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The Cryosphere Discussions

# Antiphase dynamics between cold-based glaciers in the Antarctic Dry Valleys region and ice extent in the Ross Sea during MIS 5

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## Abstract

During the interglacial and interstadials of Marine Isotope Stage 5 (MIS 5e, 5c, 5a), outlet and alpine glaciers in the Dry Valleys region, Antarctica, appear to have advanced in response to increased precipitation from enhanced open ocean conditions in the Ross Sea. We provide further evidence of this antiphase behaviour through retreat of a peripheral lobe of Taylor Glacier in Pearse Valley, a region that was glaciated during MIS 5. We measured cosmogenic $^{10}Be$ and $^{26}Al$ in three granite cobbles from thin, patchy drift (Taylor 2 Drift) in Pearse Valley to constrain the timing of retreat of Taylor Glacier. Assuming simple continuous exposure, our minimum, zero erosion, exposure ages suggest Taylor Glacier partially retreated from Pearse Valley no later than 65–74 ka. Timing of retreat after 65 ka and until the Last Glacial Maximum (LGM) when Taylor Glacier was at a minimum position, remains unresolved. The depositional history of permafrost sediments buried below Taylor 2 Drift in Pearse Valley was obtained from $^{10}Be$ and $^{26}Al$ depth profiles to ~3 metres in permafrost in proximity to the cobble sampling sites. Depth profile modelling gives a depositional age for near-surface (<1.65 m) permafrost at Pearse Valley of 180 ka $^{+20}/_{-40}$ ka, implying deposition of permafrost sediments predate MIS 5 advances of Taylor Glacier. Depth profile modelling of deeper permafrost sediments (>2.09 m) indicates a depositional

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age of >180 ka. The cobble and permafrost ages reveal Taylor Glacier advances during MIS 5 were non-erosive or mildly erosive, preserving the underlying permafrost sediments and peppering boulders and cobbles upon an older, relict surface. Our results are consistent with U/Th ages from central Taylor Valley, and suggest changes in moisture delivery over Taylor Dome during MIS 5e, 5c and 5a appear to be associated with the extent of the Ross Ice Shelf and sea ice in the Ross Sea. At a coastal, lower elevation site in neighbouring Lower Wright Valley, $^{10}Be$ and $^{26}Al$ depth profiles from a second permafrost core exhibit near-constant concentrations with depth, and indicate the sediments are either vertically mixed after deposition, or are sufficiently young and post-depositional nuclide production is negligible relative to inheritance. $^{26}Al/^{10}Be$ concentration ratios for both depth profiles range between 4.0 and 5.2 and are all lower than the nominal surface production rate ratio of 6.75 indicating that prior to deposition, these sediments experienced a complex exposure-burial history. Assuming a single cycle exposure-burial scenario, the observed $^{26}Al/^{10}Be$ ratios are equivalent to a total minimum exposure-burial history of ~1.2 Ma. Our new data corroborates antiphase behaviour between outlet and alpine glaciers in the Dry Valleys region and ice extent in the Ross Sea. We suggest a causal relationship of cold-based glacier advance and retreat that is controlled by an increase in moisture availability during retreat of sea ice and perhaps the Ross Ice Shelf, and conversely, a decrease during times of sea ice and Ross Ice Shelf expansion in the Ross Sea.

## 1 Introduction

During Plio-Pleistocene warm intervals, the West Antarctic Ice Sheet (WAIS), and marine-based sectors of the East Antarctic Ice Sheet (EAIS) underwent extensive retreat (Naish et al., 2009; Pollard & DeConto, 2009; Cook et al., 2013; Blackburn et al., 2020; Patterson et al., 2014). Warmer than present global temperatures and higher than present sea levels are also observed in recent prominent interglacial periods, i.e., MIS 31 (~1.07 Ma), MIS 11 (~400 ka), and MIS 5e (130 - 115 ka) (Dutton et al., 2015; Naish et al., 2009; Pollard & DeConto, 2009). The extent of ice sheet retreat during these recent warm intervals varied significantly within different drainage basins and through time. During the penultimate interglacial (MIS 5e), the average global temperature was ~1–2°C warmer than pre-industrial (Fischer et al., 2018; Otto-Bliesner et al., 2013), Antarctic temperatures were ~3–5°C warmer (Jouzel et al., 2007) and global mean sea levels were ~6–9 metres higher than present (Dutton & Lambeck, 2012; Kopp et al., 2009). With a global average temperature currently ~1.1°C warmer than pre-industrial levels, and predicted to be ≥1.5°C in the coming decades (IPCC, 2021), interglacial conditions, such as during MIS 5, are an important analogue for

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evaluating future ice sheet behaviour and global climate processes under future warming scenarios.

Ice sheet modelling during the Last Interglacial (MIS 5e, 130–115 ka), projected Antarctic ice loss contributed ~3.5–7.5 m GMSL (global mean sea level), primarily from WAIS retreat (DeConto & Pollard, 2016; DeConto et al., 2021; Golledge et al., 2021; Turney et al., 2020). Simulated ice sheet retreat by Golledge et al. (2021) suggested ice loss in the Thwaites and Pine Island sector of the WAIS, whereas the Ross Ice Shelf remained intact. In contrast, simulations by DeConto et al. (2016), and Turney et al. (2020) suggested retreat of the Ross Ice Shelf, followed by retreat of the WAIS interior.

The $\delta^{18}O$ ice core records from Talos Dome reveal the EAIS was relatively intact during MIS 5 (Sutter et al., 2020) and recent studies suggest partial ice sheet lowering in Wilkes Subglacial Basin but no grounding line retreat (Fig. 1; Golledge et al., 2021; Sutter et al., 2020; Wilson et al., 2018). Ice core studies reveal increased accumulation rates at Taylor Dome (Steig et al., 2000) and the Allan Hills Blue Ice Area (Yan et al., 2021) near the onset of the Last Interglacial. Yan et al. (2021) hypothesized that high accumulation rates during warm interglacials may reflect open ocean conditions in the Ross Sea, caused by reduced sea ice extent, and possibly retreat of the Ross Ice Shelf relative to its present-day position. This hypothesis is supported by a depleted $\delta^{18}O$ value (−0.175 ‰) from ice core records at Roosevelt Island, indicating high sea level and reduced ice sheets during MIS 5a (Lee et al., 2020).

Terrestrial evidence from the Dry Valleys suggests Taylor and Ferrar glaciers were larger than present during globally warm mid-Pliocene climatic optimum (3.0–3.1 Ma), MIS 31 (1.07 Ma) (Swanger et al., 2011) and MIS 5 (Brook et al., 1993; Higgins et al., 2000a). The glacier advances appear to be out of phase with WAIS retreat and ocean warming during interglacial periods. Alpine glaciers in the Dry Valleys also appear out of phase with marine based ice sheet retreat and advanced during MIS11 (Swanger et al., 2017), MIS 5 (Swanger et al., 2019), and MIS 3 (Joy et al., 2017). Glacial deposits and moraines, which can be used to reconstruct past ice extent, have been preserved where cold-based glaciers have advanced and retreated during Quaternary glaciations. The past ice volume and extent of Taylor Glacier (during interglacial periods) has been derived from cosmogenic nuclide studies and mapping drift and moraine deposits in lower Kennar Valley (Swanger et al., 2011), and lower Arena Valley (Brook et al., 1993; Marchant et al., 1994), and U/Th dating in central Taylor Valley (Higgins et al., 2000a). MIS 5 age glacial deposits in central Taylor Valley and Arena Valley are mapped as Taylor 2 Drift (Bockheim et al., 2008; Brook et al.,

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1993; Cox et al., 2012; Denton et al., 1970), termed Bonney Drift by Higgins et al. (2000b). By inference, glacial deposits on the valley floor of Pearse Valley are mapped as Taylor 2 Drift (Bockheim et al., 2008; Cox et al., 2012; Denton et al., 1970). U/Th ages of algal carbonates in central Taylor Valley suggest multiple advance / retreat cycles of the Taylor Glacier snout during MIS 5, with retreat of Taylor Glacier continuing after the MIS 5/4 transition (Higgins et al., 2000a). The $\delta^{18}O$ values measured from buried ice in northern Pearse Valley also support the advance of Taylor Glacier during MIS 5 (Swanger et al., 2019). However, the timing of advance and retreat of Taylor Glacier in central Taylor Valley and in Pearse Valley remain poorly constrained.

Previous studies investigating chronology and stability of glacial drift deposits, sediments and permafrost in the Dry Valleys and Transantarctic Mountains typically focused on high elevation sites (e.g., Bergelin et al., 2022; Bibby et al., 2016; Morgan et al., 2011; 2010; Ng et al., 2005; Schäfer et al., 2000; Sugden et al., 1995). The objective of these studies has largely been to constrain the ages and / or erosion and sublimation rates of early Pleistocene, Pliocene, and Miocene landscapes. There only appears to be one study investigating the age and stability of permafrost below 1000 m elevation (Morgan et al., 2010). Yet, understanding the depositional environment and stability of permafrost at low elevations is important for interpreting landscape evolution, geomorphic processes and polar climate change on Earth, and as a terrestrial analogue for Mars (e.g., Marchant & Head, 2007).

Here, we investigate the relationship between thin, patchy drift overlying permafrost sediments in Pearse Valley. Thin, patchy drift is the only evidence of cold-based glacier overriding, and is defined as a scattering of clasts overlying older, undisturbed desert pavements (Atkins, 2013). We present cosmogenic nuclide surface exposure ages from three cobbles in Pearse Valley to determine the age of Taylor 2 Drift, and provide constraints on the timing of retreat of a peripheral lobe of Taylor Glacier during MIS 5. To determine the relationship between the thin, patchy drift and underlying permafrost sediments at Pearse Valley, we also present companion $^{10}Be$ and $^{26}Al$ depth profiles of permafrost. Combining exposure ages of cobbles from the drift and permafrost depth profiles, we constrain a minimum age of Taylor Glacier retreat, and infer the depositional history of the permafrost sediments. These data from Pearse Valley provide insight into Taylor Glacier behaviour and associated geomorphic processes during MIS 5. Additionally, we present $^{10}Be$ and $^{26}Al$ permafrost depth profiles from a coastal, lower elevation site in the neighbouring Lower Wright Valley, and together with the Pearse Valley depth profiles, discuss long-term recycling processes of Dry Valleys sediments.

