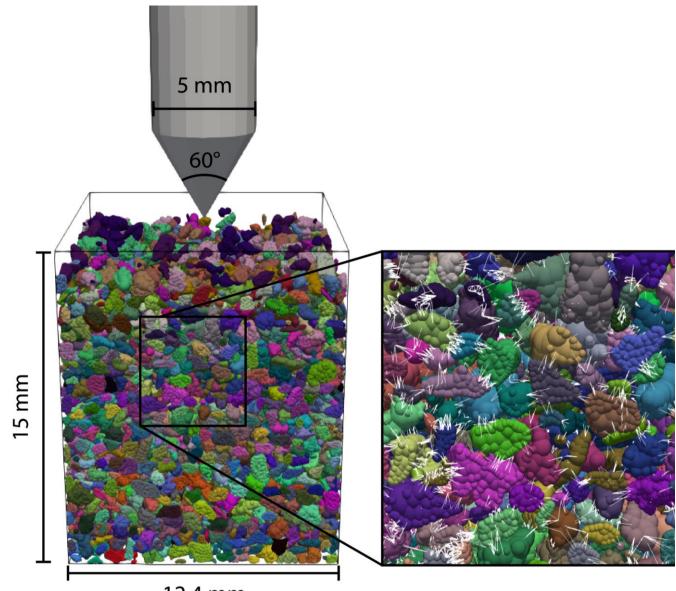
where r_{mean} (m) is a characteristic length constant for all the interactions in the numerical sample, taken as the mean sphere radius. The shear stiffness is then defined as:

$$203 K_S = \nu \times K_N \,. (5)$$

Note that due to the rather arbitrary characteristic length considered in the definition of the normal stiffness [Eq. (4)], which depends on the grain shape approximation, as well as to the simple linear relation considered for the normal force [Eq. (1)], the contact-level YADE Young's modulus *E* should not be regarded as the "true" Young's modulus of the material, but rather as a representative parameter of the elastic properties at the contacts.

209 2.2.3 Simulation setup and critical time step

210 In order to evaluate the DEM model, we have implemented a CPT configuration similar to the experimental setup used by 211 Peinke et al. (2020) (Fig. 1). The snow sample is contained in a rectangular box open at the top. The box is about 12.4 mm 212 along the x- and y-axis and about 15 mm along the z-axis. The vertical and horizontal box sizes were reduced compared to the 213 20 mm height and 20 mm diameter respectively of the sample holder used by Peinke et al. (2020). This choice has been 214 motivated by (1) simplifying the geometry with a rectangular numerical sample, (2) matching the sample height imaged with 215 μ CT and (3) reducing the computational time. A sample size sensitivity analysis has been performed to ensure that border 216 effects are not introduced by reducing the sample size (Fig. S2). The penetrometer tip displays a maximal radius R of 2.5 mm 217 and an apex angle a of 60°. Initially in a centered position at the box surface, it is displaced downwards through the sample at a constant speed of 20 mm s⁻¹. The simulation stops when the tip reaches the bottom of the box. The walls (box and tip) are 218 219 represented by facets with rigid boundary conditions. The gravity is set to 9.81 m s⁻².



12.4 mm

Figure 1: Illustration of DEM CPT modelling for the RGIr sample. The penetrometer is moving downward at a constant speed of mm s⁻¹. Snow grains (represented with different colours) are modelled by overlapping spheres clumped together. The zoomed window shows the initial cohesive interactions between the spheres of adjacent grains (white lines).

225

The stability of the explicit integration scheme is ensured by estimating the critical time step, based on the propagation speed of elastic waves in the sample (Zhao, 2017):

$$228 \qquad \Delta t_{cr} = \min\left(\frac{m_i}{K_{N,i}}\right)^{0.5} \quad , \tag{6}$$

with m_i and $K_{N,i}$ the mass and normal stiffness of the DE *i*. The mass m_i or, equivalently the material density ρ , can be artificially increased to increase the time step (Hagenmuller et al., 2015). A numerical sensitivity analysis (Fig. S3) has shown that increasing the density by a factor *f* equal to 100 does not affect the simulation results, while significantly reducing the computing time. Finally, a Cundall's non-viscous damping coefficient Λ was applied to the particle acceleration to dissipate kinetic energy and avoid numerical instabilities (Šmilauer et al. 2015). A value of 0.05 was chosen according to the results of a numerical sensitivity analysis (Fig. S4).

235 2.2.4 Input parameters

In view of the preceding paragraph, the density of the ice grains was set to $\rho = f \ge 917$ kg m⁻³. The contact law parameters were 236 237 derived from typical values measured on ice. The Poisson coefficient P was set to 0.3 (Schulson and Duval, 2009). The typical 238 Young's modulus E, the cohesion strength C and the friction coefficient $tan(\varphi)$ values for the ice are usually evaluated around 1 x 10¹⁰ Pa, 1 x 10⁶ Pa and 0.2, respectively (Gammon et al., 1983; Schulson and Duval, 2009). For this study, a sensitivity 239 240 analysis to the values of these parameters was performed to get insights into their influence and best adjust simulation results to the experimental measurements. The considered ranges were 1 x 10^{8} -1 x 10^{10} Pa for E, 5 x 10^{5} -5 x 10^{6} Pa for C and 0.2-0.5 241 242 for $tan(\varphi)$, respectively. Note that the range of the Young's modulus E ensures small grain overlaps, i.e. compliance with the 243 rigid grain assumption (Fig. S5). We must mention that, due to longer computing times, fewer parameter values could be 244 explored for large Young's modulus values. For the PP sample, no numerical simulations could be performed for a Young's modulus of 1 x 10¹⁰ Pa, as computing times were unreasonable ($E = 1 \times 10^8$ Pa, $t \sim 4$ months and $E = 1 \times 10^9$ Pa, $t \sim 10$ months 245 on a 72 cores machine with 2.6 GHz Intel Xeon processors (2.6 GHz) and 500 GB RAM. YADE scripts enable parallelisation 246 247 on up to 5 cores).

Simulation setup							
Sample width	W	13 mm					
Sample height	Н	15 mm					
Tip radius	R	2.5 mm					
Cone apex	а	60°					
Tip velocity	v	20 mm s ⁻¹					
Gravity	g	9.81 m s ⁻²					
Numerical parameters							
Time step	dt	~ 1 x 10 ⁻⁶ -1 x 10 ⁻⁸ s					
Mass factor	f	100					
Non-viscous damping coefficient	Λ	0.05					
Material properties							
Grain density	ρ	917 x 10 ² kg m ⁻³					
Poisson coefficient	Р	0.3					
Friction coefficient	$tan(\varphi)$	0.2–0.5 (default value 0.2)					
Young's modulus	Ε	1 x 10 ⁸ -1 x 10 ¹⁰ (default value 1 x 10 ⁹) Pa					
Cohesion	С	$5 \ge 10^{5} - 5 \ge 10^{6}$ (default value $2 \ge 10^{6}$) Pa					

²⁴⁹ Table 2: Input parameters used for the simulations presented in this paper.

251 **2.3 Data processing**

The outputs of the DEM simulations are the resisting force exerted by the grains on the penetrating rod and the displacement of the grains. These results can be directly compared to the experimental measurements.

254 2.3.1 Force sampling

255 The sum of the forces along the z-axis applied on all the facets constituting the penetrometer (cone and rod) is recorded at each 256 time step. The characteristics of the raw numerical force profiles depend on the numerical parameters (notably the time step), 257 and are not necessarily suited for direct comparison with experimental results. To obtain numerical profiles that can be 258 compared to their experimental counterparts, the simulated force values were averaged over windows corresponding to 259 displacement increments of 4 µm, thus matching the sampling frequency of the SMP. This averaging is also useful to smooth 260 out high-frequency fluctuations linked to the very small time steps used in DEM. Finally, numerical and experimental force 261 profiles are then re-sampled by linear interpolation over a regular grid with a step of 4 µm over the same depth. The profiles 262 span from a depth of 0 mm (initial contact between the cone and the sample surface) to the chosen maximum depth, which, in 263 our study, is set to 7 mm (i.e., 1750 points). This value corresponds to the minimum depth reached by the penetrometer during 264 the experimental CPT tests for the selected samples.

