Geographic variation and temporal trends in ice phenology in Norwegian lakes during the period 1890-2020

Jan Henning L'Abée-Lund[1], Leif Asbjørn Vøllestad[2], John Edward Brittain[1,3], Ånund Sigurd Kvambekk[1] and Tord Solvang[1]

[1] Norwegian Water Resources and Energy Directorate, Box 5091 Majorstuen, N-0301 Oslo, Norway

[2] Centre for Ecological and Evolutionary Synthesis, University of Oslo, Box 1066 Blindern, N-0316 Oslo, Norway

[3] Natural History Museum, University of Oslo, Box 1072 Blindern, N-0316 Oslo, Norway

Correspondence to: Jan Henning L'Abée-Lund (jlabeelund@gmail.com)

Abstract

Long-term observations of ice phenology in lakes are ideal for studying climatic variation in time and space. We used a large set of observations from 1890 to 2020 of the timing of freeze-up and break-up, and the length of ice-free season, for 101 Norwegian lakes to elucidate variation in ice phenology across time and space. The dataset of Norwegian lakes is unusual, covering considerable variation in elevation (4 – 1401 m a.s.l.) and climate (from oceanic to continental) within a substantial latitudinal and longitudinal gradient (58.2 – 69.9° N; 4.9 – 30.2° E).

The average date of ice break-up occurred later in spring with increasing elevation, latitude and longitude. The average date of freeze-up and the length of the ice-free period decreased significantly with elevation and longitude. No correlation with distance from the ocean was detected, although the geographical gradients were related to regional climate due to adiabatic processes (elevation), radiation (latitude) and the degree of continentality (longitude). There was a significant lake surface area effect as small lakes froze-up earlier due to less volume. There was also a significant trend that lakes were completely frozen over later in the autumn in recent years. After accounting for the effect of long-term trends in the large-scale NAO index, a significant but weak trend over time for earlier ice break-up was detected.

An analysis of different time periods revealed significant and accelerating trends for earlier break-up, later freeze-up and completely frozen lakes after 1991. Moreover, the trend for a longer ice-free period also accelerated during this period, although not significant.

An understanding of the relationship between ice phenology and geographical parameters is a prerequisite for predicting the potential future consequences of climate change on ice phenology. Changes in ice phenology will have consequences for the behaviour and life cycle dynamics of the aquatic biota.

Keywords: Lake ice, Ice phenology, Climate change, Lake characteristics, Geographical variation

1 Introduction

The surface area of lakes makes up a substantial part (15-40 %) of the arctic and sub-arctic regions of the Northern Hemisphere (Brown and Duguay 2010). Most of these lakes freeze over annually. In addition to its substantial biological importance (Prowse 2001), this annual freezing has significant repercussions for transportation, local cultural identity and religion (Magnusson et al., 2000; Prowse et al., 2011; Sharma et al., 2016; Knoll et al., 2019). The importance of freshwater and ice formation for people has resulted in the monitoring of freezing and thawing of lake ice for centuries (Sharma et al., 2016).

Lakes and their ice phenology are effective sentinels of climate change (Adrian et al., 2009) and ice phenology has been studied extensively (e.g., reviewed by Brown and Duguay, 2010). In general, freeze-up and break-up have changed over time, freeze-up occurs later and break-up appears earlier despite different timespans on global (Magnuson et al., 2000 (1846-1995); Benson et al., 2012 (1855-2005); Sharma & Magnusson, 2014 (1854-2004); Du et al., 2017 (2002-2015)), regional (Duguay et al., 2006 (1952-2000); Mishra et al., 2011 (1916-2007); Hewitt et al., 2018 (1981-2015)) and local scales (Choiński et al., 2015 (1961-2010); Takács et al., 2018 (1774-2017)). Despite these general results, the strength of the trends varies among studies. The time of freeze-up was delayed by 0.3 to 5.7 days/decade (Benson et al. 2012, Magnusson et al. 2000), whereas the timing of ice break-up was delayed by between 0.2 and 6.3 days/decade (Mishra et al., 2011; Magnusson et al., 2000). Some of this variation is a consequence of differences in the length of the study period, covering from more than a century to just a single decade. However, based on time series for 2000-2013 from 13 300 Arctic lakes, Šmejkalova et al. (2016) showed significant and more dramatic trends in earlier start and end of break-up in Northern Europe as the rate was 0.10 days/year 0.14 days/year, respectively, and that this change was significantly correlated with the 0 °C isotherm. The wide variation in time period and the particular time-period studied is important to consider when trying to compare the strength of trends in ice phenology parameters as significant associations with ice break-up and oscillations (2-67 years) have been documented (Sharma & Magnusson, 2014). Global mean temperature has changed considerably since 1880 (Hansen et al., 2006), and the change (increase) in temperature is particularly evident in later decades (IPCC, 2007; Benson et al., 2012). By dividing data from the 1976-2005 period into shorter timer periods, Newton and Mullan (2020) showed, for Fennoscandia, an increase in the magnitude of the general trend in earlier break-up in 1991-2005 compared to earlier periods. In North America the trend was for earlier break-up, but it was neither spatially nor temporally consistently explained by local or regional variation in climate (Jensen et al., 2007). In a recent study, Filazzola et al. (2020) showed that unusually shorter ice cover periods are becoming more frequent and even shorter, especially since 1990.

In Fennoscandia, recording ice phenology has long traditions due to the importance of frozen lakes and rivers for transport and recreation (Sharma et al., 2016). Data from Swedish and Finnish lakes have been studied in detail (e.g., Eklund, 1999; Kuusisto & Elo, 2000; Livingstone, 2000; Yoo & D'Odorico, 2002; Blenckner et al., 2004; Korhonen, 2006; Palecki & Barry, 1986). Based on Swedish data for the period 1710-2000, Eklund (1999) showed that ice break-up did not change from 1739 to 1909, became 5 days earlier in the period 1910-1988 and still 13 days earlier during the final period (1988-1999). Furthermore, ice freeze-up was later in the 1931-1999 period than in the 1901-1930 period. Similarly, stronger trends in both freeze-up and break-up in the last decade of the 1950-2009 time period have been shown for both Finnish and Karelian lakes (Blenckner et al., 2004; Efremova et al., 2004; Korhonen, 2006). Moreover, Blenckner et al. (2004) showed that large variability was apparent south of 62° N, indicating that lakes in southern Sweden were more influenced by large-scale climate effects (such as the North Atlantic Oscillation; NAO (Hurrell, 1995)) than northern lakes. This pattern was explained by the mountain range between Norway and Sweden affecting the regional circulation in the north. The large-scale anomaly in the NAO in the winter season shifts between strong westerly winds with warm and moist air and cold, easterly dry winds across the North Atlantic and Western Europe. The positive phases of NAO are associated with milder and rainy delayed winters and early springs in Northern Europe (Hurrell, 1995). This significantly affects the timing of ice break-up in lakes (Palecki & Barry, 1986; Livingstone, 2000; Yoo & D'Odorico, 2002). However, Yoo & D'Odorico (2002) argued that climatic forcing such as CO_2-induced regional and global warming may have a pronounced effect leading to earlier break-up. On the other hand, George et al., (2004) showed that ice correlations (freeze-up and length of the period of ice cover) differed strongly between Windermere situated close to the Irish Sea in northwest England and Pääjärvi situated some distance from the Baltic sea in southern Finland. The number of days with ice has fallen dramatically for lake Windermere, whereas no such trend was detected for Pääjärvi. They postulated that the position of the boundary between the oceanic and continental climate regimes can change and produce a significant shift in winter dynamics of lakes located near this zone. In addition to this effect between climate zones, the boundary of the 0 °C isotherm is important as it strongly affect ice formation and break-up (Brown and Duguay, 2010; Filazzola et al., 2020).

Despite the fact that registration of ice phenology has been undertaken in a large number of lakes and rivers in Norway, as early as 1818 in some lakes (www.nve.no), few lakes have been studied in detail and no country-wide analysis has been done. Trends in freeze-up and break-up have been analyzed for two subalpine lakes in Central Norway (Kvambekk and Melvold, 2010; Tvede, 2004; Solvang, 2013). Although not covering the exact same period, both freeze-up and break-up show different trends in the two lakes. Although geographically close to lakes in Sweden and Finland, Norwegian lakes demonstrate considerably more variation in topography and climate. Norway covers most of the Scandinavian north-south mountain ridge with several summits above 2300 m a.s.l. while the highest mountains in Sweden and Finland only reach 2106 and 1328 m a.s.l., respectively. This mountain ridge ensures that Sweden and Finland generally have a continental climate, contrary the complex climate in Norway with oceanic climate in the west and south, continental climate in the east along the Swedish and Finnish border, and tundra and subarctic climates in the mountain regions in the southern and northern parts. Norwegian lakes, situated in the western parts of the Scandinavian peninsula, encompass a large a variation in elevation over short distances as well as substantial latitudinal and longitudinal variation. A large and complex coast also introduces considerable climate variability. This makes Norwegian lakes well suited for testing the effect of climate change on ice phenology, also in relation to elevation.

In the present study, we have analysed long-term (1890-2020) observations of lake freeze-up, ice break-up and length of ice-free period in 101 Norwegian lakes. The lakes cover a broad range of climatic zones described by geographical parameters (elevation, latitude and longitude), as well as lake characteristics (area, water inflow and water level amplitude). The main aim of the analyses was to detect potential temporal trends in ice phenology while adjusting for both geographical parameters and lake characteristics.

2 Material and methods

2.1 Lakes studied

We collated observations from 101 Norwegian lakes, covering a wide range in latitude (58.2 – 69.9° N), longitude (4.9 – 30.2° E) and elevation (4 – 1401 m a.s.l.). The lakes are situated in three major climatic zones (boreal, subalpine, alpine) and with varying distances from the ocean. Thus, they differ widely in several geographic characteristics (Figure 1, Appendix 1). Most of the lakes are relatively small (median area 6.9 km^2), although the dataset also includes Norway's largest lake, Mjøsa (369.3 km^2). Their catchment areas vary between 7.1 and 18101.9 km^2 (median 235 km^2) and mean annual inflow to the lakes varies between 5.6 10^6 and 9935.7 10^6 m^3 $year^{-1}$ (median 256 10^6 m^3 $year^{-1}$). About 50 % of the lakes (N = 53) were developed for hydropower production with an annual water level variation varying from 1 to 30.3 m. The lake and catchment information were extracted from Norwegian Water Resources and Energy Directorate website www.nve.no.

