"Thermodynamic vs. dynamic Arctic wintertime sea ice thickness effects" authors' responses to referee #1

Thank you for the useful and constructive comments. Authors' responses are in red.

This paper documents the estimation of thermodynamic and dynamical components of sea ice volume changes during the Arctic growth season. The authors have combined 3 data sets: the AWI-SMOS weekly sea ice thickness data, the Polar Pathfinder sea ice drift data and the authors new SLICE, brightness temperature derived thermodynamic ice growth estimate data. The final results presented are quantified and thoroughly compared to ice volume change studies. Such observational derived estimates of ice volume change are challenging to produce, and the authors must be commended at the efforts they have made. However, if this study is to be of use to the wider scientific community, much more information needs to be included. Readers of this paper must then be informed of the conceptual consequences of using the data presented. At the moment very limited information is included about the most appropriate interpretation of the presented results. From this reviewers point of view, this study is as much an accuracy assessment of the SLICE data as an estimate of the components of Arctic sea ice thickness change, though this aspect of the paper is mentioned very little. A more thorough description of the issues are described below.

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Thank you for the positive comments and constructive feedback. The SLICE methodology was initially validated in Anheuser et al. (2022) and this paper is an implementation of SLICE towards understanding the relative contributions of thermodynamic and dynamic sea ice effects in the Arctic. Favorable comparison between our results and those found in literature bolster confidence in this work. Additionally, we hope the responses below support this view.

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My main issue with this paper is the incorrect definition of uncertainty used that causes misleading claims within the results and discussion section. The equational form of (6) will give the interannual variability of the budget terms, with no information of the observational uncertainty given. When uncertainty is described for these observational estimates, the reader expects to see information included on how much we can trust these estimates based on how well defined the original measurements are. What role does this uncertainty have in the final budget calculations? The discussion section does attempt to discuss the role of observational uncertainty in the calculations of the paper, though these are compared to scarcely believable claims that the SLICE data has an uncertainty of less than 1mm/ per week. While the uncertainty in AWI-SMOS is plotted, no accurate information on how this plot was calculated is given, making it difficult to interpret. Uncertainty measurements are included

in the Pathfinder data set, and I suggest that similar plots need including too. Pathfinder uncertainties can often be as high as 30-50%. As the authors are the creators of the SLICE dataset, then I expect them to also include similar plots showing the uncertainty in this data also. Possible covariances within the data are not mentioned at all. This has the potential to be the highest source of uncertainty in the final values presented. While calculating covariances may be beyond the scope of the study, the possibility of the occurrence needs to be presented.

We have replaced our standard error approach with an uncertainty propagation approach. Please see the updated section below. A key aspect to this calculation is that uncertainty is reduced through temporal averaging in creation of the climatologies, similar to regridding of satellite altimeter data as in AWI CS2SMOS. In regards to Polar Pathfinder uncertainties, DeRepentigny et al. (2016) found the weekly sea ice motion vectors to have a 7% error and the Tschudi et al. (2020) lists a maximum ice motion error of 0.7 cm s-1. Lastly, we have updated the discussion section to reflect the new uncertainty section.

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Uncertainty in the individual weekly observations of thermodynamic, dynamic, advection and deformation effect can be calculated using a general formula for uncertainty in a function of several variables (Taylor, 1982):

$$\delta_q = \sqrt{\left(\frac{\partial q}{\partial x}\delta_x\right)^2 + \dots + \left(\frac{\partial q}{\partial z}\delta_z\right)^2},\tag{1}$$

where q is the computed value; x, \dots, z are independent and random inputs to that computed value and $\delta_x, \dots, \delta_z$ are those inputs associated uncertainties. Applying Eq. 1 to the terms as described in Section 3, we have:

$$\delta_{thm} = \sqrt{\delta_{SLICE}^2 + \left(\frac{thermodynamic\ growth}{CS2SMOS}\delta_{CS2SMOS}\right)^2}$$
 (2)

$$\delta_{dyn} = \sqrt{\left(\frac{1}{\Delta t}\sqrt{2}\delta_{CS2SMOS}\right)^2 + \delta_{thm}^2} \tag{3}$$

$$\delta_{adv} = \sqrt{\left(\frac{u}{\Delta x}\sqrt{2}\delta_{CS2SMOS}\right)^2 + \left(\frac{\partial CS2SMOS}{\partial x}\delta_u\right)^2 + \left(\frac{v}{\Delta y}\sqrt{2}\delta_{CS2SMOS}\right)^2 + \left(\frac{\partial CS2SMOS}{\partial y}\delta_v\right)^2}$$
(4)

$$\delta_{def} = \sqrt{\delta_{dyn}^2 + \delta_{adv}^2},\tag{5}$$

where δ_{thm} , δ_{dyn} , δ_{adv} , δ_{def} , δ_{SLICE} , $\delta_{CS2SMOS}$, δ_u , and δ_v are uncertainties in the thermodynamic growth, dynamic effect, advection effect, deformation effect, SLICE, CS2SMOS thickness, x direction component of sea ice motion vector, and y direction component of sea ice motion vector, respectively; u is the x direction component of sea ice motion vector; v is the v direction component of sea ice motion vector; v is time step size; and v and v are the grid box size. These uncertainty formulas do not account for covariances between the input terms. Though covariances may be present across the input data, inclusion of their effects on uncertainty is outside the scope of this work.