265 **2.3.2 Statistical indicators**

Quantitatively, the DEM numerical model is evaluated by comparisons with experimental force profiles in terms of three statistical indicators: the mean macroscopic force \overline{F} (N), the amplitude of force fluctuations σ (N), and the correlation length *l* (mm). The indicator σ is calculated as the variance of the detrended force profile as follows:

269
$$\sigma = \overline{\tilde{F}}^2, \qquad \tilde{F} = \frac{F - F_{sm}}{F_{sm}}$$
 (7)

with \tilde{F} ([Eq. (5)], Peinke et al. 2019), the detrended force profile, F, the force profile and F_{sm} , the averaged force profile calculated over a rolling window $\Delta z = 3$ mm. The correlation length l (mm) is also computed on the detrended force profile (Peinke et al. 2019). In our study, the snow samples exhibit a rather homogeneous structure allowing us to consider that l is constant over the depth (Peinke et al., 2019). These three statistical indicators have been chosen because they are easily quantifiable and commonly used to describe force profiles obtained by CPT in snow (Johnsson and Schneebeli, 1999; Löwe and van Herwijnen, 2012; Peinke et al. 2019). In addition, they constitute key parameters to derive additional microstructural properties based on Poisson shot noise models (Löwe and van Herwijnen, 2012; Peinke et al. 2019).

To select the set of model mechanical parameters (*E*, *C* and $tan(\varphi)$) providing the best fit to the experimental measurements, a total error RE_{tot} is computed according to:

279
$$REtot = \sqrt{2 RE_F^2 + RE_\sigma^2 + RE_l^2}$$
 (8)

280 with RE_k the logarithmic relative error calculated for the three statistical indicators, $k = (F, \sigma, l)$, as:

$$281 RE_k = \frac{\log(\text{measured value}_k) - \log(\text{computed value}_k)}{\log(\text{measured value}_k)}$$
(9)

Given the difficulties in reproducing the correlation length with the DEM model for two out of four samples and the fact that the values of the statistical indicators vary over several orders of magnitude (see Section 3.2), the logarithmic relative errors RE_k were computed with the log of the considered values. We have attributed a weight factor of 2 to the logarithmic relative error RE_F related to the mean macroscopic force, to put more emphasis on the correct reproduction of this quantity. Hence, for each snow sample, the set of mechanical parameters minimising the total error RE_{tot} was determined.

287 **2.3.3 Grain displacement analysis**

The position of all grains was recorded every ~0.4 mm of penetration in the DEM simulations. The total displacements and the trajectories can therefore be reconstructed for each grain. Due to the thermodynamically active nature of snow, interrupted experimental tests were not feasible and only the initial (before CPT) and the final states (after CPT) of the snow sample could be imaged by μ CT. Grain tracking, applied to the micro-tomographic images, has been performed by Peinke et al. (2020), providing the total displacement of the identified grains. We thus compared the total displacement between the CPT experiments and the DEM simulations at the same penetration depth, i.e., at the maximal penetration measured experimentally. Note that grain tracking could not be performed for the PP sample due to the small size of the grains.

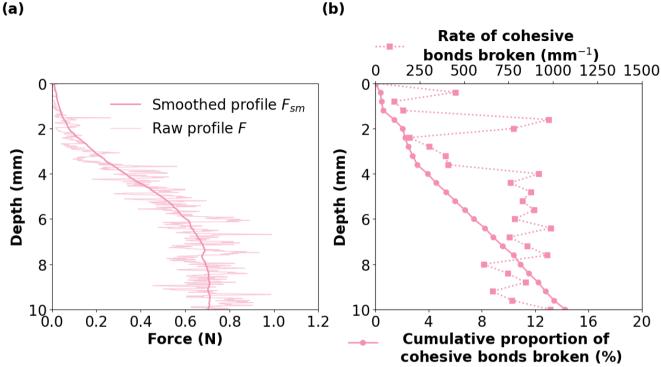
295 The profiles of vertical and radial displacements were averaged around the cone axis and over the height of an area located 296 between the top section of the cone and the sample surface. A displacement threshold of 0.03 mm was set to define the CZ 297 (Peinke et al., 2020). Only the radial profiles were compared to the experimental results, as we suspect the vertical profiles 298 derived from µCT scans might be misleading (Peinke et al. 2020). Indeed, before acquiring the post-CPT µCT scans, the tip 299 was removed from the snow. This procedure was performed about one hour after the tip penetration, to allow for bonds between 300 ice grains to re-form by sintering and limit grain displacements during tip removal. However, despite this precaution, some 301 grains in contact with the tip might have been dragged upward due to friction with the tip. Therefore, the upward component 302 of the vertical displacement might have been overestimated in the experimental results, especially for the larger grains.

303 **3 Results**

304 **3.1 Simulated Cone Penetration Tests**

This section presents an example of CPT simulation results for the case of the RG snow sample with the following mechanical parameters: $E = 1 \times 10^9$ Pa, $C = 5 \times 10^6$ Pa and $tan(\varphi) = 0.2$ (Table 3). The results for the other snow samples are shown in





309 310 Figure 2: (a) Force F as a function of penetration depth (light line) obtained for the RG sample. The superposed smoothed profile 311 (bold line) F_{sm} corresponds to the force value averaged over a rolling window of 3 mm. (b) Rate of cohesive bonds broken per unit 312 penetration depth and cumulative proportion of cohesive bonds broken (%) as a function of tip penetration depth. The initial 313 number of cohesive bonds is indicated in Table 1. The results are obtained with the mechanical parameters indicated in Table 3.

315 The simulated penetration force globally increases with depth and is characterised by high-frequency fluctuations whose 316 amplitude also tends to increase with depth (Fig. 2 (a)). The force profile displays an 'S' shape with three stages: 1) up to \sim 317 3.5 mm depth, the profile is convex, 2) between \sim 3.5 and \sim 6 mm depth, the increase of force with depth is almost linear, and 318 3) for depths larger than 6 mm, the force reaches a nearly constant value. A similar behaviour is observable for the RGIr and 319 PP samples (Fig. S6 (a) and S10 (a)), with slight variations in the transition depths between the different stages. For the DH 320 sample, the macroscopic force profile also displays stages 1 and 2, but the stabilisation at a nearly constant value is less evident 321 for the results presented in Fig. S (a). Stage 3 might be reached at greater depths for this sample. 322

- The penetration of the tip induces bond failures in the simulated samples (Fig. 2 (b)). Overall, for the RG sample, about 15%
- 323 of the cohesive interactions broke over 10 mm of penetration, corresponding to an average rate of ~650 bond failures mm⁻¹.
- 324 This average bond failure rate is variable among the samples, reaching up to 1400 bond failures mm⁻¹ for RGIr sample (Figs.
- 325 S6 (b), S8 (b), S10 (b)). In detail, for the RG sample, we notice an increase in the bond failure rate at around 3.5 mm of

penetration depth (Fig. 2 (b)), coinciding with the transition between the first and second stages observed in the force signal (Fig. 2 (a)). Bond failure intensity then remains nearly constant as the macroscopic force reaches its steady-state value. Similar characteristics are observed for the other snow types (Figs. S6, S10) except for the DH sample, for which the slope change between the first and second stages is less clear (Fig. S8 (b)).

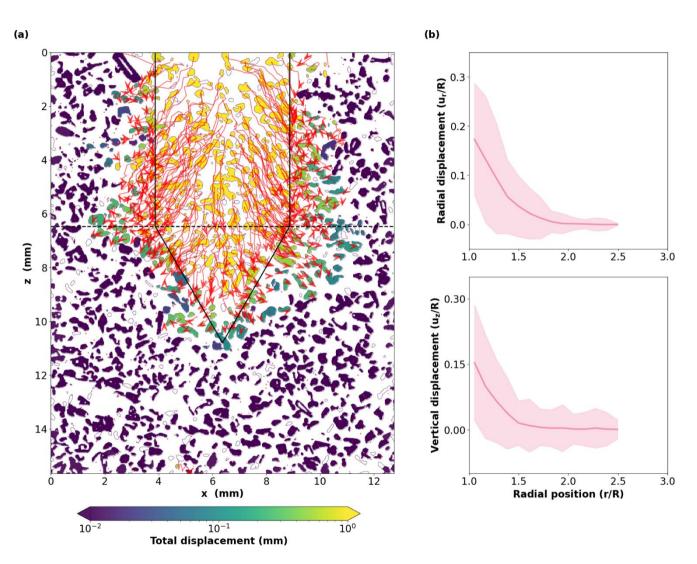




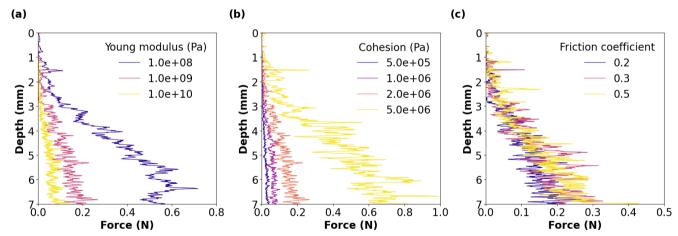
Figure 3: (a) Simulated grain displacement map for the RG sample. The red arrows indicate the grain trajectories while the tip is penetrating (sampling = 0.4 mm). White grains correspond to grains that are not represented in the DEM simulation. The final tip position is indicated by the black solid lines. The horizontal black dashed line indicates the cone top. (b) Radial (upper panel) and vertical (lower panel) displacement profiles (red curves) for the RG sample. These profiles represent averages computed from the sample surface to the cone top. By convention, downward (respectively upward) movement corresponds to positive (respectively negative) values of vertical displacement. The shadowed areas around the solid lines represent the standard deviation of grain displacements. The results are obtained with the mechanical parameters indicated in Table 3.