2.2 Ice observations

Observations of the timing of ice formation on the lakes in autumn and ice break-up in spring were undertaken visually or by fixed-location video cameras. The data were made available by the Norwegian Water Resources and Energy Directorate (NVE), the hydropower association Glommens og Laagens Brukseierforening, or by private persons. NVE operates a national hydrological database that contains information on ice conditions. The first observations are from 1818, but substantial records started in the 1890s. Video cameras have now replaced visual observations in some lakes. Satellite data is also being increasingly used to detect ice cover or open water. In our dataset, we have included lakes with more than 7 years of observations for at least one ice phenology variable in the analysis. This resulted in 101 lakes of which 76 have a registration period exceeding 30 years (Figure 2, Appendix 2). The average length of the data series was 53 years (range 11 – 149 years).

Prior to 2000, most observations of ice phenology were carried out manually by NVE observers, power company employees, farmers and landowners. Afterwards web cameras and remote sensing became increasingly important. In most years, registrations were on a daily basis. After 2000, personnel conducted weekly observations, and in these cases remote sensing were used to improve accuracy. Registrations by personnel were conducted at the shore. Thus, the accuracy is high for the date of freeze-up, whereas the setting the date of a completely ice-covered lake and break-up have an uncertainty of a couple of days.

The date of ice break-up was set when the lake was estimated to be free of ice based on the available observations. The length of the ice-free period during summer was then estimated as the difference between the day of freeze-up in the autumn and the day of ice break-up in spring. All dates are given as Julian day number during the year (1 January is day 1). For some lakes in certain years ice formation started in winter after 1 January. For these years the day number was extended past the normal 365 days. The observations were always made at the same site in each lake. The date of freeze-up was set when the first formation of ice was observed. Subsequent temporary ice-free periods, often due to mild weather combined with strong winds, did not change this date. The date when the whole lake was covered by ice was also noted, when possible. This date is more variable, and information is frequently missing. It would require extensive travel and several observation points to ascertain this date with high certainty, unless there are time-lapse cameras or satellite data. We have a total of 4371 observations on ice break-up, 3035 observations of freeze-up, 4221 observations of when the lakes were completely frozen over, and 2808 observations of the length of the ice-free period.

Some of the lakes are used as hydropower reservoirs, and thus within-year water level variation may differ from the normal annual cycle. For such lakes we have included information on the year of regulation and the maximum amplitude of water level variation. Although we do not have information on exact water level variation within a given year, maximum and minimum occurs when freeze-up and break-up normally take place, respectively.

For one particular large lake there are observations from two different locations (called Tustervatn and Røssvatn) that were partly overlapping in time. The observations of the time of ice break-up and ice freeze-up were strongly and positively correlated. The correlation between the two different estimates of time of freeze-up ($r = 0.501$, $n = 37$, $p = 0.002$) were lower than for the time of break-up ($r = 0.887$, $n = 38$, $p < 0.001$). There was no tendency for a particular temporal trend for this particular lake, so we have used the longest of the two time-series in the analyses.

2.3 Climate data

As a potential large-scale climate driver, especially impacting ice break-up, we used the North Atlantic Oscillation (NAO) index. We therefore extracted the PCA-based winter (December to March) NAO index (National Center for Atmospheric Research Staff (Eds.), last modified 10 September 2019: https://climatedataguide.ucar.edu/climate-data/hurrell-north-atlantic-oscillation-nao-index-pc-based (accessed 28 October 2020)). Variation in winter NAO is known to impact on winter temperature and precipitation, depending on location (Hurrell, 1995; Stenseth et al., 2003). An elevated index leads to mild and wet winters in Europe, while a low index leads to cold and dry winters. The PCA-based winter NAO-index covers the period from 1898 to 2020. The winter index covers the period December – February, and we used this index to test for large-scale variation in timing of ice break-up as the winter index influences both winter precipitation and temperature.

2.4 Modelling and statistical analyses

2.4.1 Average time of ice break-up and freezing and length of ice-free period

We tested for variation in timing of the different phenological events using general linear models (GLM) and model selection procedures. Based on prior knowledge, we assumed that these timing traits would vary depending on longitude (Long), latitude (Lat), and elevation above sea level (Ele, m) and that there might be interactions among these traits. Further, we assumed that distance to the sea might be important as it impacts on both precipitation and temperature. We estimated the distance from each lake to the sea as distance from the outlet of the lake to the coastal shelf (a line drawn between the outermost islands along the coast) on maps (1:1,000,000). An increasing distance from the coastal shelf line reflects an increasing importance of continental climate. As the coastline of Norway bends eastwards at increasing latitude, the coastal distance may more correctly reflect oceanic/continental climate than longitude.

Various lake and catchment characteristics may also have an impact on ice phenology. Thus, in this analysis we used total lake surface area (Area, km^2), total catchment area (Catch, km^2) and annual mean inflow (Flow, m^3) as descriptors.

We started by evaluating a full model including all relevant parameters (equation 1):

Equation 1: $Y = \mu + \alpha_1 Ele + \alpha_2 Lat + \alpha_3 Long + \alpha_4 Ele*Lat + \alpha_5 Ele*Long + \alpha_6 Long*Lat + \alpha_7 Ele*Long*Lat + \alpha_8 Distance + \alpha_9 Area + \alpha_{10} Catch + \alpha_{11} Flow + \varepsilon$.

We then compared the models using a backward selection procedure by removing the least important parameters until we ended with the "best model". Models were compared with the corrected Akaike Information Criteria (AIC_c) (Burnham and Anderson, 1998). Models with AIC_c values 2 units below that of a competing model are assumed to give the better fit to the data. When presenting the results of the model selection (se Appendix 3) we present the AIC_c values for the three best models as well as the full model and present the best model by giving parameter estimates and overall model results (in the Result chapter).

2.4.2 Temporal variation in timing of ice break-up, freeze-up and length of ice-free period

We used several different approaches to test for temporal variation in the different ice phenology traits.

Firstly, in order to identify the main parameters influencing variation in time of freeze-up, time when lakes were completely frozen over and length of the ice-free period, we used general linear mixed models (GLMM), using basically the same parameters as in our average modelling approach (Equation 2) (see Appendix 4). Year was, however, always included as a continuous variable to test for linear temporal trends. In addition, the parameters Regulated (yes/no) and water level amplitude (Amplitude, m) were always either excluded or included in parallel in the analyses. To account for temporal autocorrelation of observations from the same lake we included lake identity as a random factor (random intercept) in the analyses. We used the same model selection procedure as above, but always kept year as a fixed factor. We selected the best model based on the AIC criterion (Burnham and Anderson, 2004).

Equation 2: $Y = \mu + (\text{Lake:Random intercept}) + \alpha_1\text{Ele} + \alpha_2\text{Lat} + \alpha_3\text{Long} + \alpha_4\text{Ele*Lat} + \alpha_5\text{Ele*Long} + \alpha_6\text{Long*Lat} + \alpha_7\text{Ele*Long*Lat} + \alpha_8\text{Distance} + \alpha_9\text{Area} + \alpha_{10}\text{Catch} + \alpha_{11}\text{Flow} + \alpha_{12}\text{Year} + \alpha_{13}\text{Regulated} + \alpha_{14}\text{Amplitude} + \varepsilon$.

Secondly, to test for temporal variation in timing of ice break-up, we used the same general linear mixed models, with lake as a random variable (random intercept) and year was always included as a fixed parameter to test for temporal trends (Equation 2). Here, we also included a large-scale climate index in the modelling (see Appendix 5). We included both a linear and a non-linear (squared) effect of NAO as potential drivers of variation in the timing of ice break-up. NAO-estimates are only available starting in 1899. Thus, this analysis covers a shorter time frame than the other traits. We selected the best model based on the AIC criterion (Burnham and Anderson, 2004).

Thirdly, we wanted to investigate if there has been any non-linearity in the temporal trends. Numerous papers indicate that large-scale climatic changes have occurred mainly during recent years (Blenckner et al., 2004; Mishra et al., 2011; Post et al., 2018), especially during the last decades. We therefore selected several lakes (N = 35) with long and complete data series and analysed for temporal trends in four different 30-year periods (1901-1930, 1931-1960, 1961-1990, 1991-2020). In these analyses we applied a simplified approach. We used a general mixed modelling approach, with ice phenology as response variable, year as predictor, and lake identity as random factor. We thus assume that all lakes have the same temporal trends (same slope) within each time period. Including a random slope did not change the conclusions.

All statistical analyses were performed using JMP 12 (JMP Version 12. SAS Institute Inc., Cary, NC, 1989-2019).

3 Results

All lakes had distinct periods without ice every year. The observations of average timing of ice break-up, time of lake freeze-up, time when the lake was completely frozen and length of ice-free period were strongly correlated (Figure 3, Table 1).

3.1 Spatial variation in average ice phenology

We tested for drivers of variation in average time of ice break-up, lake freeze-up, time when a lake is completely frozen over and the length of the ice-free period. A summary of the model selection results is presented in Appendix 4.

The spatial variation in average time of ice break up was best explained by a complex model including a three-way interaction between latitude, longitude and elevation (Table 2). The best model did, however, include a weak negative effect of annual inflow to the lake, but not distance to the sea. Distance to sea was, however, included in a model within 0.4 AIC_c units of the best model. There were only small effects of the various lake characteristics, but ice break-up was later with increasing latitude (2.3 days per °N), longitude (1.5 days per °E) and elevation (3.4 days per 100 m) (Figure 4). The lakes are situated geographically such that latitude and longitude are strongly positively correlated (r = 0.825, p< 0.001), longitude and elevation are negatively correlated (r = -0.404, p< 0.001), and latitude and coastal distance are negatively correlated (r = -0.479, p< 0.001), indicating that the effects should be interpreted with caution. Furthermore, there was large within-lake variability in timing of ice break-up (Table 3), with an average coefficient of variation (CV; defined as standard deviation divided by the mean) of 8.90 %. Within-lake CV was negatively correlated with latitude, longitude, elevation and distance to the coastline. This indicates larger phenological variation in lakes in southern and western areas and at lower elevation.

