

1 Comments and responses to referees and authors.

2 ref 1

3 My response:

4
5 Thank you for your remarks.

6
7 1. My commentary is based on observations via Google Earth imagery. This makes it
8 possible for any reader to look at the field evidence and surrounding areas. Charles
9 Darwin noted, 'How odd it is that anyone should not see that observation must be for or
10 against some view if it is to be of any service' (Ayala, 2009). This quotation highlights
11 issues in the philosophy of science and the nature of evidence both of which I touch upon
12 in my responses hereafter. I have numbered the main points sequentially for the benefit
13 of the reader.

14
15 2. My original comments, and indeed my responses posed here, are intended to show
16 readers the field evidence as I see it; 'it is essential to the scientific process that any
17 hypothesis be "tested" by reference to the natural world that we experience with our
18 senses' (Ayala, 2009).

19
20 3. Although it may not be 'possible to determine the origin of rock glaciers', the reviewer
21 acknowledges that my argument is 'sufficiently convincing' to warrant using the
22 glacial model for the Dos Lunas (DL) rock glacier. My comments are based on
23 observations from various glacier-rock glacier landforms in the area. I chose to
24 illustrate it with one specific example, but I fill in some more detail in my responses to
25 others below.

26
27 4. In the responses I use the following convention to help readers identify locations on
28 Google Earth (GE) by pasting in the numbers in the GE search bar between square
29 parentheses. Thus, Dos Lunas (DL) can be identified as decimal latitude and longitude
30 [-30.24664,-69.78667]. A transect along the 'fall line' on the feature starts at the top with
31 the last term (260) being a bearing in degrees from the preceding couplet as origin: {-
32 30.24235,-69.76730,260}. This decimal degree convention is more useful to
33 georeference features at various scales and transects for recording purposes than the
34 traditional ° ' ". See Whalley (2021a, 2021b; collated references are at the end) for
35 illustrations about the notation for studying rock glaciers elsewhere.

36
37 referee 2

38
39 Thank you for your comments. I fill in some detail here in direct response to your remarks
40 (other information is provided below). I have tried to keep these succinct and directly
41 related to what is particularly pertinent.

42 5. As my comments were primarily about field observations (see 1, supra), I only included
43 two papers about the rheology of ice rock mixtures. It is the mechanical nature of the mixture
44 model (rock/ice-snow/water/air) that determines the rheology. A thin glacier (<30m thick,
45 slope angle ca 10°, with an ablation-reducing debris cover) will flow at rock glacier
46 velocities, < 1 m a⁻¹. However, talus (scree or rockfill) as an 'ice sparse' composite will not
47 flow unless the ice content is high (perhaps >60%) and in thick (≈ 20 m) deformable bands or

48 lenses. The geophysical signature of a rock glacier at any location depends upon the field-
49 mixture model, as well as the volume examined, given its inhomogeneity and anisotropy. The
50 permafrost model correlates geophysical signatures to a formational mode for all rock
51 glaciers (i.e. exclusively of non-glacier origin, see 17 below). My commentary suggests there
52 is directly observable field evidence for a glacial origin for the deforming ice at DL. But, as
53 Lliboutry noted (1990) of a comment by Haeberli (1989), 'I do not deny that many (not all)
54 rock glaciers are below melting point at depth'.