The uncertainty in SLICE is taken from Anheuser et al. (2022), who report SLICE to have a thermodynamic growth mean bias of 4×10^{-4} m d^{-1} and standard deviation bias of 2.2×10^{-3} m d^{-1} when compared against ice mass balance buoy data. Here we use this standard deviation as SLICE uncertainty. The analysis presented in Anheuser et al. (2022) does not include

the effect of uncertainty in initial sea ice thickness, so we add the second term on the right side of manuscript Eq. 3 to account for the uncertainty in CS2SMOS sea ice thickness. Tschudi et al. (2020) lists a maximum ice motion vector error of 0.7 cm s⁻¹, which we use here for the uncertainty in the ice motion vector components. The uncertainty in CS2SMOS is calculated for each week and available in the data product. Lastly, the time step is one week and grid cell size is 25,000 m. Using these inputs, we calculate uncertainty in the thermodynamic growth, dynamic effect, advection effect, deformation effect terms at each time step and grid cell location.

When the terms are averaged to form Fig. 2, the uncertainties are reduced through the averaging. Applying 1 to an averaging operation, we have the following:

$$\delta_{mean} = \sqrt{\left(\frac{1}{N}\delta_1\right)^2 + \dots + \left(\frac{1}{N}\delta_N\right)^2},\tag{6}$$

where δ_{mean} is the uncertainty of the mean; N is the number of samples; and $\delta_1, \dots, \delta_N$ are the individual uncertainties of each sample. Figure 5 shows the uncertainty of the mean for each of the processes studied here. The effect with the greatest uncertainty is deformation as it is a summation of uncertainties in the other terms due to it being calculated as a residual from those other terms.

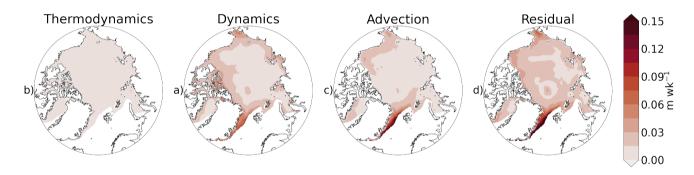


Figure 1. Uncertainty calculated per Section for each grid cell during wintertime from late 2010 through early 2021 (except the winter of 2011-2012) sea ice thickness changes due to a) thermodynamic growth, b) dynamic effect, c) advection effects and d) deformation effects. Uncertainty increases with a decrease in latitude as the number of weeks with ice cover and number of satellite overpasses decreases.

The values given in this paper with claims such as 'error is highest in the East Greenland, Barents and Kara Seas, where motion vectors are largest and most variable' are telling us only about which regions are most variable from week to week, and from year to year. Indeed it may be possible that such regions have lower observational uncertainty than regions given here that are said to have low uncertainty, when the metric supplied by this study are only indicate that these region have a low interannual variability, or indeed constant values over the growth season.

The updated uncertainty methodology should give more support for this discussion and has been updated per the new uncertainty estimates.

The first main issue leads into the second limitation of this study: the limited plotted data supplied in the paper. Arctic sea ice is widely described to have large interannual variability, for example in minimum extent and volume. The results presented and discussed in this paper are mostly from 10 year climatologies, that by definition will remove all such variability. It is thus then difficult to assess the accuracy and role of the different budget components. Each yearly total is shown in an additional plot in the appendix, and a table is supplied quantifying the key regions. No example weekly calculations are presented at all, which is a crucial omission as this is the time scale of the original calculations. To adequately document the calculations performed for this study, and to be of use to modelling studies, time series of the regions presented can be included. An example figure showing the input and output data from a single week will be a very useful inclusion. Uncertainty data for the shorted timescales needs including too. The final data presented in this paper comes from short time scale calculations. Uncertainty in these calculations comes from these timescales also.

The aim of this work is to present a mean climatology over this time period rather than measurements of individual weeks. As discussed in the uncertainty section, the temporal averaging reduces uncertainties while measurements from individual weeks are much more uncertain. In our judgement, the smallest time step containing useful data is monthly. We have added monthly regional line plots and monthly basin-wide spatial plots from the 2019-2020 winter to the manuscript to pair with the MOSAiC plot (Figs. 2, 3), which we have also changed to monthly resolution rather than weekly to remain consistent (Fig. 4). The monthly data is created by taking the mean of the weekly data found within each month, sorted by the date on the first day of the week.