The best models explaining variation in the timing of lake freeze-up, time when the lake is completely frozen, and the length of the ice-free period usually contained an interaction effect between longitude and elevation. All models also included a positive effect of lake surface area (Table 2, Appendix 3). Overall, lakes freeze-up earlier and have a shorter ice-free period with increasing longitude and elevation. Large lakes also take longer to freeze and were ice-free for longer than smaller lakes. The within-lake variation in timing of freeze-up (mean CV = 4.45 %) and when the lake was completely frozen (mean CV = 4.55 %) was less than the variation in the length of the ice-free period (mean CV = 15.04 %). The CV of these three phenological traits were negatively correlated with elevation and coastal distance (Table 3). The effect of longitude was more variable.

3.2 Temporal variation in timing of lake freeze-up, time when the lake is completely frozen and length of ice-free period

The best models, based on the AIC_c criterion, for timing of lake freeze-up, time when the lake was completely frozen and the length of the ice-free period contained geographic parameters such as elevation, latitude and longitude (Appendix 4). Lake surface area also had a positive effect on all these three phenological traits. In addition, lake regulation and the amplitudinal range in water level had an impact on all traits. There was little temporal variation in these traits on the long timescale analysed here; only for when the lake was completely frozen over, did we find a significant (p<0.001) positive temporal trend, indicating that the lakes are completely frozen later in the autumn in recent years (Table 4).

3.3 Temporal trends in timing of ice break-up

The best model for the timing of ice break-up included the effects of geography, time and climate (Appendix 5). Ice break-up occurred later during spring with increasing elevation, latitude, and longitude. These effects are complex, as indicated by the various significant interaction effects. In addition, there was a significant negative temporal trend in ice break-up, i.e. ice break-up occurred earlier in the spring (Table 5). There was also a significant climate effect, with a negative linear effect of the NAO ($p<0.001$).

3.4 Non-linear temporal trends in ice phenology

Many studies indicate that climate is changing faster during recent decades. To investigate for potential non-linear trends in ice phenology we analysed for temporal trends within four different time periods (1901-1930, 1931-1960, 1961-1990, 1991-2020). We selected 35 lakes with relatively long, and continuous data series exceeding 50 years for both date of break-up and date of completely frozen lake (Appendix 6). We used a period-specific mixed model, assuming similar temporal trends (slopes) for all lakes (random intercept only). During the three first time periods none of the slope estimates were significant (Figure 5, Table 6), whereas during the last time period (1991-2020) most temporal trends were significant. During this period ice break up happened approximately 2 days earlier per decade, whereas time of ice freeze-up and time when lake is completely frozen were on average 6 and 3 days later per decade. Furthermore, the length of the ice-free period has become 7 days longer per decade, although this effect was marginally non-significant ($p = 0.068$).

4 Discussion

Our analysis of ice phenology of 101 Norwegian lakes covering the period from the 1890s to 2020 gave two major results. Firstly, the analysis indicated significant trends in ice phenology in recent years. Ice break-up occurred earlier, ice freeze-up and completely frozen occurred later, and all trends were accelerating. This results in a longer ice-free season. Secondly, the coefficient of variation in the different ice phenology variables were larger in lakes in southern and western areas and at lower elevation, indicating that lakes in these areas are most influenced by climate change.

4.1. Geographical parameters

The investigated lakes cover a range of climatic zones in a latitudinal, longitudinal and elevational perspective. These variables clearly showed complex and significant interactions, especially for ice break-up, indicating the problems in illuminating the individual importance of the geographical parameters. The date of break-up generally occurs later with increasing latitude, modified by macro- to local scale atmospheric circulation and lake characteristics (Blenckner et al. 2004, Livingstone et al. 2009). Our results support this latitudinal trend, but we also found that longitude, elevation, and lake size had significant effect.

Ice break-up dates are shown to be 2.3 days later with each degree of higher latitude. This is considerably less than previously documented in Fennoscandia (3.3-5.4 days/°N) (Efremova et al., 2013; Blenckner et al., 2004) and in North America 3.5 days/°N (Williams et al., 2006). There is no obvious reason for this difference. One possible explanation could be that registration of ice parameters differs both within and between studies. Moreover, the oceanic effect could modify the relationship as the majority of lakes in northern Norway are situated close to the ocean in contrast to the southern lakes that are mostly continental.

Moreover, we found that ice break-up was 3.4 days delayed by a 100 m increase in elevation. This is also slightly lower than in Karelian lakes where Efremova et al. (2013) found a delay of 5 days/100 m. Although there is considerable climatic difference between Norway and Karelia as Karelian lakes in general experience a more continental climate, the Karelian lakes also covers less variation in elevation.

Although several studies have investigated ice phenology in Europe, most of them have not included longitude in their analyses. One reason could be the complexity of the parameter. In contrast to latitude which reflect insolation received, longitude reflects more the distance from the coast. However, one exception is the study of Polish lakes by Wrzesinski et al. (2015). The lakes are situated in the northern region and covered a wide longitudinal range (14 – 24° E), although a somewhat smaller range compared to the Norwegian lakes. Wrzesinski et al. (2015) found that break-up increased by 1 day/°E, compared to 1.5 days/°E in our study. The location of the Polish lakes indicate that any effect of the Baltic Sea is similar. In contrast, in Norway the climate becomes more continental when moving eastwards, especially south of 61° N where the mountain chain that runs north-south creates a distinct difference in climate from west to east. Thus, the longitudinal effect could as well be due to the climatic conditions as the proximity to the ocean renders the climate milder in the west. The longitudinal effect should therefore be treated with caution. However, the global study by Sharma et al. (2019) showed that distance to the coast was important in determining whether lakes had annual winter ice cover. In our analysis the distance from ocean did not per se have any significant effect of any of the ice phenology parameters.

Our results demonstrated a complex relationship among geographical parameters describing date of freeze-up. The best models explaining variation in the timing of lake freeze-up contained an interaction effect between longitude and elevation, in addition to a positive effect of lake surface area. This differs from the results from other studies in the region. The Karelian lakes, covering 54-68° N, freeze-up 2.3 days earlier for every degree of increasing latitude (Efremova et al., 2013), while Swedish (58-68° N) and Finnish (61-69° N) lakes freeze-up 2.8 and 4.5 days earlier for each degree of increasing latitude, respectively (Blenckner et al., 2004). The most obvious explanation for this discrepancy is due to altitudinal variation. The Norwegian lakes cover 1400 m in elevation range, whereas the lakes in Karelia are all situated lower than 204 m, in Sweden lower than 340 m and in Finland lower than 473 m. An additional complicating factor is the oceanic climate that, if anything, is more pronounced for Norwegian lakes than lakes in Sweden, Finland and Karelia.

In our model, distance from the coast does not significantly contribute neither to freeze-up nor break-up date, probably as distance to the coast was included both in the latitude and longitude variables. This in in contrast to the analyses of 41 Finnish lakes where a pronounced deflection of isolines of both freeze-up and break-up date northward near the Baltic Sea coast was documented (Palecki and Barry, 1986; Korhonen, 2006).

The predictable seasonal cycle in solar radiation is characteristic of higher latitudes. Weyhenmeyer et al. (2011) hypothesised, based on a global dataset, that lakes north of 61° N had lower inter-annual variability in seasonal cycle than lakes at latitudes lower than 61° N. The Norwegian lakes are distributed along a latitudinal gradient to test this hypothesis in a robust way. Our results lend support to this, as the within-lake coefficient of variation (CV) of ice break-up, freeze-up and length of ice-free season were negatively correlated with latitude, longitude, elevation and/or distance to coastline. This indicates larger phenological variation in lakes in southern and western areas and at lower elevation.

4.5 Temporal trends

Although many studies have documented trends in ice phenology, few studies have investigated changes across specific periods to elucidate periods with stronger trends. In a study of global datasets Benson et al. (2012) and Newton and Mullan (2020) showed that trends in ice variables were steeper over the last 30-year period. Similar increase in trends in the last two decades have been shown for Karelian lakes (Efremova et al., 2013) and the Great Lakes region (Mishra et al., 2011).

Our analyses revealed significant, accelerating trends for earlier break-up, later freeze-up and lately completely frozen lakes after 1991. Moreover, the trend for a longer ice-free period also accelerated during this period, although the trend was not significant. These trends are in accordance with an increase in air temperature in the spring and autumn, as well for the global temperature over the last decades (Benson et al., 2012; Hansen et al., 2006). Our results are in accordance with Newton and Mullan (2019), showing marked differences in ice phenology in Fennoscandian lakes (Sweden, Finland) across 30-year periods after 1931. Newton and Mullan (2020) found that break-up appeared to be earlier and trends more pronounced in southern regions during the first period. In the next period, 1961-1999, break-up trends increased in magnitude, and the lakes with negative trends in the previous period shifted to be positive. In the last period, the strength of the trends in earlier break-up increased and reached 3.9 days/decade. In our study, the trend in the 1991-2020 was 2.0 days/decade. Korhonen (2006) also showed a significantly earlier break-up of 6-9 days over a hundred years for Finnish lakes, although the data was not analysed in 30-years periods. One plausible reason for a slower trend in Norwegian lakes during this period than in the rest of Fennoscandia is the influence of the ocean. There has been considerable change in the 0 °C isotherm with a marked reduction in the area below 0 °C (see Figure 6). The change is recorded throughout Norway but at a larger scale in the continental areas between 61 and 63° N. A significant correlation between break-up and 0 °C isotherm has been documented for Arctic lakes in North America, Europe and Siberia (Šmejkalova et al. 2016). The extension of the Gulf Stream, the North Atlantic Drift, along the Norwegian coast contributes to a mild climate and reduced climate. The speed of thermal change in the ocean is less rapid and less variable than in inland waters (Woolway and Maberly, 2020).

Changes in ice phenology depend on several climatic forcing variables, such as air temperature, solar radiation, wind and snowfall (Magnusson et al. 1997). A significant increase in global air temperature during the last century is well documented (e.g. Hansen et al., 2006; Robinson, 2020). Newton and Mullan (2020) showed that rising temperature appears to be the dominant factor for the shift towards earlier break-up and later freeze-up in the Northern Hemisphere. Precipitation may also play a role in the observed trends. Nordli et al. (2007) found a significant correlation (R^2=0.58) between date of break-up in lake Randsfjorden and the mean temperature in February to April. Duguay et al. (2006) showed that trends towards later freeze-up corresponded with areas of increasing autumn snow cover, and that spatial trends in break-up were consistent with changes in spring snow cover duration. Similarly, Jensen et al. (2007) in a study of ice phenology trends across the Laurentian Great Lakes region found that variability in the strength of trends in earlier break-up were partly explained by number of snow days or snow depth. For the lake Litlosvatn, in the mountain area of western Norway, Borgstrøm (2001) found a clear relationship between spring snow depth and the date on which the lake was free of ice. The altitudinal gradient causes considerable regional difference in annual precipitation in Norway (Hanssen-Bauer, 2005). The general trend in increasing temperature and precipitation observed from 1875 to 2004, has been modelled to increase to 2100, although there will be regional differences (Hanssen-Bauer et al., 2017). Thus, our results concerning the recent trends in ice phenology probably indicate a new situation for ice formation in Norwegian lakes. However, this is in agreement with a general trend in the Northern Hemisphere shown by an increase in extreme events for lake ice (Filazzola et al., 2020).