55 6. Why don't all the slopes in the area show flow-features when, in a known permafrost area,
56 there are plentiful scree slopes? The answer is that they will do so only if there is a thick
57 enough body of ice, as a glacier in a conventional sense or with a thick snow/ice body
58 covered with debris (5). On cliffed slopes with snow avalanching, this can be achieved if
59 perennial snow accumulates (and is buried, perhaps sequentially, under debris). This is the
60 point made by reference to rheology in Whalley and Azizi (2003) and the mixture model (see
61 5). It is the creep of massive ice, not rock debris – even if this is in a permafrost area.
62 Permafrost is not necessary, but it is sufficient to keep creep rates lower than at ice pressure
63 melting point. As an illustration, the transect, 1: {-30.2423,-69.7670,260} down the centre of
64 DL rock glacier can be compared with a parallel transect, 2: {-30.24908,-69.76338,270}. The
65 latter, some 700 m to the south of 1, is representative of much of that mountainside and must
66 be under the same environmental conditions, temperature, snowfall and ablation, as the rock
67 glacier, 1. However, transect 2 shows no signs of flow. The reason must lie in the 'mixture
68 model', debris from the upper slopes has covered a perennial snowpack of a 'buried glacieret',
69 'buried debris-rich glacieret' or 'glacier enterré' (Lliboutry, 1961; Lliboutry, 1990). That there
70 is no glacier/glacieret remnant showing at 1 is because the thick ice mass necessary for flow
71 is covered with debris from above. The top of this original, small and confined, glacier would
72 have been under the cliffs in the vicinity of Google Earth locality [-30.2429,-69.7747] and
73 fed down gullies higher on the slope. Extant equivalents can be seen at the top of gullied
74 south-facing slopes in the vicinity of [-30.23512,-69.83599]. The glacier and its protecting
75 debris load have now crept downhill and formed the DL rock glacier. A short transect {-
76 30.24318,-69.77858,160} for about 150 m, i.e. some 250 m east of the Halla et al. 'root zone'
77 transect, is lower in the centre (by 5-10m) from the edges. This shows that ice had flowed out
78 of this area and has not been replaced. This effect is similar to other rock glaciers with
79 extending flow regimes (Whalley and Palmer, 1998, Whalley, 2021b).

80 7. Observations using GE brings to light further changes in surface topography of rock
81 glaciers, notably the appearance of pools that show melting of ice below the surface debris.
82 Recent coverage by GE shows meltwater pool exposures are becoming increasingly common.
83 Ridges and furrows, piled up in lower (snout) regions are the result of basically compressive
84 glacier flow with debris loads becoming increasingly thick near and at the snouts. This
85 inhibits melting further from upstream amounts (where the debris load is thinner). Glaciers
86 and rock glaciers may exhibit extending flow where, usually on steeper slopes and perhaps
87 more restricted valley sections, transverse ridges and furrows are replaced by irregular or
88 longitudinal features. Meltwater pools can form variously in them according to local
89 topography and thickness of the debris cover.

90 8. These meltwater pools can be of considerable size, that shown in my Fig 1 at [-30.2413,-
91 69.8542] has a water area of about 3 000 m² and has been in existence at least between 2006
92 – 2019 (from GE imagery). The total 'missing' volume of rock glacier is some 40 x 10³ m³,
93 suggesting that the mixture model is predominantly of high percentage (massive) ice from a
94 buried glacier tongue. This is commensurate with the sides of a 'thermokarst depression'

95 shown (Figure 4) of Trombotto-Liaudat and Bottegal (2020) at Morenas Coloradas debris-
96 covered glacier [-32.9426,-69.3988] although the exact location is not given. Other long-lived
97 meltwater pools can be seen up-valley to the exposed glacier at Morenas Coloradas, further
98 examples can be seen in some of the images in Janke et al. (2015). Whether rock glaciers
99 extend back into visible debris free and debris-covered versions (as suggested in the
100 classification of Janke et al. (2015)) depends upon the relative inputs of glacier ice and
101 weathered debris over time. The Colina Mountain example (Janke et al., 2015, Fig. 21B) [-
102 34.3428,-70.0492] has a continuum of classes of debris-covered glacier/rock glacier with
103 surface forms that include meltwater pools [-34.3437,-70.0486] & [-34.3494,-70.0583] and
104 lateral erosion of pool with an exposed glacier ice cliff [-34.3571,-70.0718].

105 **HAEBERLI**

106 Thank you for your comments Wilfried.

107 9. Please note that I said, 'The geophysical data supplied by Milana and Güell (2008) and
108 Halla et al. (2020) will be useful in the interpretation of these factors in glacier/rock glacier
109 formation ...' In other words, evaluating the nature of the 'mixture model' that should be
110 applied to the rheology (6, supra) will be helpful in establishing the geophysical properties
111 and variability in rock glaciers. I am well aware of the range of geophysical results available
112 from rock glaciers and why they can be so variable (acknowledged by Referee 2) and noted
113 this in my original comment. This is also part of the review of the mixture models provided
114 by Whalley and Azizi (1994) and I do not propose to discuss this variability here as my point
115 was, and is, to look at visible forms and how they might inform us as to the origin of rock
116 glaciers. The rheology gives the landform and its details, not the variable geophysical
117 signature.