340 Figure 3 (a) shows the total displacement of the grains as well as grain trajectories. The largest displacements (up to several 341 mm) are observed for grains initially located on the path of the tip. Around the tip, the displacements are < 1 mm and are 342 mainly localised close to the tip. Grain trajectories indicate that grains are pushed downward from each side of the tip. Grains 343 initially located on the tip axis display quasi-straight vertical trajectories. The trajectories become more radial and curved away 344 from the tip medial axis, with grains also being pushed aside. Both radial and vertical displacement profiles show a pronounced 345 decreasing trend and reach almost zero values at a radial position of about 1.7-1.8R (Fig. 3 (b)). The vertical profile attests of 346 a dominant downward movement of the grains close to the tip. Similar observations are made for the DH (Fig. S9) and PP 347 (Fig. S11) samples. In contrast, for the RGIr sample, vertical displacements are smaller and oriented slightly upward on 348 average, for the mechanical parameters chosen here (Fig. S7).

349 **3.2 Sensitivity to mechanical parameters**

The influence of the mechanical parameters (Young's modulus, cohesion, friction coefficient) involved in the contact law has been systematically explored. For the RG sample, the force profiles obtained for the different values of the parameters within the explored ranges (Table 2) are presented in Figure 4, and synthetic plots of the sensitivity of the statistical indicators to these parameters are presented in Figure 5. The results for the other snow samples can be found in Sect. S2.3. Table S3 also summarises the values of statistical indicators in all cases.

355



Force (N) Force (N)

First, it can be observed that increasing Young's modulus decreases the mean macroscopic force (Figs. 4 (a) and 5 (a)) and the correlation length (Fig. 5 (c)). The influence of Young's modulus on the amplitude of force fluctuations is more complex and displays a co-dependency with the cohesion values (Fig. 5 (b)). For low (respectively high) cohesion values, the amplitude of

force fluctuations shows a decreasing (respectively increasing) trend with Young's modulus. Regarding the influence of cohesion, it is observed that increasing this parameter increases the three statistical indicators. Finally, increasing the friction coefficient, generally also leads to an increase of the three statistical indicators. Note however that, over the range of explored friction coefficient values (0.2-0.5), the sensitivity to this parameter is less important than for the other two mechanical parameters (where *E* is varied over two orders of magnitude and *C* is varied over one order of magnitude). Despite changes in absolute force values, the evolution of the force profiles (Figs. S14, S18 and S22) and statistical indicators (Figs. S15, S19 and S23) with the mechanical parameters follow similar trends for all the samples.

371

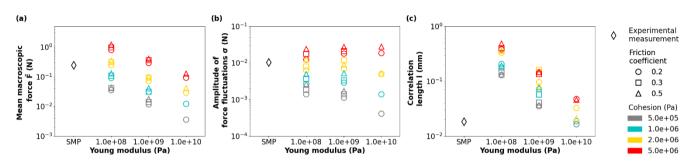


Figure 5: Evolution of statistical indicators as functions of Young's modulus, cohesion and friction coefficient: (a) Mean macroscopic force \overline{F} , (b) amplitude of force fluctuations σ , and (c) correlation length *l*. The experimental results (black diamonds) are also represented in the plots. The results presented here correspond to the RG sample.

376

372

The number of broken bonds per increment of tip penetration depth appears rather insensitive to Young's modulus (Figs. S12 (a), S16 (a), S20 (a), S24 (a)) and is only slightly reduced when cohesion increases (Figs. S12 (b), S16 (b), S20 (b), S24 (b)). Conversely, this quantity is significantly affected by the friction coefficient, with an increase of the average bond failure rate when $tan(\varphi)$ increases (Figs. S12 (c), S16 (c), S20 (c), S24 (c)).

Finally, it is observed that the influence of the mechanical parameters on the radial grain displacement profiles is negligible (Figs. S13, S17, S21, S25). Young's modulus shows no influence on the vertical grain displacement either. Cohesion appears to play a role in the vertical displacement profile for the RGIr sample, by enhancing upward movements. Larger friction coefficients tend to increase the downward movement of the grains close to the tip for all the snow types.

385 **3.3** Comparison of DEM results with experimental measurements

A first noticeable observation is that, for the values of the mechanical parameters tested, the orders of magnitude of the statistical indicators obtained numerically are consistent with the experimental results in most of the cases (Figs. 5, S15, S19, S23, Table S2, Table S3). This demonstrates that the DEM model is indeed capable of reproducing the main characteristics of the CPT force profile (Fig. S26, Table S2). However, we highlight the difficulty of matching the three statistical indicators at once for a given combination of the three mechanical parameters studied. Hence, for the RG sample (Fig. 5), the DEM simulation can reproduce the experimental mean macroscopic force and the amplitude of force fluctuations but tends to overestimate the correlation length by a factor of 8 for the best combination of mechanical parameters. For the RGIr and DH samples (Figs. S15, S18), all the experimental statistical indicators can be reproduced individually, but not for one single combination of the mechanical parameters. For the PP sample, the experimental mean macroscopic force and the amplitude of force fluctuations can be reproduced numerically, but the correlation length is systematically overestimated by a factor of at least 8 (Fig. S23).

397

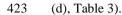
Sample	E (Pa)	C (Pa)	tan(q)	REF	REσ	RE	REtot
RG	1 x 10 ⁹	5 x 10 ⁶	0.2	1.2 x 10 ⁻¹	1.2 x 10 ⁻¹	5.2 x 10 ⁻¹	5.6 x 10 ⁻¹
RGlr	1 x 10 ⁹	1 x 10 ⁶	0.3	5.5 x 10 ⁻²	-4.6 x 10 ⁻¹	1.1 x 10 ⁻¹	4.8 x 10 ⁻¹
DH	1 x 10 ¹⁰	5 x 10 ⁶	0.2	1.2 x 10 ⁻¹	-1.1 x 10 ⁻¹	-2.3 x10 ⁻¹	3.1 x 10 ⁻¹
PP	1 x 10 ⁹	2 x 10 ⁶	0.5	-1.3 x 10 ⁻¹	-1.6 x 10 ⁻¹	6.5 x 10 ⁻¹	6.9 x 10 ⁻¹

Table 3: Selected combination of mechanical parameters for RG, RGIr, DH and PP samples. The indicated values of Young's modulus *E*, cohesion *C* and friction coefficient $tan(\varphi)$ correspond to the combinations that yield the lowest total error RE_{tot} on the statistical indicators (mean macroscopic force \overline{F} , amplitude of force fluctuations σ , correlation length *l*) measured experimentally. Logarithmic relative error RE_k for all the mechanical parameter combinations tested are indicated in Table S3.

