1 Significant underestimation of peatland permafrost along the

2 Labrador Sea coastline in northern Canada

- 3 Yifeng Wang¹, Robert G. Way¹, Jordan Beer¹, Anika Forget¹, Rosamond Tutton^{1,2}, Meredith C. Purcell³
- 4 ¹Northern Environmental Geoscience Laboratory, Department of Geography and Planning, Kingston, K7L 3N6, Canada
- 5 2Global Water Futures, Wilfrid Laurier University, Yellowknife, X1A 2P8, Canada
- 6 ³Torngat Wildlife, Plants, and Fisheries Secretariat, Happy Valley-Goose Bay, A0P 1E0, Canada
- 7 Correspondence to: Yifeng Wang (yifeng.wang@queensu.ca)
- Abstract. Northern peatlands cover approximately four million km², and about half of these peatlands are estimated to contain 8 permafrost and periglacial landforms, like palsas and peat plateauxplateaus. In northeastern Canada, peatland permafrost is predicted to be spatially concentrated in the western interior of Labrador and but is assumed to be largely absent along the 10 11 Labrador Sea and Gulf of St. Lawrence coastline. However, the paucity of observations of peatland permafrost in the interior, coupled with ongoing traditional and ongoing use of perennially frozen peatlands along the coast by Labrador Inuit and Innu, 12 suggests a need for re-evaluation of -easts doubt on the reliability of existing maps of peatland permafrost distribution estimates 13 in for the region. In this study, we develop a multi-stage consensus-based point inventory of peatland permafrost complexes 14 in coastal Labrador and adjacent parts of Quebec using high-resolution satellite imagery, and we validate it with extensive 15 field visits and low-altitude aerial photography and videography. A subset total of 1885 2092 wetland complexes that 16 17 potentially contained peatland permafrost were inventoried were inventoried, of which 1023 1119 were interpreted asclassified 18 as likely containing peatland permafrost. Likely peatland permafrost complexes were mostly found in lowlands within 40-22 19 km of the coastline where mean annual air temperatures often exceeded up to +1.2 °C are recorded. Evaluation of the 20 geographic distribution of peatland permafrost complexes reveals aA clear gradient in peatland permafrost distribution exists from the outer coasts, where peatland permafrost is more abundant, to inland peatlands, where permafrost is generally absent. 21 22 This coastal gradient may be attributed to a combination of climatic and geomorphological influences which lead lead to lower 23 insolation, thinner snowpacks, and more poorly drained, frost-susceptible materials along the coast. The results of this study also-suggest that existing maps estimates of permafrost distribution for southeastern Labrador require adjustments -to better 24 25 reflect thethe abundance of peatland permafrost complexes which are located to the south of the regional sporadic discontinuous permafrost limit. This study constitutes the first dedicated peatland permafrost inventory for Labrador, and our 26 results provides an important baseline for future mapping, modelling, and climate change adaptation strategy development in 27 28 the region.

1 Introduction

Near the southern boundary of latitudinal permafrost zonation, lowland perennially frozen ground is primarily restricted to peatlands wetlands as in the form of palsas (peat mounds with a frozen core of mineral and organic material) and peat plateauxplateaus (fields of frozen peat elevated above the general surface of the surrounding peatlandlarge, elevated fields of frozen peat) (Payette, 2004; International Permafrost Association Terminology Working Group, 2005; Zoltai, 1972; Zoltai and Tarnocai, 1975). Persistence of these cryotic landforms at the extreme limits of their viability is facilitated by a large temperature offset between the ground surface and the top of permafrost, caused by the thermal properties of thick layers of overlying peat and the buffering effect of ground ice (Burn and Smith, 1988; Williams and Smith, 1989). In recent years, many studies have shown that peatland permafrost can be very sensitive to climate warming and ecosystem modifications (Beilman et al., 2001; Borge et al., 2017; Thibault and Payette, 2009). Understanding the distribution of these ice-rich, thaw-sensitive periglacial environments is important for predicting assessing thermokarst potential (Gibson et al., 2021; Olefeldt et al., 2016), local hydrological and vegetation change (Zuidhoff and Kolstrup, 2005), regional infrastructure and or land-use planning, and global carbon stores and carbon cycling activities (Hugelius et al., 2014).

Palsas and related landformspeat plateaus are primarily thought to occur in continental locations (Fewster et al., 2020; Hustich, 1939) where colder winters allow deeper frost penetration and drier summers promote less thaw. For example As such, palsas and peat plateauxplateaus have been described in many continental locations in Canada, including Yukon Territory, the Northwest Territories, and the Prairie provinces (e.g., Beilman et al., 2001; Coultish and Lewkowicz, 2003; Mamet et al., 2017; Thie, 1974; Zoltai, 1972). However, but these landforms they have also been documented in coastal locations including the Hudson Bay Lowlands in northern Manitoba, Ontario, and Quebec (e.g., McLaughlin and Webster, 2014; Ou et al., 2016; Pironkova, 2017). In the Labrador region of northeastern Canada, continental- to hemispheric-scale studies have suggested that depicted peatland permafrost is as present in the region's continental interior but is as far less abundant or completely absent along most of the Labrador Sea coastline (Fewster et al., 2020; Hugelius et al., 2020; Olefeldt et al., 2021). However, historic and contemporary use of coastal peatland permafrost environments by Labrador Inuit and Innu is well documented (Anderson et al., 2018), and published field-based observations (e.g., Anderson et al., 2018; Andrews, 1961; Brown, 1975, 1979; Davis et al., 2020; Dionne, 1984; Elias, 1982; Hustich, 1939; Seguin and Dionne, 1992; Smith, 2003; Way et al., 2018; Wenner, 1947) suggest that peatland permafrost is more abundant along some sections of the coast than in the interior. This recurring ongoing undermisestimation of peatland permafrost in the region has led to has an impact on predictions of low ground ice content (O'Neill et al., 2019), thermokarst potential (Olefeldt et al., 2016), and carbon content (Hugelius et al., 2014) in the region.

Locally, preservation of peatland permafrost complexes is also-relevant to Labrador Inuit and Innu because these areas are frequented for traditional activities such as bakeapple (cloudberry; Inuttitut: appik; Innu-aimun: shikuteu; Rubus chamaemorus) berry-picking (Anderson et al., 2018; Karst and Turner, 2011; Norton et al., 2021), goose hunting, and fox trapping (Way et al., 2018). Improvements to our understanding of regional peatland permafrost distribution will provide an

important baseline for local and regional climate change adaptation strategy development, while better representation of the distribution of thaw-sensitive terrain will inform future development of linear and built infrastructure in and around Labrador's coastal communities coastal Labrador (Way et al., 2021b; Bell et al., 2011).

62 63

64 65

66

67

68 69

70 71

72

73 74

75

76 77

78

79

80

81

82

83

84

Previous peatland permafrost mapping efforts in Labrador hasve been limited to scattered observations of palsa bogs through from the National Topographic Database (Natural Resources Canada, 2005) and the Ecological Land Classification (Environment Canada, 1999), withand no dedicated comprehensive peatland permafrost inventorying efforts have been completed to date (Way et al., 2018). In this study, we develop a multi-stage, consensus-based point inventory of contemporary peatland permafrost complexes within 100 km of the along the Labrador Sea and part of the Gulf of St. Lawrence coastline (Figure 1), comprising the region of Nunatsiavut and surrounding areas, including the land claims agreement-in-principle of the Labrador elaimed by the Innu Nation (Nitassinan) and coastal areas claimed by the NunatuKavut Community Council (NunatuKavut). The goal of this inventory is to map and contextualize the contemporary distribution of peatland permafrost complexes throughout coastal Labrador, using extensive validation efforts from a combination of field visits and low-altitude image and videory acquisitions methods. We hypothesize that this point-based inventory will reinforce the local understanding of a high abundance of peatland permafrost landforms in coastal locations, which will be relevant for This point-based inventory will refine our understanding of peatland permafrost distribution at local to regional scales and will be relevant for carbon modelling, land use planning, infrastructure development, and climate change adaptation strategy developmenties at local to regional scales in northeastern Canada. This contribution will also provide insights into the reliability of existing relevant peatland permafrost and permafrost distribution maps-productsin eastern Canada, which currently claim an absence or low abundance of both peatland permafrost and permafrost along the Labrador Sea coastline. Based on our results, we also propose amendments to the current limits of the sporadic discontinuous and isolated patches of permafrost distribution zones in southeastern Labrador. This point-based inventory is a first step towards understanding the distribution of peatland permafrost in Labrador and will contribute to refined which will help to refine regional and global estimates of ground ice content, thermokarst potential, and carbon storage in northern Canada.

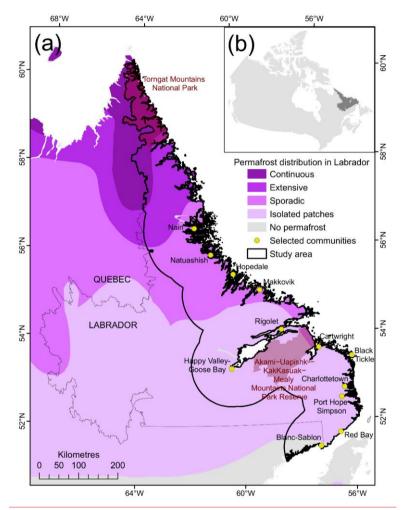


Figure 1. (a) Permafrost zonation in Labrador (Heginbottom et al., 1995) with the boundary for the inventory study area (black line) corresponding to areas within 100 km of the Labrador Sea coastline. Map is annotated with locations of the Torngat Mountains National Park, Akami-Uapishku-KakKasuak-Mealy Mountains National Park Reserve, and selected communities; (b) Inset map showing Labrador's position in Canada.

2 Study area

2.1 Bioclimatic setting

Labrador's climate is strongly influenced by atmosphere-ocean interactions from the adjacent Labrador Sea (Barrette et al., 2020; Way and Viau, 2015). Labrador's-In coastal Labrador, long, cold winters and short, cool summers are largely dictated by the Labrador Current that carries cold Arctic waters down the eastern coast of mainland Canada (Banfield and Jacobs, 1998; Foster, 1983; Roberts et al., 2006; Way et al., 2017). Mean annual air temperatures (1980-2010) decrease with continentality and latitude, ranging from -1+.9-2°C in parts of the Torngat Mountains National Park to +1.5-°C near the community of Blanc-Sablon (Karger et al., 2017, 2021). Labrador is also characterized by some of the highest precipitation amounts in the North American boreal zone (Banfield and Jacobs, 1998; Hare, 1950) due to its varying relief, high moisture availability from the adjacent Atlantic Ocean, and high frequency of passing winter storm systems (Brown and Lemay, 2012). Precipitation totals as high as ~2700 mm per year are estimated for some locations at high elevations along the coast (Karger et al., 2017, 2021), with solid precipitation fractions increasing with both latitude and elevation (~0.35 at Blanc-Sablon; ~0.5 at Nain) (Environment and Climate Change Canada, 2022).