Biological consequences

Shifts in ice phenology have major repercussions for the biota of lakes and rivers (Prowse, 2001; Prowse et al., 2011; Caldwell et al., 2020), as ice cover changes the aquatic environment, not only in terms of light penetration, but also the physical characteristics of the environment such as temperature. Of special interest is that the trend in earlier ice break-up and the loss of ice will stimulate biological production. In late autumn, solar insulation is restricted and thus, a prolonged period without ice has limited consequences for aquatic production. Caldwell et al. (2020) tested a conceptual model that expressed how earlier break-up affected aquatic ecosystems. The effect differed between and within tropic levels. Whereas contrasting effects were found between littoral and pelagic zooplankton production, the modelled brook trout (*Salvelinus fontinalis*) did not profit from the increased zooplankton production and experienced reduced fitness. A review of the long-term dynamics of fish species in Europe (Jeppesen et al. 2011), revealed a shift towards higher dominance of eurythermal species. Loss of ice cover increased resting metabolism by approximately 30 % in an Atlantic salmon (*Salmo salar*) population (Finstad et al., 2004), and the recruitment of an alpine brown trout (*Salmo trutta*) population was strongly affected by accumulated snow depth and thereby the timing of ice-break (Borgstrøm and Museth, 2005). Moreover, the outcome of competition in sympatric populations of brown trout and Arctic charr (*Salvelinus alpinus*) is strongly dependent on the duration of ice-cover as high charr abundance is correlated with low trout population growth rate only in combination with long winters (Helland et al., 2011). In addition, aquatic insects, such as Ephemeroptera and Plecoptera may change their voltinism and their emergence timing in a warmer climate (Brittain 1978, 2008; Sand & Brittain 2009). We still have limited knowledge about how climate change in general may have impacts on Arctic and Alpine fishes and fish populations (Reist et al., 2006). This is also the case with changes in ice phenology. The biological consequences of changes in ice phenology will occur first and be most marked in lakes with high coefficient of variation in the ice phenology parameters; that is, in lakes situated in the coastal lowlands, in the southernmost part, and in the eastern part of southern Norway. We suggest that biological consequences will be small in these areas as ubiquitous species with wide environmental limits often dominate, although those species dependent on regular ice conditions will be replaced by species with a more flexible life history (Brittain 2008; Brittain et al. 2020). In the long-term this replacement is also likely to occur in lakes at higher elevation as ice cover duration decreases and becomes more variable at the same time as winter temperature increases.

5 Conclusions

Ice phenology is complex and determined by the interaction of a range of parameters. This study shows that elevation, latitude and longitude all significantly affect ice phenology in Norwegian lakes. Overall, the length of ice-free season becomes longer with increasing values of each parameter. Lake characteristics are of minor importance, although lake size had a significant effect. In addition, a significant temporal effect of changing climate was detected during 1991-2020 but not in the three earlier 30-year periods). During this latter period ice break up happened approximately 2 days earlier per decade, whereas timing of ice freeze-up and time when lakes are completely frozen was on average 6 and 3 days, respectively, later per decade. Furthermore, the length of the ice-free period has become 7 days longer per decade. These trends are shown to happen concomitantly with a considerable reduction in the area with annual mean air temperature below 0 °C. The reduction is most pronounced in continental areas between 61and 63 °N. An understanding of the relationship between ice phenology and geographical and climate parameters is a prerequisite for predicting the potential consequences of climate change on ice phenology and lake biota.

Data availability. All ice phenology data are available at doi:10.5061/dryad.bk3j9kd9x.

Author contributions. JHL-L designed this study. JHL-L, LAV and JEB led the writing of this paper. LAV conducted the formal analysis. Data curation was conducted by JHL-L, ÅSK and TS. JHL-L collated basic characteristics for individual lakes.

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgements. We would like to acknowledge Glommens og Laagens Brukseierforening hydropower company for giving access to ice phenology of 13 lakes. Halvor Lien provided observation of ice phenology of lake Møsvatn which was carried out by Halvor Hamaren until 1987, and himself afterwards. We acknowledge Julio Pereira and Henrik L'Abée-Lund for technical assistance. Ole Einar Tveito, The Norwegian Meteorological Institute, kindly supplied maps of the zero-degree isotherm.

References

Adrian, R., O'Reilly, C.M., Zaragese, H., Baines, S.B., Hessen, D.O., Keller, W., Livingstone, D.M., Sommaruga, R., Straile, D., Van Donk, E., Weyhenmeyer, G.A., and Winder, M.: Lakes as sentinels of climate change, Limnol. Oceanogr., 56, 2283-2297, 2009.

Benson, B.J., Magnusson, J.J., Jensen, O.P., Card, V.M., Hodgkins, G., Korhonen, J., Livingstone, D.M., Stewart, K.M., Weyhenmeyer, G.A., and Granin, N.G.: Extreme events, trends, and variability in Northern Hemisphere lake-ice phenology (1855-2005), Clim. Change, 112, 299-323, https://doi.org/10.007/s10584-011-0212-8, 2012.

Blenckner, T., Järvinen, M., and Weyhenmeyer, G.A.: Atmospheric circulation and its impact on ice phenology in Scandinavia, Boreal Env. Res., 9, 371-380, 2004.

Borgstrøm, R.: Relationship between spring snow depth and growth of brown trout, *Salmo trutta*, in an Alpine lake: Predicting consequences of climate change, Arct., Antarct. Alp. Res., 33, 476-480, https://doi.org/10.1080/15230430.2001.12003457, 2001.

Feltkode endret

Borgstrøm, R. and Museth, J.: Accumulated snow and summer temperature – critical factors for recruitment to high mountain populations of brown trout (*Salmo trutta* L.), Ecol. Freshw. Fish, 14, 375-384, https://doi.org/10.1111/j.1600-0633.2005.00112.x, 2005.

Brittain, J.E.: Semivoltinism in mountain populations of *Nemurella pictetii* (Plecoptera), Oikos, 30, 1-6, 1978.

Brittain, J.E.: Mayflies, biodiversity and climate change, in International advances in The ecology, zoogeography, and systematics of mayflies and stoneflies edited by Hauer, F.R., Stanford, J.A., and Newell, R.L., University of California Publications in Entomology, 128, 1-14, 2008.

Brittain, J. E., Heino, J., Friberg, N., Aroviita, J., Kahlert, M., Karjalainen, S. M., Keck, F., Lento, J., Liljaniemi, P., Mykrä, H., Schneider, S.C., and Ylikörkkö, J.: Ecological correlates of riverine diatom and macroinvertebrate alpha and beta diversity across Arctic Fennoscandia, Freshwat. Biol, https://doi.org/10.1111/fwb.13616, 2020.

Brown, L.C. and Duguay, C.R.: The response and role of ice cover in lake-climatic interactions, Progr. Phys. Geogr., 34, 671-704, https://doi.org/10.1177/0309/333/0375653, 2010.

Burnham, K. P. and Anderson, D. R.: Model selection and inference: a practical information-theoretic approach. Springer Verlag, New York, 1998.

Burnham, K. P. and Anderson, D. R.: Multimodel inference. Understanding AIC and BIC in model selection, Sociol. Methods Res., 33, 261-304, https://doi.org/10.1177/0049124104268644, 2004.

Caldwell, T.J., Chandra, S., Feher, K., Simmons, J.B., and Hogan, Z.: Ecosystem response to earlier ice break-up date: Climate-driven changes to water temperature, lake-habitat-specific production, and trout habitat and resource use, Glob. Change Biol., 26, 5475-5491, https://doi.org/10.1111/gcb.15258, 2020.

Choiński, A., Ptak, M., Skowron, R., and Strzelczak, A.: Changes in ice phenology on polish lakes from 1961 to 2010 related to location and morphometry, Limnologica 53, 42-49, http://dx.doi.org/10.1016/j.limn0.2015.05.005, 2015.

Du, J.D., Kimball, J.S., Duguay, C., Kim, Y., and Watts, J.D.: Satellite microwaves assessment of Northern Hemisphere lake ice phenology from 2002 to 2015, The Cryosphere, 11, 47-63, https://doi.org/10.5194/tc-11-47-2017, 2017.

Duguay, C.R., Prowse, T.D., Bonsal, B.R., Brown, R.D., Lacroix, M.P., and Menard, P.: Recent trends in Canadian lake ice cover, Hydrol. Process., 20, 781-801, https://doi.org/10.1002/hyp.6131, 2006.

Efremova, T., Palshin, N., and Zdorovennov, R.: Long-term characteristics of ice phenology in Karelian lakes, Estonian J. Earth Sci., 62, 33-41, https://doi.org/10.3176/earth.2013.04, 2013.

Eklund, A.: Islägging och islossning i svenska sjöar. SMHI Hydrologi, no. 81, 24 pp, 1999.

Finstad, A.G., Forseth, T., Næsje, T.F. and Ugedal, O.: The importance of ice cover for energy turnover in juvenile Atlantic salmon, J. Animal Ecol., 73, 959-966, https://doi.org/10.1111/j.0021-8790.2004.00871.x, 2004.

Filazzola, A., Blagrave, K., Imrit, M.A., and Sharma, S.: Climate changes drives increases in extreme events for lake ice in the Northern Hemisphere, Geophys. Res. Letters 47, e2020GL089608, https://doi.org/10.1029/2020GL089608, 2020.

George, D.G., Järvinen, M., and Arvola, L.: The influence of the North Atlantic Oscillation on the winter characteristics of Windermere (UK) and Pääjärvi (Finland), Boreal Environ. Res. 9, 389-399, 2004.