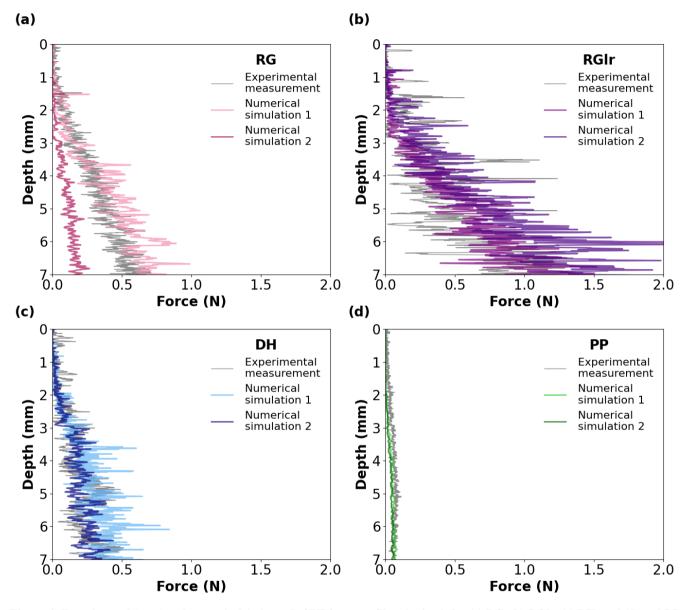
402

403 Based on the sensitivity analysis (Sect. 3.2.3.), we selected for each sample the combination of the three mechanical parameters that minimises the total error RE_{tot} (Tables 3, S3). The corresponding simulated force profiles (referred to as 'Numerical 404 405 simulation 1') are compared with the experimental profiles in Fig. 6. Note that the error values quoted in the text below 406 correspond to relative errors calculated without the logarithmic function, as they are easier to grasp. These values therefore 407 differ from the logarithmic relative errors shown in Tables 3 and S3 and used for the parameter selection. From a qualitative 408 point of view, a good overall agreement is observed between these numerical and experimental force profiles. For the RG 409 sample, the experimental mean macroscopic force is overestimated by $\sim 20\%$ by the numerical result, the amplitude of force 410 fluctuation is overestimated by ~70% and the correlation length is largely overestimated by a factor of 8 (Figs. 5, 6 (a), Table 411 3). Both the experimental and numerical force profiles reach a quasi-steady-state value at about the same depth (~6 mm, S27). 412 For the RGIr sample, the experimental mean macroscopic force is fairly reproduced with a relative error of 6%, the amplitude 413 of force fluctuations is underestimated by $\sim 60\%$ and the correlation length is overestimated by 35% (Figs. S15 and 6 (b), 414 Table 3). We note that the slope change between 2.5 and 3 mm penetration depth is reproduced numerically. However, it 415 appeared difficult to reproduce numerically the amplitude of force fluctuations in the upper section (from 0 to 4 mm) of the 416 experimental profile. For the DH sample, the experimental mean macroscopic force is overestimated by ~25%. The 417 experimental amplitude of force fluctuations is underestimated by 28% and the correlation length is about half of the 418 experimental value (Figs. S19, 6 (c), Table 3). The numerical results minimise the force peaks observed in the upper part of 419 the experimental profile (above 3 mm) but reproduce fairly well the main features of the amplitude of force fluctuations, 420 especially the force "jump" at 3 mm depth. Finally, for the PP sample, the experimental mean macroscopic force is 421 underestimated by \sim 30%, while the experimental amplitude of force fluctuations is underestimated by \sim 60%. In this case, the

422 experimental correlation length could not be reproduced at all, with values overestimated by a factor of 20 (Figs. S23 and 6



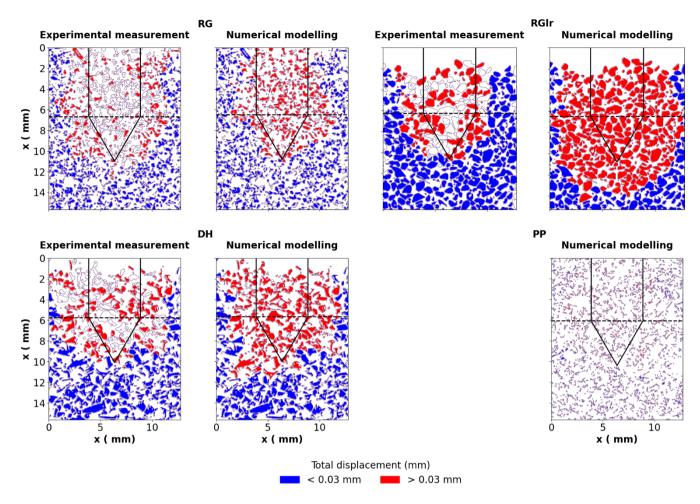




425

Figure 6: Experimental (grey) and numerical (coloured) CPT force profiles obtained for (a) RG, (b) RGlr, (c) DH, and (d) and PP samples. The "Numerical simulation 1" profiles correspond to the best fit of the mechanical parameters determined for each sample (Table 3), while "Numerical simulation 2" profiles correspond to an overall best fit of the mechanical parameters for the four samples ($E = 1 \ge 10^{9}$ Pa, $C = 2 \ge 10^{6}$ Pa and $tan(\varphi) = 0.2$, Table S3).

431 For comparison, we also selected the single set of mechanical parameters that minimises the combined total error RE_{tot} on RG, RGlr. DH and PP samples. Corresponding values are: $E = 1 \times 10^9$ Pa, $C = 2 \times 10^6$ Pa and $tan(\varphi) = 0.2$. The respective 432 433 logarithmic relative errors for each sample can be found in Table S3. As before, the error values presented in the text below 434 correspond to the relative errors without the logarithmic function applied to the values. In general, the corresponding simulated 435 force profiles (referred to as 'Numerical simulation 2' in Fig. 6) also show a fair agreement with the experimental results. For 436 the RG sample, however, the experimental mean macroscopic force is significantly underestimated by ~70% (Figs. 5, 6 (a), 437 Table S3). The numerical amplitude of force fluctuations is underestimated by \sim 35%, while the correlation length is 438 significantly overestimated by a factor of 5. For the RGIr sample, the agreement is acceptable for the three statistical indicators 439 with relative errors around 50%. For the DH sample, the experimental mean macroscopic force is reproduced at 90%, while 440 the experimental amplitude of force fluctuations is underestimated by 60% and the experimental correlation length is 441 overestimated by a factor of ~ 2 . Finally, for the PP sample, the experimental mean macroscopic force is underestimated by 442 \sim 80%, the amplitude of force fluctuations by \sim 85% and the experimental correlation length is again strongly overestimated by 443 a factor of 20 (Figs. S23 and 6 (d), Table S3).



445

Figure 7: Total displacement maps obtained experimentally with μ CT (left panels) and numerically with DEM simulation (right panels) for the RG, RGIr, DH and PP samples. A displacement threshold of 0.03 mm has been set to define the deformation zone (Peinke et al., 2020). White grains correspond to non-trackable grains in μ CT scans (Peinke et al., 2020) and grains not represented in the DEM simulations. The final tip position is indicated with black solid lines. The horizontal black dashed line indicates the cone top. Displacement profiles shown in Fig. 8 are computed from the sample surface to the cone top. Numerical results are obtained with the mechanical parameters indicated in Table 3. The experimental displacement field could not be determined for the PP sample.

As shown in Fig. 7, the DEM simulations also proved capable of reproducing, at least qualitatively, the experimental grain displacement patterns derived from μ CT scans for the four snow types. Essentially similar results are obtained with the individual best-matching sets of mechanical parameters indicated in Table 3 (Fig. 7), and with the globally-matching set of parameters introduced in the previous paragraph (Fig. S28). For the RG sample, the overall shape and size of the deformation zone are well reproduced by the simulations. For the DH sample, the radial extension of the deformation zone is well reproduced by the simulations, but the vertical extension tends to be overestimated. The largest discrepancies are observed for the RGlr sample, for which the radial and vertical extensions of the deformation zone are overestimated compared to the





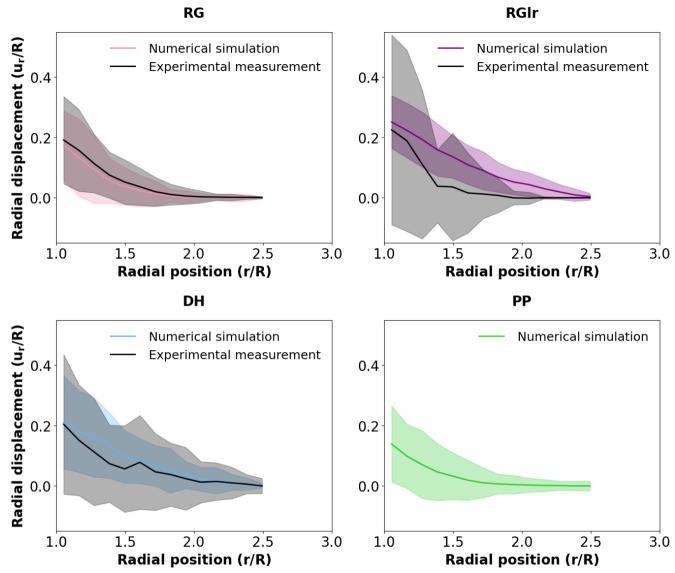


Figure 8: Radial displacement profiles (solid lines) obtained experimentally (black) and numerically (coloured) for the RG, RGlr,
DH and PP samples. The shadowed areas around the solid lines correspond to the standard deviation of grain displacement and
exhibit the variability of the radial displacement of grains. The numerical results are obtained with the mechanical parameters
indicated in Table 3.