Ecologically, Labrador exhibits a combination of is characterized by taiga forests in the interior, tundra in the north, and wind-swept coastal barrens along the coastline of the Labrador Sea (Roberts et al., 2006). Tree cover is sparse in the coastal barrens due to a combination because of climatic and physiographic limitations, but dense patches of black spruce (*Picea mariana*), white spruce (*Picea glauca*), tamarack (*Larix laricina*), and balsam fir (*Abies balsamea*), interspersed with deciduous trees, like paper birch (*Betula papyrifera*) and trembling aspen (*Populus tremuloides*), do exist in sheltered locations and on some_slopes (Roberts et al., 2006). Wetlands are found throughout Labrador, but total wetland abundance is difficult to assess given widespread disagreement between existing estimates of wetland and peatland extents for this region (Supplement Sect. S1). Generally, wetlands in Labrador tend to decrease in abundance but increase in size as latitude increases. Most wetlands along the southern Labrador coast are classified as raised bogs, while inland, most wetlands are string and blanket bogs (Foster and Glaser, 1986).

114 2.2 Physical environment

Labrador is mostly underlain by igneous and metamorphic bedrock (Roberts et al., 2006). Extensive blankets of glacial till were deposited during and following—the retreat of the Laurentide Ice Sheet (12-6 k years BP) (Bell et al., 2011; Dyke, 2004), along with thin layers of medium—to fine-grained marine and glaciomarine sediments in coastal lowland areas below the marine limit (Fulton, 1995). The post-glacial marine limit decreases with latitude, from ~150 m a.s.l. in southeastern Labrador and along the Quebec Lower North Shore to 0 m a.s.l. at the northernmost tip of Labrador in the Torngat Mountains (Dyke et al., 2005; Occhietti et al., 2011; Vacchi et al., 2018). The broad distribution of near-surface bedrock and hardpans (Smith, 2003) results in poor drainage that has facilitated peatland development across large areas of southern Labrador,

particularly in depressions and over flat deposits, following a general gradient of string and blanket bogs in the interior to raised bogs along the coast (Foster and Glaser, 1986).

2.3 Permafrost distribution

122

123

124

125

127

129

130

131

132

133

134

135

136

137

138

139

140

141

142

143

145

146

147

148

149

150

151

While permafrost conditions in Labrador, including the presence of peatland permafrost landforms, have been noted 126 during ecological, palynological, glaciological, and archaeological surveys and studies (Anderson et al., 2018; Andrews, 1961; Hustich, 1939; Smith, 2003; Wenner, 1947), permafrost-specific field investigations are limited to R.J.E. Brown's (1975) 128 helicopter survey of permafrost conditions in the late 1960s and the Labrador Permafrost Project from 2013 to 2017 that began only in 2013 (Way, 2017). Compared to other parts of Canada, where permafrost focused studies are more concentrated, oOur understanding of permafrost distribution in Labrador has relied on extensive extrapolation of limited field observations and broad assumptions of the interactions between air temperature, vegetation cover, snow cover, and permafrost presence (Ives, 1979). According to the Permafrost Map of Canada (Heginbottom et al., 1995), the area underlain by permafrost in Labrador is less extensive than comparable regions in northern Canada like Yukon Territory or the Northwest Territories. permafrost area in Labrador is relatively low, with a Approximately two-thirds of Labrador is classified in the isolated patches of permafrost zone (<10 % permafrost by area), but the distribution of permafrost does become more widespread farther north (Figure 1). zone. Permafrost distribution in Labrador follows a latitudinal gradient, with extensive discontinuous (50 90% permafrost by area) and sporadic discontinuous (10-50% permafrost by area) permafrost zones in the north, and the isolated patches of permafrost zone in the south (Figure 1). Along the Labrador coastline, the sporadic discontinuous permafrost zone (10-50 % permafrost by area) extends slightly further south along the outer edge of the Akami-Uapishku-KakKasuak-Mealy Mountains National Park Reserve than in the interior, though the rationale justification for this departure is unclearnot clarified in published literature. Continuous permafrost (>90 % permafrost by area) is expected to persist only at high elevations and latitudes, mostly in the Torngat Mountains (Heginbottom et al., 1995).

144 2.4 Inventory extent

> This study is focused on the coastal areas of Labrador and Quebec, within 100 km of the Labrador Sea coastlines (Figure 1). This area of interest was informed by knowledge gained from prior works in the region (Anderson et al., 2018; Andrews, 1961; Brown, 1975, 1979; Davis et al., 2020; Dionne, 1984; Elias, 1982; Hustich, 1939; Seguin and Dionne, 1992; Smith, 2003; Way, 2017; Way et al., 2018; Wenner, 1947) that indicated a greater abundance of peatland permafrost landforms along the coast as compared to the interior of Labrador. Exhaustive descriptions of records of peatland permafrost and other periglacial landforms in Labrador have been presented by Brown (1979) and Way (2017), both of whom found limited evidence of peatland permafrost in Labrador's interior.

Formatted: Indent: First line: 1.27 cm

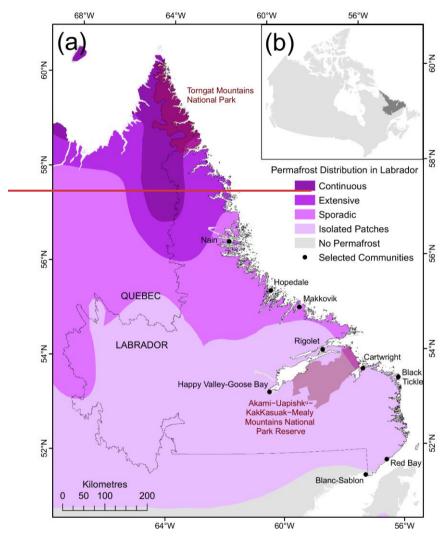


Figure 1. (a) Permafrost zonation in Labrador according to the Permafrost Map of Canada (Heginbottom et al., 1995) and locations of the Torngat Mountains National Park, Akami-Uapishk"—KakKasuak—Mealy Mountains National Park Reserve, and communities mentioned in the text; (b) Location of Labrador in relation to Canada.

3 Methods

156 157

158

159

160

161

162

163

164

165

167

168

169

170

171

172

173

174

175

176

177

178

179 180

181 182

183

Palsas and peat plateaux plateaux are typically found in bogs and in the region may measure up to 4 m higher than their surrounding wetlands, so large peatland permafrost landforms can be identified and mapped from high-resolution satellite imagery (Borge et al., 2017; Gibson et al., 2020, 2021). Our point inventory, which includes only the largest and most visually apparent of large peatland permafrost complexes within 100 km of the along the Labrador Sea and Gulf of St. Lawrence coastline, was generated through a multi-stage mapping and consensus-based review process, supported by extensive validation efforts mostly completed between 2013-2017 and 2021-2022. Mapping and identification activities were informed by existing wetland and peatland distribution products (Supplement Sect. S1), but significant disagreement between these products limited their direct application and utility during the inventorying process. An initial inventory of wetlands of interest (WOIs) was developed as a subset of the wetlands in coastal Labrador deemed potentially suitable (e.g., bogs and fens) for the development and persistence of peatland permafrost landforms. Prospective peatland permafrost complexes were identified 166 from high resolution satellite imagery, resulting in an initial database of wetlands of interest (WOIs). The presence of peatland permafrost landforms within these WOIs was then evaluated through a consensus-based review process-of high-resolution satellite imagery involving by three mappers with permafrost-specific field experience in the region. Final interpretation of peatland permafrost presence or absence within the WOIs was based on reviewer agreement and was informed by field- and imagery-based validation of peatland permafrost landform presence or absences in the region.

3.1 Data sources

WOIs were identified and evaluated using Maxar (Vivid) optical satellite imagery, available as the World Imagery basemap via ArcGIS Online (0.5 m ground sampling distance; 5 m absolute spatial accuracy) (Esri, 2022). These satellite imagery mosaics consisted of summer imagery with minimal cloud and snow cover, with and acquisition dates for Labrador that rangedranging from 2010 to 2020.

Topographic data from Natural Resources Canada covering the WOIs were extracted from the Canadian Digital Elevation Data (CDED; 50 m spatial resolution), with a small gap near the provincial border between Labrador and the Quebec Lower North Shore that was filled in using the Canadian Digital Surface Model (CDSM), Gridded mean annual air temperature (MAAT) and mean annual thawing degree days (TDD) for the 1981 to 2010 climate normal were extracted from CHELSA V2.1 (~1 km spatial resolution) (Karger et al., 2017, 2021) at the WOI locations. Mean annual freezing degree days (FDD) for the WOI locations for 1981 to 2010 were calculated from MAAT and TDD over the same climate normal, following prior work in the region (Way et al., 2017; Way and Lewkowicz, 2018).

3.2 Inventorying peatland permafrost complexes

3.2.1 Identifying wetlands of interest (WOIs)

184 185

186

187

188

189

190

191

192

193

194

195

196

197

198

199

200

201

202

203

204

205

206

207

208

209

210

211

212

213

A team of three mappers used ArcGIS Online to identify and place point features within WOIs throughout coastal Labrador (Supplement Sect. S1Figure 2). The point-based nature of the inventorying process allowed for evaluation of the entire study area by incorporating field- and imagery-based validation for many WOIs over a large study area, as opposed to detailed validation of peatland permafrost areal coverage within a given WOI. These is point based mapping approach aimed to generate a conservative sample of some of the largest peatland permafrost complexes in the region. Mapping and identification activities-WOIs were mainly restricted to include only those WOIs that contained prospective peatland permafrost landforms that exceeded exceeding 2 m in length or width (-2 m in length or width 4 m²-), which was determined to be the smallest detectable feature based ondue to the 0.5 m spatial resolution of the satellite imagery. This point based mapping approach aimed to generate a conservative sample of some of the largest peatland permafrost complexes in the region. Some prospective peatland permafrost locations in interior Labrador and along the Labrador Quebec interior border were included in the inventory (Brown, 1955, 1975; Way, 2017; Way and Lewkowicz, 2014), but mapping activities were primarily concentrated along the coast. This focus on the coastal barrens ecozone was based on an evaluation of existing literature-based observations (Anderson et al., 2018; Andrews, 1961; Brown, 1975, 1979; Davis et al., 2020; Dionne, 1984; Elias, 1982; Hustich, 1939; Seguin and Dionne, 1992; Smith, 2003; Way, 2017; Way et al., 2018; Wenner, 1947) and records of palsa bogs from the National Topographic Database (Natural Resources Canada, 2005) that favoured the coastline, and a lack of identified features during extensive field activities in the interior (Way, 2017; Way and Lewkowicz, 2016, 2018) (Supplement Sect. S2). Mappers were instructed to identify WOIs based on local geomorphology, local hydrology and drainage patterns, the presence of a white or grey lichen surface cover corresponding to Cladonia and/or Ochrolechia spp. lichens, evident shadows indicative of elevated landform edges and surface uplift, and the presence of thermokarst pondings or exposed peat that may indicate indicative of ongoing thaw processes. The inventory sought to only include active contemporary peatland permafrost landforms, so WOIs with extensive thermokarst pondsing but no evident peatland permafrost landforms were not included in the database. Individual WOIs ranged in size from ~0.2 km² to larger than ~3.5 km². However, the total area underlain by peatland permafrost within each WOI was not able to be reliably evaluated using satellite imagery. WOIs near one another were sometimes difficult to discern due to potential connectivity between adjacent systems, but contiguous WOIs could generally be identified by differences in drainage, vegetation, and morphology, or because of separation by linear infrastructure like roads, airstrips, and trails (Figure 2). Mappers also assigned each WOI a ranking of 1 (low confidence) to 3 (high confidence) to reflect their self-assessed score to reflect their confidence in their interpretation of permafrost presence within the wetland complex (1 = low confidence, 2 = medium confidence, 3 = high confidence).