Hansen, J., Sato, M., Ruedy, R., Lo, K., Lea, D.W., and Medina-Elizade, M.: Global temperature change, PNAS 103, 14288-14293, www.pnas.org/cgi/doi/10.1073/pnas.0606291103, 2006.

Hanssen-Bauer, I.; Regional temperature and precipitation series for Norway: Analyses of time-series updated to 2004, Norwegian Meteorological Institute, report no. 15, 2005.

Hanssen-Bauer, I., Førland, E.J., Haddeland, I., Hisdal, H., Mayer, S., Nesje, A., Nilsen, J.E.Ø., Sandven, S., Sandø, A.B., Sorteberg, A., and Ådlandsvik, B.: Climate in Norway 2100 – a knowledge base for climate adaptation, NCCS report no. 1, 2017.

Helland, I.P., Finstad, A.G., Forseth, T., Hesthagen T., and Ugedal, O.: Ice-cover effects on competitive interactions between two fish species, J. Anim. Ecol., 80, 539–547, https://doi.org/ 10.1111/j.1365-2656.2010.01793.x, 2011.

Hewitt, B.A., Lopez, L.S., Gaibisels, K.M., Murdoch, A., Higgins, S.N., Magnusson, J.J., Paterson, A.M., Rusak, J.A., Yao, H., and Sharma, S.: Historical trends, drivers, and future projections of ice phenology in small north temperate lakes in the Laurentian Grate Lakes watershed, Water, 10, 70, https://doi.org/10.3390/w10010070, 2018.

Hurrell, J.W.: Decadal trends in the North Atlantic Oscillation: regional temporal and precipitation, Science, 269, 676-679, 1995.

IPPC. Climate change 2007: The physical science basis, ed. by Solomon, S., Dahe, Q., Manning, M., Marquis, M., Averyt, K., Tignor, M.M.B., Miller, H.L. jr., and Chen, Z., Cambridge University Press, 2007.

Jensen, O., Benson, B.J., Magnuson, J.J., Card, V.M., Futter, M.N., Soranno, P.A., and Stewart, K.M.: Spatial analysis of ice phenology trends across the Laurentian Great Lakes region during a recent warming period, Limnol. Oceanogr., 52, 2013-2026, https://doi.org/10.2307/4502353, 2007.

Jeppesen, E., Mehner, T., Winfield, I.J., Kangur, K., Sarvala, J., Gerdeaux, D., Rask, M., Malmquist, H.J., Holmgren, K., Volta, P., Romo, S., Eckmann, R., Sandström, A., Blanco, S., Kangur, A., Stabo, H.R., Tarvainen, M., Ventelä, A.-M., Søndergaard, M., Lauridsen T.L., and Meerhoff, M.: Impacts of climate warming on the long-term dynamics of key fish species in 24 European lakes, Hydrobiologia, 694, 1-39, https://doi.org/10.1007/s10750-012-1182-1, 2012.

Knoll L.B., Sharma, S., Denfeld, B.A., Flaim, G., Hori, Y., Magnuson, J.J., Straile, D., and Weyhenmeyer, G.A.: Consequences of lake and river ice loss on cultural ecosystem services, Limnol. Oceanogr. Letters 4, 119-131, https://doi.org/10.1002/lol2.10116, 2019.

Korhonen, J.: Long-term changes in lake ice cover in Finland, Nordic Hydrology 37, 347–363, 2006.

Kuusisto, E. and Elo, A.-R.: Lake and rive ice variables as climate indicators in Northern Europe, Verh. Internat. Verein. Limnol. 27, 2761-2764, 2000.

Kvambekk, Å.S. and Melvold, K.: Long-term trends in water temperature and ice cover in the subalpine lake, Øvre Heimdalsvatn, and nearby lakes and rivers, Hydrobiologia, 642, 47-60, https://doi.org/10.1007/s10750-010-0158-2, 2010.

Livingstone, D.M., Large-scale climatic forcing detected in historical observations of lake ice break-up, Verh. Internat. Verein. Limnol. 27, 2775-2783, 2000.

Livingstone, D.M., Adrian, R., Blenckner, T., George, G., and Weyhenmeyer, G.A.: Lake ice phenology, in The impact of climate change in European lakes edited by George, G., Aquat. Ecol. Ser., 4, 51-61, https://doi.org/10.1007/978-90-481-2945-4_4, 2009.

Magnuson, J.J., Webster, K.E., Assel, R.A., Bowser, C.J., Dillon, P.J., Eaton, J.G., Fee, E.J., Hall, R.I., Mortsch, L.R., Schindler, D.W., and Quinn, F.H.: Potential effects of climate changes on aquatic systems: Laurentian Great Lakes and Precambrian shield region., Hydrol. Processes, 11, 825-871, 1997.

Magnuson, J.J., Robertson, B.M., Benson, B.J., Wynne, R.H., Livingstone, D.M., Arai, T., Assel, R.A., Barry, R.G., Card, V., Kuusisto, E., Granin, N.G., Prowse, T.D., Stewart, K.M., and Vuglinski, V.S.: Historical trends in lake and river ice cover in the Northern Hemisphere, Science, 289, 1743-1746, 2000.

Mishra, V., Cherkauer, K.A., Bowling, L.C., and Huber, M.: Lake ice phenology of small lakes: impact of climate variability in the Great Lakes region, Glob. Planet. Change, 76, 166-185, https://doi.org/10.1016/j.gloplacha.2011.01.004, 2011.

Newton, A.M.W. and Mullan, D.: Climate change and Northern Hemisphere lake and river ice phenology, The Cryosphere, https://doi.org/10.5194/tc-2020-172, 2020.

Nordli, Ø., Lundstad, E., and Ogilvie, A.E.J.: A late-winter to early-spring temperature reconstruction for southeastern Norway from 1758 to 2006., Annales Glaciol., 46, 404-408, https://doi.org/10.3189/172756407782871657, 2007.

Feltkode endret

Palecki, M.A. and Barry, R.G.: Freeze-up and break-up of lakes as an index of temperature changes during the transition seasons: a case study for Finland, J. Appl. Met., 25, 893-902, https://doi.org/10.1175/1520-0450(1986)025<0893:FUABUO>2.0.CO;2, 1986.

Post, E., Steinman, B.A., and Mann, M.E.: Acceleration of phenological advance and warming with latitude over the past century, Scientific Reports 8, 3927, https://doi.org/10.1038/s41598-018-22258-0, 2018.

Prowse, T.D.: River ice ecology. II: Biological aspects, J. Cold Reg. Engineering, 15, 17-33, https://doi.org/10.1061/(ASCE)0887-381X(2001)15:1(17), 2001.

Prowse, T., Alfredsen, K., Beltaos, S., Bonsal, B.R., Bowden, W.B., Duguay, C.R., Korhola, A., McNamara, J., Vincent, W.F., Vuglinsky, V., Walter, K.M.A., and Weyhenmeyer, G.A.: effects of changes in Arctic lake and river ice, Ambio 40, 63-74, https://doi.org/10.1007/s13280-011-0217-6, 2011.

Reist, J. D., Wrona, F.J., Prowse, T.D., Power, M., Dempson, J.B., Beamish, R.J., King, J.R., Carmichael, T.J., and Sawatzky, C.D.: General effects of climate change on Arctic fishes and fish populations, Ambio, 35(7), 370-380, https://doi.org/10.1579/0044-7447, 2006.

Robinson, S.-A.: Climate change adaptation in SIDS: A systematic review of the literature pre and post the IPCC Fifth Assessment Report, Climate Change 11(3), https://doi.org/10.1002/wcc.653, 2020.

Sand, K. and Brittain, J.E.: Life cycle shifts in *Baetis rhodani* (Ephemeroptera) in the Norwegian mountains, Aquat. Insects, 31, 283-291, 2009.

Sharma, S. and Magnusson, J.J.: Oscillatory dynamics do not mask linear trends in the timing of ice breakup for Northern Hemisphere lakes from 1855-2004, Clim. Change 124, 835-847, https://doi.org/10.1007/s10584-014-1125-0, 2014.

Sharma, S., Magnuson, J.J., Batt, R.D., Winslow, L.., Korhonen, J., and Aono, Y.: Direct observations of ice seasonality reveal changes in climate over the past 350-570 years, Science Rep. 6, 25061, https://doi.org/10.1038/srep25061, 2016.

Sharma, S., Blagrave, K., Magnusson, J.J., O'Reilly, C.M., Oliver, S., Batt, R.D., Magee, M.R., Straile, D., Weyhenmeyer, G.A., Winslow, L., and Woolway, R.I.: Widespread loss of lake ice around the Northern Hemisphere in a warming world, Nat. Clim. Change, 9, 227-231, https://doi.org/10.1038/s41558-018-0393-5, 2019.

Šmejkalova, T., Edwards, M.E., and Dash, J.: Arctic lakes show strong decadal trend in earlier spring ice-out, Scientific Reports 6, 38449, https://doi.org/10.1038/srep38449, 2016.

Solvang, T.: Historical Trends in Lake and River Ice Cover in Norway: signs of a changing climate, Master Thesis, Department of Geosciences, University of Oslo, http://urn.nb.no/URN:NBN:no-39915, 2013.

Stenseth, N. C., Ottersen, G., Hurrell, J.W., Mysterud, A., Lima, M., Chan, K.-S., Yoccoz, N.G., and Ålandsvik, B.: Studying climate effects on ecology through the use of climate indices: The North Atlantic Oscillation, El Nino Southern Oscillation and beyond. Proc. R. Soc. London, Series B, 270, 2087-2096, https://doi.org/10.1098/rspb.2003.2415, 2003.

Takács, K., Kern, Z., and Pásztor, L.: Long-term ice phenology records from eastern-central Europe, Earth System Science Data, 10, 391-404, https://doi.org/10.5194/essd-10-391-2018, 2018.

Tvede, A.M.: Hydrology of Lake Atnsjøen and River Atna, Hydrobiologia, 521, 21-23, 2004.

Weyhenmeyer, G.A., Westoo, A.-K., and Willen, E.: Increasingly ice free winters and their effect on water quality in Sweden's largest lakes, Hydrobiologia, 599, 111-118, https://doi.org/10.1007/978-1-4020-8379-2_13, 2008.

Weyhenmeyer, G.A., Livingstone, D.M., Meili, M., Jensen, O., Benson, B., and Magnusson, J.J.: Large geographical differences in the sensitivity of ice-covered lakes and rivers in the Northern Hemisphere to temperature changes, Glob. Change Biol., 17, 268-275, https://doi.org/10.1111/j.1365-2486.2010.02249.x, 2011.