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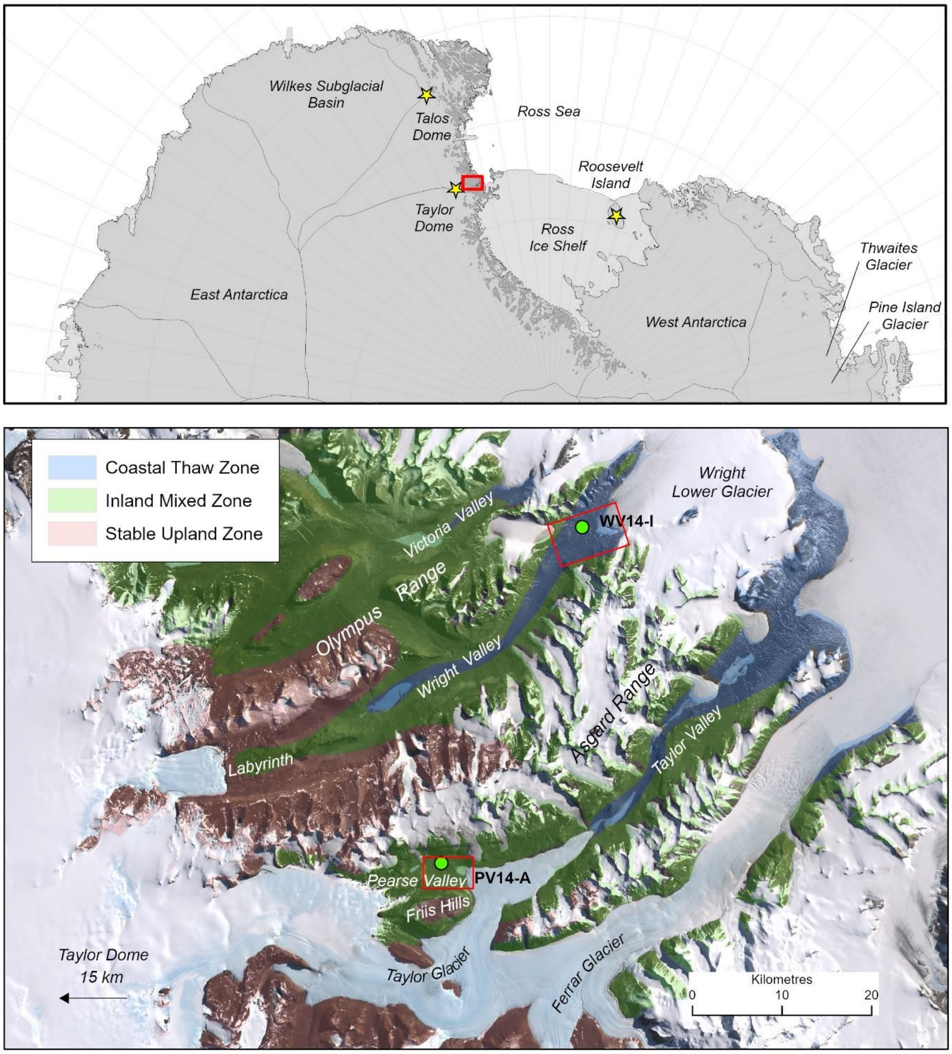


**Figure 1.** Study area and location of Dry Valleys. Yellow stars show ice core sites discussed in the text. The green circles show the locations of the Pearse Valley and Lower Wright Valley sites where permafrost cores were recovered. The three microclimatic zones are the stable upland zone (brown), inland mixed zone (green), and coastal thaw zone (blue). Modified from Marchant and Head (2007); and Salvatone and Levy (2021). Red rectangles in the lower diagram show the locations of Pearse Valley in Fig. 2 and Lower Wright Valley in Fig. 3.

## 2 Geologic setting and study area

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The Dry Valleys are a hyperarid, cold polar desert and can be subdivided into three geographic zones (stable upland, inland mixed, and coastal thaw zones), which are defined by their microclimatic parameters of atmospheric temperature, soil moisture, and relative humidity (Fig. 1; Marchant & Denton, 1996; Marchant & Head, 2007). The stability and evolution of geomorphic features and permafrost are controlled by subtle variations within each microclimatic zone. The active-layer in permafrost is defined as soil horizons where the ground temperature fluctuates above and below 0°C seasonally (Davis, 2001; Yershov, 1998). Antarctic permafrost soils along the floors and flanks of ice-free valleys are vertically mixed, initially through deposition of reworked sediments, and secondarily through active-layer cryoturbation up to 70 cm depth of the surface (Bockheim et al., 2007; 2008). Cryoturbation is defined as soil movement due to repeated freeze–thaw, generally within the active-layer of permafrost (French, 2017). Active-layers can be distinguished by the presence (wet active-layer) or absence (dry active-layer) of water. Soils in the coastal thaw zone are seasonally moist and comprise wet active-layers, whereas soils in the inland mixed zone are dry and comprise dry active-layers (Marchant & Head, 2007). Our study sites focused on two different microclimatic zones (Fig. 1); Pearse Valley in the inland mixed zone, and Lower Wright Valley in the coastal thaw zone, which differ in age, elevation, and distance from the coast.

### 2.1 Pearse Valley

Pearse Valley is an ice-free valley that is bounded by the Friis Hills in the south, the Asgard Range in the north and opens onto peripheral lobes of Taylor Glacier in the east and west (Fig. 1). Taylor Glacier flows east from Taylor Dome of the EAIS, terminating in Taylor Valley. At the eastern end of Pearse Valley, a lobe of Taylor Glacier terminates into Lake Joyce, a closed-basin proglacial lake (Fig. 2). Taylor Glacier and local alpine glaciers have advanced in the present interglacial and occupy their maximum position since the Last Glacial Maximum (LGM) (Higgins et al., 2000a). At the head of Pearse Valley, glacially incised bedrock sits at a similar elevation to the Labyrinth platform in upper Wright Valley, likely formed by a network of subglacial drainage channels beneath wet-based glacial conditions during the Miocene Climate Transition (Fig. 1; Lewis & Ashworth, 2016; Chorley et al., 2022). The northern valley wall comprises gelifluction lobes, buried snowpack deposits, meltwater channels derived from ephemeral streams, and fans fed by the meltwater channels in front of the lobes (Heldmann et al., 2012; Swanger et al., 2019). The valley floor consists of a lower elevation area on the southern side, and a higher elevation area on the northern side of the valley. The PV14-A core and cobble samples are located on the central northern side of the valley floor (Fig. 2).

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The local bedrock comprises basement granites and Ferrar dolerite intrusives (Cox et al., 2012; Gunn & Warren, 1962). Glacial deposits on the valley floor are mapped as Taylor 2 Drift (Bockheim et al., 2008; Denton et al., 1970). These sediments were inferred as waterlain and melt-out tills following the penultimate down-valley advance of the Taylor Glacier during MIS 5 (70 – 130 ka) (Cox et al., 2012; Higgins et al., 2000a; Swanger et al., 2019). The valley floor landscape is characterized by hummocky moraines with a combination of glacigenic, and fluvial deposits, and aeolian sediments. Variably weathered granite boulders (up to 3 m in diameter) form a lag deposit on the drift surface, inferred as a till deflation or a separate younger depositional unit (Higgins et al., 2000b). The northern and southern Pearse Valley walls comprises extensive rock glaciers (Swanger et al., 2019).

### 2.1.1 Modern climate

Pearse Valley is situated in the inland mixed zone of the Dry Valleys (Marchant & Denton, 1996). The valley has a mean annual temperature of -18°C (Marchant et al., 2013) and precipitation rates of 20–50 mm/yr (water equivalent), and 100–200 mm/yr in the adjacent Asgard Range, the source region for the local alpine glaciers (Fountain et al., 2010). Mean summer air temperatures (December through February) in Pearse Valley are -2 to -7°C (Marchant et al., 2013). Ground surface temperatures measured at the Pearse Valley meteorological station between 27–28 November, 2009, recorded a peak temperature of 10°C due to solar heating (Heldmann et al., 2012). Winds in Pearse Valley are strong enough to mobilise sand grains and form aeolian surface features such as sand dunes (Heldmann et al., 2012).

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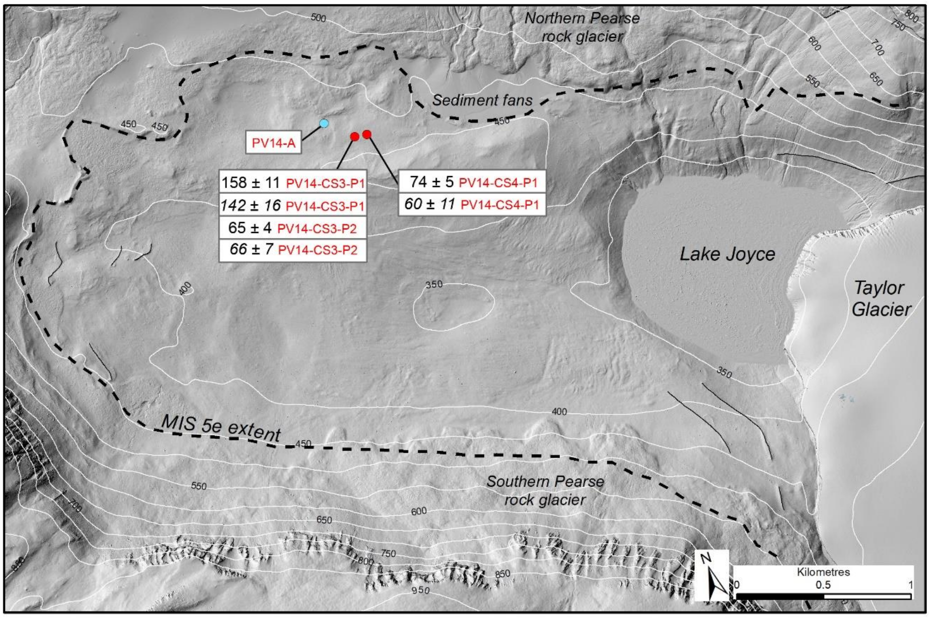


**Figure 2.** Map of Pearse Valley with MIS 5e extent of Taylor Glacier (black dashed line; Cox et al. 2012), sample locations and PV14-A permafrost drill site. Thin black lines trace undated moraines. PV14-A drill site (blue circle) and measured $^{10}$Be and $^{26}$Al (italics) ages of cobbles residing on boulders are shown in kiloyears with 1σ uncertainties (red circles). Lidar image from Fountain et al. (2017).