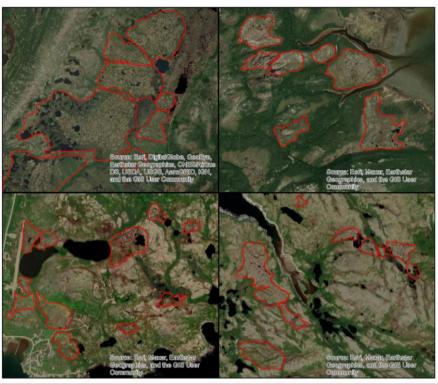


Figure 2. Examples of wetland complexes of interest (WOIs) in Labrador that were identified by the mapping team using high-resolution satellite imagery available via Esri ArcGIS Online. Examples of WOI boundaries are shown in red and were determined based on differences in drainage or vegetation from adjacent WOIs or based on separation following linear infrastructure, such as roads, airstrips, or trails. Identification was restricted to WOIs that contained prospective peatland permafrost landforms.

3.2.2 Quality control of WOI database

214

216 217 218

219

220

221

222 223 The WOI inventory was subjected to a quality control check, during which each complexes wasere reviewed and duplicates or points clearly not corresponding to wetlands were removed. In some cases, non-wetland locations may have been retained because of difficulties discerning peat plateauxplateaus from surface peat overpeat covered bedrock or flat coastal tundra

Formatted: Font: 9 pt, Bold, English (United States)

Formatted: Space After: 10 pt, Line spacing: single

Formatted: Font: 9 pt, Bold

Formatted: Font: 9 pt, Bold, English (United States)

Formatted: Font: 9 pt, Bold

3.2.3 Consensus-based review of WOI database

The quality-controlled WOI inventory was sent back to the mappers for a consensus-based review, following a similar method—used—byto—Way et al.'s (2021a) approach—for rock glacier inventorying in northern Labrador. Each WOI was independently reviewed by two team members, both of whom had access to the mapper's initial confidence rating, and one of whom had access to a field-validated dataset of WOIs (see Sect. 3.3 Validation of subset of WOI database). Both team members were asked to indicate whether each WOI contained peatland permafrost landforms. WOIs that were evaluated by both reviewers as containing peatland permafrost were considered likely to contain palsas or peat plateauxplateaus, while WOIs that were evaluated by both reviewers as not containing peatland permafrost were considered unlikely to contain palsas or peat plateauxplateaus. WOIs with conflicting classifications were considered to possibly contain palsas or peat plateauxplateaus. This consensus-based review process resulted in a full inventory of WOIs that were classified as likely, possibly, or unlikely to contain peatland permafrost landforms.

3.3 Validation of subset of WOI database

The full, consensus-based inventory results were compared with a field-<u>and imagery</u>-validated dataset of 285571 WOIs, with and without contemporary peatland permafrost landforms. From July to September 2021 <u>and 2022</u>, field evaluations of WOIs were undertaken via in-person field visits, remotely piloted aircraft (RPA) image acquisitions (DJI Mini 2 <u>microdrone</u>, weighing less than 250 g), video clip acquisition from a helicopter survey, and image acquisitions from commercial Twin Otter aircraft flights. Interpretation of the presence or absence of permafrost landforms within each WOI that was visited or aerially surveyed was <u>also determined through consensus</u> <u>determined following a consensus based approach</u> between two mappers. Any <u>WOIs with</u> disagreements in interpretation <u>were re-evaluated and discussed</u> <u>were discussed on a wetland complex by complex basis until consensus could be reached between the two mappers.</u>

Field visits to peatland complexes WOIs were undertaken at road-accessible locations within 500 m of the Trans-Labrador Highway and other accessible side roads via truck or ATV and at coastal locations via speedboat from the nearby communities of Black Tickle, Cartwright, Rigolet, and Nain. The number of WOIs that could be visited for field validation was restricted by weather conditions, tides, the availability of local guides and boat drivers with location-specific expertise, and other logistical and operational constraints. During field visits, team members probed the soil to the depth of refusal (maximum of 125 cm). The nature of refusal, interpreted as frozen ground, compact sediment, clasts, rock, or not applicable (N/A; >125 cm), was noted and used to assess permafrost presence or absence. Where the cause of probe refusal was unclear, instantaneous ground temperature measurements were collected using vertically arranged thermistors connected to an Onset Hobo UX120-006M 4-Channel Analog Data Logger (accuracy ±0.15 °C) (Davis et al., 2020; Holloway and Lewkowicz, 2020; Way et al., 2021b; Way and Lewkowicz, 2015). Ground temperatures were recorded at the base of within the probed hole for a minimum of 10 minutes to allow for thermal equilibration. Frost probing and instantaneous ground temperature

measurements <u>were targeted towards locations considered aimed to sample locations that were</u>-most likely to contain frozen ground and thus mostly occurred on elevated peat-covered microtopography within each <u>WOleomplex</u>.

Low-altitude RPA imagery of prospective peatland permafrost complexes was-were collected using a DJI Mini 2 minicrodrone when weather conditions were suitable (i.e., no rain, no fog, low wind). speedboat.

Low-altitude georeferenced video footage was collected using a GoPro Hero9 camera that was mounted onto a helicopter during a fuel cache eleaning mission throughout in northern Labrador in July and August 2021, led by the Torngat Wildlife, Plants, and Fisheries Secretariat. The camera was set to record real-time video (1080 p, 60 fps, wide) at an oblique angle (~45°). The flight altitude was between 90 m and 120 m a.g.l., similar to coastal Nunavik transects performed by Boisson and Allard (2018), and the flight plan between the Goose Bay Airport and the Torngat Mountains National Park was designed to fly over WOIs in coastal locations north of the community of Hopedale-Makkovik (55.50° N) (Supplement Sect. S2).

Low-altitude georeferenced aerial images were also collected using handheld digital cameras (Nikon Coolpix W300 or Olympus Tough TG-6) during commercial Air Borealis Twin Otter flight segments between Cartwright and Black Tickle and between Goose Bay, Rigolet, Makkovik, Postville, Hopedale, Natuashish, and Nain. The Twin Otter flights only crossed over WOIs along existing commercial flight routes.

271 3.4 Compilation of final WOI database

The final WOI database of likely, possible, or unlikely peatland permafrost complexes was developed following the incorporation of the field-validated dataset. WOIs that were classified as likely or possibly to contain peatland permafrost were subject to a final round of review in which the peatland permafrost landforms were identified as palsas, peat plateaus, or both palsas and peat plateaus (mixed).

276 3.4-5 Statistical analyses of final WOI database

ANOVA (analysis of variance) and *post hoc* Tukey's HSD (honest significant difference) tests were performed to determine whether the mean latitude, distance from coastline, elevation, MAAT, TDD, and FDD were statistically significantly different between the final classes of likely, possibly, and unlikely peatland permafrost complexes. Statistical analyses were performed in R 4.0.3 (R Core Team, 2020).

4 Results

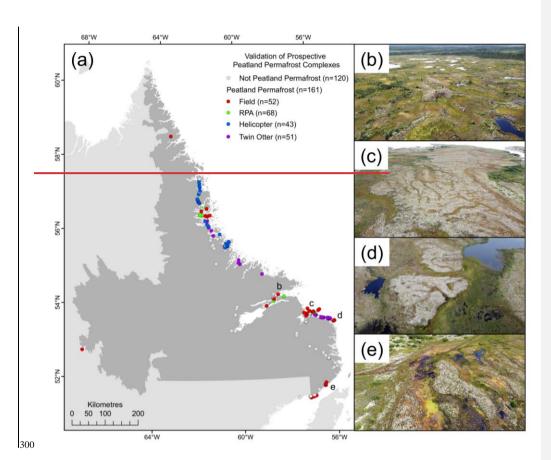
4.1 Peatland permafrost complex identification and review

A total of 1885-2092 unique WOIs, limited to the largest and most visually apparent prospective peatland permafrost complexes within the study area, were included in the full inventory. In the consensus building review process, rReviewer

agreement was very high (8889 %) during the consensus-building review process, with 1016-1116 complexes classified by both reviewers as likely containing peatland permafrost and 643-750 complexes classified by both reviewers as not likely containing unlikely to contain peatland permafrost, and only 226 complexes with conflicting classifications of permafrost presence or absence (1211 %) (Supplement Sect. S23).

4.2 Validation of peatland permafrost complexes

In Summer 2021 and 2022, in-person field visits (n=60-63 WOIs), RPA visits (n=97-141 WOIs), helicopter video clips (n=69 WOIs), and Twin Otter images (n=97-314 WOIs) were combined to evaluate peatland permafrost presence at 271 531 WOIs, 47-49 of which were cross-validated using multiple methods (Figure 23; Supplement Sect. S24). Previous work from 20137 to 2020, including field visits (n=7-23 WOIs) and RPA image collection (n=10-19 WOIs), were also used to validate palsa or peat plateau presence at an additional 40-19 complexes and peatland permafrost absence at an additional seven complexes (Anderson et al., 2018; Way, 2017). Out of the 281-557 WOIs evaluated via field and/or imagery validation methods, 161311 were interpreted as containing to contain peatland permafrost landforms. Comparison between the validation dataset and the consensus-based inventory resulted in re-classification of 34-39 of the 226 possible peatland permafrost complexes (157 %) to either likely (n=37) or unlikely (n=2736) peatland permafrost complexes.



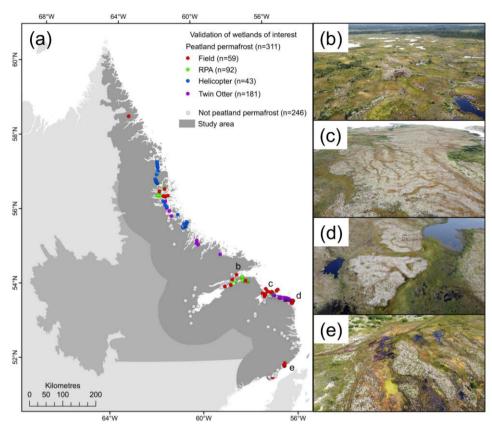
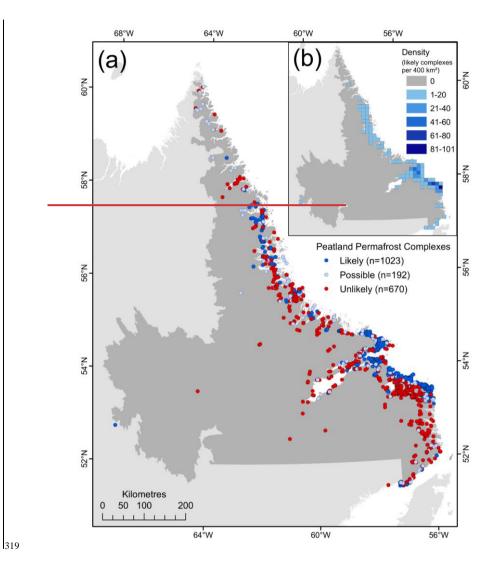


Figure 23. (a) Locations of validated peatland permafrost complexes in coastal Labrador from field-based activities and imagery acquisition using RPA, helicopter, and Twin Otter from 2013-2017 to 2021-2022; Example of peatland complexes containing palsas and/or peat plateauxplateaus near (b) Rigolet, (c) Cartwright, (d) Black Tickle-Domino, and (e) Red Bay.

