

We thank Dr. Marin for their constructive and thorough feedback on our manuscript. The comments were very helpful, and have led to improvements in the paper. Please see our following responses **in bold** and our proposed alterations.

Pag 7 line 173, It would be interesting to know how much the percentage of infilled areas is.

**The percentage of infilled areas was available in the Supplementary Material. We now moved it into the main manuscript for easier reference.**

Pag 7 line 173, It is not clear if the signature always developed also in the cases where the dates were outside of two and a half standard deviation.

**The SAR snowmelt signature varies within the infilled areas (i.e., those outside two and a half standard deviation). Backscatter does not decrease during ablation or increase after snowmelt has occurred in many infilled pixels. This reduced fidelity becomes increasingly pronounced as slope and elevation increases. Backscatter values within some infilled pixels drop during the ablation period and increase after melt has occurred (i.e., have a snowmelt signature), especially those on low slopes in open or glacierized areas, however. For universal application over the Lajoie, we used standard deviation to remove variable SAR onset estimates stemming from noise or insensitivities in backscatter values. Future research will develop a time series analyses approach to reduce error in melt onset estimates.**

Pag 8 line 198, The identification of the end of the season using sparsely acquired high resolution data is a very challenging task. Some methods have been presented in the literature that try to address this problem using HR multi-source (optical-optical and optical-SAR) data and it is worth to mentioning them

**To improve the detection of snow disappearance, we used Landsat 8 and Sentinel-2 images (optical-optical), however additional preprocessing, or more complex snow detection algorithms, may improve results. Due to the large volumes of data we processed, additional optical image processing, or optical-radar fusion, was beyond the scope of the study. Line 405 has been amended and the following passage has been added:**

**“The inclusion of optical data with a more frequent revisit interval or the use of complex snow detection algorithms could improve this analysis. Support vector machine algorithms have yielded promising results for multi-source snow detection with SAR-optical data in mountainous environments (He et al. 2015, Lui et al. 2020). Further, downscaling and gap filling algorithms can improve snow cover detection using multi-source optical data (Premier et al. 2021, Revuelto et al. 2021). Revuelto et al. (2021) produced 20 m snow cover products by downscaling observations from the high-frequency, low-resolution Moderate Resolution Imaging Spectroradiometer with Sentinel-2 observations. In future studies, downscaling or machine learning methods can be adopted to reduce error in snow disappearance estimates.”**

Table 2, It is not clear, at least to me in the present form, what the image frequency is. For example 1-7 days refer to the day in which there is an acquisition?

The image frequency here refers to the number of days between image acquisitions. For example, in 2021 from the descending orbit there was a minimum gap between acquisitions of one day, and a maximum gap of 12 days. The chart has been changed to 'Acquisition Frequency'.

Table 3, It would be interesting to know the percentage of image for which the cloud cover is less than a given low threshold e.g., 30%.

For optical images, we used approximately half of all available Sentinel-2 and Landsat-8 images. In 2018, 217 images were available with 53% falling below the 60% cloud cover threshold. In 2019, 2020, and 2021, 51%, 56%, and 55% of images were used, respectively. We added the following sentence (line 184):

“Images were further filtered to remove scenes with more than 60% cloud cover, resulting in the use of 53%, 51%, 56% and 55% of all available images from 2018-2021, respectively.”

Figure 2, It is not clear why the shaded blue representing the melt period is not stopping at SWE = 0 for some years. I think this is the rule to be applied once the onset is identified.

In Figure 2 there was an error in the graph, the shaded blue area for the Green Mountain station was displayed on the graph for the Downton Lake Upper station. The shaded area has now been corrected and updated in the manuscript. We chose to use the final breakpoint in the piecewise linear regression to mark the end of the melt period as SWE does not reach zero. After the end of snowmelt, the observed SWE often ranges from 5 mm to 15 mm due to error in the snow pillows. The final breakpoint was used instead as it can be applied universally between the stations and study years.

Interestingly the onset for the runoff is derived in simplified snow model by considering the average temperature (and the radiation) i.e., degree day model. If air temperatures are available for the Lajoie basin, it would be interesting to discuss the difference between Sentinel-1 in identifying the runoff onset (temperature can be spatialized at high resolution and thresholded accordingly).