468

Similarly, the radial displacement profiles obtained from the DEM numerical simulations are overall in good agreement with
 their experimental counterparts (Figs. 8 and S29). Consistently with the displacement maps, the largest discrepancy is observed

471 for the RGIr sample. In particular, the abrupt slope break seen in the experimental profile at a radial position of about 1.5 is 472 not reproduced in the numerical profile. Note however that, due to a relatively low number of trackable grains (Fig. 7), the 473 standard deviation of the grain radial displacements is larger in the experimental measurements, which may result in a larger 474 uncertainty on the average profile. In contrast, simulations on the RG and DH samples show a very good agreement with the 475 experiments. The CZ (defined with displacement threshold set at 0.03 mm) obtained from numerical simulations extends 476 radially up to 1.6R, 2.2R, 2.0R and 1.5R for the RG, RGIr, DH and PP samples, respectively. In comparison, the CZ derived 477 from uCT scans extends radially up to 1.7R, 1.5R and 1.9R for the RG, RGlr and DH samples, respectively (no measurement 478 for PP sample).

479 4 Discussion

480 **4.1 Evaluation of the DEM model**

We used three mechanical parameters, namely Young's modulus, the cohesion and the friction coefficient, to adjust the simulated force profiles to the experimental results. Overall, the numerical model could reproduce relatively well the mechanical response of all studied numerical samples with a single set of mechanical parameters ($E = 1 \ge 10^9$ Pa, $C = 2 \ge 10^6$ Pa and $tan(\phi) = 0.2$) (Fig. 6), indicating that the differences in the force profiles among the samples are mainly dependent of the snow microstructure.

486 It should also be noted that the values of the mechanical parameters obtained by adjusting the model on the experimental data 487 (either globally for all samples or for each sample individually, Table 3) are reasonably close to the mechanical properties of ice. Young's modulus of ice is measured between 9 x 10⁹ Pa and 10 x 10⁹ Pa (Gammon et al., 1983), while our selected values 488 range between 1 x 10^9 Pa and 1 x 10^{10} Pa. Recall that, in YADE, the Young's modulus is a numerical parameter used to define 489 490 the normal contact stiffness, and is not expected to necessarily correspond to the physical Young's modulus of the material 491 (Sect. 2.2.2). Nevertheless, the fact that the numerical value of E is in the same range of magnitude as the elastic properties of 492 ice provides confidence that the DEM model and the used contact law ([Eqs. (1)-(5)]) correctly capture the physical processes 493 at play. Similarly, the numerical cohesion values, ranging between 1×10^6 Pa and 5×10^6 Pa, are in agreement with typical 494 cohesion values measured on ice (in the range 2×10^6 Pa to 6×10^6 Pa, Schulson and Duval, 2009). Finally, numerical friction 495 coefficients appear to be on the order of 0.2-0.5, while values measured experimentally generally range from 0.02 to 1 (Fish 496 and Zaretsky, 1997; Maneno and Arakawa, 2004). All these results reinforce the confidence in the relevance of the DEM 497 model.

We acknowledge that the mechanical parameters obtained from minimising the logarithmic relative errors on the statistical indicators do not necessarily represent optimal values, in the sense that only a limited number of parameter sets could be tested. Based on the sensitivity analysis, a more proper inversion procedure could be developed to retrieve true optimal values of the mechanical parameters. This would certainly provide more robust elements as to whether a single set of mechanical parameters can be used to represent the experimental results of all snow types, or whether these mechanical parameters differ according

- to the snow type. Our current analysis cannot provide a conclusive answer to this question. Note that ice is a polycrystalline material, whose mechanical behaviour can be strongly anisotropic depending on the ice structure (Fish and Zaretsky 1997; Thorsteisson, 2001; e.g. Maeno and Arakawa, 2004). Therefore, it is not unlikely that ice bonds between grains could be characterised by different mechanical properties depending on the specific conditions of snow formation and evolution.
- As further proof of DEM predictive capabilities, we could also observe that the grain displacement fields measured for the different snow types were overall well reproduced by the simulations (Figs. 7 and 8). In particular, the model captures the radial extent of the deformation zone, which is on the order of 1.5*R*-2.2*R*. A discrepancy between the numerical and experimental radial displacement profiles was observed for the RGIr sample. However, it can be noted that these experimental radial displacement profiles for the RGIr sample also show the largest divergence with the prediction of the cavity expansion model (CEM) (Yu and Carter, 2002), as shown by Peinke et al. (2020). In fact, the radial profile predicted by the CEM for this sample is similar to the radial profile obtained numerically in this study.

514 **4.2 Interpretation**

515 **4.2.1 Sensitivity to the mechanical parameters**

516 The sensitivity analysis revealed a strong influence of the mechanical parameters on the simulation results. In particular, a 517 clear dependence of the mean macroscopic force with Young's modulus E was observed, suggesting that a significant part of 518 the sample undergoes elastic deformation, while brittle failures are confined in a region close to the tip. Note that a similar 519 dependence to E with a cohesive contact law has been observed in DEM modelling of soil compression (De Pue et al., 2019) 520 and snow compression (Bobillier et al., 2020). The mean macroscopic force, the amplitude of force fluctuations and the 521 correlation length all increase with the cohesion C and, to a smaller extent, with the friction coefficient $tan(\varphi)$. This can be 522 related to the fact that increasing cohesion and friction between grains increase bond strength. It was also observed that 523 cohesion tends to prevent bond failures and to favour the upward movement of grains for samples with a large initial density, 524 such as RGIr. In contrast, increasing the friction coefficient enhances the bond failure rate and the downward movement of 525 grains (Figs, S12, S16, S20, S24). When sliding between grains is inhibited, a grain dragged by the tip movement will entrain 526 surrounding grains more easily, thus enlarging the deformation zone and triggering additional bond failures. Finally, radial 527 grain displacements and the radius of the deformation zone appeared to be mostly insensitive to the mechanical parameters, 528 indicating that these features are mainly controlled by CPT configuration and snow microstructure.

529 4.2.2 Compaction zone development

For all snow types, the force profiles computed numerically display a 'S' shape (Figs. 1, S6, S8, S10). We attribute this shape to the development of a compaction zone (CZ) in front of the tip during its penetration into the numerical sample. More specifically, the first stage of the force profiles (slope increase) is presumably caused by the progressive entry of the cone into the sample. The second stage (constant slope) is attributed to the development of the CZ in front of the tip. The third stage (quasi-constant force value) suggests that a steady-state regime, with a fully-developed CZ, is reached. Depending on the snow type, the numerical results indicate that full development of the CZ occurs for 6 mm to 8 mm of penetration depth. These results agree with the experimental profiles for the RG, DH and PP samples. Globally, we can highlight that the DEM simulations are able to reproduce fairly well the global shape of the experimental profiles, and thus to correctly capture the development of the CZ.

539 Nevertheless, in another experimental study, the CZ has been reported to be fully developed only for around 40 mm of depth 540 penetration (Herwijnen, 2013), which is significantly deeper than the experimental and numerical results obtained in this study. 541 A first hypothesis to explain this discrepancy is that since the maximum depth of our CPT force profiles is 10 mm, we might 542 miss information on the full CZ development. A second explanation could be related to the differences in the experimental 543 setups. Indeed, Peinke et al. (2020) performed CPT on snow samples contained in cylinders of 20 mm diameter and 20 mm 544 height, which is significantly smaller than the decimetric snow samples considered by Herwijnen (2013). Boundary effects 545 might thus play a role in limiting the development of the CZ. Finally, the tip geometry also differs between the two studies. 546 Peinke et al. (2020) used a plain tip, while Herwijnen (2013) used the original SMP tip geometry with a cone radius larger 547 than the rod. A sensitivity analysis comparing the two geometries showed an influence over the upper 12 mm of the force 548 profiles (Peinke, 2020). The plain tip geometry resulted in larger values of the mean macroscopic force and the amplitude of 549 force fluctuations values. This effect might also influence the characteristics of the CZ development, which could be studied 550 in the future using the presented numerical model.