4.3 Peatland permafrost complex inventory

 A total of 1023-1119 out of 1885-2092 WOIs were classified as likely containing peatland permafrost landforms, with an additional 192-187 wetland complexes classified as possibly containing peatland permafrost landforms (Figure 43). The largest clusters of likely and possible peatland permafrost complexes were located along the coastline between Makkovik (55.0° N) and Black Tickle (53.5° N) (Figure 43; Figure 4A5A). Of the 1023-The likely peatland permafrost complexes, 1022 were at low elevations (mean elevation of 30-29 m a.s.l.) (Figure 54C) within 41-22 km of the coastline (mean distance from

coastline of 4.22.6 km) (Figure 34; Figure 4B5B), and one was located in alpine tundra, near the Labrador Quebec interior border (Brown, 1979). The 1022 coastal complexes that were 1Likely to contain peatland permafrost complexes were distributed along the coastline from 51.4° N near Blanc-Sablon to 58.6° N in the Torngat Mountains National Park (Figure 34; Figure 4A5A), with most complexes located in southeastern Labrador (mean latitude of 54.1° N) (Supplement Sect. S34). Comparison against gridded climate products showed that the MAAT of at peatland permafrost complexes ranged from -7.5 °C to +1.2 °C, with corresponding ranges for FDDs of 11266 degree days to 3466-3471 degree days and TDDs of 736-733 degree days to 1704-1704 degree days (Figure 4D5D-F). Despite the wide range in MAAT, the majority of the coastal likely peatland permafrost complexes (9087%) were found in locations with MAATs between -2 °C and ±1 °C (Figure 4D5D).



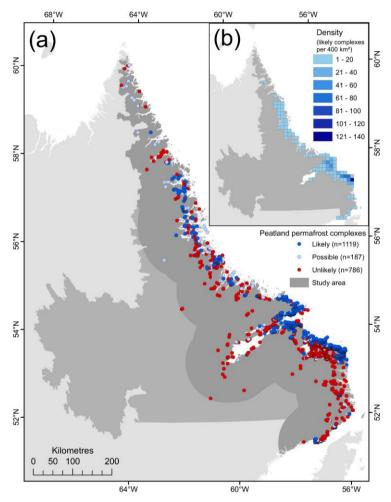
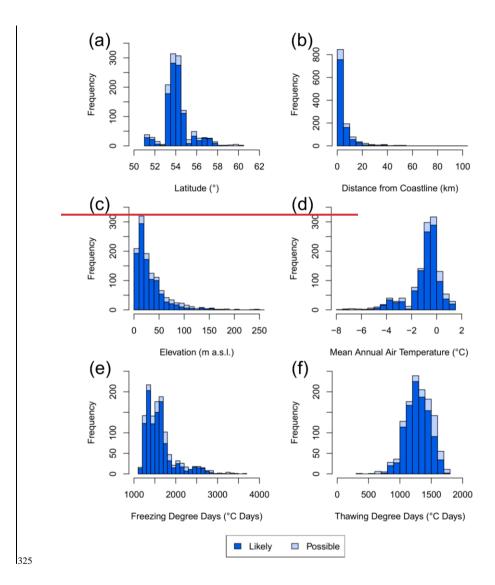


Figure 43. (a) Spatial distribution of inventoried peatland complexes (n=18852092) classified as likely containing peatland permafrost landforms (n=192187), possibly containing peatland permafrost landforms (n=192187), and unlikely to contain peatland permafrost landforms (n=670786); (b) Inset map showing density of Number of wetlands of interest that are likely to contain peatland permafrost landforms complexes within 20 by 20 km (400 km²) grid cells.



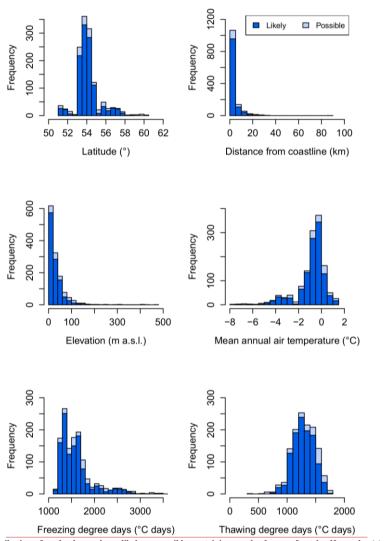


Figure 54. Distribution of wetland complexes likely or possibly containing peatland permafrost landforms by (a) latitude; (b) distance from the coastline; (c) elevation; and (d) mean annual air temperature; (e) mean annual freezing degree days; and (f) mean annual thawing degree days for the 1981 to 2010 climate normal.

ANOVA and *post hoc* Tukey's HSD tests revealed that the mean distance from coastline, elevation, MAAT, FDD, and TDD were statistically different between the likely, possibly, and unlikely peatland permafrost complexes at the 95 % confidence level. When compared with the complexes that likely contained peatland permafrost, the 192-187 complexes that possibly contained peatland permafrost were similarly distributed all along the coastline but were skewed further north (mean latitude of 54.5° N) and extended as far as 60.2° N (Supplement Sect. S34). These less certain features were at greater distances from the coastline (mean distance from coast of 11.77.8 km) and at higher elevations (mean elevation of 68-66 m a.s.l.). The 670-786 complexes that were unlikely to contain peatland permafrost were well distributed between 51.4° N and 60.2° N (Figure 43) but were located further from the coastline (mean distance from coastline of 18.910.7 km), at higher elevations (mean elevation of 79-78 m a.s.l.), and at warmer-higher MAATs (mean MAAT of -0.6-5 °C) than the complexes that likely or possibly contained peatland permafrost (Supplement Sect. S34).

Likely and possible peatland permafrost complexes were also classified according to the type of peatland permafrost landforms found within the wetland complex. Complexes that were exclusively comprised of palsas accounted for half of the likely and possible peatland permafrost complexes (50 %) and were distributed along the entire study area. Complexes with exclusively peat plateaus were less common (29 %) and were spatially concentrated between ~53° N and ~55° N. The remaining 21 % of the likely and possible peatland permafrost complexes were interpreted to contain a combination of palsas and peat plateaus, but it is possible that many of these complexes contain dissected and heavily degraded peat plateaus that now resemble palsas. Further field-based investigations would be required to differentiate these degradational landforms.

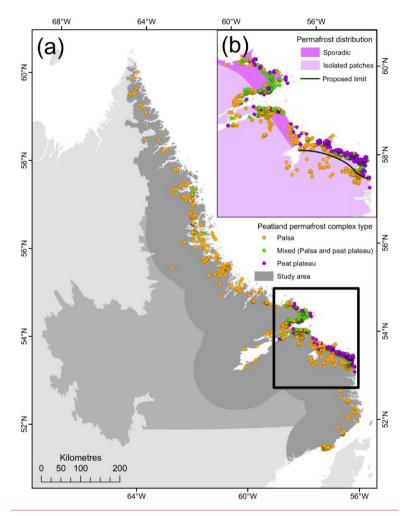


Figure 6. (a) Spatial distribution of likely and possible peatland permafrost complexes classified by peatland permafrost landform type as palsas, peat plateaus, or a mix of both palsas and peat plateaus for coastal Labrador. (b) Inset map showing existing permafrost distribution zones (Heginbottom et al., 1995) for a subsection of coastal Labrador and the location of a new proposed location for the southern limit of the sporadic discontinuous permafrost zone.

Formatted: Centered

5 Discussion

354

5.1 Distribution of peatland permafrost in Labrador

356 In this study, we demonstrated that pPeatland permafrost complexes in Labrador and adjacent portions of 357 Quebec are concentrated abundant in lowlands within 40-22 km of the Labrador Sea or Gulf of St. Lawrence coastline (Figure 5B). A geographic gradient is especially apparent between Rigolet (54.2° N) and Black Tickle (53.5° N), where peatland 358 359 permafrost complexespeat plateaus are abundant along the coast but generally absent from wetlands farther inland (Figure 4). The higher density of peatland permafrost complexes along the coast could be linked to climatic factors like persistent fog and 360 361 cloud cover leading to less incoming solar radiation (Way et al., 2018) or thinner and denser snowpacks (Seppälä, 1994; Vallée 362 and Payette, 2007) in the wind-exposed barrens along the coast (Way et al., 2018). Further work should focus on exploring the 363 role of local climate conditions in the formation and persistence of peatland permafrost in coastal Labrador and similar northern 364 coastal locations. In Labrador, pPeatland permafrost was found across a large range of MAATs, spanning identified for a 365 range of MAATs from -7.5 °C to +1.2 °C. Permafrost persistence at MAATs above +1 °C in southeastern Labrador was previously noted in a field study at five palsa complexes (Way et al., 2018). Peatland permafrost complexes in Labrador were 366 367 located at higher MAATs than is predicted for other northern coastal regions like northern This compares favourably with a 368 previous study from the region, where palsas and peat plateaux were identified in five locations in southeastern Labrador with 369 corresponding MAATs of up to +1.0 °C (Way et al., 2018). Finland, Norway, and Sweden (approximately (Fewster et al., 370 2020)+0.4 °C) (Parviainen and Luoto, 2007). By contrast, the distribution of MAATs in peatland permafrost locations in 371 Labrador is warmer than the upper MAAT thresholds of +0.4 °C and +0.2 °C that have been modelled for northern Finland, 372 Norway, and Sweden (Parviainen and Luoto, 2007) and for the rest of Canada and Alaska (Fewster et al., 2020), respectively. 373 Our results also suggest that the upper-MAAT threshold of +0.2 °C for peatland permafrost areas previously applied to that is 374 estimated for North America (Fewster et al., 2020) is too low for Labrador and adjacent parts of Quebec where peatland 375 permafrost landforms continue to persist, due to their relict and resilient nature (Dionne, 1984; Way et al., 2018) and predicted 376 resilience (Way et al., 2018) of many peatland permafrost landforms. Large thermal offsets (up to and often exceeding 2.0 °C 377 in southeastern Labrador) (Way and Lewkowicz, 2018) are typical of organic-rich landscapes like peatlands and may promote 378 continued landform persistence despite The large thermal offset that is typical of peatland permafrost (Burn and Smith, 1988; 379 Williams and Smith, 1989)(Way and Lewkowicz, 2018) is expected to promote continued landform persistence under the 380 context of a warming climate (Jorgenson et al., 2010). This may further exacerbate discrepancies between peatland permafrost 381 observations and regional estimates, calling into question the utility of simplified threshold-based approaches when modelling 382 with future climate scenarios, leading to exceedance of existing MAAT based thresholds. Information on the timing of 383 peatland initiation following deglaciation (Gorham et al., 2007), rates of peat deposition (Tarnocai, 2009; Gorham, 1991), and 384 corresponding peat thicknesses should also be considered in studies of peatland permafrost distribution, as thicker peat deposits 385 may influence permafrost development and protect permafrost persistence through a larger thermal offset (Smith and Riseborough, 2002). 386