Thank you for this suggestion – using a simple degree day model (Eq. 1), we now compare SAR estimates of melt onset with modelled estimates based on extrapolated air temperatures from the Downton Lake Upper snow pillow.

$$M = C_m(T_a - T_b) \quad \text{Eq. 1}$$

where M is snowmelt in mm d<sup>-1</sup>, C<sub>m</sub> is a degree-day coefficient in mm day<sup>-1</sup> °C<sup>-1</sup>,

$T_a$  is the mean daily air temperature in  $^{\circ}\text{C}$ , and  $T_b$  is the threshold temperature in  $^{\circ}\text{C}$ . For this analysis, we used a degree-day coefficient of  $3.2 \text{ mm day}^{-1} \text{ }^{\circ}\text{C}^{-1}$ , based on reported values from North American and maritime snowpacks (DeWalle et al. 2002, Hock 2003). We selected a threshold temperature of  $0^{\circ}\text{C}$ , which is commonly used in snowpack studies (Zhou et al. 2021). We extrapolated air temperatures from the Downton Lake Upper station using a lapse rate of  $6.5^{\circ}\text{C km}^{-1}$ . We identified melt onset per pixel as the first date in the ablation season where accumulated melt is only positive (i.e., the date after the last negative or zero melt value).

On average, SAR predicts earlier melt onset compared to results from the degree day model (Figure 1). These discrepancies may stem from the melt model and its assumptions, the inability of temperature-based models to account for aspect, slope, and terrain cover, or errors in the SAR-derived melt onset estimates.

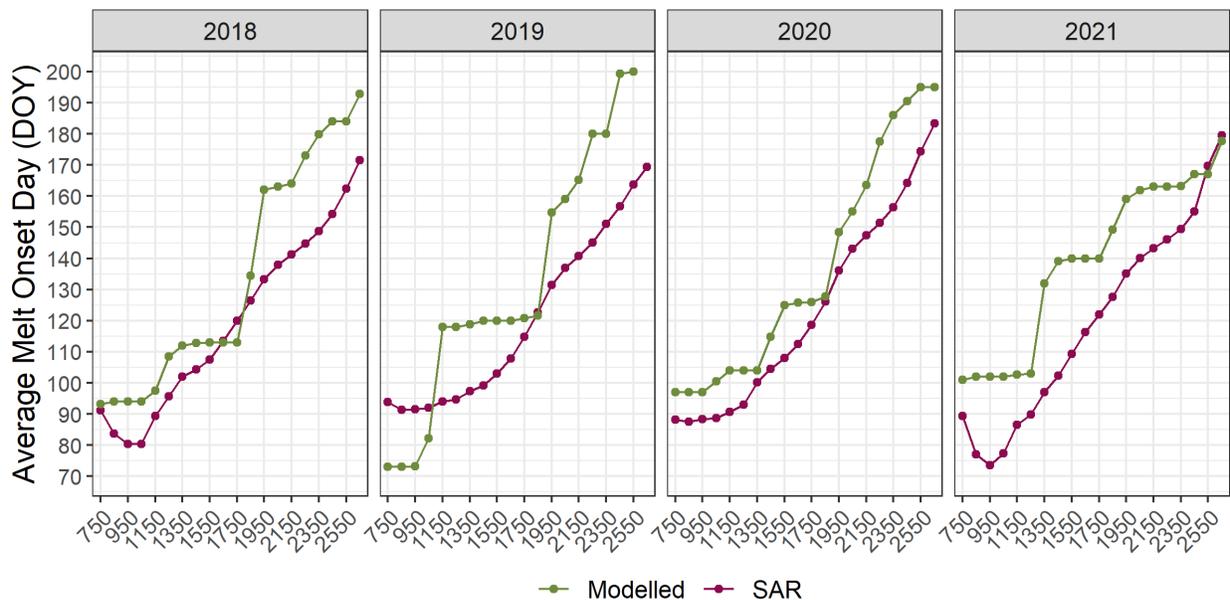


Figure 1. Comparisons of average melt onset dates from SAR and modelled estimates. Modelled estimates are obtained from a simple degree day model using temperature data from the Downton Lake Upper snow pillow.

At low elevations the Lajoie is forested. In warm maritime environments, forest cover can enhance snowmelt at low elevations due to increased longwave radiation and warmer air temperatures (Lundquist et al. 2013). Forest cover can also prevent longwave radiative cooling at night, leading to winter snowpacks that are closer to  $0^{\circ}\text{C}$ . In these low-elevation forests, the earlier onset estimates from SAR, when compared to the degree-day model, may capture the effects of forests on melt onset. These effects would not be accounted for in a temperature-based model. At high elevations in the Lajoie, snow melt may be accelerated due to the greater prevalence of steep slopes. As steep slopes tend to accumulate less snow, they may melt earlier compared to surrounding areas which a temperature-based model can not account for. Further, surface melt can occur when temperatures are below  $0^{\circ}\text{C}$ , leading to the underestimation of melt when using  $0^{\circ}\text{C}$  as the threshold temperature.

Additional analyses are needed to determine how model calibration, variation in lapse rates, and topography impact results, and is beyond the scope of the project. Future work will compare SAR and modelled estimates of snowmelt onset in greater depth.