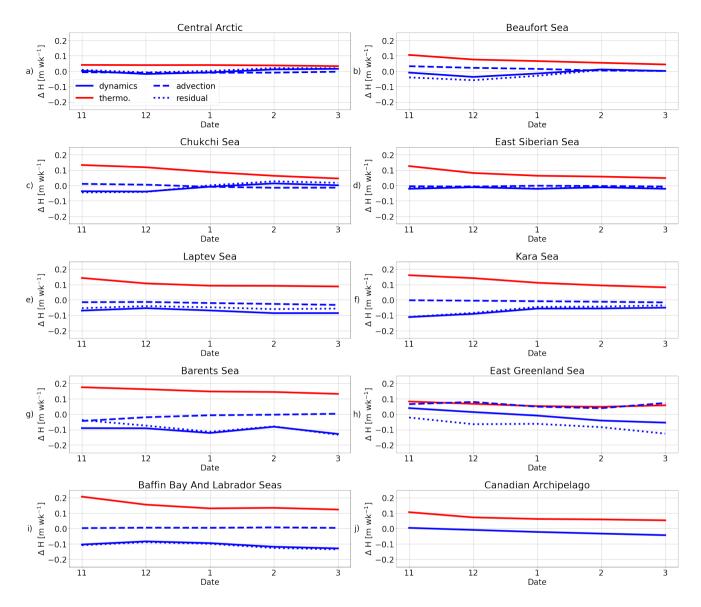


Figure 2. Mean monthly time series of dynamic effect (blue), thermodynamic growth (red), advection effect (blue dash dot) and deformation effect (blue dotted) for a) the East Greenland Sea, b) the Barents Sea, c) the Kara Sea, d) the Laptev Sea, e) the East Siberian Sea, f) the Chukchi Sea, g) the Beaufort Sea, h) the Canadian Islands, i) the Central Arctic and j) the entire Arctic. Thermodynamic growth typically declines through the season as thickness increases and dynamic effect patterns are variable from region to region.

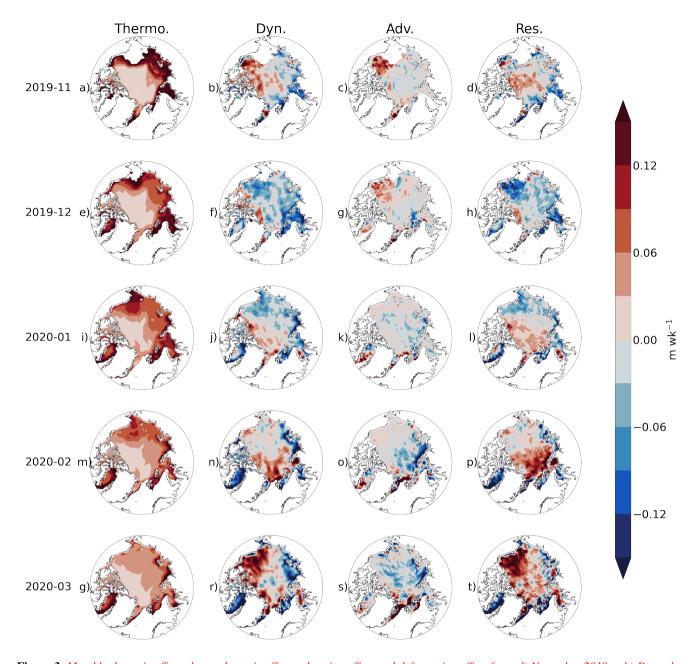


Figure 3. Monthly dynamic effect, thermodynamic effect, advection effect and deformation effect for a-d) November 2019, e-h) December 2019, i-l) January 2020, m-p) February 2020 and q-t) March 2020. Thermodynamic effect decreases while dynamic and deformation effects increase through the growth season.

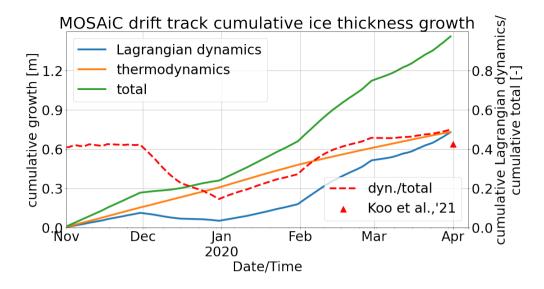


Figure 4. Cumulative Lagrangian dynamics (residual effects), thermodynamic and total sea ice thickness growth (primary vertical axis) and cumulative Lagrangian dynamics over cumulative total sea ice thickness growth (secondary vertical axis) along the MOSAiC drift track determined using the methodology described here and averaged over each month. The red triangle represents cumulative Lagrangian dynamics over cumulative total growth over a similar area reported by Koo et al. (2021) who used ICESat-2 to determine dynamics vs. thermodynamics along the MOSAiC drift track. Lagrangian dynamics accounts for over half of all thickness growth by 1 April 2020.