The regional distribution of fine-grained sediments and local depositional history are expected to play an important role in landscape suitability for peatland permafrost landforms (O'Neill et al., 2019; Seppälä, 1986; Zoltai, 1972). For example, differences in the distribution of palsa versus peat plateau landforms have previously been attributed to varying thicknesses and extents of the underlying sediment, with thicker sediment deposits leading to the development of palsas and thinner sediment deposits linked to the development of peat plateaus (Allard and Rousseau, 1999). Differences in sediment grain size may also influence the thickness of the ice lenses and the depth at which they form, with thicker ice lenses developing deeper in finer sediments and thinner ice lenses forming at shallower depths in coarser sediments (Allard and Rousseau, 1999). Further examination of how these variables could influence peatland permafrost formation and persistence in coastal Labrador is challenged by the paucity of information on surficial materials and marine limits along most of the Labrador Sea coastline (Hagedorn, 2022; Occhietti et al., 2011). To date, local marine limits have been identified at some individual locations and study sites in coastal Labrador (e.g., Bell et al., 2011; Dyke et al., 2005; Occhietti et al., 2011; Vacchi et al., 2018), but widespread mapping of marine sediments has only been completed for a small section of northern coastal Labrador from Goose Bay to Hopedale (Hagedorn, 2022). Based on the information that is currently available, we can qualitatively link the distribution of the largest clusters of peatland permafrost complexes, particularly peat plateau complexes, to locations where post-glacial marine invasions had occurred, such as along the lowland-dominated coastline between Makkovik (55.0° N) and Black Tickle (53.5° N), where frost-susceptible, glaciomarine surficial materials are generally widespread (Fulton, 1989, 1995; Hagedorn, 2022; Occhietti et al., 2011). Meanwhile, fewer peatland permafrost complexes were mapped between Makkovik (55.0° N) and Hopedale (55.5° N), where the elevated topography resulted in limited marine invasions and post-glacial marine deposition along the coast. Significant and coordinated advances in surficial mapping will be required before similar links between peatland permafrost distribution and surficial material type, sediment grain size, and elevation relative to the marine limit can be made for other parts of the coastline. As peatland permafrost landforms form from the epigenetic development of segregated ice, it would be expected that the regional distribution of fine grained sediments and local depositional history would also play an important role in the landscape suitability of these complexes (O'Neill et al., 2019; Zoltai, 1972; Seppälä, 1986).(Allard and Rousseau, 1999) (Allard and Rousseau, 1999)We identified the largest clusters of peatland permafrost complexes in locations where post glacial marine invasions had occurred, such as along the lowland dominated coastline between Makkovik (55.0° N) and Black Tickle (53.5° N), where frost susceptible, glaciomarine surficial materials are widespread (Fulton, 1989, 1995; Hagedorn, 2022; Occhietti et al., 2011). Fewer peatland permafrost complexes were mapped between Makkovik (55.0° N) and Hopedale (55.5° N), where the elevated topography resulted in limited marine invasions and post-glacial marine deposition along the coast (Hagedorn, 2022; Occhietti et al., 2011).

387 388

389

390

391

392

393

394

395

396

397

398

399

400

401

402

403

404

405

406

407

408

409

410

411

412

413

414

415

416

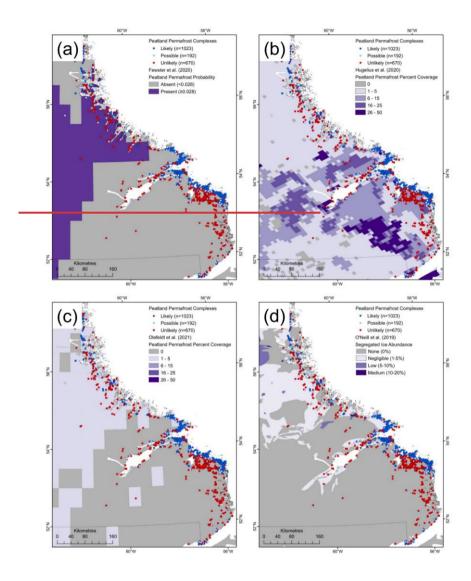
417

418

5.2 Implications for peatland permafrost and permafrost distribution in northeastern Canada

Comparisons between our inventory results and several recent national to global wetland, peatland (Supplement Sect. S1: Supplement Sect. S45), and peatland permafrost distribution products (e.g., Fewster et al., 2020; Hugelius et al., 2020;

420 Olefeldt et al., 2021) (Figure 57) provide compelling evidence that peatland permafrost along the Labrador coast has been is 421 poorly represented by existing datasets. While differences in scale may explain some of this discrepancy, the general pattern 422 presented in most previous datasets, showing relatively greater peatland permafrost in the continental interior and less along 423 the coast, is directly contradicted by the results of this study. Although these products can have inconsistencies with one another 424 (Figure 5A C), they do consistently model more abundant peatland permafrost in the continental interior and little to no 425 peatland permafrost along much of the Labrador Sea and Gulf of St. Lawrence coastline. Model predictions showing more 426 peatland permafrost in the interior compared to along the coast This reversed pattern could reflect inaccurate assumptions on 427 the climate limits of peatland permafrost and/or may reflect the absence may be due to a lack of field data from other many 428 northern coastal peatland permafrost environments (Borge et al., 2017). Inclusion of physiographic variables, like soil 429 conditions, and frost-susceptibility of sediments, and more detailed surficial deposit maps are glacial depositional information 430 is-likely necessary for an improved representation of peatland permafrost in northern coastal regions. Recent work by O'Neill 431 et al. (2019), for example, has demonstrated that segregated ice can be more reliably modelled along sections of the Labrador 432 Sea coastline (Figure 75D) by incorporating paleogeographic variables like vegetation cover, surficial geology, and glacial 433 lake and marine limits.



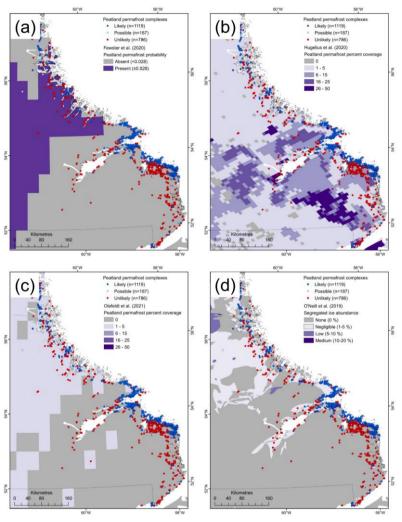


Figure 57. Comparison of inventoried peatland permafrost complexes with peatland permafrost presence and percent coverage as modelled by (a) Fewster et al. (2020); (b) Hugelius et al. (2020); and (c) Olefeldt et al. (2021) and with segregated ice content as modelled by (d) O'Neill et al. (2019).

The results of our inventory also suggest that some amendments to existing representations of permafrost distribution maps may be required for coastal Labrador. For example, the highest density of peatland permafrost complexes (Figure 4B) along the Labrador Sea and Gulf of St. Lawrence coastline was found near the community of Black Tickle (53.5° N) on the Island of Ponds (2 palsa complexes, 19 mixed palsa and peat plateau complexes, and 59 peat plateau complexes within 94 km²) (Figure 6) (Figure 3B), which is currently classified in the isolated patches of permafrost zone on the Permafrost Map of Canada (Heginbottom et al., 1995) and the no permafrost zone on the 2000-2016 Northern Hemisphere Permafrost Map (Obu et al., 2019) (Supplement Sect. S56). The identification of large swaths of likely peatland permafrost complexes, including more than 150 peat plateaus, between Cartwright (53.7° N) and Black Tickle (53.5° N) suggest that the physiography-based Permafrost Map of Canada's limit for the sporadic discontinuous zone along the Labrador coast (Heginbottom et al., 1997) (Supplement Sect. S5), could reasonably be extended south by ~110 km from its current position of (~53.7° N-to ~53.1° N) (Figure 6B). This southerly extension of the sporadic discontinuous permafrost zone-limit has previously been suggested by Allard and Seguin (1987) and Payette (2001) who indicated that , based on regional vegetation and geomorphology favoured permafrost along much of this coastline (Payette, 1983). Unexpectedly, large cClusters of likely peatland permafrost complexes were also identified near the communities of Red Bay (Supplement Sect. S67) and Blanc-Sablon, both of which are considered to be underlain by little to no permafrost (Heginbottom et al., 1995; Obu et al., 2019) which are respectively classified in the isolated patches and no permafrost zones on the Permafrost Map of Canada (Heginbottom et al., 1995) and in the no permafrost zone on the 2000-2016 Northern Hemisphere Permafrost Map (Obu et al., 2019) (Supplement Sect. S65). A 15 kmn extension of the southern limit of the Permafrost Map of Canada's isolated patches permafrost zone to include the Blanc-Sablon region would better reflect contemporary permafrost conditions in this area, especially given that of permafrost zone by 15 km to include Blanc Sablon may better reflect regional permafrost conditions, as permafrost has been previously detected in both

440 441

442

443

444

445

446 447

448 449

450

451

452

453

454

455

456

457

458

459

460

461

462 463

464

465

466

467

468

469

470

471

5.3 Challenges and limitations of a point-based inventory of peatland permafrost complexes in coastal Labrador point-based inventorying of peatland permafrost complexes in coastal Labrador

mineral soils in the community and in surrounding peatlands below the marine limit (Dionne, 1984).

The most challenging aspects of the inventorying process involved interpreting peatland permafrost presence in isolated WOIs containing small landforms, while in the case of more obvious peatland permafrost features, there were at times difficulties in determining distinct wetland boundaries (Figure 2) (Tarnocai et al., 2011)Differences in peatland permafrost landform size, shape, and vegetation coverage across a large, heterogeneous study area like Labrador (Beer et al., 2021) can lead to difficulties in feature mapping and identification, especially when performed by a single mapper. However, we believe that these issues were mitigated through t The inclusion of multiple mappers, which in the inventorying process facilitated the development of a large initial database that and reduced the potential for omission of prospective WOIs. The consensus-based review process that followed was designed to minimize the inclusion of false positives in the final dataset of 1023-1119 likely peatland permafrost complexes, but we recognize that this conservative approach may have resulted in the exclusion of some

Formatted: French (Canada)

Field Code Changed

Field Code Changed

Formatted: French (Canada)

Field Code Changed

Field Code Changed

complexes. At the northern end of the study area, where other types of periglacial landforms become more common, misclassification of palsas for other elevated periglacial landforms may have contributed to the designation of a higher number of possible peatland permafrost complexes. It is certainly possible that some segregated ice mounds with less than 40 cm of overlying peat (i.e., lithalsas) may have been included in the inventory, particularly near the northern end of the study area where wetlands are less abundant and peat deposits may be thinner (Supplement Sect. S1). This suggests that the definition of peatlands, as wetlands containing at least 40 cm of surface peat (Tarnocai et al., 2011), and its application to palsas and lithalsas, can introduce some ambiguity during inventorying. The most challenging aspect of the inventorying process was interpreting WOIs containing small landforms in relatively isolated wetlands, while in the case of more obvious features, there were at times difficulties in determining distinct wetland boundaries. At the northern end of the study area, where other types of periglacial landforms become more common, misclassification of palsas and peat plateaux for other elevated landforms (i.e., lithalsas) may have contributed to the designation of a higher number of possible peatland permafrost complexes in certain subregions.