## 2.2 Lower Wright Valley

Lower Wright Valley is ice-free and is bounded by the Asgard Range in the south, and the Olympus Range in the north (Fig. 1). The mouth of the valley at the eastern end is cut off from the Ross Sea by the Wright Lower Glacier, a lobe of the Wilson Piedmont Glacier. Lake Brownworth, a proglacial lake fed by the Wright Lower Glacier, supplies the westward flowing Onyx River. The WV14-I core is located on the northern side of Lower Wright Valley (Fig. 3). Radiocarbon dates of lacustrine algae from glaciolacustrine deposits suggest Lake Brownworth is a small remnant of a much larger lake that existed during the LGM and early Holocene (Hall et al., 2001). The post-glacial, Holocene age landscapes form hummocky moraines, with a combination of deltas, shorelines and glaciolacustrine sediments (Hall et al., 2001). Glacial meltwater streams drain into Lake Brownworth and the Onyx River from the north and south valley walls. The local bedrock comprises basement metasediments and granites, and Ferrar dolerite intrusives (Cox et al., 2012). Metasediments, granite, dolerite and occasional basalt sediments in the Lower Wright Valley have accumulated since the last

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deglaciation by lacustrine, fluvial and aeolian processes (Hall et al., 2001; Hall & Denton, 2005).

### 2.2.1 Modern climate

Lower Wright Valley is situated in the coastal thaw zone of the Dry Valleys (Marchant & Denton, 1996) and has a mean annual temperature of -21°C (Doran et al., 2002) and precipitation rates of 26–51 mm/yr (water equivalent) (Fountain et al., 2010). Mean summer air temperatures (December through February) in Lower Wright Valley are -5 to -7°C, and can exceed 0°C for >6 days per year (Doran et al., 2002). Meltwater forms during summer months (December and January) when temperatures can rise to as much as 10°C at some locations (Hall et al., 2001).

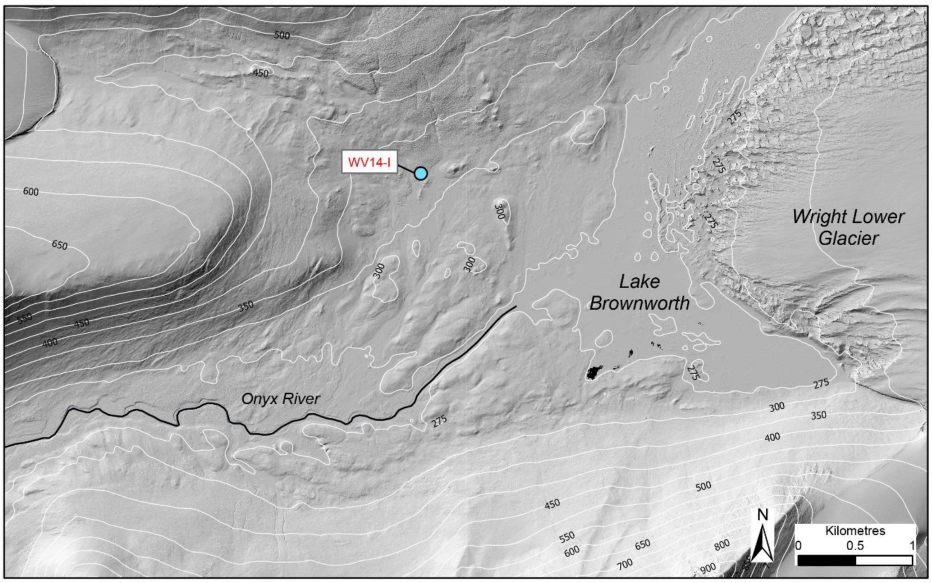


**Figure 3.** Map of Lower Wright Valley and WV14-I permafrost drill site (blue circle). Lidar image from Fountain et al. (2017).

## 3 Methods

### 3.1 Surface exposure sample collection

Three granite cobble samples were collected for surface exposure analysis from Pearse Valley (Table 1; Fig. 2). We targeted perched cobbles, resting on larger flat boulders to

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minimise the possibility of post-depositional disturbance and hence best reflect deposition from retreating glacier ice or from surface deflation through sublimation. Samples that showed minimal weathering or fracturing were selected. The three cobbles were perched on larger host boulders (>1 m diameter) which were elevated above the local surface permafrost valley deposits (Fig. 4). Two samples (PV14-CS3-P1 and PV14-CS3-P2) are small cobbles perched on the same host boulder, while the third sample (PV14-CS4-P1) is a slightly larger cobble perched on a different host boulder less than 80 metres away.


**Figure 4.** Boulders and cobbles from Taylor 2 Drift on the central northern side of Pearse Valley. (a) PV14-CS3-P1 and PV14-CS3-P2 cobbles perched on a dolerite boulder. (b) Close view of PV14-CS3-P2. (c) PV14-CS4-P1 cobble hosted on dolerite boulder. (d) A granite boulder, hosting a dolerite boulder.

### 3.2 Permafrost core locations and characteristics

During the 2014/15 austral field season, permafrost cores were recovered from Pearse Valley and Lower Wright Valley using a gasoline powered dry drilling technique (Fig. 1). These two cores were sampled for sedimentological and for cosmogenic nuclide analysis. After extraction, the core sections were divided into ~10 cm portions for sub-sampling and analysis. Permafrost sediments were collected in a combination of Whirl-Pak bags and PVC core liners.

### 3.2.1 Pearse Valley borehole core

The PV14-A core is located on an elevated bench that extends along the northern side of the valley floor at 450 masl (77.7062°S, 161.5467°E), ~3 km north-west of the present position of the Taylor Glacier lobe (Fig. 2). The core was recovered to a depth of 3.16 m (Fig 5a; Table 2). The active-layer above the ice-cemented permafrost consists of a thin armoured surface layer of desert pavement (~0.02 m thick), which caps a layer of loose dry sand (~0.35 m thick). Recovered sediments from beneath the armoured desert pavement comprise a dry active-layer of loose sand and pebbles down to 0.37 m depth. Below 0.37 m depth the recovered sediments comprise ice-cemented permafrost. The $^{10}Be$ and $^{26}Al$ depth profiles start below the 0.02 m thick surface armoured pavement. The first three samples were collected from the dry active-layer and nine from the ice-cemented permafrost (Fig. 5a). Sediments within the permafrost core comprise gravelly sands derived from weathered Beacon Supergroup, granite, granodiorite, diorite, and dolerite origins, which are structureless, or weakly bedded. Between 0.73–0.86 m depth, the core comprises several ice lenses indicative of ice accumulation below a paleosublimation unconformity. Several small ice lenses were also recovered between 1.57–1.87 m depth. Only two of the three active-layer samples, and six of the nine permafrost core samples were successful in providing paired $^{10}Be$ and $^{26}Al$ concentrations.

### 3.2.2 Lower Wright Valley borehole core

The WV14-I core is located in eastern Wright Valley at 326 masl (77.4252°S, 162.6664°E), ~2 km west of Wright Lower Glacier (Fig. 3). The core was recovered to a depth of 2.91 m (Fig. 5b; Table 2). The active-layer above the ice-cemented permafrost consists of a thin armoured surface layer of desert pavement (~0.02 m thick), which caps a layer of loose sand and pebbles (~0.26 m thick). Below 0.28 m depth the recovered sediments comprised ice-cemented permafrost. The $^{10}Be$ and $^{26}Al$ depth profiles start on the armoured desert pavement. Two samples were collected from the active-layer and 10 from the ice-cemented permafrost (Fig. 5b). The permafrost sediments are structureless, to thinly laminated, fine to coarse, and pebbly granular sands, which we interpret to be fluvial and aeolian deposits. Sediments within the core are derived from weathered granite, metasedimentary, dolerite and basalt origins. From 0–0.98 m depth, core sections were broken and loose sediment was recovered. Sediments recovered from 0.98–2.91 m were ice-cemented, except when encountering ice lenses. Several small ice lenses were recovered between 1.80–2.03 m

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depth. Hall et al. (2001) suggested sediments at Lower Wright Valley are delta, shoreline and glaciolacustrine deposits associated with a large proglacial lake at the LGM and in the early Holocene (25–7 ka). Only four of the 10 permafrost core samples were successful in providing paired $^{10}Be$ and $^{26}Al$ concentrations.

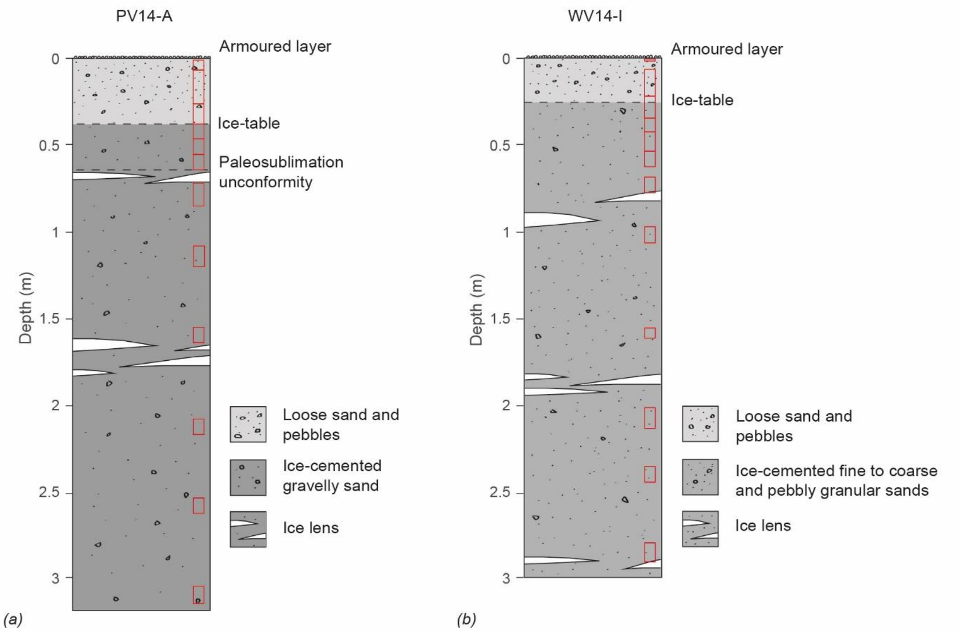


**Figure 5.** (a) Pearse Valley (PV14-A) permafrost core, and (b) Lower Wright Valley (WV14-I) permafrost core sedimentology. Locations of cosmogenic nuclide samples shown in red boxes.