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Another concern is due to the definition of the budget equations and the interpretation of this definition. There are three main sources of volume change of sea ice used: rate of change of thickness data, the advection of sea ice thickness, and the new brightness temperature derived thermodynamic ice growth. The final deformation-based estimates of ice volume change is given as the product of the three input data sets. This is conceptually fine, and the results of this method will be of use to the scientific community. However, the authors must tell the reader that it is their own interpretation that gives the difference between the input data as deformation. There are many points within the paper where the deformation values are listed alongside the other components as an additional data value. What is a more accurate definition is thus:

The rate of change in volume and advection estimates are compared to a novel source of data on the thermodynamic change in sea ice thickness. There are relatively large differences between these two data, which are presented here. When considering the source in this difference as from sea ice deformation, these results are the consequence.

This is a more accurate and conceptually more useful description for the wider community. As mentioned above, the final deformation values are thus highly dependent on the uncertainty in the input data. The magnitude of the deformation estimates must be compared to the magnitude of any uncertainty.

The authors agree with this critique on a fundamental level. We have added emphasis to the fact that defining the residual as deformation is our own interpretation rather than a direct measurement of deformation itself. One point of disagreement here is that the differences between the overall volume change, advection, and thermodynamic growth are large. All are of the same order of magnitude.

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The previous concern brings up two notable omissions from the data presented in this paper. First the lack of ice concentration data to convert sea ice thickness to sea ice volume, and then the use of the ice drift data to create an estimation of sea ice divergence. The two data sources can then be used to capture wider scale deformation data and compared to the emergent deformation. Indeed, divergence rates are discussed in this paper, but without presented calculations on the deformation rates then this is purely supposition. The authors comment that the use of Pathfinder for vector calculus applications is not appropriate, if this is so then how can the results of this study be accurate?

The SLICE methodology is only viable in areas with 95% or greater sea ice concentration due to open water contamination of the passive microwave brightness temperatures and thus our analysis only applies to times when a given grid cell contains higher than this threshold sea ice concentration. While this was mentioned in the discussion and in Anheuser et al. (2022), it was an unintentional omission on our part to not mention this in the data and methodology sections. We have added the below figure showing the percent of the total study time each grid cell is found to meet this criteria and add a 50% threshold in this metric for each grid cell above which results will not be reported. We certainly do discuss divergence as we suspect that to be the cause of the deformation fields and have added a statement that this is a speculative explanation on our part rather than a measurement of divergence itself in the flow field. Long term divergence in the Polar Pathfinder data is found to be noisy, whereas a long term average of the flow itself (and thus advection) does not exhibit the same noise.

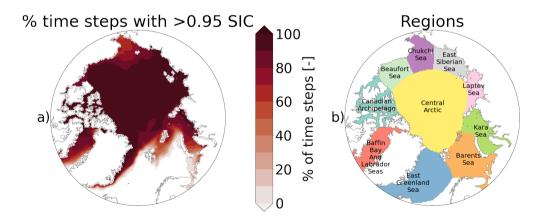


Figure 5. Plots showing a) percent of total time steps with 95% or greater sea ice concentration and b) location, extent and corresponding name of regions used in Section 4.

Specific points: L3 and 67 This is not the first such study, see Ricker et al. (2021)

Thank you for this reference. It is highly applicable and we added it to the introduction and discussion.

L 80 Previous studies using similar methods to those presented in this paper (Ricker et al., 2021; Holland and Kimura, 2016) both also use an ice concentration data set too. This data is crucial in high divergence/advection areas (say the Greenland Sea). Ice thickness products using the Radar Freeboard method (as presented here) will not capture the loss in ice volume due to ice lead opening.

We have added these references. The sea ice concentration issue is discussed above and we have added this aspect in the 150 methodology.

L 153 What is meant exactly here as 'Deformation effect'? Is it the divergence term that is listed as due to deformation in the previous sentence?

155 Correct.

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Equations 4 and 5 and the list on lines 174 - 178 Please reformat 'dynamic effect' and others using the normal text rather than equation text format.

160 Done.

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L 206 This measure of uncertainty appears to be the interannual variability - the standard deviation of measurements over 10 years of weekly data. Check this. Are interannual variability and uncertainty in the mean related?

The uncertainty and associated discussion have been updated per above.

L 211 This is a correct reason for uncertainty, but it is equated to a map of variability.

This statement is still true even in the new uncertainty methodology and has been updated to reference the new plots.

Figure 1. It will be good to see the mean total change in volume on this plot too as it is the main input data for these calculations.

We have added a plot of this.

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L 239 It is important to express for this experimental data set, that the results obtained in this paper suggest that 'it experiences a decrease in thickness due to deformation via lead formation'. The data presented here does not observe lead formation, so this must be presented as the authors explanation of the observed results. This is especially true for the interpretation of a 10 year climatology, and even more true for budget calculations that do not incorporate ice concentration data, and thus will not capture the spatial divergence of ice within the time derivative section of equation (2). This rephrasing is crucial to perform throughout the paper.