While other inventorying approaches, including grid-based approaches methods (Ramsdale et al., 2017; Gibson et al., 2020, 2021; Borge et al., 2017), were considered, a point-based inventory was ultimately developed for this study. The implementation of a grid-based approach with delineation of individual landforms for each WOI could have been useful for estimating ground ice content, thermokarst potential, carbon content, and overall permafrost coverage, but the purpose of this study was to generate an initial inventory to guide future research that will facilitate quantitative assessments of peatland permafrost distribution and coverage in these regions. Our field experience in the region suggests that areal delineations of peatland permafrost complexes in coastal Labrador will require extensive validation, and it is unlikely that even experienced permafrost mappers could accurately map the extents of permafrost throughout some complexes without extensive field investigations.

have been applied in peatland permafrost mapping studies in the Northwest Territories (Gibson et al., 2020, 2021) and parts

of Norway (Borge et al., 2017), Despite the above limitations, the point based nature of our studyour inventory allowed for the incorporation of dedicated, co-located field- and imagery-based validation information. Post-validation adjustments to the inventory, including Despite the high agreement during the review stage, reclassification of 34.39 WOIs following integration of field—and imagery-based validation—highlights the importance of ground-truthing in remote sensing—or modelling-based periglacial landform inventories.

Owing to a lack of prior field-based assessments of permafrost conditions in Labrador, it was also difficult to independently validate our peatland permafrost inventory results. However, a detailed aerial photograph- and field-based Ecological Land Classification (ELC) survey undertaken in the late 1970s did cover a subset of our study area in southeastern Labrador (Environment Canada, 1999). The ELC identified a total wetland area of 666 km² that contained which was at least partly covered by inventoried peatland permafrost landforms (Figure 86). Comparison with the present study showed that mMappers successfully-identified likely-peatland permafrost complexes in 2+23 of the 24 contiguous ELC wetland areas indicated out of the 24 ELC wetland polygons that were identified as containing palsas. Re-eExamination of the three-one

remaining polygons ELC peatland permafrost-containing wetland area revealed the presence of wetland complexes with irregular ponding patterns indicative of thermokarst development and elevated landforms that could be peatland permafrost but, due to their small size, would require in situ field visits for validation. Some of the inventoried likely peatland permafrost complexes that were not captured as part of the peatland permafrost areas from the ELC were instead classified in other wetlands, like string bogs, and in raised marine terrain units. Overall, the results of our inventory are in good agreement with the limited previous overlapping field investigations and inventorying efforts from the ELC.

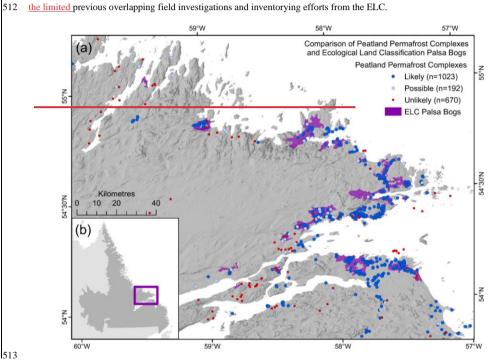
507

508

509

510

511



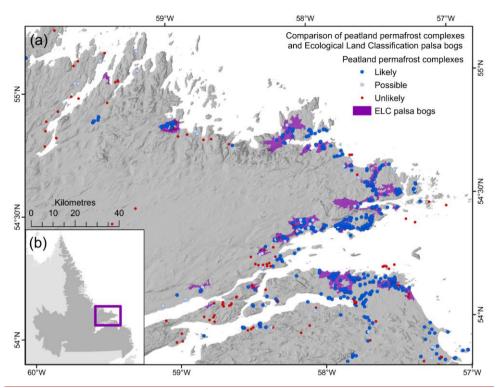


Figure 68. (a) Comparison of inventoried peatland permafrost complexes with palsa bog regions identified in the Ecological Land Classification (ELC) survey (Environment Canada, 1999); (b) Inset map showing the extent of the peatland permafrost area that was mapped in the Location of ELC palsa bogs in relation to Labrador.

6 Conclusions

This study provides the first detailed <u>point</u> inventory of peatland permafrost landforms along the <u>understudied</u> Labrador Sea <u>and Gulf of St. Lawrence</u> coastline. Using high-resolution satellite imagery and extensive field- and imagery-based validation efforts, we applied a multi-stage, consensus-based inventorying approach to identify <u>a total of 1885 wetlands</u> of interest, 1023 of which were classified as 1119 likely to contain peatland permafrost landformscomplexes. Likely <u>pPeatland</u> permafrost complexes were primarily found in <u>lowlands on outer coasts</u>, <u>coastal</u>, <u>lowland locations</u>-spanning from 51.4° N to 58.6° N, with the largest clusters <u>of complexes</u> occurring <u>just~110 km</u> south of the <u>previously mapped</u> limit of sporadic discontinuous permafrost in northeastern Canada (Heginbottom et al., 1995).

Comparisons between our point inventory results and existing wetland, peatland, and peatland permafrost distribution products reveal major discrepancies between this study and in prior estimates of peatland permafrost in Labrador with implications for ground ice content (O'Neill et al., 2019), thermokarst potential (Olefeldt et al., 2016), and carbon content (Hugelius et al., 2014). Significant advances in the development of relevant datasets on surficial materials, marine limits, peatland distribution, and peat ages and thicknesses, along with field-based advances in climate monitoring for cloud cover, fog, and snow, are critically needed to better characterize northern coastal regions like Labrador. Our results highlight the importance of field-based validation for periglacial landform mapping and modelling, particularly when mapping small, dynamic features like palsas and peat plateaux, and of considering physiography and geomorphology in for accurate representations of peatland permafrost in larger scale spatial products. This study provides an important baseline for future peatland permafrost mapping and modelling efforts along the Labrador Sea coastline and will support local to regional infrastructure and climate change adaptation strategy development. The significant underestimation of peatland permafrost along the Labrador Sea and Gulf of St. Lawrence coastline identified shown in this study should inform future regional to global permafrost, peatland permafrost, and carbon content mapping efforts, infrastructure and climate change adaptation strategy development, and wildlife management considerations for Labrador and other northern coastal locations regions.

Data availability. Likely and possible peatland permafrost locations from the coastal Labrador peatland permafrost complex inventory are freely available for download from Nordicana D (Wang et al., 2022).

Author contribution. YW and RW designed the study and drafted the manuscript. YW led the raw data collection and the data analysis. RW contributed to raw data collection and data analysis and was the PI for the NSERC Discovery Grant supporting peatland permafrost research activities in Labrador. JB contributed to raw data collection and data analysis. AF and RT contributed to raw data collection. MP coordinated the collection of helicopter video footage. JB, AF, RT, and MP reviewed and contributed edits to the manuscript.

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgements. The authors would like to acknowledge the Nunatsiavut Government (Rodd Laing), the Nunatsiavut Research Centre (Carla Pamak, Michelle Saunders), the NunatuKavut Community Council (Bryn Wood, George Russell Jr., Charlene Kippenhuck), and the Innu Nation (Jonathan Feldgajer, Jack Penashue) for their support guidance ofn research conducted on traditional Inuit and Innu lands. We thank Caitlin Lapalme, George Way, and Amy Norman of Goose Bay-i Freeman Butt and Tanya Barney of Red Bay-: Jeffrey and Wendy Keefe of Black Tickle-: Barbara Mesher of Cartwright Jane and Jack Shiwak, -and-Tyler and Harvey Palliser, and Sandi and Karl Michelin of Rigolet, and Caroline Nochasak, Liz Pijogge, Carla Pamak, Kayla Wyatt, Frédéric Dwyer-Samuel, and Patricia Johnson-Castle of Nain for their logistical and in-kind support. We acknowledge Caitlin Lapalme, Victoria Colyn, Kayla Wyatt, Frédéric Dwyer-Samuel, and Adrian Earle,

- 560 Patrick Lauriault, Michaela Smitas-Kraas, and Tara Ryan for their valuable field assistance, and we are grateful to Derrick
- 561 Pottle, Tyler Palliser, Martin Shiwak, Reginald Maggo, Jeremy Ivany, Martin Andersen, Eldred Andersen, Jeffrey Keefe, and
- 562 Gary Bird, Pat Davis, and Anthony Elson for their boat operation and guiding services. Funding for this research was provided
- 563 by the Natural Sciences and Engineering Research Council of Canada, the Northern Scientific Training Program, and Queen's
- 564 University, We thank Dr. Antoni Lewkowicz and Dr. Luise Hermanutz for many insightful conversations about permafrost
- 565 and peatlands in Labrador. The manuscript has also benefitted from helpful comments and suggestions from Michelle
- 566 Saunders, Dr. H. Brendan O'Neill, and Dr. Hanna Lee, and we thank two anonymous referees and Dr. Steve Kokelj for their
- 567 <u>constructive reviews.</u>

568 References

- 569 Allard, M. and Rousseau, L.: The international structure of a palsa and a peat plateau in the Rivière Boniface region, Québec:
- 570 Inferences on the formation of ice segregation mounds, Géographie Phys. Quat., 53, 373–387,
- 571 https://doi.org/10.7202/004760ar, 1999.
- 572 Allard, M. and Seguin, M. K.: Le pergélisol au Québec nordique : bilan et perspectives, Géographie Phys. Quat., 41, 141-152,
- 573 https://doi.org/10.7202/032671ar, 1987.
- 574 Anderson, D., Ford, J. D., and Way, R. G.: The impacts of climate and social changes on cloudberry (bakeapple) picking: A
- 575 case study from southeastern Labrador, Hum. Ecol., 46, 849–863, https://doi.org/10.1007/s10745-018-0038-3, 2018.
- 576 Andrews, J. T.: The glacial geomorphology of the northern Nain-Okak section of Labrador, M.Sc. Thesis, McGill University,
- 577 Montreal, Canada, 301 pp., 1961.
- 578 Banfield, C. E. and Jacobs, J. D.: Regional patterns of temperature and precipitation for Newfoundland and Labrador during
- 579 the past century, Can. Geogr. Géographe Can., 42, 354–64, 1998.
- 580 Barrette, C., Brown, R., Way, R. G., Mailhot, A., Diaconescu, E. P., Grenier, P., Chaumount, D., Dumont, D., Sévigny, C.,
- Howell, S., and Senneville, S.: Nunavik and Nunatsiavut regional climate information update, in: Nunavik and Nunatsiavut:
- 582 From science to policy, an integrated regional impact study (IRIS) of climate change and modernization, second iteration,
- edited by: Ropars, P., Allard, M., and Lemay, M., ArcticNet Inc., Quebec City, Canada, 62, 2020.
- 584 Beilman, D. W., Vitt, D. H., and Halsey, L. A.: Localized permafrost peatlands in western Canada: Definition, distributions,
- 585 and degradation, Arct. Antarct. Alp. Res., 33, 70–77, https://doi.org/10.1080/15230430.2001.12003406, 2001.
- 586 Bell, T., Putt, M., and Sheldon, T.: Landscape hazard assessment in Nain, Phase I: Inventory of surficial sediment types and
- infrastructure damage, Final Report to Nunatsiavut Government and Nain Inuit Community Government, 2011.
- 588 Boisson, A. and Allard, M.: Coastal classification of Nunavik and dynamics of the Arctic/Subarctic coastal environments,
- 589 Vulnerabilities of the Quebec's Arctic Territory in the Context of Climate Change, Kuujjuaq, Canada, 2018.
- 590 Borge, A. F., Westermann, S., Solheim, I., and Etzelmüller, B.: Strong degradation of palsas and peat plateaus in northern
- 591 Norway during the last 60 years, The Cryosphere, 11, 1–16, https://doi.org/10.5194/tc-11-1-2017, 2017.