### 3.3 Analytical methods

Each core sample processed for cosmogenic nuclide analysis was heated at 100°C overnight to remove ice and dry the sediment. Dried core samples, and cobble surface samples were crushed and sieved to obtain the 250 – 500 µm fraction. Quartz was separated and purified using the hot phosphoric acid method (Mifsud et al., 2013) and beryllium and aluminium were extracted from quartz via conventional HF dissolution and ion exchange chromatography (Child et al., 2000). Isotope ratios were measured by Accelerator Mass Spectrometry on the SIRIUS accelerator at the Australian Nuclear Science and Technology Organisation (Wilcken et al., 2019).

Measured $^{10}Be/^{9}Be$ ratios were normalised to the 07KNSTD (KN-5.2) standard of Nishiizumi et al. (2007) with a nominal $^{10}Be/^{9}Be$ ratio of 8560 x $10^{-15}$. Measured $^{26}Al/^{27}Al$ ratios were

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normalised to the KNSTD (KN-4.2) standard of Nishiizumi (2004) with a nominal $^{26}Al/^{27}Al$ ratio of 30960 x $10^{-15}$. The nuclide concertation data for the perched cobbles, and Pearse Valley and Lower Wright Valley depth profiles are shown in Tables 1 and 2, respectively. Full procedural $^{10}Be/^{9}Be$ blanks were obtained using a solution processed from dissolved beryl mineral with a known $^{9}Be$ concentration (1068 and 1048 µg/g (solution)) and resulted in ratios of 1.9 ± 0.4 x$10^{-15}$ and 1.3 ± 0.3 x$10^{-14}$. Blank corrections to measured $^{10}Be/^{9}Be$ ratios amounted to <2%. Procedural $^{26}Al/^{27}Al$ blanks were processed from standard reference ICP aluminium solutions (1000 µg/ml ±1%) and resulted in ratios 3.6 ± 1.7 x $10^{-14}$ and 1.3 ± 0.6 x$10^{-15}$. Blank corrections to measured $^{26}Al/^{27}Al$ ratios amounted to 4 to 35% for Pearse Valley erratics and <1% for all other samples. Final errors in $^{10}Be$ and $^{26}Al$ concentrations are obtained by quadrature addition of the final AMS analytical error (the larger of the total statistical or standard mean error), a reproducibility error based on the standard deviation of the set of standard reference samples measured during the run (typically 1-2% for either $^{10}Be$ or $^{26}Al$), a 1% error in Be spike concentration and a representative 3% error for ICP Al concentration of the native $^{27}Al$ in the final purified quartz powder. Unless otherwise stated, all analytical uncertainties are 1σ.

Surface exposure ages for the cobble samples were calculated using version 3 of the CRONUS-Earth calculator (http://hess.ess.washington.edu/; Balco et al., 2008) using the LSDn scaling scheme (Lifton et al., 2014) and the primary default calibration data set of Borchers et al. (2016) (Table 1). Complete analytical data for all measurements are shown in Table S1, and data from surface samples are archived on the ICE-D Antarctica database (http://antarctica.ice-d.org).

### 3.4 Dual nuclide depth profile models and parameters

$^{10}Be$ and $^{26}Al$ data from core samples were used to model the surface exposure age of permafrost sediments at Pearse and Lower Wright valleys via the depth profile technique (Anderson et al., 1996). We implemented a modified version of the Monte Carlo-based code of Hidy et al. (2010) that allows profiles of both $^{10}Be$ and $^{26}Al$ to be modelled jointly (after Hidy et al. (2018)). For shallow profiles in sediments, where non-unique solutions for exposure age and erosion rate are likely, this approach allows estimation of exposure age and pre-depositional nuclide concentration (i.e., inheritance) given reasonable observation-based constraint on erosion rate or net erosion (e.g., Bergelin et al., 2022; Hidy et al., 2010, 2018; Mercader et al., 2012; Morgan et al., 2010).

The simplest assumptions are that all depth profile sediments have the same inherited nuclide concentration at the time of deposition and that post depositional sediment mixing is

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absent. The former is a reasonable assumption for our core samples given that these sediments comprise a combination of well mixed, thick glacial tills, fluvial and aeolian sediments that were deposited at a given time when the ice retreated from each valley. As described in Sect. 3.2 above, the upper ~0.3 m of both cores consists of loose sandy sediment that is mobile or active. Fig. 6 shows the evolution of a cosmogenic nuclide depth profile over time with the added feature of a near-constant $^{10}$Be concentration in a cryoturbated active-layer above ice-cemented permafrost sediments. Any post-depositional nuclide production is unknown, but the inheritance determined by the best-fit depth profile asymptote can be subtracted from the measured values for each sample (Hidy et al., 2018).

To ensure consistency with the cobble exposure ages, we obtain production rates applied in the depth profile model from the CRONUS-Earth calculator. For the PV14-A core, we use a site-specific spallation $^{10}$Be surface production rate of 8.40 atoms $^{10}$Be g$^{-1}$ (quartz) yr$^{-1}$, and a $^{26}$Al surface production rate of 59.7 atoms $^{26}$Al g$^{-1}$ (quartz) yr$^{-1}$. For the WV14-I core, we use a site-specific spallation $^{10}$Be surface production rate of 7.47 atoms $^{10}$Be g$^{-1}$ (quartz) yr$^{-1}$, and a $^{26}$Al surface production rate of 53.2 atoms $^{26}$Al g$^{-1}$ (quartz) yr$^{-1}$. These production rates were calculated using LSDn scaling (Lifton et al., 2014) and the primary calibration data set of Borchers et al. (2016). These production rates yield $^{26}$Al/$^{10}$Be surface production rate ratios of 7.11 and 7.12 for Pearse Valley and Lower Wright Valley, respectively. We assume a neutron attenuation length of 140 ± 5 g cm$^{-2}$, as used in previous Antarctic studies for $^{10}$Be and $^{26}$Al (Bergelin et al., 2022; Borchers et al., 2016). Spallogenic production rate uncertainty has not been included in the modelling. Muogenic production with depth, including an assumed 8% uncertainty, followed Model 1A from Balco (2017). We assume bulk density to be constant with depth but sampled from a normal distribution of 1.7 ± 0.1 g cm$^{-3}$ based on bulk density measured from two core samples. Erosion rate and net erosion were constrained between 0–0.4 cm/ka and 400 cm, respectively, based on field observations described in Sect. 4.2. Within these constraints, exposure age, surface erosion rate, and inheritance for $^{10}$Be and $^{26}$Al were simulated with uniform distributions, and model output was based on n=100,000 acceptable depth profile solutions.

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**Table 1.** Cosmogenic $^{10}$Be and $^{26}$Al concentrations and apparent exposure ages from Pearse Valley

| Sample name | Latitude (DD) | Longitude (DD) | Elevation (masl) | Sample thickness (cm) | Topographic shielding | $^{10}$Be conc. ($10^5$ atoms $g^{-1}$)[a] | $^{26}$Al conc. ($10^5$ atoms $g^{-1}$)[b] | Apparent $^{10}$Be exposure age (ka)[c,d] | Apparent $^{26}$Al exposure age (ka)[c,d] | $^{26}$Al/$^{10}$Be ratio | Erosion-corrected $^{10}$Be exposure age (ka)[e] |
|---|---|---|---|---|---|---|---|---|---|---|---|
| PV14-CS3-P1 | -77.70737 | 161.55283 | 451 | 6 | 0.993 | 12.40 ± 0.39 | 76.57 ± 4.48 | 158 ± 11 (5) | 142 ± 16 (9) | 6.18 ± 0.41 | 174 ± 13 (6) |
| PV14-CS3-P2 | -77.70737 | 161.55283 | 451 | 3 | 0.993 | 5.36 ± 0.15 | 37.99 ± 1.54 | 65 ± 4 (2) | 66 ± 7 (3) | 7.09 ± 0.35 | 68 ± 5 (2) |
| PV14-CS4-P1 | -77.70747 | 161.55582 | 451 | 5 | 0.993 | 5.94 ± 0.16 | 33.71 ± 5.14 | 74 ± 5 (2) | 60 ± 11 (9) | 5.68 ± 0.88 | 77 ± 5 (2) |

All samples are granite cobbles and have a density of 2.65 g $cm^{-3}$.

[a] Normalised to the 07KNSTD (KN-5.2) standard of Nishiizumi et al. (2007).

[b] Normalised to the KNSTD (KN-4.2) standard of Nishiizumi (2004).

[c] Exposure ages calculated using the CRONUS-Earth calculator (http://hess.ess.washington.edu/math/), using the LSDn scaling scheme.

[d] Both internal and external uncertainties (shown at the 1σ level). Internal uncertainties (given in parentheses) are analytical uncertainties only and external uncertainties are absolute uncertainties and include production rate and scaling errors.

[e] Calculated using an erosion rate of 0.65 mm/ka.