We have added an emphasis to the fact that this explanation and others like it are the interpretation of the authors.

L 243 'as the flow pattern deposits ice and leads to ridging' again this is not directly observed by the data presented and must be presented as the authors interestation.

We have added an emphasis to the fact that this explanation and others like it are the interpretation of the authors.

L 248 'negative effects from both advection and deformation' again misleading. The advection is an emergent signal present in the data. The deformation is entirely from the resultant difference in the data, and its role as deformation is entirely the authors interpretation. This must be expressed as such.

We have added an emphasis to the fact that this explanation and others like it are the interpretation of the authors.

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L 249 'location and strength of the Beaufort Gyre' has this been quantified anywhere? Or is it up to the reader to interpret this entirely from the plotted arrow vectors? If it is from the arrow vectors, please indicate this in the text, as this is a highly interpretive method and not at all accurate. There are other studies that presented data on the Beaufort Gyre and may be a useful inclusion here.

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While making quantitative connections between the strength or location of the Beaufort Gyre and the effects demonstrated here is out of scope of this study, we thought it reasonable to mention this potential connection. We have replaced "the location and strength of the Beaufort Gyre" with "interannual variability".

L 251 More information on figure 4 is required. What time scale are these plots calculated? The ration taken at weekly time scales, or over the whole 10 years?

The caption states this is a wintertime mean for 2010-2021 but we have added this to the body of the text. Our first submission shows the ratio is the ten year mean of dynamic effect over the ten year mean of thermodynamic growth, though it may be more useful to the take the ratio at weekly time scales and then average.

Figure 6: More information is required for this figure caption. Which data come from this study and which are from the MOSAiC campaign? What timescale are the various data on?

All of the data is from the weekly results demonstrated by this paper except the red triangle with is the Koo et al. (2021) result. We have added "determined using the methodology described here" to the first sentence.

L 284 Again, ridging is not observed in this study. What is shown here, is entirely an observed change in sea ice volume that cannot be accounted for using advection or thermodynamics growth estimates. This needs to be written as such before the results can be of use.

We have added an emphasis to the fact that this explanation and others like it are the interpretation of the authors.

L 292, If divergence is discussed, then it really needs to be calculated too.

The divergence field calculated using the Polar Pathfinder vectors is much more noisy than the vectors themselves.

L 296. It is the other way around. Positive deformation emerges, that can be accounted for through ridging.

230 We have updated this.

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235

L 305, why are they not suitable? Has this been tried? Are the results less certain than the estimation presented in this paper?

We have tried this approach, but noise in the divergence field makes the results less useful.

L 356 fix equation number.

We have updated this.

L 366 Can this claim be backed up? I see no discussion on the uncertainty of the SLICE data. As this is a new data sourced, then it needs to be thoroughly analysed for its accuracy and reliability.

	The uncertainty discussion has been updated per above.
245	L 384 the observational uncertainty here is compared the data variability. These are not the same thing.
	The uncertainty discussion has been updated per above.
250	Figure 7. How are these values calculated? Is it the mean uncertainty value taken from the data product?
	The uncertainty discussion has been updated per above.
	L 393 I find an uncertainty of 1mm/per week for this data source to be somewhat to good to be true.
255	The uncertainty discussion has been updated per above.
	L 407. So ice concentration data was used in this study? What data is this?
260	See our response on the sea ice concentration issue above.
	L 417 A calculation of the resultant uncertainty in advection can thus be calculated. This needs to be quantified.
	The uncertainty discussion has been updated per above.
265	L 442, Needs to be: where the SLICE thermodynamic growth data can only account for (a third possibly) of, on average, the total observed ice thickness change.
270	We have added emphasis to the fact that we use residuals here and our explanations are based on our own interpretation of these residuals.
	L 447 – this data is not yearly. This is the week to week data surely?
	Yes, we have updated this to "weekly".
275	L 453 This claim cannot be made as this is the only uncertainty shown. The uncertainty in Pathfinder and SLICE needs to be shown in similar depth.

The uncertainty discussion has been updated per above.

280 L 455 this is a comparison between uncertainty and variability and highly misleading.

The uncertainty discussion has been updated per above.

Appendix - there is no text in the appendix. Add this figure to the main body of text.

285

We have moved this figure to the results section where it is briefly introduced.