Formatted: French (Canada)

- 592 Brown, R. and Lemay, M.: Chapter 2: Climate variability and change in the Canadian Eastern Subarctic IRIS region (Nunavik
- and Nunatsiavut), in: Nunavik and Nunatsiavut: From science to policy. An Integrated Regional Impact Study (IRIS) of climate
- change and modernization, ArcticNet Inc., https://doi.org/10.13140/2.1.1041.7284, 2012.
- 595 Brown, R. J. E.: Permafrost investigations in Quebec and Newfoundland (Labrador), National Research Council of Canada,
- 596 Division of Building Research, Technical Paper 449, https://doi.org/10.4224/20374659, 1975.
- 597 Brown, R. J. E.: Permafrost distribution in the southern part of the discontinuous zone in Québec and Labrador, Géographie
- 598 Phys. Quat., 33, 279–289, https://doi.org/10.7202/1000364ar, 1979.
- 599 Burn, C. R. and Smith, C. A. S.: Observations of the "thermal offset" in near-surface mean annual ground temperatures at
- 600 several sites near Mayo, Yukon Territory, Canada, ARCTIC, 41, 99-104, https://doi.org/10.14430/arctic1700, 1988.
- 601 Coultish, T. L. and Lewkowicz, A. G.: Palsa dynamics in a subarctic mountainous environment, Wolf Creek, Yukon Territory,
- 602 Canada, Proc. 8th Int. Conf. Permafr., 6, 2003.
- 603 Davis, E., Trant, A., Hermanutz, L., Way, R. G., Lewkowicz, A. G., Siegwart Collier, L., Cuerrier, A., and Whitaker, D.:
- 04 Plant-environment interactions in the low Arctic Torngat Mountains of Labrador, Ecosystems, 24, 1038-1058,
- 605 https://doi.org/10.1007/s10021-020-00577-6, 2020.
- 606 Dionne, J.-C.: Palses et limite méridionale du pergélisol dans l'hémisphère nord : Le cas de Blanc-Sablon, Québec, Géographie
- Phys. Quat., 38, 165–184, https://doi.org/10.7202/032550ar, 1984.
- 608 Dyke, A. S.: An outline of North American deglaciation with emphasis on central and northern Canada, in: Developments in
- 609 Quaternary Sciences, vol. 2, Elsevier, 373–424, https://doi.org/10.1016/S1571-0866(04)80209-4, 2004.
- 610 Dyke, A. S., Dredge, L. A., and Hodgson, D. A.: North American deglacial marine- and lake-limit surfaces, Géographie Phys.
- 611 Quat., 59, 155–185, https://doi.org/10.7202/014753ar, 2005.
- 612 Elias, S. A.: Paleoenvironmental interpretation of Holocene insect fossils from northeastern Labrador, Canada, Arct. Alp. Res.,
- 613 14, 311, https://doi.org/10.2307/1550794, 1982.
- 614 Environment and Climate Change Canada: Canadian Climate Normals, 2022.
- 615 Environment Canada: Audio Tape Transcript of the East-Central Labrador Ecological Land Survey,
- 616 https://ftp.maps.canada.ca/pub/nrcan_rncan/archive/vector/labrador/, 1999.
- 617 Esri: World Imagery, 2022.
- 618 Fewster, R. E., Morris, P. J., Swindles, G. T., Gregoire, L. J., Ivanovic, R. F., Valdes, P. J., and Mullan, D.: Drivers of Holocene
- 619 palsa distribution in North America, Quat. Sci. Rev., 240, 106337, https://doi.org/10.1016/j.quascirev.2020.106337, 2020.
- 620 Foster, D. R.: The history and pattern of fire in the boreal forest of southeastern Labrador, Can. J. Bot., 61, 2459-2471,
- 621 https://doi.org/10.1139/b83-269, 1983.
- 622 Foster, D. R. and Glaser, P. H.: The raised bogs of south-eastern Labrador, Canada: Classification, distribution, vegetation and
- 623 recent dynamics, J. Ecol., 74, 47, https://doi.org/10.2307/2260348, 1986.
- 624 Fulton, R. J. (Ed.): Quaternary Geology of Canada and Greenland, Minister of Supply and Services Canada, Ottawa, Canada,
- 625 846 pp., 1989.

Formatted: French (Canada)

- 626 Fulton, R. J.: Surficial Materials of Canada, "A" Series Map 1880A, Natural Resources Canada, Geological Survey of Canada,
- 627 https://doi.org/10.4095/205040, 1995.
- 628 Gibson, C., Morse, P. D., Kelly, J. M., Turetsky, M. R., Baltzer, J. L., Gingras-Hill, T., and Kokelj, S. V.: Thermokarst mapping
- 629 collective: Protocol for organic permafrost terrain and preliminary inventory from the Taiga Plains test area, Northwest
- 630 Territories, N. W. T. Open Rep., 2020, 29, 2020.
- 631 Gibson, C., Cottenie, K., Gingras-Hill, T., Kokelj, S. V., Baltzer, J. L., Chasmer, L., and Turetsky, M. R.: Mapping and
- 632 understanding the vulnerability of northern peatlands to permafrost thaw at scales relevant to community adaptation planning,
- 633 Environ. Res. Lett., 16, 055022, https://doi.org/10.1088/1748-9326/abe74b, 2021.
- 634 Gorham, E.: Northern peatlands: Role in the carbon cycle and probable responses to climatic warming, Ecol. Appl., 1, 182–
- 635 195, https://doi.org/10.2307/1941811, 1991.
- 636 Gorham, E., Lehman, C., Dyke, A., Janssens, J., and Dyke, L.: Temporal and spatial aspects of peatland initiation following
- 637 deglaciation in North America, Quat. Sci. Rev., 26, 300–311, https://doi.org/10.1016/j.quascirev.2006.08.008, 2007.
- 638 Hagedorn, G. W.: Preliminary delineation of marine sediments in east-central Labrador: Parts of NTS map areas 13F, -G, -I, -
- 639 J, -K, -N AND -O, Gov. Nfld. Labrador Dep. Ind. Energy Technol. Curr. Res., 22, 189-201, 2022.
- 640 Hare, F. K.: Climate and zonal divisions of the boreal forest formation in eastern Canada, Geogr. Rev., 40, 615,
- 641 https://doi.org/10.2307/211106, 1950.
- 642 Heginbottom, J. A., Dubreuil, M. A., and Harker, P. T.: Canada, Permafrost, National Atlas of Canada, Natural Resources
- 643 Canada, Geomatics Canada, https://doi.org/10.4095/294672, 1995.
- 644 Heginbottom, J. A., Brown, J., Melnikov, E. S., and Ferrians Jr., O. J.: Circumarctic map of permafrost and ground ice
- 645 conditions, in: Permafrost: Sixth International Conference, Sixth International Conference on Permafrost, Guangzhou, China,
- 646 1132-1136, 1997.
- 647 Holloway, J. E. and Lewkowicz, A. G.: Half a century of discontinuous permafrost persistence and degradation in western
- 648 Canada, Permafr. Periglac. Process., 31, 85–96, https://doi.org/10.1002/ppp.2017, 2020.
- 649 Hugelius, G., Strauss, J., Zubrzycki, S., Harden, J. W., Schuur, E. A. G., Ping, C.-L., Schirrmeister, L., Grosse, G., Michaelson,
- 50 G. J., Koven, C. D., O'Donnell, J. A., Elberling, B., Mishra, U., Camill, P., Yu, Z., Palmtag, J., and Kuhry, P.: Estimated stocks
- 651 of circumpolar permafrost carbon with quantified uncertainty ranges and identified data gaps, Biogeosciences, 11, 6573-6593,
- 52 https://doi.org/10.5194/bg-11-6573-2014, 2014.
- 653 Hugelius, G., Loisel, J., Chadburn, S., Jackson, R. B., Jones, M., MacDonald, G., Marushchak, M., Olefeldt, D., Packalen, M.,
- 654 Siewert, M. B., Treat, C., Turetsky, M., Voigt, C., and Yu, Z.: Large stocks of peatland carbon and nitrogen are vulnerable to
- permafrost thaw, Proc. Natl. Acad. Sci., 117, 20438–20446, https://doi.org/10.1073/pnas.1916387117, 2020.
- 656 Hustich, I.: Notes on the coniferous forest and tree limit on the east coast of Newfoundland-Labrador, Acta Geogr., 7, 81,
- 657 1939.
- 658 International Permafrost Association Terminology Working Group: Multi-language glossary of permafrost and related ground-
- 659 ice terms, edited by: van Everdingen, R. O., 2005.
- 660 Ives, J. D.: A proposed history of permafrost development in Labrador-Ungava, Géographie Phys. Quat., 33, 233-244,
- 661 https://doi.org/10.7202/1000360ar, 1979.

- 662 Jorgenson, M. T., Romanovsky, V., Harden, J., Shur, Y., O'Donnell, J., Schuur, E. A. G., Kanevskiy, M., and Marchenko, S.:
- Resilience and vulnerability of permafrost to climate change, Can. J. For. Res., 40, 1219–1236, https://doi.org/10.1139/X10-
- 664 060, 2010.
- 665 Karger, D. N., Conrad, O., Böhner, J., Kawohl, T., Kreft, H., Soria-Auza, R. W., Zimmermann, N. E., Linder, H. P., and
- 666 Kessler, M.: Climatologies at high resolution for the earth's land surface areas, Sci. Data, 4, 170122.
- 667 https://doi.org/10.1038/sdata.2017.122, 2017.
- 668 Karger, D. N., Conrad, O., Böhner, J., Kawohl, T., Kreft, H., Soria-Auza, R. W., Zimmermann, N. E., Linder, H. P., and
- 669 Kessler, M.: Climatologies at high resolution for the earth's land surface areas V2.1, 2021.
- 670 Karst, A. L. and Turner, N. J.: Local ecological knowledge and importance of bakeapple (Rubus chamaemorus L.) in a
- 71 southeast Labrador Métis community, Ethnobiol. Lett., 2, 6–18, https://doi.org/10.14237/ebl.2.2011.28, 2011.
- 672 Mamet, S. D., Chun, K. P., Kershaw, G. G. L., Loranty, M. M., and Kershaw, G. P.: Recent increases in permafrost thaw rates
- 673 and areal loss of palsas in the western Northwest Territories, Canada: Non-linear palsa degradation, Permafr. Periglac. Process.,
- 674 28, 619–633, https://doi.org/10.1002/ppp.1951, 2017.
- McLaughlin, J. and Webster, K.: Effects of climate change on peatlands in the far north of Ontario, Canada: A synthesis, Arct.
- 676 Antarct. Alp. Res., 46, 84–102, https://doi.org/10.1657/1938-4246-46.1.84, 2014.
- 677 Natural Resources Canada: Standards and Specifications of the National Topographic Data Base, Edition 3.1, 2005.
- 678 Norton, C. H., Cuerrier, A., and Hermanutz, L.: People and plants in Nunatsiavut (Labrador, Canada): Examining plants as a
- 679 foundational aspect of culture in the Subarctic, Econ. Bot., 75, 287–301, https://doi.org/10.1007/s12231-021-09530-7, 2021.
- 680 Obu, J., Westermann, S., Bartsch, A., Berdnikov, N., Christiansen, H. H., Dashtseren, A., Delaloye, R., Elberling, B.,
- Etzelmüller, B., Kholodov, A., Khomutov, A., Kääb, A., Leibman, M. O., Lewkowicz, A. G., Panda, S. K., Romanovsky, V.,
- 682 Way, R. G., Westergaard-Nielsen, A., Wu, T., Yamkhin, J., and Zou, D.: Northern Hemisphere permafrost map based on
- 683 TTOP modelling for 2000–2016 at 1 km2 scale, Earth-Sci. Rev., 193, 299–316
- 684 https://doi.org/10.1016/j.earscirev.2019.04.023, 2019.
- 685 Occhietti, S., Parent, M., Lajeunesse, P., Robert, F., and Govare, É.: Late Pleistocene-early Holocene decay of the Laurentide
- 686 Ice Sheet in Québec-Labrador, in: Developments in Quaternary Sciences, vol. 15, Elsevier, 601-630,
- 687 https://doi.org/10.1016/B978-0-444-53447-7.00047-7, 2011.
- 688 Olefeldt, D., Goswami, S., Grosse, G., Hayes, D., Hugelius, G., Kuhry, P., McGuire, A. D., Romanovsky, V. E., Sannel, A.
- 689 B. K., Schuur, E. A. G., and Turetsky, M. R.: Circumpolar distribution and carbon storage of thermokarst landscapes, Nat.
- 690 Commun., 7, 13043, https://doi.org/10.1038/ncomms13043, 2016.
- 691 Olefeldt, D., Hovemyr, M., Kuhn, M. A., Bastviken, D., Bohn, T. J., Connolly, J., Crill, P., Euskirchen, E. S., Finkelstein, S.
- 592 A., Genet, H., Grosse, G., Harris, L. I., Heffernan, L., Helbig, M., Hugelius, G., Hutchins, R., Juutinen, S., Lara, M. J.,
- 693 Malhotra, A., Manies, K., McGuire, A. D., Natali, S. M., O'Donnell, J. A., Parmentier, F.-J. W., Räsänen, A., Schädel, C.,
- 694 Sonnentag, O., Strack, M., Tank, S. E., Treat, C., Varner, R. K., Virtanen, T., Warren, R. K., and Watts, J. D.: The Boreal-
- 695 Arctic Wetland and Lake Dataset (BAWLD), Earth Syst. Sci. Data, 13, 5127–5149, https://doi.org/10.5194/essd-13-5127-
- 696 2021, 2021.
- 697 O'Neill, H. B., Wolfe, S. A., and Duchesne, C.: New ground ice maps for Canada using a paleogeographic modelling approach,
- 698 The Cryosphere, 13, 753–773, https://doi.org/10.5194/tc-13-753-2019, 2019.