118 10. I am also aware of Gruben glacier/rock glacier and its ice-dammed lakes and the so-called
119 'periglacial part'. But readers should note that an interpretation of that rock glacier landsystem
120 suggests that the rock glacier *does* have a glacier ice core (Whalley, 2020). It is no different
121 from the observations of glacier ice cores in rock glaciers that have been recorded over the
122 years from many parts of the world, for example; Kesseli (1941), Potter et al. (1998) and
123 more recently Whalley (2021b). No amount of geophysical pleading can refute these
124 observations. It is for time, as more meltwater pools are exposed, and readers to evaluate. A
125 rough calculation (see 8, supra) shows that such meltwater pools are from the decay of
126 massive glacier ice – which is what was the case at Gruben (Whalley, 2020).

127 11. It is certainly true that boreholes and exposures do show the complex nature of ice and
128 debris in rock glaciers, see for example Janke et al. (2015) and Jones et al. (2019), especially
129 near rock glacier snouts. Because of the increasing surface debris loads down-valley, ice
130 exposures tend to be hidden by debris. However, some snout collapses can be seen in GE,
131 such as at Glockturmferner (Austria) [46.89846,10.65058], compared with earlier views
132 (Kerschner, 1983). Lliboutry described a section in the one of the four 'glaciers enterrés'
133 below the west face of Cerro Negro (Andes of Santiago). The exact location is unknown but
134 is in the vicinity of [-33.1484,-70.2367] (Lliboutry, 1961, Fig. 1). The section (Lliboutry,
135 1961 Fig. 4) and (Lliboutry, 1965 Fig.17.21) shows complex relationships between ice;
136 young, old bubbly and bubble free ice together with silt and pebbled bands. This is more
137 complex than the section shown by Trombotto-Liaudat and Bottegal (2020). Figure 8 of
138 Janke et al. (2015) shows section of a meltwater pool showing banding, similar to Gruben
139 rock glacier's drained lakes (Whalley, 2020). There is clearly much to be gained about the

140 structures of glaciers as they become exposed at the snouts of rock glaciers. This will help in
141 matching geophysical attributes to structural glaciology and debris content.

142 12. Although there have been descriptions of rock glaciers since the early 20th C, the paper
143 by Wahrhaftig and Cox (1959) has become particularly important in discussion about these
144 features (Stine, 2013). Indeed, it has become the 'Urtext' for those believing the 'permafrost'
145 origin of rock glaciers promoted by Wahrhaftig and Cox. The book by Barsch (1996)
146 provides the stated dogma of the permafrost viewpoint. This text is followed by Barsch
147 (1987) who denigrates many observations of glacier ice cores. Subsequently, sins of omission
148 have followed by disregarding any other possibilities than the permafrost dogma, e.g. Swift et
149 al. (2021). Please see Whalley (2021a) where some of these wrongs are addressed.

150 13. Professor Haeberli, as a true believer in the Urtext and permafrost dogma, has always
151 maintained that rock glaciers cannot have glacier ice cores (i.e. be glacial). For him, this
152 means that not only do glacier ice cores not exist but that any continuum or equifinality does
153 not occur (pace Referee 2). Yet there are many reports of glacier ice in rock glaciers, as well
154 as the well-established work of Potter at Galena Creek that cannot be denied (although I leave
155 it to readers to adjudicate). Quoting many references that support a permafrost viewpoint
156 amounts to 'affirming the consequent' (modus tollens). In terms of swans and rock glaciers,
157 all swans are not white and at least some rock glacier swans are black and contain glacier ice
158 cores. Thus, supposition and following a particular point of view is insufficient to replace
159 valid contra-observations. In a Popperian sense therefore we might have to wait for contra-
160 indications of permafrost, or affirmation of the appearance of glacier ice by meltwater ponds.