References

- Anheuser, J., Liu, Y., and Key, J.: A simple model for daily basin-wide thermodynamic sea ice thickness growth retrieval, The Cryosphere, 16, 4403–4421, https://doi.org/10.5194/tc-16-4403-2022, 2022.
 - DeRepentigny, P., Tremblay, L. B., Newton, R., and Pfirman, S.: Patterns of Sea Ice Retreat in the Transition to a Seasonally Ice-Free Arctic, J. Climate, 29, 6993 7008, https://doi.org/10.1175/JCLI-D-15-0733.1, 2016.
 - Holland, P. R. and Kimura, N.: Observed Concentration Budgets of Arctic and Antarctic Sea Ice, Journal of Climate, 29, 5241–5249, https://doi.org/10.1175/jcli-d-16-0121.1, 2016.
- Koo, Y., Lei, R. B., Cheng, Y. B., Cheng, B., Xie, H. J., Hoppmann, M., Kurtz, N. T., Ackley, S. F., and Mestas-Nunez, A. M.: Estimation of thermodynamic and dynamic contributions to sea ice growth in the Central Arctic using ICESat-2 and MOSAiC SIMBA buoy data, Remote Sensing of Environment, 267, https://doi.org/10.1016/j.rse.2021.112730, 2021.
 - Ricker, R., Kauker, F., Schweiger, A., Hendricks, S., Zhang, J. L., and Paul, S.: Evidence for an Increasing Role of Ocean Heat in Arctic Winter Sea Ice Growth, Journal of Climate, 34, 5215–5227, https://doi.org/10.1175/jcli-d-20-0848.1, 2021.
- Taylor, J. R.: An introduction to error analysis: The study of uncertainties in physical measurements, University Science Books, 20 Edgehill Rd., Mill Valley, CA 94941, 1982.
 - Tschudi, M. A., Meier, W. N., and Stewart, J. S.: An enhancement to sea ice motion and age products at the National Snow and Ice Data Center (NSIDC), Cryosphere, 14, 1519–1536, https://doi.org/10.5194/tc-14-1519-2020, 2020.

"Thermodynamic vs. dynamic Arctic wintertime sea ice thickness effects" authors' responses to referee #2

Thank you for the useful and constructive comments. Authors' responses are in red.

Review of "A climatology of thermodynamic vs. dynamic Arctic wintertime sea ice thickness effects during the CryoSat-2 era" by Anheuser et al.

This paper aims to quantify the components of dynamic and thermodynamic sea ice growth in the Arctic during the winter season from 2010 to 2021. The authors make use of three products: 1) The SLICE model (Stefan's Law Integrated Conducted Energy) providing thermodynamic ice growth and introduced in an earlier paper ("A climatology of thermodynamic vs. dynamic Arctic wintertime sea ice thickness effects during the CryoSat-2 era") by the authors of this study. 2) The AWI CS2SMOS sea ice thickness data set providing weekly sea ice thickness grids for the Arctic, and here used to derive total sea ice thickness growth. And 3) NSIDC Pathfinder sea ice motion data to derive advection of sea ice.

The topic is relevant and the approach using the SLICE model is interesting. In general, I think this paper could be interesting and a benefit for the sea ice and climate community. But from my point of view the study is lacking further information on methods (and may be required corrections), and a sound uncertainty estimation. I have a few major concerns:

Thank you for the positive comments.

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The SLICE thermodynamic ice growth is subtracted from the total growth in sea ice thickness using the CS2SMOS product. Are the authors aware that the CS2SMOS sea ice thickness for each grid cell does not include open water? So hypothetically assuming that within a grid cell (with pure level ice) sea ice diverges, forming leads, the averaged ice thickness will be the same in CS2SMOS (especially for the CryoSat-2 domain). In other words, thickness=0, is not used for averaging. But I cannot see that this is considered in the current approach.

The SLICE methodology is only viable over 95% or greater sea ice concentration due to open water contamination of the passive microwave brightness temperatures and thus our analysis only applies to times when a given grid cell contains higher than this threshold sea ice concentration. While this was mentioned in the discussion and in the SLICE paper, it was an unintentional omission on our part to not mention this in the data and methodology sections. This condition means that even without considering the effects of changing sea ice concentration, the maximum error in a given term will be 5%. We have also

added the below figure showing the percent of the total study time each grid cell is found to meet this criteria and add a 50% threshold in this metric for each grid cell above which results will not be reported.

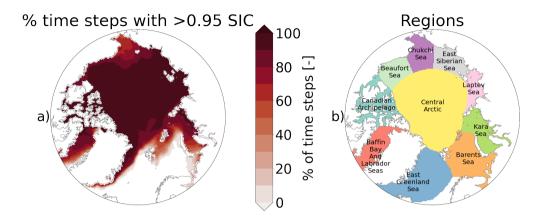


Figure 1. Plots showing a) percent of total time steps with 95% or greater sea ice concentration and b) location, extent and corresponding name of regions used in Section 4.

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Partly related to 1): I wonder how new ice formation in leads is affecting the overall findings in this study. It is not clear to me how this is handled in this study. SLICE does not seem to consider new ice formation in leads or am I wrong?

You are correct, we have no way of capturing new ice within a gridcell using the methodology of this paper. With a 95% sea ice concentration threshold, this will likely not be a significant source of volume generation. Nonetheless, have added a description of this issue to the discussion.