Williams, S.G. and Stefan, H.G.: Modelling lake ice characteristics in North America using climate, geography, and lake bathymetry, J. Cold Reg. Engineering, 20, 140-167, https://doi.org/10.1061/(ASCE)0887-381X(2006)20:4(140), 2006.

Woolway, R.I. and Maberly, S.C.: Climate velocity in inland standing waters, Nature Climate Change, https://doi.org/10.1038/s41558-020-0889-7, 2020.

Wrzensinski, D., Choinski, A., Ptak, M., and Rajmund, S.: Effect of North Atlantic Oscillation on the pattern of lake ice phenology in Poland, Acta Geophys., 63, 1664-1684, https://doi.org/10.1515/acgea-2015-0055, 2015.

Yoo, J.C., P., D'Odorica, P.: Trends and fluctuations in the dates of ice break-up of lakes and rivers in Northern Europe: the effect of the North Atlantic Oscillation, J. Hydrol., 268, 100-112, 2002.

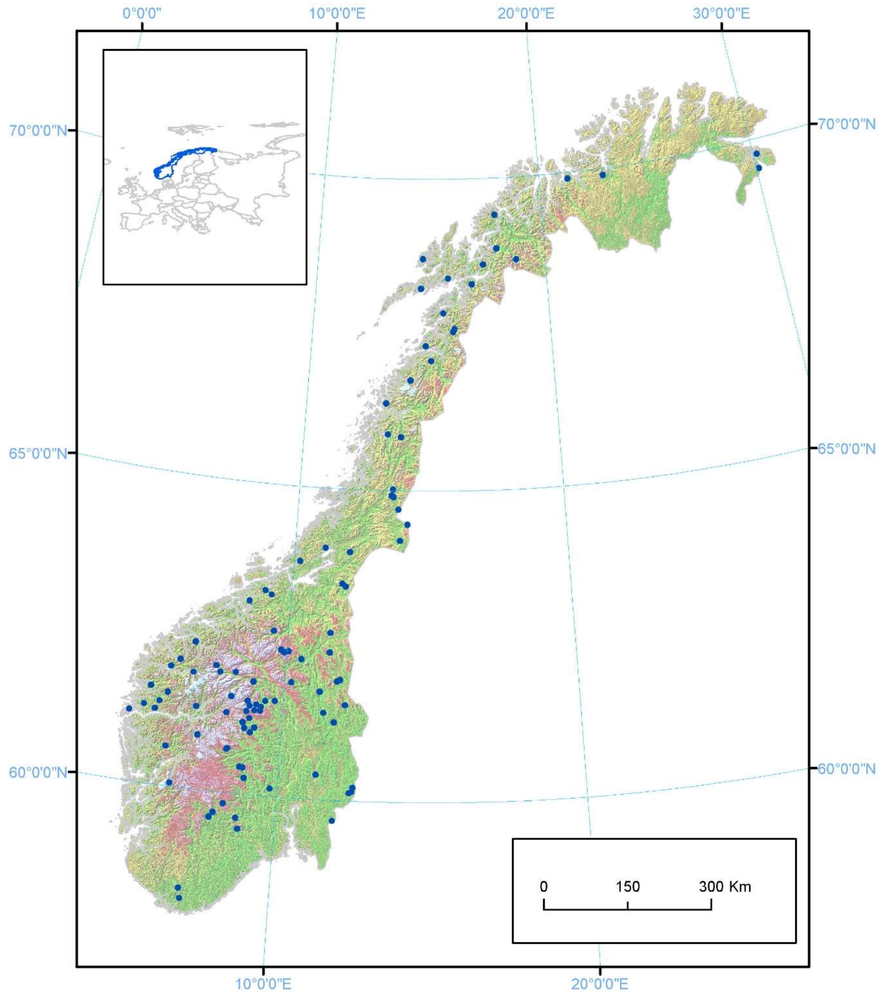

Figure 1. Topographic map of Norway with the 101 lakes included in the analysis. Information on the locations and names of the lakes is given in Table S1 in the online Supplement.

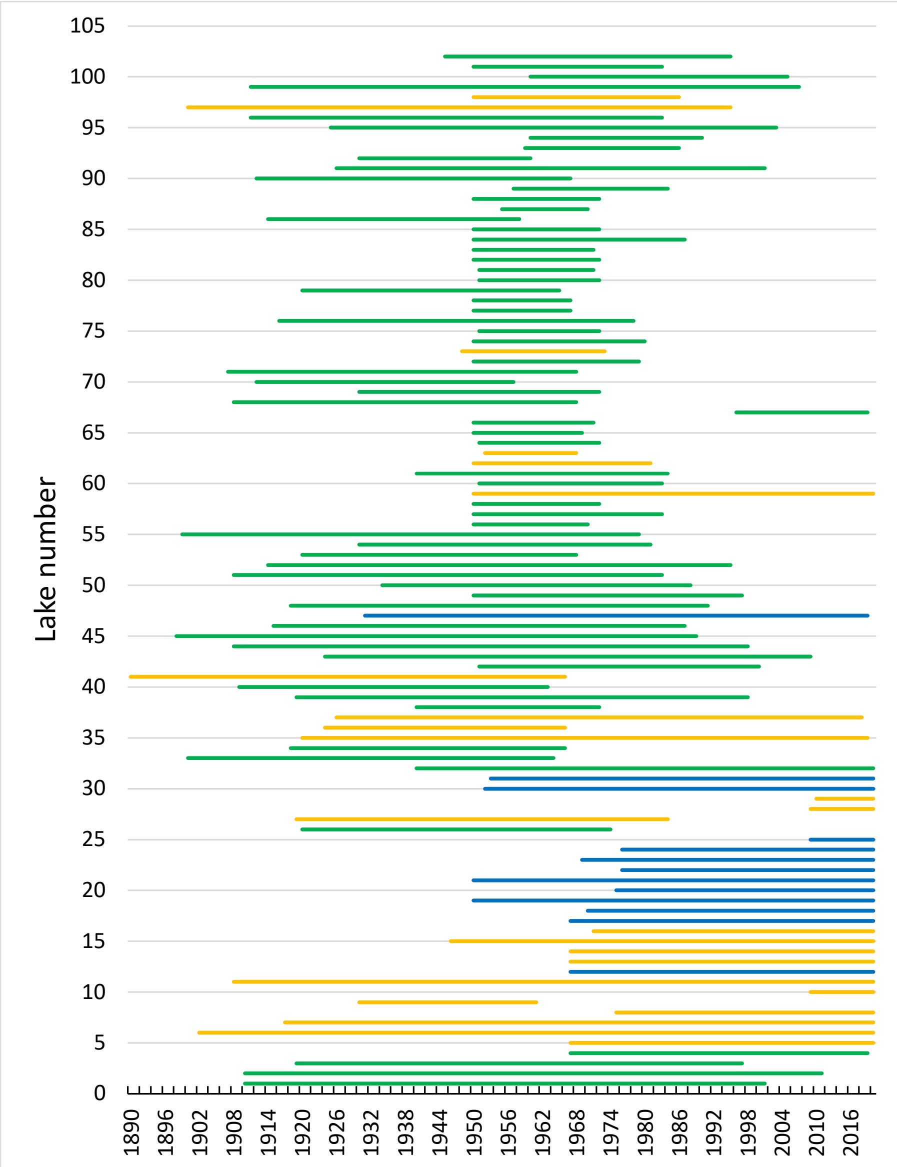

Figure 2. Chart showing the registration periods (1890-2020) for ice phenology (ice freeze-up, frozen lake and ice break-up) for individual lakes. For Lake 41, registration started in 1818 but was not continuous. In several data series there are years with missing registration of variables. The colour indicate elevation for each lake (green: <500 m a.s.sl.; yellow: 500-1000 m a.s.l; blue: >1000 m a.s.l.). For information on each lake see Appendix 1 and 2.