161 14. I have mentioned the work of the late Professor Louis Lliboutry in reporting 'glacier
162 enterré' and in particular the complexities of snout stratigraphy. He also said (Lliboutry,
163 1990); 'I do not wish to enter into a public controversy with W. Haeberli about the origin of
164 rock glaciers; he has always been deaf to my arguments. Nevertheless, the readers of his
165 passionate assertions (Haeberli, 1989) must be aware that he intentionally omits to quote my
166 detailed observations in the dry Andes (Lliboutry, 1955, 1965, 1986).' Further, 'Nevertheless,
167 for the advancement of science, the essential point is not "must rock glaciers be left to
168 scientists claiming to be permafrost specialists" but "what can we learn from the existence of
169 rock glaciers in a given area"? I maintain that the geographical study of rock glaciers as an
170 extreme case of glacier fluctuations, as an indicator of favourable mass balances in the past,
171 or of past surges, would be much more rewarding than to consider them as a mere case of
172 standard permafrost, or of creeping regolith.' (Lliboutry, 1990).

173 **Halla et al**

174 Dear Authors. Thank you for your comments

175 Regarding your first point, I appreciate that your detailed work refers to a single feature. By
176 implication however, your findings refer to the general study of water storage in glaciers and
177 rock glaciers. Thus, your study becomes a part of an overall appreciation of water content in
178 South America and needs to accommodate a variety of findings under slightly different
179 climatic conditions – as you are arguing for a zonal (or morphoclimatic) interpretation.

180 15. I appreciate your view (third point) that, 'the assessment and discussion of the origin of a
181 distinct rock glacier or landform should be based on on-site specific geomorphological
182 characteristics (form, process, and material) of the landform. Indeed, I recently (Whalley,

183 2021a) I suggested that it was necessary (though geomorphological mapping) 'to recognise
184 and link materials (M), 'processes' (P, that is mechanisms integrated over time) and visual
185 categorization and geometrical information (G). In principle, this information, i.e. site
186 metadata, can be collected and a database interrogated to maximise geomorphological
187 knowledge'. I suggest above (points 6 and 9) that it is the rheological (dynamic) properties of
188 a feature and related to the materials, that account for the forms seen. In this it is necessary to
189 look at the connectivity of material movement downslope and the origin of both water/ice
190 and solids. Further, that other examples in the literature, which can be seen on Google Earth,
191 do show rheological properties that are consistent with a glacier ice core (for valley floor rock
192 glaciers) or a substantial snow/ice mass that has been buried by copious debris supplies from
193 above – which is the case at DL. As mentioned above (6) ice that collected in the vicinity of
194 [-30.2429,-69.7747] has moved downslope and now lies buried under the debris in the snout
195 lobes. That there are no 'glacial deposits, like moraines' as 'traces of a former glacier' is rather
196 easily explained; the rock glacier deposits *are* the moraines. A transect {-30.24316,-
197 69.77959,255} shows a distinct (right) lateral moraine of a former small debris-covered
198 glacier, with its main ice collection area at about [-30.2429,-69.7784]. This small glacier was
199 clearly overwhelmed by the ice and sediments of the ice rock glacier of DL.

200 16. It is arguable whether science should be conducted according to inductive or deductive
201 principles (see Ayala (2009) for basic discussion related to Darwin). Goudie and Viles (2010)
202 argue for an abductive view in the construction of ideas and models but in order to overcome
203 'prejudices and conditioning' the 'critical rationalist approach' of Karl Popper should be used
204 to 'attempt to disprove rather than verify our hypotheses' (Schumm, 1991). In other words,
205 and in this case, alternative viewpoints are not only acceptable but to be welcomed (12,
206 supra). Thus, my observations of meltwater pools in a wide variety of instances in the
207 literature, which show that ice melting is not 'iso-volumetric' supports a massive ice origin. A
208 theory should make predictions that can be tested. I suggest that meltwater pools will be seen
209 on DL around [-30.2479,-69.7850] in the next ten years to become like [-30.2413,-69.8542]
210 to which it is topographically similar and functionally related.