The way uncertainties are considered in this study does not seem sound, or at least needs further explanation. I would assume that the uncertainty of the climatological mean as calculated here (Eq. 6) is mostly affected by temporal (interannual and seasonal) variability. This needs some improvement from my point of view. Moreover, the SLICE uncertainty for the weekly thermodynamic ice growth in most of the Central Arctic is close to 0, which does not seem realistic, also considering the comparison with independent data sets in Anheuser et al. (2022).

A comprehensive uncertainty analysis is important here, since different input products are used, where either of them adds to the uncertainty budget. Uncertainty of the sea ice motion product is barely mentioned, but especially in the Fram Strait I believe this can lead to significant errors in the final retrievals, also since ice thickness is very heterogenous there.

These are good constructive critiques. We have replaced our standard error approach with an uncertainty propagation approach. Please see the updated section below. A key aspect to this calculation is that uncertainty is reduced through temporal averaging in creation of the climatologies, similar to regridding of satellite altimeter data as in AWI CS2SMOS. In regards to

Polar Pathfinder uncertainties, DeRepentigny et al. (2016) found the weekly sea ice motion vectors to have a 7% error and the Tschudi et al. (2020) lists a maximum ice motion error of 0.7 cm s-1. Lastly, we will update the discussion section to reflect the new uncertainty section.

55 Uncertainty in the individual weekly observations of thermodynamic, dynamic, advection and deformation effect can be calculated using a general formula for uncertainty in a function of several variables (Taylor, 1982):

$$\delta_q = \sqrt{\left(\frac{\partial q}{\partial x}\delta_x\right)^2 + \dots + \left(\frac{\partial q}{\partial z}\delta_z\right)^2},\tag{1}$$

where q is the computed value; x, \dots, z are independent and random inputs to that computed value and $\delta_x, \dots, \delta_z$ are those inputs associated uncertainties. Applying Eq. 1 to the terms as described in Section 3, we have:

$$\delta_{thm} = \sqrt{\delta_{SLICE}^2 + \left(\frac{thermodynamic\ growth}{CS2SMOS}\delta_{CS2SMOS}\right)^2}$$
(2)

$$\delta_{dyn} = \sqrt{\left(\frac{1}{\Delta t}\sqrt{2}\delta_{CS2SMOS}\right)^2 + \delta_{thm}^2} \tag{3}$$

$$\delta_{adv} = \sqrt{\left(\frac{u}{\Delta x}\sqrt{2}\delta_{CS2SMOS}\right)^2 + \left(\frac{\partial CS2SMOS}{\partial x}\delta_u\right)^2 + \left(\frac{v}{\Delta y}\sqrt{2}\delta_{CS2SMOS}\right)^2 + \left(\frac{\partial CS2SMOS}{\partial y}\delta_v\right)^2}$$
(4)

$$\delta_{def} = \sqrt{\delta_{dyn}^2 + \delta_{adv}^2},\tag{5}$$

where δ_{thm} , δ_{dyn} , δ_{adv} , δ_{def} , δ_{SLICE} , $\delta_{CS2SMOS}$, δ_u , and δ_v are uncertainties in the thermodynamic growth, dynamic effect, advection effect, deformation effect, SLICE, CS2SMOS thickness, x direction component of sea ice motion vector, and y direction component of sea ice motion vector; respectively; u is the x direction component of sea ice motion vector; v is the v direction component of sea ice motion vector; v is time step size; and v and v are the grid box size. These uncertainty formulas do not account for covariances between the input terms. Though covariances may be present across the input data, inclusion of their effects on uncertainty is outside the scope of this work.

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The uncertainty in SLICE is taken from Anheuser et al. (2022), who report SLICE to have a thermodynamic growth mean bias of 4×10^{-4} m d^{-1} and standard deviation bias of 2.2×10^{-3} m d^{-1} when compared against ice mass balance buoy data. Here we use this standard deviation as SLICE uncertainty. The analysis presented in Anheuser et al. (2022) does not include the effect of uncertainty in initial sea ice thickness, so we add the second term on the right side of manuscript Eq. 3 to account for the uncertainty in CS2SMOS sea ice thickness. Tschudi et al. (2020) lists a maximum ice motion vector error of 0.7 cm s⁻¹, which we use here for the uncertainty in the ice motion vector components. The uncertainty in CS2SMOS is calculated for each week and available in the data product. Lastly, the time step is one week and grid cell size is 25,000 m. Using these inputs, we calculate uncertainty in the thermodynamic growth, dynamic effect, advection effect, deformation effect terms at each time step and grid cell location.