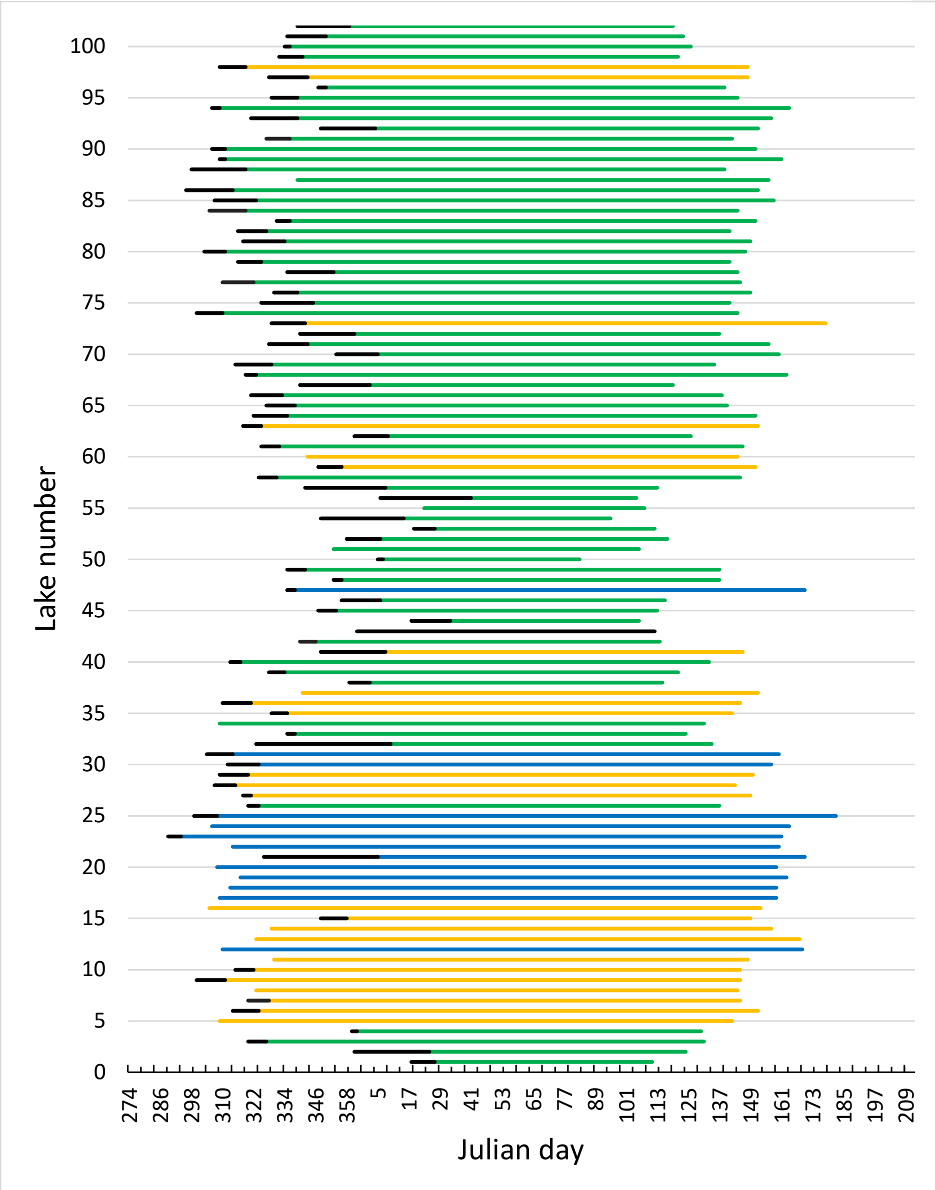

Figure 3. Median date of freeze-up (black line), frozen lake and break-up (coloured lines) for 101 Norwegian lakes during 1890-2020. The X-axis start at 274 (October 1) and end at 212 (July 31). The colour indicate elevation for each lake (green: <500 m a.s.sl.; yellow: 500-1000 m a.s.l; blue: >1000 m a.s.l.). For information on each lake see Appendix 1 and 2.

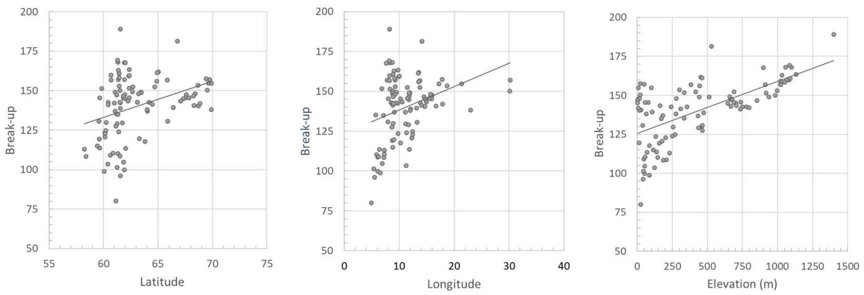

Figure 4. The correlation between the average timing of ice break-up and latitude, longitude and elevation of 101 Norwegian lakes during the period 1890-2020. The lines represent best linear fit. Left panel r = 0.345, p < 0.001. Center panel r = 0.329, p < 0001. Right panel r = 0.630, p < 0.001.

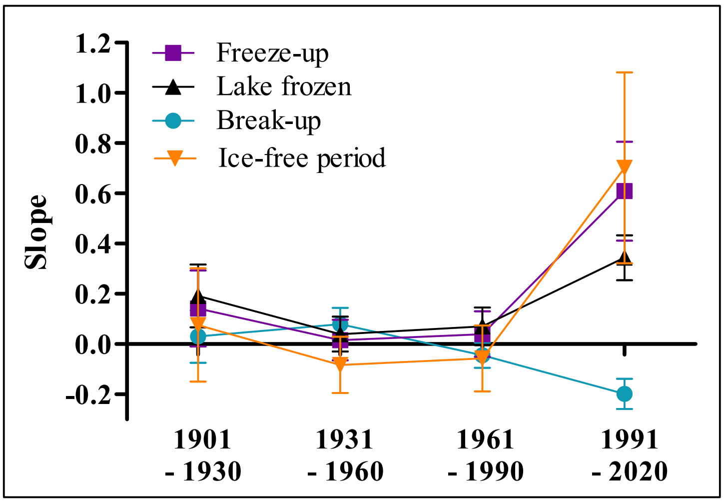

Figure 5. Estimated slopes from general linear mixed models with aspects of ice phenology as response variables (parameter estimates and significance level are given in Table 5). Means and standard error are given.

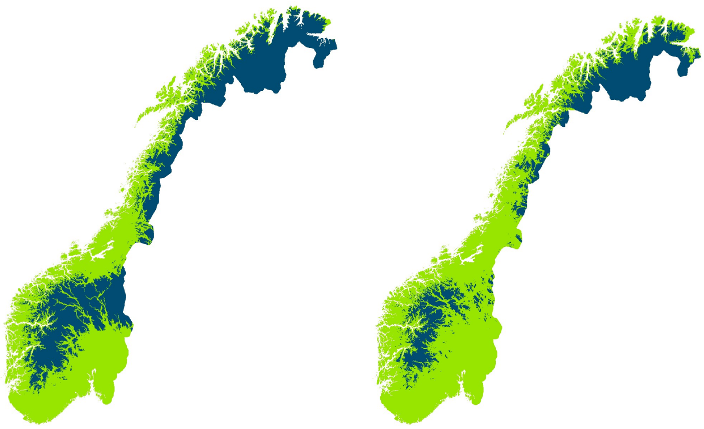

Figure 6. Areal maps of the annual mean air temperature below 0 °C (dark colour) in 1961-1990 (left panel) and 1991-2020 (right panel). (https://rmets.onlinelibrary.wiley.com/doi/full/10.1002/qj.3208)

Table 1.

Correlation between timing of ice-break-up, lake freeze-up, time when the lake was completely frozen and length of ice-free period for 101 Norwegian lakes. All correlations coefficients are significant at P<0.001.

	Lake freeze-up	Lake completely frozen	Length of ice-free period
Ice break-up	-0.741	-0.692	-0.829
Lake freeze-up		0.934	0.868
Lake completely frozen			0.829

Table 2. Model summary. Testing for temporal variation in time of ice break-up, time of lake freeze-up, time when the lake is completely frozen, and length of ice-free period for 99 lakes in Norway. Parameter estimates for the best model are given (see Appendix table 1 for results from the model selection). Significant parameter estimates are given in bold.

Time of ice break-up: Summary statistics with parameter estimates (β ± S.E.), *t*-values and significance level (P). Model *F*-ratio = 91.46 (d.f. = 8, 92), total N = 101, P < 0.0001, R^2 = 0.888.

Parameter	β	S.E.	*t*-value	P
Intercept	-222.39	39.32	-5.66	**<0.001**
Latitude	5.58	0.69	8.08	**<0.001**
Longitude	-0.22	0.53	-0.41	0.684
Elevation	0.36	0.004	9.41	**<0.001**
Latitude * Longitude	0.10	0.15	0.65	0.515
Latitude * Elevation	0.008	0.002	3.65	**<0.001**
Longitude * Elevation	-0.008	0.002	-4.44	**<0.001**
Latitude * Longitude * Elevation	0.001	0.001	2.89	**0.005**
Annual inflow	-0.001	0.001	-1.77	**0.080**

Time of lake freeze-up: Summary statistics with parameter estimates (β ± S.E.), *t*-values and significance level (P). Model *F*-ratio = 23.14 (d.f. = 6, 80), total N = 87, P < 0.0001, R^2 = 0.634.

Parameter	β	S.E.	*t*-value	P
Intercept	394.04	64.25	6.13	**<0.001**
Latitude	-0.32	1.08	-0.30	0.767
Longitude	-3.28	0.73	-4.48	**<0.001**
Elevation	-0.03	0.007	-4.28	**<0.001**
Latitude*Longitude	0.32	0.12	2.60	**0.011**
Latitude*Elevation	0.005	0.003	1.79	0.077
Lake surface area	0.14	0.03	4.05	**<0.001**

Time when lake is completely frozen: Summary statistics with parameter estimates (β ± S.E.), *t*-values and significance level (P). Model *F*-ratio = 42.57 (d.f. = 3, 96), total N = 100, P < 0.0001, R^2 = 0.570.

Parameter	β	S.E.	*t*-value	P
Intercept	389.92	5.66	68.84	**<0.001**
Longitude	-3.08	0.39	-7.87	**<0.001**
Elevation	-0.04	0.005	-9.42	**<0.001**
Lake surface area	0.15	0.04	4.12	**<0.001**

Length of ice-free period: Summary statistics with parameter estimates (β ± S.E.), *t*-values and significance level (P). Model *F*-ratio = 34.06 (d.f. = 6, 80), total N = 87, P < 0.0001, R^2 = 0.719.

Parameter	β	S.E.	*t*-value	P
Intercept	301.63	106.90	2.82	**0.006**
Latitude	-0.10	1.80	-0.06	0.954
Longitude	-6.43	1.22	-5.29	**<0.001**
Elevation	-0.08	0.01	-6.84	**<0.001**
Latitude * Longitude	0.62	0.21	3.07	**0.003**
Latitude * Elevation	0.01	0.005	1.88	**0.064**
Lake surface area	0.15	0.06	2.73	**0.008**

Table 3. Summary statistics for the coefficient of variation (mean, median and range), and correlation between coefficient of variation (CV) and various geographic traits for each lake (elevation, latitude, longitude and distance to the coastline).

	CV			**Correlation coefficient**			
	mean	**median**	**range**	**elevation**	**latitude**	**longitude**	**Coastal distance**
Ice break-up	8.94	6.87	3.94 – 29.93	-0.477 (<0.001)	-0.238 (0.018)	-0.361 (<0.001)	-0.297 (0.003)
Lake freeze-up	4.45	4.16	1.94-10.18	-0.228 (0.034)	-0.092 (0.397)	-0.229 (0.033)	-0.237 (0.027)
Lake completely frozen	4.60	4.31	2.82-9.35	-0.445 (<0.001)	0.159 (0.117)	0.249 (0.808)	-0.367 (<0.001)
Length of ice-free period	15.04	11.55	5.73-42.83	-0.225 (0.036)	0.542 (<0.001)	0.324 (0.002)	-0.427 (<0.001)

Table 4. Model summary. Testing for temporal variation in time of lake freeze-up, time when the lake is completely frozen, and length of ice-free period for 99 lakes in Norway. Lake identity is modelled as a random factor, and year is always included in the model as a fixed effect. Summary statistics with parameter estimates (β ± S.E.), *t*-values and significance level (P) for the best model are given (see Appendix table 2 for results from the model selection). Significant parameter estimates are given in bold.