551 4.2.3 Grain-tip interaction

552 The sensitivity analysis to the grain shape representation (Sect. S1.1) provides interesting insights into the interpretation of 553 force profiles. In particular, the study highlighted that the grain shape representation could be relatively coarse (high volumetric error E_V) but still produce a force profile with an acceptable mechanical error E_M compared to a reference profile obtained for 554 555 a fine grain shape representation ($E_V < 10\%$) (Fig. S1, Table S1). This is notably the case for the RG sample, for which the 556 selected grain shape representation (L = 5, S = 0.3) corresponds to a value of E_V of about 40%. Large values of E_V often imply 557 grain loss, as the smallest grains identified in the μ CT scans cannot be represented by the DEM with coarse spherical elements. 558 Yet, the similarity of the force profile to the reference force profile indicates the limited contribution of these smallest grains 559 to the macroscopic force, compared to the largest grains with stronger bonds. The loss of grains and bonds might nevertheless directly affect the force fluctuations, providing a potential explanation as to why the DEM model underestimates the correlation 560 561 length obtained experimentally for the samples with the smallest grain sizes (RG and PP) (Figs. 5, S23).

562 **4.2.3 Scaling relation for the mean macroscopic force**

To try and synthesise the large number of simulation results obtained in this study, scaling relations describing the evolution of the statistical indicators as a function of the main simulation parameters can be looked for. We focused in particular on the mean macroscopic force \overline{F} , which was observed to depend both on the mechanical parameters *E*, *C* and $tan(\varphi)$, as well as on

- sample microstructure. Since the range of friction coefficient values (between 0.2-0.5) that we could explore remained limited compared to the ranges of *E* and *C*, the parameter $tan(\varphi)$ was not included in this analysis and the results presented below
- 568 correspond to a single value $tan(\varphi) = 0.3$.
- First, inspection of our results (see Figs. 5 (a), S15 (a), S19 (a), S23 (a)) indicates that the dependencies of the mean macroscopic force \overline{F} to the Young's modulus *E* and cohesion *C* appear to be consistent across the four tested samples (see also Table S4). More precisely, \overline{F} scales with *E* according to a power law of the form $\overline{F} \sim C^{-\alpha}$, with an exponent α on the order of 1/2. Similarly, \overline{F} scales with *C* according to a power law of the form $\overline{F} \sim C^{\beta}$, with β on the order of 3/2.
- Second, we can expect \overline{F} to be also related to the rate of cohesive broken bonds per unit penetration depth. In particular, it is observed (see Figs. S12, S16, S20, S24) that the slope λ of the cumulative proportion of broken bonds as a function of depth is essentially independent of the Young's modulus and cohesion. Conversely, as shown in Fig. 9 (a), this slope λ is linearly related to the initial contact density ν defined as:
- $577 \quad \nu = z\Phi \tag{10}$
- with *z* the coordination number (number of initial cohesive interactions between grains divided by the number of grains, see Table 1) and Φ the volume fraction of the sample (ice density = 917 kg m⁻³, see Table 1).
- 580
- From these different observations, the following scaling law for the mean macroscopic force \overline{F} can be proposed:

582
$$\overline{F} = B T C \left(\frac{c}{E}\right)^{\alpha} f(v)$$
 (11)

- with *B* a dimensionless constant, T (m²) the surface area of the cone (with a radius *R* and a cone apex *a*, Table 2) in contact with the sample, and *f* a function to be determined. Figure 9 (b) shows the dimensionless quantity $\overline{F}T^{-1}E^{1/2}C^{-3/2}$ plotted against the initial contact density ν . We observe that all the simulation results for the four snow types and the different values of Young's modulus and cohesion nicely merge on a unique logarithmic trend. Note, however, that a relatively larger dispersion is observed for RGlr ($\nu = 1.63$) compared to the other samples.
- Equation (11) encapsulates in a single relation the main physics controlling the mean macroscopic force recorded by the penetrometer. In particular, this relation indicates that the influence of snow microstructure can be captured, at least as a first approximation, by the initial contact density ν . Former studies already showed that this parameter plays a key role in the mechanical behaviour of cohesive granular materials (Gaume et al. 2017). Looking for similar relations describing the other statistical indicators (amplitude of force fluctuations and correlation length) constitutes an interesting prospect for future analyses, although we can anticipate these indicators to display more complex dependencies. Further analyses will also be required to explore the influence of the friction coefficient on these relations.
- 595

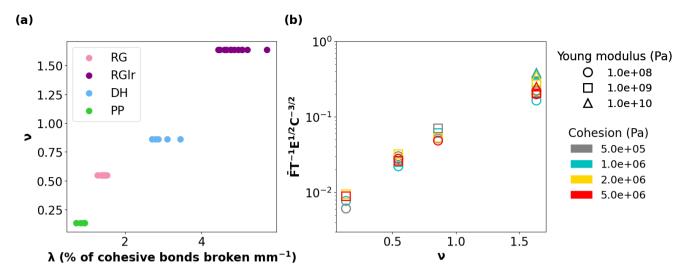


Figure 9: (a) Initial contact density ν versus the slope λ of the proportion of cohesive bonds broken per unit depth (mm⁻¹) for each snow type. The values of initial contact density ν were computed with Eq. (10) and the values indicated in Table 1. The slopes λ were computed from the evolution of the cumulative proportion of cohesive bonds broken (Figs. S12, S16, S20, S24) over a window of 7 mm depth. (b) Dimensionless quantity $\overline{F}T^{-1}E^{1/2}C^{-3/2}$ (see Eq. (11)) versus the initial contact density ν for all simulation results. All the results are provided for a friction coefficient $tan(\varphi)$ of 0.3.

596

603 5 Conclusion

We have evaluated a numerical model based on DEM that reproduces the mechanical behaviour of snow in the brittle regime. The DEM model takes into account the ice properties and the snow microstructure captured by tomography. The experimental configuration of the CPT measurements conducted on different snow types by Peinke et al. (2020) has been reproduced with the DEM model. Three parameters namely, the mean macroscopic force, the amplitude of force fluctuations and the correlation length, were used to quantify the similitude of the numerical and experimental profiles. The grains displacement field was computed and compared to the experimental displacement field derived from μ CT scans acquired before and after the CPT. The DEM model has demonstrated overall a good capability to reproduce the mechanical responses of CPT performed in

611 different snow types. The computed force profiles satisfactorily reproduce the main characteristics of the experimental force 612 profiles. The results revealed that the force profile characteristics are strongly dependent on the microstructure. A sensitivity 613 analysis also demonstrated the dependence of the mechanical response to the mechanical parameters of the contact law. In 614 particular, a simple scaling law could be derived relating the mean macroscopic force computed by the DEM to the mechanical 615 parameters *E* (Young's modulus) and *C* (cohesion) and to the microstructure characteristics captured by the initial contact 616 density. The displacement fields are also well reproduced by the model, except for the RGlr sample showing a larger extent 617 for the numerical results. The agreement in terms of radial displacement profiles is very good. The grains are mainly travelling

- downward during the CPT, although for the RGlr sample, the upward movements close to the surface are not negligible. The
- 619 CPT implies a complex deformation field with a compression zone around the apex and an expansion zone close to the surface 620 (Peinke et al., 2020). Therefore being able to reproduce the force profiles (including high-frequency fluctuations) and 621 displacement fields for this mechanical test constitutes a strong validation of the reliability of the DEM model.
- 622 However, a downside of the DEM method is its high computational cost (simulation times ranging between 1 week to several 623 months depending on the physical and numerical parameters for the chosen CPT configuration), which limited the range of 624 mechanical parameters that could be explored for all snow types. The developed DEM model nonetheless constitutes a versatile 625 approach that could be applied to various materials and configurations in future studies. In particular, it will be possible to use 626 the model to gain more physical insights into the interaction between the tip and the grains, in order to better interpret the CPT 627 force profiles. Such analyses will provide ways to test and derive relevant macro- and micromechanical parameters to 628 characterise the microstructure properties from the CPT force signal solely. In particular, the validity of the assumptions made 629 by the HPP-NHPP method, as well as the influence of the CZ development, will be assessed. Future studies may also consider 630 refining the used contact laws to investigate, e.g. the influence of sintering processes on CPT results.

631 Code availability

632 Codes can be provided by the corresponding author upon request.

633 Data availability

All data can be provided by the corresponding author upon request.

635 Author contribution

- 636 CH, PH and GC developed the numerical model, CH performed simulations and evaluated the numerical model, IP, PH, GC,
- 637 JR designed experiment, IP acquired experimental data, IP processed and analysed experimentation measurements, CH
- analysed and interpreted numerical results, CH wrote the manuscript draft, PH and GC reviewed and edited the manuscript.

639 **Competing interests**

- 640 GC is a member of the editorial board of The Cryosphere. The peer-review process was guided by an independent editor, and
- 641 the authors have no other competing interests to declare.