The sampling time provided by Sentinel-1 seems to be not adequate, in Shannon sense, to properly sampling the melting which has probably a temporal resolution less than one day. That means that the error could be potentially of several days. How is this uncertainty propagating in the case of snow melt duration analysis when different years are compared? What is the ideal revisit time needed for this kind of analysis?

Snowmelt durations are impacted by both errors in SAR onset estimates and optical snow disappearance estimates. Between study years SAR acquisitions are available every five to twelve days, averaging 6 days in each year. Optical observations are more variable between study years, and are summarized in Table 1.

Table 1. Revisit intervals for optical imagery from Sentinel 2 and Landsat 8 for the Lajoie Basin.

Year	Minimum Revisit Interval	Maximum Revisit Interval	Average Revisit Interval
2018	1 day	17 days	4 days
2019	1 day	15 days	4 days
2020	1 day	18 days	3 days
2021	1 day	8 days	3 days

We calculated errors in snowmelt durations ( $\sigma_D$ ) using Equation 2.

$$\sigma_D = \sqrt{\sigma_{ONS}^2 + \sigma_{END}^2} \quad \text{Eq. 2}$$

where error in the onset ( $\sigma_{ONS}$ ) is six days, and error in snow disappearance ( $\sigma_{END}$ ) is the average revisit interval from the annual Landsat 8 and Sentinel 2 image collections. Rounding to the nearest day, this yields an average error in the duration estimates of  $\pm 7$  days for all study years. Ideally, a revisit time for SAR observations of three days would reduce this error to  $\pm 4$  days. Daily SAR observations could reduce error in durations to  $\pm 3$  days; however, for reduced error, or operational applications, high-resolution optical imagery is needed at more frequent revisit intervals.

Table 3 (now Table 4 due to revisions) has been updated to include the revisit frequency of optical observations. Line 401 has been amended, and the following passage has been added.

“Inaccuracies in snowmelt durations are attributed to errors in SAR onset estimates and optical snow disappearance estimates. Between study years SAR acquisitions are available every five to twelve days, averaging six days in each year. Optical observations are variable between study years (Table 4); however, average revisit intervals are consistent at four days in 2018 and 2019

and three days in 2020 and 2021. We calculated errors in snowmelt durations ( $\sigma_D$ ) using Equation 1.

$$\sigma_D = \sqrt{\sigma_{ONS}^2 + \sigma_{END}^2} \quad \text{Eq. 1}$$

where error in the onset ( $\sigma_{ONS}$ ) and disappearance dates ( $\sigma_{END}$ ) are taken as the average revisit interval for each data type and study year. Rounding to the nearest day, this yields an average error in the duration estimates of  $\pm 7$  days for all study years. Ideally, a revisit time for SAR observations of three days would reduce this error to  $\pm 4$  days. Daily SAR observations could reduce error in durations to  $\pm 3$  days; however, for reduced error, or operational applications, high-resolution optical imagery is needed at more frequent revisit intervals.”

## References

- DeWalle, D. R., Z. Henderson and A. Rango. 2002. Spatial and temporal variations in snowmelt degree-day factors computed from snotel. In Proceedings of the 70th Annual Meeting of the Western Snow Conference, Granby, CO, USA: 73–81.
- He, G., P. Xiao, X. Feng, X. Zhang, Z. Wang and N. Chen. 2015. Extracting snow cover in Mountain areas based on SAR and optical data. *IEEE Geoscience and Remote Sensing Letter*, 12(5): 1136-1140.
- Hock, R. 2003. Temperature index melt modelling in mountain areas. *Journal of Hydrology*, 282: 104-115.
- Lui, Y., X. Chen, JS. Hao and L. Li. 2020. Snow cover estimation from MODIS and Sentinel-1 SAR data using machine learning algorithms in the western part of the Tianshan Mountains. *Journal of Mountain Science*, 17(4): 884-897.
- Lundquist, J. D., S. E. Dickerson-Lange, J. A. Lutz and N. C. Cristea. 2013. Lower forest density enhances snow retention in regions with warmer winters: A global framework developed from plot-scale observations and modeling. *Water Resources Research*, 49: 6356-6370.
- Premier, V., C. Marin, S. Steger, C. Notarnicola & L. Bruzzone. 2021. A novel approach based on a hierarchical multiresolution analysis of optical time series to reconstruct the daily high-resolution snow cover area. *IEE Journal of Selected Topics in Applied Earth Observation and Remote Sensing*, 14: 9223-9240.
- Revuelto, J., E. Alonso-Gonzalez, S. Gascoin, G. Rodriguez-Lopez and J. I. Lopez-Moreno. 2021. Spatial downscaling of MODIS snow cover observations using Sentinel-2 snow products. *Remote Sensing* 13(22): 4513.