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**Table 2.** Depth profile data from Pearse Valley and Lower Wright Valley

| Sample name | Sample depth (m) | $^{10}$Be conc. ($10^6$ atoms $g^{-1}$)[a] | $^{26}$Al conc. ($10^6$ atoms $g^{-1}$)[b] | $^{26}$Al/$^{10}$Be ratio |
|---|---|---|---|---|
| **Pearse Valley** | | | | |
| PV14-SS-5 | 0.02 - 0.07 | 4.24 ± 0.095 | - | - |
| PV14-A-01 | 0.07 - 0.27 | 4.37 ± 0.097 | 18.67 ± 0.73 | 4.27 ± 0.19 |
| PV14-A-02 | 0.27 - 0.37 | 4.35 ± 0.097 | 17.97 ± 0.71 | 4.13 ± 0.19 |
| PV14-A-03 | 0.37 - 0.47 | 4.42 ± 0.098 | 19.63 ± 0.82 | 4.44 ± 0.21 |
| PV14-A-04 | 0.47 - 0.56 | - | 19.94 ± 0.78 | - |
| PV14-A-05 | 0.56 - 0.65 | 4.40 ± 0.098 | 18.28 ± 0.69 | 4.16 ± 0.18 |
| PV14-A-07 | 0.73 - 0.86 | 3.96 ± 0.089 | 17.95 ± 0.70 | 4.53 ± 0.20 |
| PV14-A-10 | 1.09 - 1.21 | - | 16.38 ± 0.64 | - |
| PV14-A-15 | 1.56 - 1.65 | 3.80 ± 0.085 | 15.09 ± 0.59 | 3.97 ± 0.18 |
| PV14-A-20 | 2.09 - 2.18 | 3.98 ± 0.080 | 17.50 ± 0.66 | 4.40 ± 0.19 |
| PV14-A-25 | 2.55 - 2.64 | 3.85 ± 0.086 | 16.70 ± 0.66 | 4.33 ± 0.20 |
| PV14-A-30 | 3.06 - 3.16 | - | 16.76 ± 0.66 | - |
| **Lower Wright Valley** | | | | |
| WV14-SS-01 | 0 - 0.02 | 4.10 ± 0.092 | 22.89 ± 0.89 | 5.58 ± 0.25 |
| WV14-I-01 | 0.07 - 0.23 | 3.73 ± 0.175 | 19.04 ± 0.75 | 5.10 ± 0.31 |
| WV14-I-02 | 0.23 - 0.35 | 3.92 ± 0.088 | 18.43 ± 0.72 | 4.70 ± 0.21 |
| WV14-I-03 | 0.35 - 0.43 | 4.00 ± 0.089 | 20.38 ± 0.77 | 5.09 ± 0.22 |
| WV14-I-04 | 0.43 - 0.54 | - | 22.72 ± 0.89 | - |
| WV14-I-05 | 0.54 - 0.63 | - | 21.66 ± 0.85 | - |
| WV14-I-07 | 0.69 - 0.78 | - | 19.99 ± 0.79 | - |
| WV14-I-10 | 0.98 - 1.07 | 4.09 ± 0.091 | 20.54 ± 0.81 | 5.02 ± 0.23 |
| WV14-I-14 | 1.56 - 1.62 | - | 20.62 ± 0.81 | - |
| WV14-I-20 | 2.02 - 2.14 | 4.22 ± 0.094 | 21.80 ± 0.86 | 5.17 ± 0.23 |
| WV14-I-23 | 2.36 - 2.45 | - | 21.41 ± 0.84 | - |
| WV14-I-29 | 2.80 - 2.91 | - | 13.60 ± 0.53 | - |

We assume a constant bulk density of 1.7 ± 0.1 g $cm^{-3}$ based on bulk density measurements made on two core samples.

Topographic shielding is 0.9932 for Pearse Valley, and 0.9968 for Lower Wright Valley, respectively.

[a] Normalised to the 07KNSTD (KN-5.2) standard of Nishiizumi et al. (2007).

[b] Normalised to the KNSTD (KN-4.2) standard of Nishiizumi (2004).

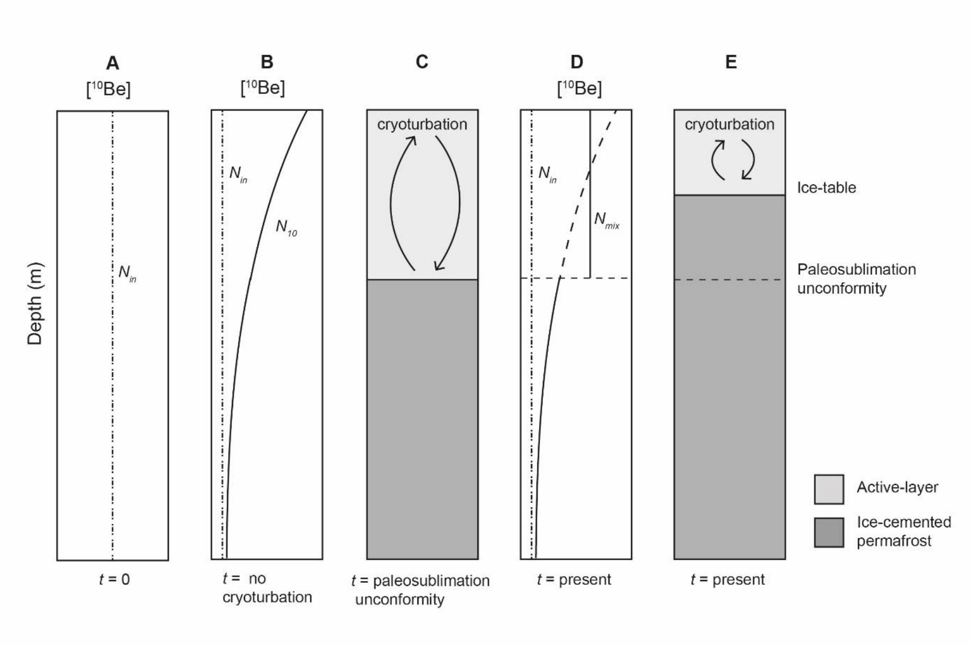


**Figure 6.** Schematic representation of a $^{10}Be$ depth profile in permafrost modified by active-layer cryoturbation. (a) Initial $^{10}Be$ profile (constant with depth) in well-mixed glacial till or sediment. All quartz grains are assumed to have been deposited with a common nuclide inheritance ($N_{in}$). (b) After prolonged exposure and in the absence of sediment mixing, an exponentially decreasing nuclide depth profile is obtained. (c) Permafrost profile during an interval where air temperature is warmer than present allowing near surface sediments to form an active-layer above the paleo-sublimation depth. Sediments below the unconformity are perennially frozen. (d) Vertical mixing via active-layer cryoturbation results in an average $^{10}Be$ value ($N_{mix}$). A decreasing $^{10}Be$ profile remains below the unconformity. (e) Present-day permafrost profile with shallower active-layer and ice-table than shown in (c).

## 4 Results

### 4.1 Surface exposure ages and erosion rates at Pearse Valley

Boulders and cobbles of granite, gneiss, Beacon sandstone and dolerite pepper the Pearse Valley floor, forming a thin, patchy drift overlying an older, well-weathered relict drift surface. Some boulders lodged in the relict drift host smaller perched boulders, cobbles, and pebbles on their surfaces, indicating deposition of perched clasts occurred after the most recent retreat of Taylor Glacier (Fig. 4).

Our surface exposure chronology is based on three granitic cobbles on the northern side of the central valley floor (Table 1, Fig. 2). Two samples (PV14-CS3-P2 and PV14-CS4-P1)

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yielded minimum zero erosion $^{10}$Be exposure ages of 65 ± 4 ka and 74 ± 5 ka (1σ external errors), respectively, whereas the third sample (PV14-CS3-P1) yielded an older age of 158 ± 11 ka, presumably affected by inheritance (Table 1). The three $^{26}$Al/$^{10}$Be concentration ratios range from 5.7 to 7.1 and when plotted on $^{10}$Be-$^{26}$Al/$^{10}$Be diagram, largely indicates a simple exposure within their 1σ error ellipses without any prior complex history (Fig. 7). While this assumption of zero erosion makes negligible difference for LGM and younger ages, we evaluate the influence of surface erosion on the exposure ages above using known erosion rates reported from Antarctica and geological evidence from the sites. Bedrock and regolith erosion rates in the Dry Valleys range from 0.1–4 mm/ka (Putkonen et al., 2008; Summerfield et al., 1999). A compiled study across Antarctica showed that granite populations have a mean erosion rate of 0.13 mm/ka, and in the Dry valleys, a max erosion rate of 0.65 mm/ka (Marrero et al., 2018). Applying the max erosion rate (0.65 mm/ka) from granite surfaces in the Dry Valleys, erosion corrected $^{10}$Be exposure ages of our granitic cobbles resulted in 174 ± 13 ka (PV14-CS3-P1), 68 ± 5 ka (PV14-CS3-P2) and 77 ± 5 ka (PV14-CS4-P1) (1σ external errors; Table 1). The cobble sample PV14-CS3-P2 displays minimal edge rounding which suggests negligible erosion and is unlikely to be much older than the zero-erosion age.

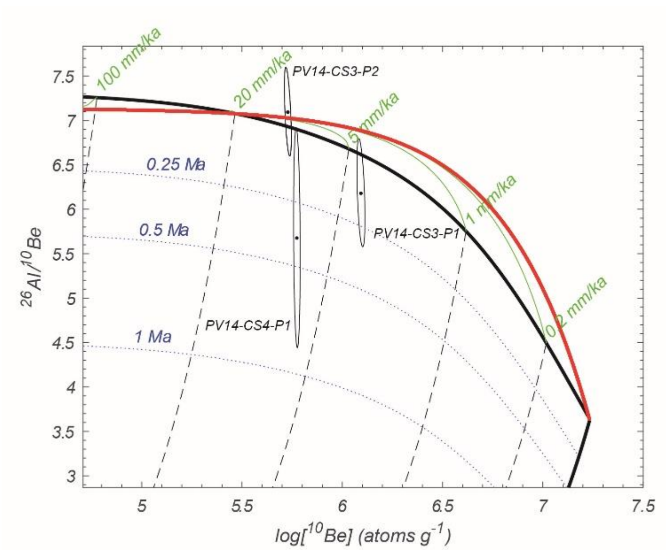


**Figure 7.** Two-isotope plot of Pearse Valley cobbles. Nuclide concentrations with 1σ uncertainties, using the time-dependent LSDn scaling scheme of Lifton et al. (2014). Burial isochrons (dotted lines), decay trajectories (dashed), the exposure-erosion region (bounded by black and red lines), and steady-state erosion loci (green) are shown.