211 17. I shall not argue about your geophysical results – which was not my intention in the first
212 place – and referee 2 (supra) has already commented on these. However, you state that DL
213 should be considered as a 'talus rock glacier'. I have no difficulty with the terminology only
214 that it must necessarily be 'creeping permafrost'. Some authors e. g. Evin et al. (1997) have
215 argued for 'hybrid models' and Monnier and Kinnard (2015) have discussed 'glacier-rock
216 glacier transitions' and Jones et al. (2019) present water content evidence from a variety of
217 rock glacier models. More investigations are clearly required.

218 18. With respect to 'surface texture, the geomorphological characteristics and spatial
219 connection of the rock glacier to the upslope are recommended proxies for visual
220 observations' (IPA, 2020) I have here outlined some reasons for considering the
221 characteristics at DL (and elsewhere) as indicative of glacier flow. However, the IPA
222 document presents a major misunderstanding of the nature of rock glaciers by concentrating
223 on kinematics rather than dynamics (rheological properties). Any flow mechanisms, i.e.
224 dynamics not just kinematics, needs to consider the full implications of the materials
225 involved. In other words, the IPA statement follows the pure Urtext (12) with not even
226 alternatives such as hybrid or equifinality possibilities.

227 19. I do not have space to argue my case about the IPA (2020) publication but rather point
228 out that in stating that 'rock glacier (or permafrost) creep has to be understand (sic) here as a

229 generic term' (p. 6) and 'Rock glaciers, as landforms resulting from a permafrost creep
230 process, should not be confused with debris-covered glaciers'. (p. 11) it follows the
231 'exclusive' approach (5 supra). In particular, by assuming the dogma associated with the
232 permafrost Urtext (12) and by ignoring the glacial/glacigenic model for which there is good
233 evidence, it has engendered 'belief perseverance' in some sectors of the geoscience
234 community where there is also 'confirmation bias' that has not been assuaged by showing
235 falsifiers (black swans). That I have generated some discussion is a good thing, although I
236 return to my original quotation from Charles Darwin on observations. But thank you for your
237 paper and its valuable measurements.

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240 Ayala, F.J.: Darwin and the scientific method. *Proc. Nat. Acad. Sci.*, 106, Supp. 1, 10033-
241 10039, 2009.

242 Barsch, D.: The problem of ice-cored rock glacier. In: J.R. Giardino, J.F. Shroder, J.D. Vitek
243 (Eds.), *Rock glaciers*. Allen and Unwin, London, pp. 45-53, 1987.

244 Barsch, D.: *Rockglaciers. Indicators for the present and former geocology in high mountain*
245 *environments*. Springer, Berlin, 331 pp. 1996.

246 Evin, M., Fabre, D. and Johnson, P.G.: Electrical resistivity measurements on the rock
247 glaciers of Grizzly Creek, St Elias Mountains, Yukon. *Permafrost Periglac.*, 8(2), 179-
248 189, 1997.

249 Goudie, A. and Viles, H.: *Landscapes and geomorphology: a very short introduction*. OUP,
250 Oxford, 137 pp. 2010.

251 Haerberli, W.: Glacier ice-cored rock glaciers in the Yukon Territory, Canada? *J. Glaciol.*,
252 35(120), 294-295, 1989.

253 Halla, C., Blöthe, J. H., Tapia Baldis, C., Trombotto, D., Hilbich, C., Hauck, C., and Schrott,
254 L.: Ice content and interannual water storage changes of an active rock glacier in the
255 dry Andes of Argentina, *The Cryosphere Discussions*, doi.org/10.5194/tc-2020-29,
256 2020.

257 IPA, International Permafrost Association: IPA Action Group, *Rock glacier inventories and*
258 *kinematics*. Available at:

259 https://bigweb.unifr.ch/Science/Geosciences/Geomorphology/Pub/Website/IPA/Guidelines/V4/200507_Baseline_Concepts_Inventorying_Rock_Glaciers_V4.1.pdf. 2020.

261 Janke, J.R., Bellisario, A.C. and Ferrando, F.A.: Classification of debris-covered glaciers and
262 rock glaciers in the Andes of central Chile. *Geomorphology*, 241, 98-121, 2015.

263 Jones, D. B., Harrison, S., Anderson, K., and Whalley, W. B.: Rock glaciers and mountain
264 hydrology: A review, *Earth-Sci. Rev.*, 193, 66-90,
265 <https://doi.org/10.1016/j.earscirev.2019.04.001>, 2019.