642 Acknowledgements

- 643 This work benefited from financial supports from the Centre National de la Recherche Scientifique (CNRS), the Centre
- National de la Recherche Météorologique, the Agence Nationale de la Recherche (ANR project MiMESis-3D ANR-19-CE01-645 0009). IGE and CNRM-CEN are parts of LabEx OSUG (ANR-10-LABX-0056). IGE is part of LabEx TEC21 (ANR-11-
- 646 LABX-0030). We thank the two reviewers, Richard Parson and Henning Löwe, for their constructive feedback that enabled 647 us to significantly improve the quality of our manuscript.

648 References

- 649 Bartelt, P., and M. Lehning.: A physical SNOWPACK model for the Swiss avalanche warning: Part I-Numerical model, 650 Cold Reg. Sci.Technol., 35(3), 123–145, doi: 10.1016/S0165-232X(02)00074-5, 2002.
- 651 Bishop, R. F., Hill, R., and F Mott, N.: The theory of indentation hardness tests, Proc. Phys. Soc. 57:321, doi: 10.1088/0959-652 5309/57/3/301, 1945.
- 653 Bobillier, G., Bergfeld, B., Capelli, A., Dual, J., Gaume, J., Herwijnen, A., and Schweizer, J.: Micromechanical modeling of 654 snow failure, The Cryosphere, 14, 39–49, doi:10.5194/tc-14-39-2020, 2020.
- 655 Bobillier, G., Bergfeld, B., Dual, J., Gaume, J., Herwijnen, A., and Schweizer, J.: Micro-mechanical insights into the dynamics
- 656 of crack propagation in snow fracture experiments, Sci Rep 11, 11711, doi:10.1038/s41598-021-90910-3, 2021.
- 657 Bolton, M. D., Gui, M. W., and Phillips, R.: "Review of miniature soil probes for model tests," in Proceedings of the 11th 658 Southeast Asian Geotechnical Conference (Singapore), 85–90, 1993.
- 659 Brun, E., David, P., Sudul, M., and Brunot, G.: A numerical model to simulate snow-cover stratigraphy for operational 660 avalanche forecasting, J. Glaciol., 38(128), 13–22, doi: 10.3189/S0022143000009552, 1992.
- 661 Calonne, N., F. Flin, C. Geindreau, B. Lesaffre and S. Rolland du Roscoat: Study of a temperature gradient metamorphism of
- 662 snow from 3-D images: time evolution of microstructures, physical properties and their associated anisotropy, The Cryosphere,
- 663 8, 2255-2274, doi: 10.5194/tc-8-2255-2014, 2014.
- 664 Calonne, N., F. Flin, B. Lesaffre, A. Dufour, J. Roulle, P. Puglièse, A. Philip, F. Lahoucine, C. Geindreau, J.-M. Panel, S.
- 665 Rolland du Roscoat and P. Charrier, CellDyM : a room temperature operating cryogenic cell for the dynamic monitoring of
- 666 snow metamorphism by time-lapse X-ray microtomography, Geophys. Res. Lett., 42, doi: 10.1002/2015GL06354, 2015.
- 667 Coléou, C., B. Lesaffre, J.-B. Brzoska, W. Ludwig and E. Boller: Three dimensional snow images by X-ray microtomography,
- 668 Ann. Glaciol., 32, 75-81, doi : 10.3189/172756401781819418, 2001.

- 669 Coeurjolly, D., Montanvert, A. and Chassery, J.-M.: Descripteurs de forme et moments géométriques. Géométrie discrète et
- 670 images numériques, Hermès, 2007.
- 671 De Pue, J., Di Emidio, G., Verastegui Flores, R. D., Bezuijen, A., and Cornelis, W. M.: Calibration of DEM material parameters
- to stimulate stress-strain behaviour of unsaturated soils during uniaxial compression, Soil and Tillage Research, 194, 104303,
- 673 doi: 10.1016/j.still.2019.104303, 2019.
- Dowd, T. and Brown, R.L.: A new instrument for determining strength profiles in snow cover, Journal of Glaciology, 32(111):
- 675 299–301, doi: 10.3189/S0022143000015628, 1986.
- Fierz, C., Armstrong, R. L., Durand, Y., Etchevers, P., Greene, E., and McClung, D. M.: "The international classification for
 seasonal snow on the ground," in Tech. Doc. Hydrol. 83 Paris: UNESCO, 2009.
- Fish, A. M. and Zaretsky, Y. K.: Ice strength as a function of hydrostatic pressure and temperature, CRREL report, 38207814,
 1997.
- 680 Floyer, J.A., and Jamieson, J.B. Rate-effect experiments on round-tipped penetrometer insertion into uniform snow, Journal
- 681 of Glaciology, 56(198): 664–672. doi:10.3189/002214310793146322, 2010.
- 682 Freitag, J., Wilhelms, F. and Kipfstuhl, S.: Microstructure dependent densification of polar firn derived from X-ray
- 683 microtomography, J. Glaciol., 50(169), 243–250, doi: 10.3189/172756504781830123, 2004.
- Gammon, P. H., Kiefte; H., Clouter, M. J. and Denner, W. W.: Elastic constants of artificial and natural ice samples by Brillouin
 spectroscopy, Journal of Glaciology, 29(103), 433-460, doi: 10.3189/S0022143000030355, 1983.
- Gaume, J., van Herwijnen, A., Chambon, G., Birkeland, K. W., and Schweizer, J.: Modeling of crack propagation in weak
- snowpack layers using the discrete element method, The Cryosphere, 9, 1915–1932. doi: 10.5194/tc-9-1915-2015, 2015.
- Gaume, J., van Herwijnen, A., Chambon, G., Wever, N., and Schweizer, J.: Snow fracture in relation to slab avalanche release:
- 689 Critical state for the onset of crack propagation, The Cryosphere, 11(1), 217–228, doi: 10.5194/tc-11-217-2017, 2017.
- 690 Gaume, J., Löwe, H., Tan, S., and Tsang, L.: Sacaling laws for the mechanics of loose and cohesive granular materials based
- on Baxter's sticky hard spheres, Physical Review E, 96, 032914, doi: 10.1103/PhysRevE.96.032914, 2017.
- Gubler, H. U.: On the ramsonde hardness equation, IAHS-AISH Publ. 114, 110–121, 1975.
- Gubler, H. U.: Determination of the mean number of bonds per snow grain and of the dependence of the tensile strength of
 snow on stereological parameters, J. Glaciol. 20, 329–341, doi: 10.3189/S0022143000013885, 1978.
- Hagenmuller, P., Chambon, G., Lesaffre, B., Flin, F., and Naaim, M.: Energy-based binary segmentation of snow
 microtomographic images, J. Glaciol. 59 (217), 859–873, doi: 10.3189/2013JoG13J035, 2013.
- 697 Hagenmuller, P., Calonne, N., Chambon, G., Flin, F., Geindreau, C., and Naaim, M.: Characterization of the snow
- microstructural bonding system through the minimum cut density, Cold Reg. Sci. Technol. 108, 72–79, doi:
 10.1016/j.coldregions.2014.09.002, 2014.
- Hagenmuller, P., Chambon, G., Flin, F., Morin, S., and Naaim, M.: Snow as a granular material: assessment of a new grain
- 701 segmentation algorithm, Gran. Matter 16 (4), 421–432, doi: 10.1007/s10035-014-0503-7, 2014.