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### 4.2 Cosmogenic nuclide depth profiles at Pearse Valley

The $^{10}$Be and $^{26}$Al depth profiles from the permafrost core and overlying active-layer at Pearse Valley, and associated modelled nuclide concentrations from a best-fit to all samples are shown in Fig. 8. No acceptable depth profile model fit was obtained for all measured $^{10}$Be and $^{26}$Al depth profile samples (see Fig. 8). However, the model appears to have performed better for the deeper samples >2.09 m, than for the shallower samples <1.65 m. This result suggests that the depth profile is of a composite structure. This is supported by the observation that ice lenses appear at ~0.7 m, and at ~1.70–1.80 m, which are also associated with distinct changes in $^{10}$Be and $^{26}$Al concentrations. Between 0.65 m and 1.65 m, both $^{10}$Be and $^{26}$Al cosmogenic nuclide concentrations display attenuation with depth, whilst below 1.65 m, the attenuation is interrupted by a considerable increase in nuclide concentrations as shown in the sample at 2.09 m depth. We attempt a model best-fit only to the samples above 1.65 m in order to determine the younger depositional phase. The five $^{10}$Be and five $^{26}$Al nuclide concentrations from 0.02–0.65 m exhibit a uniform concentration with depth with averages of $4.36 \pm 0.10 \times 10^{6}$ atoms g$^{-1}$ and $1.89 \pm 0.07 \times 10^{7}$ atoms g$^{-1}$, respectively, with no attenuation, indicating that these upper sediments have been vertically mixed (or possibly deposited sufficiently recently so that nuclide depth profiles effectively reflect only inheritance without significant post-depositional production). To accommodate the vertically-mixed, uniform $^{10}$Be and $^{26}$Al concentrations in the upper 0.65 m we use the mean $^{10}$Be and $^{26}$Al concentrations from these samples to best represent the process that resulted in the uniform profile (i.e., a vertically mixed cryoturbated layer or the most recent deposition) as shown in Fig. 9. The best-fit modelled nuclide concentrations for the PV14-A depth profile when restricted to samples from 0.02 to 1.65 m depth, falls within the 25$^{th}$ to 75$^{th}$ percentile of the measured concentrations. We constrained the erosion rate of the depth profiles using information from surface cobble PV14-CS3-P2 which sits ~10–20 cm above the desert pavement and has a minimum exposure age of 65 ka (Fig. 4a). Based on this observation we can assume a maximum surface lowering rate of ~0.3 cm ka$^{-1}$. Using this field observation, we applied a conservatively high erosion rate limit of 0.4 cm ka$^{-1}$ for our depth profile modelling. The solutions yield most probable $^{10}$Be and $^{26}$Al inheritance concentrations of $3.59 \times 10^{6}$ and $1.42 \times 10^{7}$ atoms g$^{-1}$, respectively (Fig. 9; Fig S1) and constrain the depositional age of the permafrost (<1.65 m depth) at 180 $^{+20}$ / $_{-40}$ ka (Fig. 10), and an erosion rate of 0.24 $^{+0.10}$ / $_{-0.09}$ cm ka$^{-1}$ (Fig. S2). By inference, the lower part of the profile (>2.09 m depth) predates the sediments above and must be deposited before ~180 ka.

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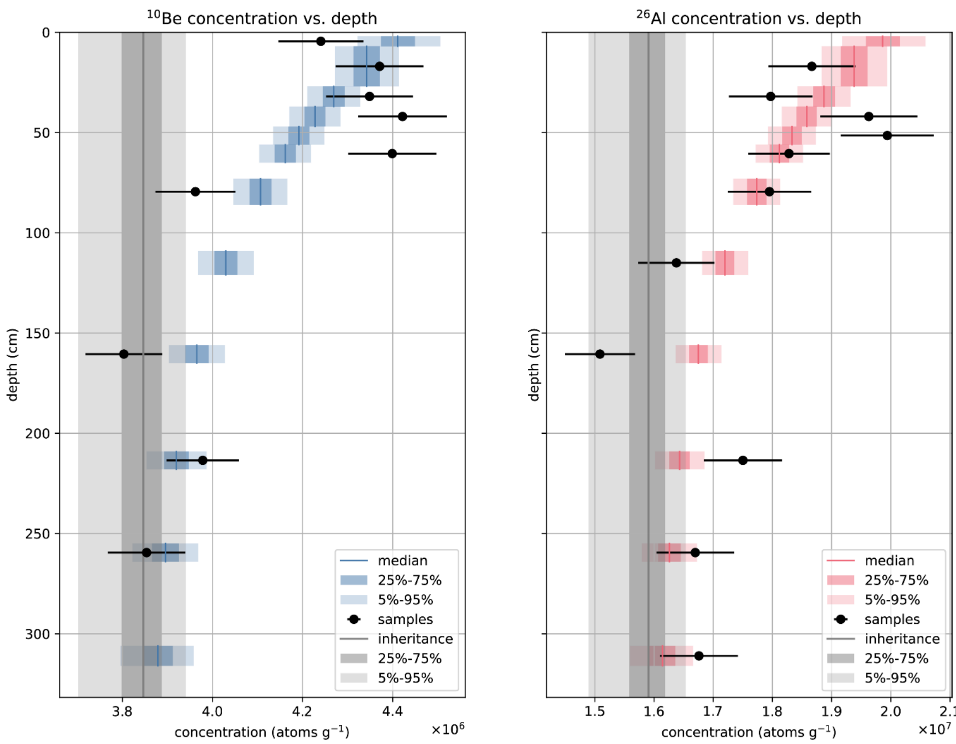


**Figure 8.** Pearse Valley (PV14-A) permafrost core depth profiles with measured $^{10}$Be and $^{26}$Al concentrations (black data points) with 1σ uncertainties for all samples between 0.02–3.16 m depth. Blue ($^{10}$Be) and red ($^{26}$Al) boxes show simulated nuclide concentrations at each sample depth.

### 4.3 Cosmogenic nuclide depth profiles at Lower Wright Valley

The $^{10}$Be and $^{26}$Al depth profiles from the permafrost core and overlying active-layer used for depth profile modelling at Lower Wright Valley is shown in Fig 11. The Lower Wright Valley $^{10}$Be and $^{26}$Al concentration profiles exhibit near-constant concentrations with depth with average values of 4.01 ± 0.10 x10$^{6}$ atoms g$^{-1}$ and 2.08 ± 0.08 x10$^{7}$ atoms g$^{-1}$, respectively. The absence of a discernible exponential attenuation indicates all sediments in the depth profile are either vertically mixed after deposition, or are sufficiently young so that post-depositional nuclide production is negligible relative to inheritance.

The depth profile model does not work well for non-attenuating profiles and usually fails to give well-constrained results. The modelled nuclide concentration depth profiles do not fit within the 5$^{th}$ to 95$^{th}$ percentile for our measured concentrations in the Lower Wright Valley depth profile (Fig. 11). The solutions yield most probable $^{10}$Be and $^{26}$Al inheritance concentrations of 4.03 x

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$10^6$ and 2.06 x $10^7$ atoms $g^{-1}$, respectively (Fig. 11; Fig. S3). Our simulations yield the depositional age of the permafrost at 4.4 $^{+8.2}$ / $_{-4.2}$ ka (5th to 95th percentile), and an erosion rate of 0.2 $^{+0.18}$ / $_{-0.18}$ cm $ka^{-1}$ (Fig. S3).

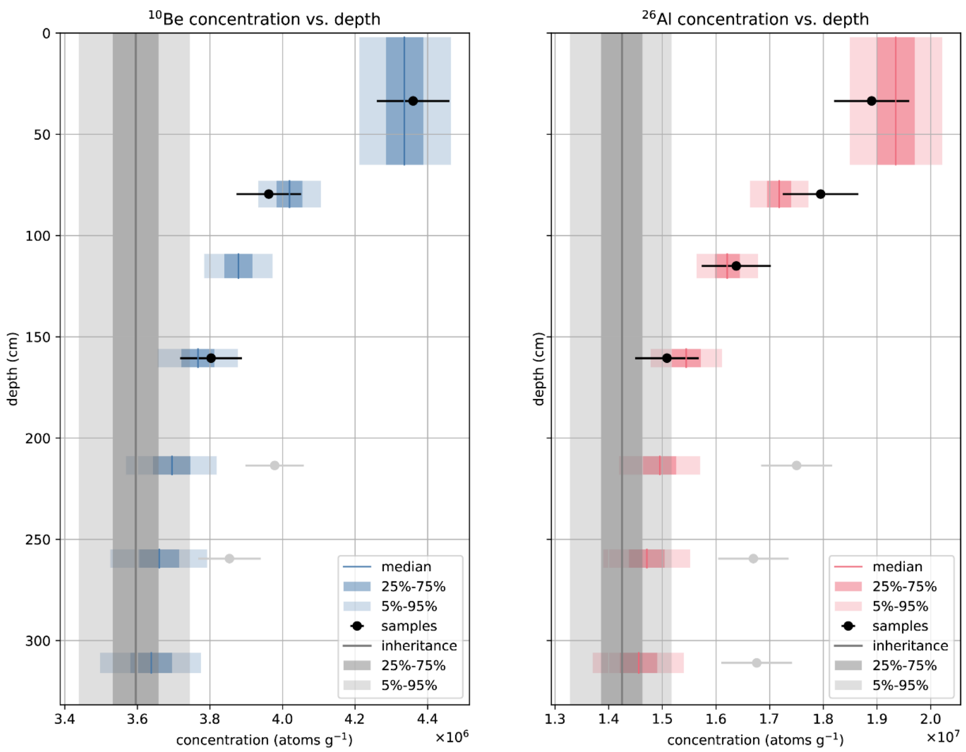


**Figure 9.** Pearse Valley (PV14-A) permafrost core depth profiles with measured $^{10}Be$ and $^{26}Al$ concentrations (black data points) with 1σ uncertainties. For all samples between 0.02–0.65 m depth, we used the average concentration of all five $^{10}Be$ and $^{26}Al$ measurements to represent the effect of cryoturbation of sediments in the active-layer. Blue ($^{10}Be$) and red ($^{26}Al$) boxes show simulated nuclide concentrations at each depth. $^{10}Be$ and $^{26}Al$ concentrations (grey data points) below 2.09 m were not included in the model.

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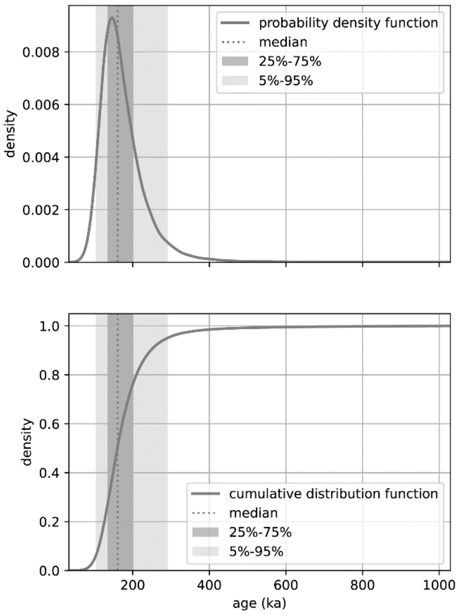


**Figure 10.** Probability density function, and cumulative distribution function for exposure age, using dual-nuclide depth profile modelling for PV14-A.