When the terms are averaged to form Fig. 2, the uncertainties are reduced through the averaging. Applying 1 to an averaging operation, we have the following:

$$\delta_{mean} = \sqrt{\left(\frac{1}{N}\delta_1\right)^2 + \dots + \left(\frac{1}{N}\delta_N\right)^2},\tag{6}$$

where δ_{mean} is the uncertainty of the mean; N is the number of samples; and $\delta_1, \dots, \delta_N$ are the individual uncertainties of each sample. Figure 2 shows the uncertainty of the mean for each of the processes studied here. The effect with the greatest uncertainty is deformation as it is a summation of uncertainties in the other terms due to it being calculated as a residual from those other terms.

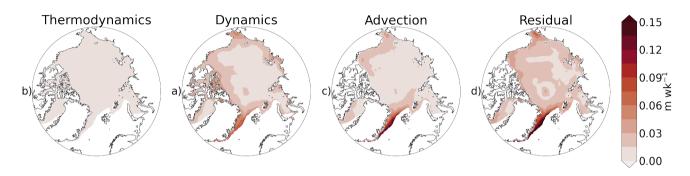


Figure 2. Uncertainty calculated per Section for each grid cell during wintertime from late 2010 through early 2021 (except the winter of 2011-2012) sea ice thickness changes due to a) thermodynamic growth, b) dynamic effect, c) advection effects and d) deformation effects. Uncertainty increases with a decrease in latitude as the number of weeks with ice cover and number of satellite overpasses decreases.

Given these points, I suggest major revisions are needed.

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Specific comments: L51 & 55: AEM (airborne electromagnetic) sounding measures the sea ice thickness, not freeboard (can be retrieved only indirectly).

Thank you for pointing this out, we have made this change.

L66: There are some recent studies that already investigated the ice growth components and should be mentioned and cited here: e.g.: Petty et al. (2018), Ricker et al. (2021). The latter also compared observational ice growth retrievals with model outputs.

We have added these references and included a comparison with our results where appropriate.

L73-75: Why is CS2SMOS not mentioned here (and in the entire introduction) as it is used in thus study?

This is a good point. We have updated the section to reference the CS2SMOS product rather than only CryoSat-2.

L98: I suggest providing numbers here for the footprint (e.g. 300 m (along track) x 1600 m (across track)).

105 We have added this information.

L155: "so to can" ... rewording

We have updated this.

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L 174: Would it not be correct to use a three-point linear regression over [i-1,i+1], centered at "i"? Otherwise, the gradient is not centered on the target week "i", but in between "i" and "i+1".

We have tried this methodology and though it doesn't appear to significantly change the results, we have implemented this in the next submission.

L193-194: "The most significant negative advection effects, less than 0.04 m wk-1, ...". It is misleading when speaking of negative effects but then stating a "less than" a positive number, which could be again a positive value.

We have updated this wording.

L197: Is negative deformation = lead formation?

Yes, in our interpretation a negative deformation effect is explained by lead formation and new ice filling those leads, thereby reducing the thickness of the ice within a grid box.

L465: Please state the version number of CS2SMOS. For most recent data and version history: https://spaces.awi.de/display/CS2SMOS/CryoSat-SMOS+Merged+Sea+Ice+Thickness

130 We have added this information.

Figures 1&2: The upper and lower limits of the color map seem to be saturated in some areas. I suggest to adjust the limits.

We have changed the limits on these plots.

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Figure 6: Increase the resolution of the figure.

We have updated this.

References

- Anheuser, J., Liu, Y., and Key, J.: A simple model for daily basin-wide thermodynamic sea ice thickness growth retrieval, The Cryosphere, 16, 4403–4421, https://doi.org/10.5194/tc-16-4403-2022, 2022.
 - DeRepentigny, P., Tremblay, L. B., Newton, R., and Pfirman, S.: Patterns of Sea Ice Retreat in the Transition to a Seasonally Ice-Free Arctic, J. Climate, 29, 6993 7008, https://doi.org/10.1175/JCLI-D-15-0733.1, 2016.
- Petty, A. A., Holland, M. M., Bailey, D. A., and Kurtz, N. T.: Warm Arctic, Increased Winter Sea Ice Growth?, Geophysical Research Letters, 45, 12 922–12 930, https://doi.org/10.1029/2018gl079223, 2018.
 - Ricker, R., Kauker, F., Schweiger, A., Hendricks, S., Zhang, J. L., and Paul, S.: Evidence for an Increasing Role of Ocean Heat in Arctic Winter Sea Ice Growth, Journal of Climate, 34, 5215–5227, https://doi.org/10.1175/jcli-d-20-0848.1, 2021.
 - Taylor, J. R.: An introduction to error analysis: The study of uncertainties in physical measurements, University Science Books, 20 Edgehill Rd., Mill Valley, CA 94941, 1982.
- Tschudi, M. A., Meier, W. N., and Stewart, J. S.: An enhancement to sea ice motion and age products at the National Snow and Ice Data Center (NSIDC), Cryosphere, 14, 1519–1536, https://doi.org/10.5194/tc-14-1519-2020, 2020.