- Hagenmuller, P., Chambon, G., and Naaim, M.: Microstructure-based modeling of snow mechanics: A discrete element
 approach, Cryosphere, 9(5), 1969–1982, doi: 10.5194/tc-9-1969-2015, 2015.
- 704
- Heggli, M., B. Köchle, M. Matzl, B. R. Pinzer, F. Riche, S. Steiner, D. Steinfeld, and M. Schneebeli: Measuring snow in 3-D
- using X-ray tomography: Assessment of visualization techniques, Ann. Glaciol., 52(58), 231–236,
- 707 doi:10.3189/172756411797252202, 2011.
- Herwijnen, A. V.: Experimental analysis of snow micropenetrometer (SMP) cone penetration in homogeneous snow layers,
- 709 Can. Geotech. J. 50, 1044–1054. doi: 10.1139/cgj-2012-0336, 2013.
- 710 Jamieson, J. B., and Johnston, C. D.: Snowpack characteristics associated with avalanche accidents, Canadian Geotechnical
- 711 Journal, 29(5), 862–866, doi: 10.1139/t92-093, 1992.
- Johnson, J., and Schneebeli, M.: Characterizing the microstructural and microchemical properties of snow, Cold Reg. Sci.
- 713 Technol. 30, 91–100. doi: 10.1016/S0165232X(99)00013-0, 1999.
- 714 Johnson, J. B. and Hopkins, M. A.: Identifying microstructural deformation mechanisms in snow using discrete-element
- 715 modeling, J. Glaciol., 51, 432–442. doi: 10.3189/172756505781829188, 2005.
- 716 Kozak, M. C., Elder, K., Birkeland, K., and Chapman, P.: Variability of snow layer hardness by aspect and prediction using
- meteorological factors, Cold Regions Science and Technology, 37(3): 357–371. doi:10.1016/S0165-232X(03)00076-4, 2003.
- Lunne, T., Robertson, P. K., and Powell, J. J. M.: Cone penetration testing in geotechnical practice. Blackie Academic, EF
 Spon/Routledge, New York, 1997.
- LeBaron, A., Miller, D., and van Herwijnen, A.: Measurements of the deformation zone around a split-axis snow micropenetrometer tip, Cold Reg. Sci. Technol. 97, 90–96. doi: 10.1016/j.coldregions.2013.10.008, 2014.
- 722 Lowe, H., and van Herwijnen, A.: A Poisson shot noise model for micropenetration of snow, Cold Regions Science and
- 723 Technology, 70: 62–70. doi: 10.1016/j.coldregions.2011.09.001, 2012.
- McCallum, A.: A brief introduction to cone penetration testing (CPT) in frozen geomaterials, Ann. Glaciol. 55, 7–14. doi:
 10.3189/2014 AoG68A005, 2014.
- 726 Mackenzie, R., and Payten, W.: A portable, variable-speed, penetrometer for snow pit evaluation, In Proceedings of the 2002
- 727 International Snow Science Workshop, Penticton, B.C. pp. 294–300, 2002.
- 728 Maeno, N. and Arakawa, M.: Adhesion shear theory of ice friction at low sliding velocities, combined with ice sintering,
- 729 Journal of Applied Physics, 95(1), 134-139, doi: 10.1063/1.1633654, 2004.
- 730 Marshall, H.P., and Johnson, J.B.: Accurate inversion of high-resolution snow penetrometer signals for microstructural and
- micromechanical properties, Journal of Geophysical Research: Earth Surface, 114(F4): F04016. doi: 10.1029/2009JF001269,
 2009.
- 733 Mede, T., Chambon, G., Hagenmuller, P., and Nicot, F.: A medial axis based method for irregular grain shape representation
- 734 in DEM simulations, Granular Matter, 20(1), 1-11, doi: 10.1007/s10035-017-0785-7, 2018a.

- 735 Mede, T., Chambon, G., Hagenmuller, P., and Nicot, F.: Snow failure modes under mixed loading, Geophys. Res. Lett. 45 736 (24), 13–351, doi: 10.1029/2018GL080637, 2018b.
- 737 Mede, T., Chambon, G., Nicot, F., and Hagenmuller, P.: Micromechanical investigation of snow failure under mixed-mode 738 loading, International Journal of Solids and Structures, 199, 95–108. doi:10.1016/j.ijsolstr.2020.04.020, 2020.
- 739 Montagnat, M., Löwe, H., Calonne, N., Schneebeli, M., Matzl, M. and Jaggi, M.: On the birth of structural and crystallographic
- 740 fabric signals in polar snow: A case study from the EastGRIP snowpack, Frontiers in Earth Science, 8:365.

741 doi:10.3389/feart.2020.00365, 2020.

- 742 Narita, H.: An experimental study on tensile fracture of snow, Contribut. Inst. Low Temperat. Sci. A32, 1–37, 1983.
- 743 Peinke, I., Hagenmuller, P., Chambon, G., and Roulle, J.: Investigation of snow sintering at microstructural scale from micro-
- 744 penetration tests, Cold Reg. Sci. Technol. 162, 43-55. doi: 10.1016/j.coldregions.2019.03.018, 2019.
- 745 Peinke, I., Hagenmuller, P., Andò, E., Chambon, G., Flin, F. and Roulle, J.: Experimental Study of Cone Penetration in Snow 746
- Using X-Ray Tomography, Front. Earth Sci. 8:63. doi: 10.3389/feart.2020.00063. 2020.
- 747 Proksch, M., Löwe, H., and Schneebeli, M.: Density, specific surface area, and correlation length of snow measured by high-
- 748 resolution penetrometry, J. Geophys. Res. Earth Surf. 120, 346–362. doi: 10.1002/2014JF003266, 2015.
- 749 Reuter, B., Schweizer, J., and van Herwijnen, A.: A process-based approach to estimate point snow instability, Cryosphere, 9, 750 837-847. doi: 10.5194/tc-9-837-2015, 2015.
- 751 Reuter, B., Proksch, M., Löwe, H., Van Herwijnen, A., and Schweizer, J.: Comparing measurements of snow mechanical 752 properties relevant for slab avalanche release, J. Glaciol. 65, 55–67. doi: 10.1017/jog.2018.93, 2019.
- 753 Ruiz, S., Straub, I., Schymanski, S. J., and Or, D.: Experimental evaluation of earthworm and plant root soil penetration-cavity 754 expansion models using cone penetrometer analogs, Vadose Zone J. 15, 1–14. doi: 10.2136/vzj2015. 09.0126, 2016.
- 755 Ruiz, S., Capelli, A., van Herwijnen, A., Schneebeli, M., and Or, D.: Continuum cavity expansion and discrete
- 756 micromechanical models for inferring macroscopic snow mechanical properties from penetration data. Geophys. Res.
- 757 Lett. 44, 8377-8386. doi: 10.1002/2017GL074063, 2017.
- 758 Shapiro, L.H., Johnson, J. B., Sturm, M., and Blaisdell, G. L.: Snow mechanics: review of the state of knowledge and 759 applications, CRREL Rep. 97-3, 1997.
- 760 Schaap, L. H. J. and Fohn, P. M. B.: Cone penetration testing in snow, Canadian Geotechnical Journal 24(3):335-341. 761 doi:10.1139/t87-044.2011.
- 762 Schneebeli, M.: Numerical simulation of elastic stress in the microstructure of snow, Annals of Glaciology, 38. doi: 763 10.3189/172756404781815284, 2004.
- 764 Schneebeli, M., and Johnson, J. B.: A constant-speed penetrometer for high resolution snow stratigraphy, Annals of 765 Glaciology, 26: 107-111, doi: 10.3189/1998AoG26-1-107-111, 1998.
- 766 Schneebeli, M., and Sokratov. S. A.: Tomography of temperature gradient metamorphism of snow and associated changes in
- 767 heat conductivity, Hydrol. Process., 18(18), 3655–3665, doi: 10.1002/hyp.5800, 2004.
- 768 Schulson, E. M. and Duval, P.: Creep and Fracture of Ice, Cambridge University Press, 2009.

- Schweizer, J., Jamieson, J. B., and Schneebeli, M.: Snow avalanche formation, Rev. Geophys., 41(4), 1016.
 doi:10.1029/2002RG000123, 2003.
- 771 Šmilauer, V., Catalano, E., Chareyre, B., Dorofeenko, S., Duriez, J., Gladky, A., Kozicki, J., Modenese, C., Scholtès, L.,
- 772 Sibille, L., Stransky, J., and Thoeni, K.: Yade reference documentation. In V. Šmilauer (Ed.), Yade Documentation (Vol. 474).
- 773 Retrieved from http://yadedem.org/doc, 2010/
- 774 Thorsteinsson, T.: An analytical approach to deformation of anisotropic ice-crystal aggregates, Journal of Glaciologiy, 47
- 775 (158), 507-516, doi; 10.3189/172756501781832124, 2001.
- Vionnet, V., Brun, E., Morin, S., Boone, A., Faroux, S., Le Moigne, P., Martin, E. and Willemet, J.-M.: The detailed snowpack
- scheme Crocus and its implementation in SURFEX v7.2, Geosci. Model Dev., 5, 773-791. doi: 10.5194/gmd-5-773-2012,
 2012.
- 779 Wautier, A., Geindreau, C., and Flin, F.: Linking snow microstructure to its macroscopic elastic stiffness tensor: A numerical
- homogenization method and its application to 3-D images from x-ray tomography, Geophysical Research Letters, 42, 8031-
- 781 8041. https://doi.org/10.1002/2015GL065227, 2015.
- 782 Yu, H. S., and Carter, J.: Rigorous similarity solutions for cavity expansion in cohesive-frictional soils, Int. J. Geomech., 2,
- 783 233–258. doi:10.1061/(ASCE)1532-3641(2002)2:2(233), 2002.