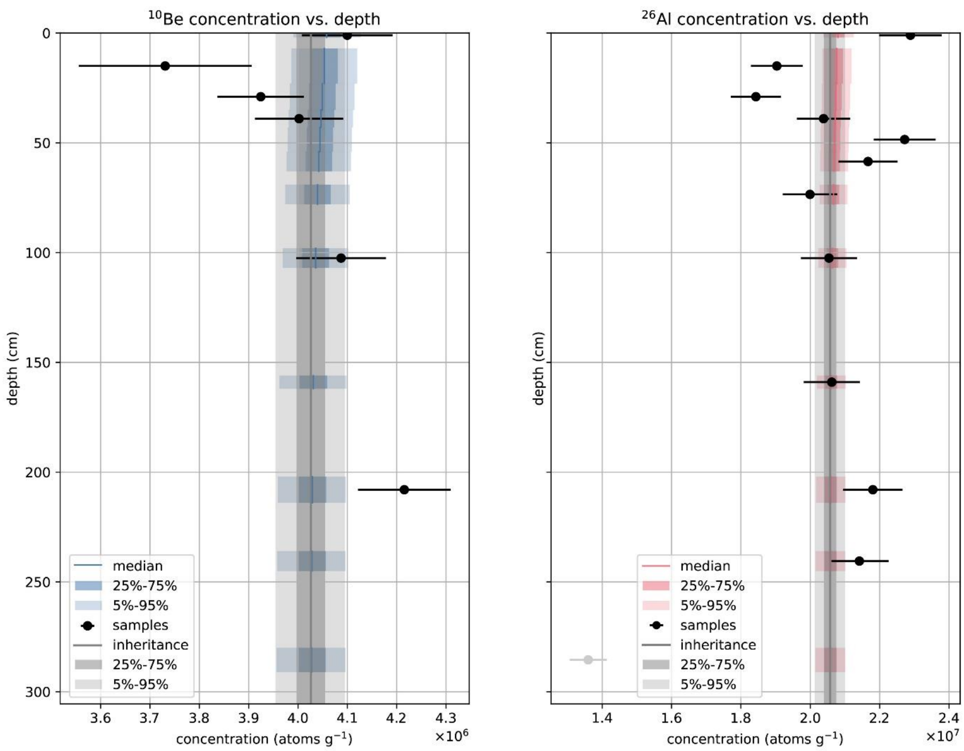

**Figure 11.** Lower Wright Valley (WV14-I) permafrost core depth profiles with measured $^{10}Be$ and $^{26}Al$ concentrations (black data points) with 1σ uncertainties. Blue ($^{10}Be$) and red ($^{26}Al$) boxes show simulated nuclide concentrations at each depth.

## 5 Discussion

### 5.1 Depositional and permafrost processes at Pearse Valley

Thin, patchy drift at Pearse Valley is a discontinuous peppering of boulders and cobbles superimposed on older loose sandy sediments, reworked clasts, and underlying permafrost sediments (Fig. 4). Surface cobble exposure ages confirm that this thin, patchy drift was deposited by a retreating cold-based Taylor Glacier during MIS 5a, and the MIS 5 / 4 transition, and corresponds with Taylor 2 Drift in central Taylor Valley. Depth profile modelling confirms that the sediments underlying Taylor 2 Drift, predate MIS 5. Undisturbed preservation of these relict surfaces is consistent with cold-based glacier activity described by Atkins (2013).

At the PV14-A permafrost core site, the present-day active-layer comprises a desert pavement surface and layer of loose vertically mixed sediments to a depth of ~0.37 m, positioned above ice-cemented permafrost sediments. The interface between this active-layer and the ice-cemented permafrost represents a sublimation unconformity. However, there is a discernible decrease in $^{10}Be$ concentration in the permafrost at ~0.65 m depth alongside an ice horizon. Such ice horizons are indicative of a paleosublimation unconformity, and the presence of a paleosublimation unconformity suggests the sediments experienced intervals that are warmer than present-day during or after deposition. This $^{10}Be$ offset cannot be explained by active-layer cryoturbation, as the present-day active-layer is only 0.37 m deep. Lapalme et al. (2017) suggested that in the upper ~0.5 m of a soil profile, ice can accumulate and sublimate due to changing ground surface temperature and humidity conditions. Below ~0.5 m depth, ice will progressively increase over time. Therefore, a paleosublimation unconformity can be inferred by the increase in ice content from 60 to 40 cm depth, which records the maximum predicted ice table depth (Lapalme et al., 2017). Therefore, we suggest the $^{10}Be$ offset between the sediments above and below 0.65 m represent a paleosublimation unconformity (Fig. 5a, 8) which probably occurred when the active-layer was thicker than present. Our depth profile model indicates that the upper section of the Pease Valley permafrost sediments (<1.65 m) was likely deposited at 180 $^{+20}/_{-40}$ ka, which does not contradict the exposure ages of the thin, patchy drift (~65–74 ka). Our measured nuclide concentrations at >2.09 m depth largely differ from the upper section and do not fit the simulated depth profile (Fig. 9). The higher nuclide concentrations in these samples, alongside the presence of several small ice lenses between

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1.57–1.87 m depth suggest these sediments were deposited during an earlier depositional event before ~180 ka.

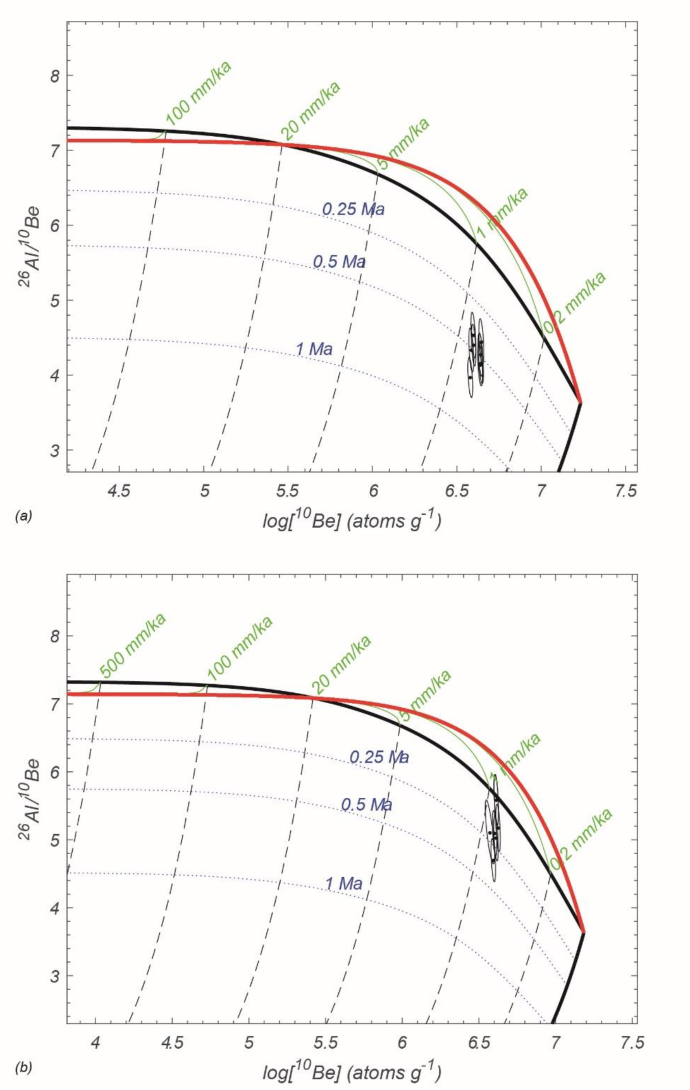


**Figure 12.** Two-isotope plot of Pearse Valley (a) and Lower Wright Valley (b) depth profiles. Nuclide concentrations with 1σ uncertainties, using the time-dependent LSDn scaling scheme of Lifton et al. (2014). Burial isochrons (dotted lines), decay trajectories (dashed), the exposure-erosion region (bounded by black and red lines), and steady-state erosion loci (green) are shown.

### 5.2 Exposure-burial history of sediments in Pearse Valley and Lower Wright Valley

While nuclide depth profiles indicate the most recent depositional history of the permafrost sediment, $^{26}Al/^{10}Be$ ratio data provides an additional insight regarding the total history of the sediment. When $^{26}Al/^{10}Be$ is plotted against $^{10}Be$ concentration on a two-isotope diagram (Fig. 12), a minimum total exposure-burial period can be inferred on the assumption that the sample experienced only one cycle of continuous exposure followed by continuous deep burial. At the Pearse Valley site, the two-isotope plot indicates that all sediments, regardless of their depth, have $^{26}Al/^{10}Be$ ratios ranging from 3.97 to 4.53, resulting in a minimum ~800 ka simple exposure (at zero erosion), and minimum ~400 ka burial, with a total exposure history of at least 1.2 Ma. At the Lower Wright Valley site, $^{26}Al/^{10}Be$ ratios for all samples range from 4.70 to 5.58, resulting in a minimum ~900 ka simple exposure, and minimum ~300 ka burial, with a total exposure history of at least 1.2 Ma.

Depth profile ages at both permafrost core sites represent the most recent phase of their depositional histories. For Pearse Valley this occurred at ~180 ka, and for Lower Wright Valley, where $^{10}Be$ and $^{26}Al$ concentrations do not attenuate, we estimate a maximum deposition age of <25 ka. This age represents the time required to change $^{10}Be$ and $^{26}Al$ above the initial inheritance level for near-surface samples by 5% - a change outside AMS $^{10}Be$ and $^{26}Al$ measurement error. However, our $^{26}Al/^{10}Be$ ratios at both sites suggest that these sediments have much longer total exposure-burial histories of at least 1.2 Ma, which most likely involves multiple recycling episodes of exposure, deposition, burial, and deflation prior to deposition at their current locations. Million-year exposure-burial recycling periods of sediments in the Dry Valleys was also observed in shallow (<1 m) pits from the Packard Dune fields in Victoria Valley (Fink et al., 2015).