- 699 Ou, C., LaRocque, A., Leblon, B., Zhang, Y., Webster, K., and McLaughlin, J.: Modelling and mapping permafrost at high
- 700 spatial resolution using Landsat and Radarsat-2 images in Northern Ontario, Canada: Part 2 regional mapping, Int. J. Remote
- 01 Sens., 37, 2751–2779, https://doi.org/10.1080/01431161.2016.1151574, 2016.
- 702 Parviainen, M. and Luoto, M.: Climate envelopes of mire complex types in fennoscandia, Geogr. Ann. Ser. Phys. Geogr., 89,
- 703 137–151, https://doi.org/10.1111/j.1468-0459.2007.00314.x, 2007.
- 704 Payette, S.: The forest tundra and present tree-lines of the northern Québec-Labrador peninsula, Proc. North. Qué. Tree-Line
- 705 Conf., 47, 3-23, 1983.
- 706 Pavette, S.: Chapitre 9: Les processus et les formes périglaciaires, in: Écologie des tourbières du Ouébec-Labrador, Presses de
- 707 l'Université Laval, 42, 2001.
- 708 Payette, S.: Accelerated thawing of subarctic peatland permafrost over the last 50 years, Geophys, Res. Lett., 31, L18208,
- 709 https://doi.org/10.1029/2004GL020358, 2004.
- 710 Pironkova, Z.: Mapping palsa and peat plateau changes in the Hudson Bay Lowlands, Canada, using historical aerial
- '11 photography and high-resolution satellite imagery, Can. J. Remote Sens., 43, 455–467
- 712 https://doi.org/10.1080/07038992.2017.1370366, 2017.
- 713 R Core Team: R, 2020.
- 714 Ramsdale, J. D., Balme, M. R., Conway, S. J., Gallagher, C., van Gasselt, S. A., Hauber, E., Orgel, C., Séjourné, A., Skinner,
- 715 J. A., Costard, F., Johnsson, A., Losiak, A., Reiss, D., Swirad, Z. M., Kereszturi, A., Smith, I. B., and Platz, T.: Grid-based
- 716 mapping: A method for rapidly determining the spatial distributions of small features over very large areas, Planet. Space Sci.,
- 717 140, 49–61, https://doi.org/10.1016/j.pss.2017.04.002, 2017.
- 718 Roberts, B. A., Simon, N. P. P., and Deering, K. W.: The forests and woodlands of Labrador, Canada: Ecology, distribution
- 719 and future management, Ecol. Res., 21, 868–880, https://doi.org/10.1007/s11284-006-0051-7, 2006.
- 720 Seguin, M. K. and Dionne, J. C.: Modélisation géophysique et caractérisation thermique du pergélisol dans les palses de Blanc-
- 721 Sablon, Quebec, Geol. Surv. Can. Curr. Res., E, 207–216, 1992.
- 722 Seppälä, M.: The origin of palsas, Geogr. Ann., 68, 141–147, 1986.
- 723 Seppälä, M.: Snow depth controls palsa growth, Permafr. Periglac. Process., 5, 283-288,
- 724 https://doi.org/10.1002/ppp.3430050407, 1994.
- 725 Smith, J. S.: Shifting sites and shifting sands: A record of prehistoric human/landscape interactions from Porcupine Strand,
- 726 Labrador, M.Sc. Thesis, Memorial University of Newfoundland, St. John's, Canada, 288 pp., 2003.
- 727 Smith, M. W. and Riseborough, D. W.: Climate and the limits of permafrost: A zonal analysis, Permafr. Periglac. Process., 13,
- 728 1–15, https://doi.org/10.1002/ppp.410, 2002.
- 729 Tarnocai, C.: The impact of climate change on Canadian peatlands, Can. Water Resour. J., 34, 453-466,
- 730 https://doi.org/10.4296/cwrj3404453, 2009.
- 731 Tarnocai, C., Kettles, I. M., and Lacelle, B.: Peatlands of Canada, Natural Resources Canada, Geological Survey of Canada,
- 732 Open File 6561, https://doi.org/10.4095/288786, 2011.

Formatted: French (Canada)

Formatted: French (Canada)

- 733 Thibault, S. and Payette, S.: Recent permafrost degradation in bogs of the James Bay area, northern Quebec, Canada, Permafr.
- 734 Periglac. Process., 20, 383–389, https://doi.org/10.1002/ppp.660, 2009.
- 735 Thie, J.: Distribution and thawing of permafrost in the southern part of the discontinuous permafrost zone in Manitoba,
- 736 ARCTIC, 27, 189–200, https://doi.org/10.14430/arctic2873, 1974.
- Vacchi, M., Engelhart, S. E., Nikitina, D., Ashe, E. L., Peltier, W. R., Roy, K., Kopp, R. E., and Horton, B. P.: Postglacial
- 738 relative sea-level histories along the eastern Canadian coastline, Quat. Sci. Rev., 201, 124–146,
- 739 https://doi.org/10.1016/j.quascirev.2018.09.043, 2018.
- Vallée, S. and Payette, S.: Collapse of permafrost mounds along a subarctic river over the last 100 years (northern Ouébec).
- 741 Geomorphology, 90, 162–170, https://doi.org/10.1016/j.geomorph.2007.01.019, 2007.
- 742 Wang, Y., Way, R. G., and Beer, J.: Coastal Labrador peatland permafrost inventory, v. 1.0. Nord. D98.
- 743 https://doi.org/10.5885/45762XD-1DB498A49B864CFB, 2022.
- 744 Way, R. G.: Field and modelling investigations of permafrost conditions in Labrador, northeast Canada, Ph.D. Thesis,
- 745 University of Ottawa, Ottawa, Canada, 293 pp., 2017.
- 746 Way, R. G. and Lewkowicz, A. G.: Investigations of discontinuous permafrost in coastal Labrador with DC electrical resistivity
- 747 tomography, Proc. 68th Can. Geotech. Conf. 7th Can. Permafr. Conf., 8, https://doi.org/10.13140/RG.2.1.1647.8803, 2015.
- 748 Way, R. G. and Lewkowicz, A. G.: Environmental controls on ground temperature and permafrost in Labrador, northeast
- 749 Canada, Permafr. Periglac. Process., 29, 73–85, https://doi.org/10.1002/ppp.1972, 2018.
- 750 Way, R. G. and Viau, A. E.: Natural and forced air temperature variability in the Labrador region of Canada during the past
- 751 century, Theor. Appl. Climatol., 121, 413–424, https://doi.org/10.1007/s00704-014-1248-2, 2015.
- 752 Way, R. G., Lewkowicz, A. G., and Bonnaventure, P. P.: Development of moderate-resolution gridded monthly air temperature
- 753 and degree-day maps for the Labrador-Ungava region of northern Canada: High-resolution air temperature and degree-day
- 754 maps for Labrador, Int. J. Climatol., 37, 493–508, https://doi.org/10.1002/joc.4721, 2017.
- 755 Way, R. G., Lewkowicz, A. G., and Zhang, Y.: Characteristics and fate of isolated permafrost patches in coastal Labrador,
- 756 Canada, The Cryosphere, 12, 2667–2688, https://doi.org/10.5194/tc-12-2667-2018, 2018.
- 757 Way, R. G., Wang, Y., Bevington, A. R., Bonnaventure, P. P., Burton, J. R., Davis, E., Garibaldi, M. C., Lapalme, C. M.,
- 758 Tutton, R., and Wehbe, M. A.: Consensus-Based Rock Glacier Inventorying in the Torngat Mountains, Northern Labrador,
- 759 Proc. 2021 Reg. Conf. Permafr. 19th Int. Conf. Cold Reg. Eng., 130–141, 2021a.
- 760 Way, R. G., Lewkowicz, A. G., Wang, Y., and McCarney, P.: Permafrost investigations below the marine limit at Nain,
- 761 Nunatsiavut, Canada, Proc. 2021 Reg. Conf. Permafr. 19th Int. Conf. Cold Reg. Eng., 38–48, 2021b.
- Wenner, C.-G.: Pollen diagrams from Labrador, Geogr. Ann., 29, 137–374, 1947.
- 763 Williams, P. J. and Smith, M. W.: The Frozen Earth: Fundamentals of Geocryology, Cambridge University Press, 1989.
- 764 Zoltai, S. C.: Palsas and peat plateaus in central Manitoba and Saskatchewan, Can. J. For. Res., 2, 291-302,
- 765 https://doi.org/10.1139/x72-046, 1972.

Zoltai, S. C. and Tarnocai, C.: Perennially frozen peatlands in the western Arctic and Subarctic of Canada, Can. J. Earth Sci.,
 12, 28–43, https://doi.org/10.1139/e75-004, 1975.

Zuidhoff, F. S. and Kolstrup, E.: Palsa development and associated vegetation in northern Sweden, Arct. Antarct. Alp. Res.,
 37, 49–60, https://doi.org/10.1657/1523-0430(2005)037[0049:PDAAVI]2.0.CO;2, 2005.