266 Kerschner, H.: Zeugen der Klimageschichte im Oberen Radurschltal. *Alpenvereinsjahrbuch*,
267 1982-83 DÖAV, 23-28, 1983.

268 Kesseli, J.E.: Rock streams in the Sierra Nevada, California. *Geogr. Rev.*, 31(2), 203-227,
269 1941.

270 Lliboutry, L.: Origine et évolution des glaciers rocheux. *R. Acad. Sci. (Paris)*, 240, 1913-
271 1915, 1955.

272 Lliboutry, L.: Les glaciers enterrés et leur rôle morphologique, *Symp. Helsinki Int. Ass. Sci.*
273 *Hydrol. Pub.* 54, pp. 272-280, 1961.

274 Lliboutry, L.: *Traité de Glaciologie*, 2. Masson & Cie, Paris, 707 pp. 1965.

275 Lliboutry, L.: About the origin of rock glaciers. *J. Glaciology*, 36(122), 125-125, 1990.

276 Monnier, S. and Kinnard, C.: Internal structure and composition of a rock glacier in the
277 Andes (upper Choapa valley, Chile) using borehole information and ground-
278 penetrating radar. *Ann. Glaciol.*, 54, 61–72, <https://doi.org/10.3189/2013aog64a107>,
279 2013.

280 Potter, J., N, Steig, E., Clark, D., Speece, M., Clark, G.T. and Updike, A.B.: Galena Creek
281 rock glacier revisited—New observations on an old controversy. *Geogr. Ann. A*,
282 80(3-4), 251-265, 1998.

283 Schumm, S.A.: To interpret the earth: ten ways to be wrong. Cambridge University Press,
284 Cambridge, p. 133, 1991.

285 Stine, M.: Clyde Wahrhaftig and Allan Cox (1959) Rock glaciers in the Alaska Range.
286 *Bulletin of the Geological Society of America* 70 (4): 383–436. *Prog. Phys. Geog.*,
287 37(1), 130-139, 2013.

288 Swift, D.A., Cook, S., Heckmann, T., Gärtner-Roer, I., Korup, O., Moore, J.: Ice and snow as
289 land-forming agents, *Snow and Ice-Related Hazards, Risks, and Disasters*. Elsevier,
290 pp. 165-198, 2021.

291 Trombotto-Liaudat, D., Bottegal, E.: Recent evolution of the active layer in the Morenas
292 Coloradas rock glacier, Central Andes, Mendoza, Argentina and its relation with
293 kinematics. *Cuadernos Investigación Geográfica*, 46(1), 159-185, 2020.

294 Wahrhaftig, C., Cox, A.: Rock glaciers in the Alaska Range. *Geol. Soc. Am. Bull.*, 70(4),
295 383-436, 1959.

296 Whalley, W.B., Palmer, C. F.: A glacial interpretation for the origin and formation of the
297 Marinet Rock Glacier, Alpes Maritimes, France. *Geogr. Ann.*, A, 80, 3/4 221-236,
298 1998.

299 Whalley, W. B.: Gruben glacier and rock glacier, Wallis, Switzerland: glacier ice exposures
300 and their interpretation, *Geogr. Annaler: A*, 102, 141-161, 2020.

301 Whalley, W.B.: Geomorphological information mapping of debris-covered ice landforms
302 using Google Earth: an example from the Pico de Posets, Spanish Pyrenees.
303 *Geomorphology*, <https://doi.org/10.1016/j.geomorph.2021.107948>, 2021a.

304 Whalley, W.B.: The Glacier – Rock Glacier Mountain Landsystem: an example from North
305 Iceland. *Geogr. Ann.*, B, 2021b.

306 Whalley, W. B., and Azizi, F.: Rheological models of active rock glaciers: evaluation,
307 critique and a possible test, *Permafrost Periglac.*, 5, 37-51, 1994.

308 Whalley, W. B., and Azizi, F.: Rock glaciers and protalus landforms: Analogous forms and
309 ice sources on Earth and Mars, *J. Geophys. Res.: Planets* (1991–2012), 108, 2003.

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