### 5.3 Fluctuations of Taylor Glacier in Pearse Valley during MIS 5

Surface exposure ages of the cobbles perched on large boulders together with constraints from a best-fit depth profile age indicate Taylor 2 Drift sediments were deposited ~65–74 ka, during MIS 5a, and the MIS 5 / 4 transition, on the northern valley floor of central Pearse Valley, whereas the underlying permafrost sediments were deposited at ~180 ka or earlier. Our surface cobble geochronology is in agreement with the minimum U/Th ages for the extent of proglacial Lake Bonney, which suggest retreat of Taylor Glacier following MIS 5c and 5a advance (Fig. 13; Higgins et al., 2000a), and the tentatively dated western section of the rock glacier derived from $\delta^{18}O$ in buried ice northern Pearse Valley (Swanger et al., 2019). These data suggest Pearse Valley was largely or partially glaciated throughout MIS 5c and 5a.

Retreat of the Taylor Glacier lobe in Pearse Valley possibly continued after 65 ka. Timing of retreat after 65 ka, until the Last Glacial Maximum, where Taylor Glacier was at a minimum

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position, remains unknown. Advance and retreat cycles during MIS 5, the final retreat of Taylor Glacier during MIS 5a, and between the MIS 5 / 4 transition and the LGM for Taylor Glacier, could be better constrained by exposure dating more drift deposits with larger spatial coverage from Pearse Valley.

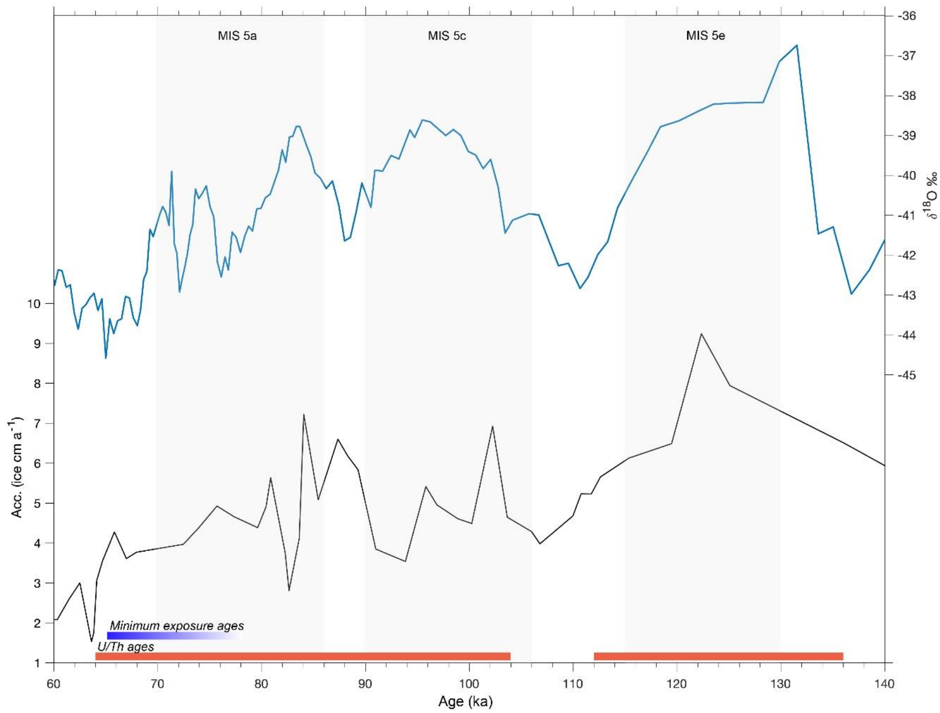


**Figure 13**. Snow accumulation rate determined from $^{10}Be$ (Acc. (ice cm $a^{-1}$)) and $\delta^{18}O$ record from Taylor Dome during MIS 5 (Steig et al., 2000). U/Th ages from algal carbonates (red bands, Higgins et al., 2000a) coincide with warm MIS substages 5e, 5c and 5a with increased accumulation rates at Taylor Dome. This is consistent with our minimum exposure ages (blue band) which show retreat of Taylor Glacier in Pearse Valley during MIS 5a, and the MIS 5 / 4 transition.

### 5.4 Advance and retreat of outlet and alpine glaciers during interglacial periods

Our new data has implications regarding the relationship between outlet and alpine glacier behaviour, regional paleoclimate and the extent of sea ice and open ocean conditions in the Ross Sea. Snow accumulation rate, atmospheric temperature, and duration of precipitation appear to be the major controls governing the advance and retreat of Taylor Glacier during previous warm intervals (Fig. 13). In central Taylor Valley, substage 5a and 5c sediments bury 5e sediments suggesting Taylor Glacier responds to regional changes over millennial

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timescales (Higgins et al., 2000a). The Taylor Glacier advances in central Taylor Valley during substages 5e, 5c and 5a correspond with increased accumulation in Taylor Dome (Higgins et al., 2000a; Steig et al., 2000). Our exposure ages indicate the retreat of Taylor Glacier in Pearse Valley occurred at ~65–74 ka during the MIS 5 / 4 transition, and is consistent with the retreat in central Taylor Valley. The presence of a lobe of Taylor Glacier in Pearse Valley throughout MIS 5 is likely linked to prolonged interglacial climate conditions. The interglacial-mode climate, where austral westerlies are in a poleward-shifted position for prolonged periods during MIS 5, is associated with periods where $CO_2$ concentrations were above ~230 ppm, the glacial-interglacial $CO_2$ threshold proposed by Denton et al. (2021).

Yan et al. (2021) suggested that peak accumulation rates occurred at ~128 ka in Southern Victoria Land and are associated with reduced sea ice and possibly retreat of the Ross Ice Shelf. The study suggested by ~125 ka, the Ross Ice Shelf had returned to a configuration comparable to present day. However, a reduction of sea ice may have enabled increased moisture delivery over Taylor Dome during MIS 5c and 5a. As Higgins et al. (2000a) suggested, increased precipitation over Taylor Dome during MIS 5a and 5c appears to have caused a subsequent readvance of Taylor Glacier. We acknowledge, this hypothesis is speculative and requires further testing of temperature, and atmospheric circulation in response to reduced sea ice extent and perhaps a reduction of the Ross Ice Shelf by climate models.

The duration of a warm interval which governs the extent of sea ice cover or open water in the Ross Sea, may in turn, influence moisture transport and accumulation on Taylor Dome and the Antarctic plateau. With temperatures predicted to be similar to the last interglacial in coming decades, on a multicentennial to millennial scale, anti-phased MIS 5 feedbacks may provide important analogues for future Antarctic ice loss (DeConto & Pollard, 2016; DeConto et al., 2021). While our geochronology suggests retreat of Taylor Glacier in Pearse Valley occurred during the MIS 5 / 4 transition, probably by a change in moisture regime and drying during MIS 5a, several uncertainties regarding advancing and retreating ice and associated processes in the Dry Valleys region and Ross Sea need to be addressed:

- The timing of advance and retreat cycles of outlet and alpine glaciers during substages 5e, 5c and 5a remain poorly constrained.
- The duration of warm intervals, bringing warm moist air to enable glacier advance is not well understood.
- The paucity of data in the Ross Sea regarding sea ice, ice shelf, and open ocean conditions during MIS 5 makes interpretation of the antiphase behaviour between advanced outlet and alpine glaciers in the Dry Valleys region, and increased open

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ocean in the Ross Sea difficult to quantify.

## 6 Conclusions

We applied cosmogenic nuclide analysis to surface cobbles and permafrost depth profiles to obtain the age of Taylor 2 Drift, and determine landscape evolution and associated processes in Pearse Valley. Our $^{10}$Be and $^{26}$Al derived surface exposure ages from cobbles emplaced on large boulders embedded in the valley floor of Pearse Valley located ~3 km from Taylor Glacier lobe give a minimum zero erosion age of ~65 to 74 ka for deposition of the thin, patchy drift, indicating that Taylor Glacier retreated from Pearse Valley during MIS 5 / 4 transition. Companion $^{10}$Be and $^{26}$Al depth profile modelling reveals a ~180 ka deposition age for near-surface permafrost deposits to 1.65 m depth that predates MIS 5. These data support antiphase behaviour between outlet and alpine glaciers in the Dry Valleys region and ice extent in the Ross Sea, and suggest a causal mechanism where cold-based glacier advance and retreat is controlled by moisture availability and drying, respectively due to ice retreat and expansion in the Ross Sea. Our work is consistent with geochronology from central Taylor Valley, supporting advance and retreat cycles of Taylor Glacier during MIS substages 5c and 5a (Higgins et al., 2000a), corresponding with increased accumulation at Taylor Dome (Steig et al. 2000). Our study highlights the need for better age constraints of alpine and outlet glaciers that advanced during MIS 5 in the Dry Valley region. In particular, for assessing the relationship between accumulation rate at Taylor Dome, and sea ice, ice shelf, and open ocean conditions in the Ross Sea during MIS 5.

The offset in $^{10}$Be concentrations at ~0.65 m depth in the Pearse Valley permafrost core, and presence of increased ice content reveals a paleosublimation unconformity, and suggests that these upper sediments have undergone active-layer cryoturbation. The permafrost >0.65 m depth in central Pearse Valley has been frozen for at least 65 ka, and perhaps ~180 ka based on our depth profile model, whereas, >2.09 m depth the depositional age of the sediment must be earlier than ~180 ka.

To compare processes of sediment evolution at Pearse Valley with a lower elevation, and more coastal environment, we also applied $^{10}$Be and $^{26}$Al nuclide analysis to permafrost depth profiles at Lower Wright Valley. While the current deposition at the latter site occurred more recently (<25 ka), total exposure-burial histories from the two sites consistently show these sediment repositories have experienced multiple glacial-interglacial cycles achieved through the recycling of sediments for at least 1.2 Ma.

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**Code availability**

The code used for depth profile modelling is available by request from the corresponding author.

**Data availability**

All data described in the paper are included in the Supplement.

**Author contributions**

JTHA, GSW, AA, and ND conducted the field work and sample collection. JTHA, DF, TF, and KW conducted the sample preparation and AMS analysis. AJH and JTHA developed the depth profile models. JTHA prepared the manuscript with contributions from all authors.

**Competing interests**

The authors declare that they have no conflict of interest.

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