Timing of lake freeze-up: Total N = 3035, R^2 = 0.676, P < 0.0001**.** The random lake effect accounts for 44.0% of total variance**.**

Parameter	β	S.E.	*t*-value	P
Intercept	491.30	62.00	7.92	**<0.001**
Year	-0.006	0.016	-0.35	0.724
Latitude	-1.82	0.92	-1.97	0.052
Longitude	-2.10	0.60	-3.53	**<0.001**
Elevation	-0.04	0.005	-8.10	**<0.001**
Lake surface area	0.12	0.03	3.60	**<0.001**
Regulated (no)	0.66	0.96	0.69	0.491
Amplitude	0.53	0.18	2.98	**0.003**

Time when lake is completely frozen: Total N = 4084, R^2 = 0.697, P < 0.0001**.** The random lake effect accounts for 50.6% of total variance**.**

Parameter	β	S.E.	*t*-value	P
Intercept	301.62	65.86	4.58	**<0.001**
Year	0.06	0.01	4.68	**<0.001**
Latitude	-0.65	1.05	-0.62	**0.537**
Longitude	-2.68	0.67	-4.01	**<0.001**
Elevation	-0.05	0.005	-9.89	**<0.001**
Lake surface area	0.15	0.04	3.93	**<0.001**
Regulated (no)	-0.53	0.84	-0.63	0.526
Amplitude	0.24	0.15	1.55	0.122

Length of ice-free period: Total N = 2807, R^2 = 0.663, P < 0.0001**.** The random lake effect account for 34.4% of total variance**.**

Parameter	β	S.E.	*t*-value	P
Intercept	433.89	108.63	3.99	**<0.001**
Year	0.02	0.03	0.52	0.606
Latitude	-2.80	1.50	-1.87	0.065
Longitude	-6.05	1.26	-4.78	**<0.001**
Elevation	-0.10	0.009	-10.90	**<0.001**
Latitude * Longitude	0.45	0.19	2.37	**0.020**
Lake surface area	0.16	0.06	2.87	**0.005**
Regulated (no)	4.79	1.91	2.51	0.012
Amplitude	0.60	0.36	1.65	0.098

Table 5. Model summary. Temporal and climate effects on in time of ice break-up 98 lakes in Norway. Lake identity is modelled as a random factor, and year is always included in the model as a fixed effect. NAO is included as the climate effect. Summary statistics with parameter estimates (β ± S.E.), *t*-values and significance level (P) for the best model are given (see Appendix table 3 for results from the model selection). Significant parameter estimates are given in bold.

Total N = 4194, R^2 = 0.726, P < 0.0001**.** The random lake effect account for 22.3 % of total variance.

Parameter	β	**S.E.**	***t*-value**	**P**
Intercept	-205.98	46.00	-4.42	**<0.001**
NAO	-3.26	0.20	-16.61	**<0.001**
Year	-0.03	0.01	-2.86	**0.004**
Latitude	6.21	0.76	8.19	**<0.001**
Longitude	-0.64	0.59	-1.08	0.283
Elevation	0.04	0.003	13.99	**<0.001**
Latitude * Longitude	-0.30	0.07	-4.35	**0.004**
Latitude * Elevation	0.008	0.002	3.59	**<0.001**
Longitude * Elevation	-0.008	0.002	-4.25	**0.004**

Table 6. Parameters estimates (slope ± se) from general linear mixed models with ice phenology estimates as response variables, year as predictor and lake identity as random effect. The time series are sorted into 30-year periods (1901-1930, 1931-1960, 1961-1990, 1991-2020). Significant estimates are given in bold, with number of observations in parenthesis. The lakes included is given in Appendix 7.

	Break-up	**Freeze-up**	**Lake frozen**	**Ice-free period**
1901-1930	0.030±0.104 N=390	0.142±0.151 N=325	0.192±0.125 N=435	0.076 ±0.226 N=254
1931-1960	0.080±0.064 N=739	0.016±0.081 N=637	0.040±0.069 N=734	-0.083±0.112 N=586
1961-1990	-0.044±0.050 N=772	0.040±0.091 N=502	0.071±0.075 N=754	-0.057±0.1309 N=475
1991-2020	**-0.198±0.060** **N=411**	**0.609±0.197** **N=116**	**0.344±0.089** **N=391**	0.702±0.380 N=107

Appendix 1.

Lake characteristics of the 101 Norwegian lakes used in the analyses. Regulated show the year when the lake was developed for hydropower, and the figure in the brackets is the annual water level amplitude. No indicate a pristine lake.

Lake no	Lake	North	East	Coastal distance (km)	Elevation (m a.s.l.)	Lake area (km^2)	Mean annual inflow (10^6 m^3)	Catchment area (km^2)	Regulated
1	Mjøsa (Hamar)	60.397	11.234	350	123	369.32	9953.72	16555.36	1920 (3.61 m)
2	Storsjø	61.392	11.363	357	251	48.1	1027.59	2293.6	1968 (3.64 m)
3	Lomnessjøen	61.732	11.202	329	255	3.67	511.93	1164.41	no
4	Osensjøen	61.246	11.739	385	437	43.37	665.79	1174.36	1941 (6.6 m)
5	Olstappen	61.514	9.402	231	668	3.2	1188.82	1305.11	1954 (13 m)
6	Aursunden	62.680	11.462	196	690	46.11	629.99	848.44	1923 (5.9 m)
7	Atnsjøen	61.852	10.226	217	701	5.01	323.1	463.2	no
8	Savalen	62.232	10.519	189	708	15.29	29.93	102.48	1973 (4.7 m)
9	Narsjø	62.364	11.477	238	737	1.95	70.67	118.86	no
10	Gålåvatn	61.530	9.717	270	778	3.04	9.72	23.1	no
11	Tesse	61.814	8.941	182	854	12.84	102.24	225.37	1942 (12 m)
12	Aursjø	61.934	8.327	140	1098	6.7	41.61	106.31	1967 (14.5 m)
13	Breidalsvatn	62.008	7.630	123	900	6.9	177.02	127.22	1944 (13 m)
14	Raudalsvatn	61.911	7.796	109	913	7.48	209.08	146.93	1952 (30.3 m)
15	Gjende	61.495	8.810	196	984	15.61	497.31	376.2	no
16	Veslevatn	61.416	9.273	224	998	4.22	33.98	44.11	1960 (2 m)
17	Kaldfjorden	61.350	9.263	245	1019	19.18	655.29	559.88	1956 (4.9 m)
18	Fundin	62.324	9.915	161	1022	10.4	155.13	252.86	1968 (11 m)
19	Vinstern	61.352	9.069	238	1032	28.19	573.95	466.3	1951 (4 m)
20	Nedre Heimdalsvatn	61.446	9.108	238	1052	7.25	134.72	129.2	1959 (2.2 m)
21	Bygdin	61.328	8.799	235	1057	40.03	398.02	305.59	1934 (9.15)
22	Marsjø	62.343	10.049	165	1064	2.68	13.95	23.39	1910 (4 m)
23	Øvre Heimdalsvatn	61.418	8.893	203	1089	0.78	26.89	24.94	no
24	Elgsjø	62.361	9.798	154	1132	2.38	22.16	33.75	1914 (5.35 m)
25	Leirvatnet	61.547	8.250	168	1401	1.04	170.31	154.72	no
26	Volbufjorden	61.080	9.110	238	434	3.94	446.88	675.85	1916 (3 m)
27	Øyangen	61.221	8.924	231	677	6.64	238.64	246.19	1918 (8.3 m)
28	Vasetvatnet	60.996	8.985	231	796	1.03	47.81	82.9	no
29	Midtre Syndin	61.058	8.782	224	937	2.73	15.68	21.47	no
30	Rødungen	60.696	8.256	193	1022	7.4	51.01	61.79	1943 (23 m)
31	Bergsjø	60.709	8.275	193	1082	1.68	5.58	28.09	1943 (11 m)
32	Vangsmjøsa	61.149	8.701	231	466	17.4	22.97	487.6	1963 (3 m)
33	Krøderen	60.123	9.783	270	133	43.91	3701.57	5091.06	1960 (2.6 m)
34	Fønnebøfjorden	60.256	8.914	217	460	0.75	455.12	687.29	no
35	Tunhovdfjorden	60.426	8.833	221	734	25.55	1141.64	1857.98	1920 (18.15 m)
36	Pålsbufjorden	60.433	8.733	215	749	19.64	1063.35	1645.84	1946 (24.5 m)
37	Møsvatn	59.824	8.317	182	918	78.51	1573.04	1509.77	1903 (18.5)
38	Seljordvatn	59.434	8.854	214	116	16.49	428.07	724.97	1943 (1 m)
39	Hjartsjå	59.608	8.763	210	158	1.07	185.76	214.35	1957 (1.8 m)
40	Vinjevatn	59.582	7.926	158	465	3.32	1249.03	905.89	1960 (3.5 m)
41	Totak	59.664	8.026	168	687	36.59	1005.39	863.22	1958 (7 m)
42	Eptevatn (Færåsen)	58.236	7.291	34	232	1.16	51.82	33.49	1921 (10 m)

43	Lygne	58.397	7.221	53	185	7.71	525.4	272.2	no
44	Sandvinvatn	60.053	6.555	91	87	4.37	1288.75	470.22	no
45	Vangsvatn	60.630	6.277	88	47	7.65	2225.36	1091.51	no
46	Vassbygdvatn	60.876	7.264	147	55	1.85	1136.22	760.47	1982 (1.4 m)
47	Tyin	61.275	8.139	189	1084	33.21	241.97	183.45	1942 (10.3 m)
48	Veitastrondvatn	61.322	7.110	133	171	17.46	895.59	386.46	1982 (2.5 m)
49	Rørvikvatn	61.208	5.761	62	336	7.14	59.9	20.69	1920 (1 m)
50	Hersvikvatn	61.135	4.929	17	24	1.37	13.53	7.06	no
51	Nautsundvatn	61.252	5.379	39	44	0.676	595	218.87	no
52	Hestadfjorden	61.335	5.887	67	146	3.24	1351.35	507.94	no
53	Jølstervatn	61.492	6.113	77	207	39.24	928.16	384.54	1952 (1.25m)
54	Blåmannsvatn	61.562	5.517	44	43	0.24	624.99	225.49	no
55	Lovatn	61.860	6.890	98	52	10.7	479.49	234.88	no
56	Hornindalsvatn	61.916	6.109	58	53	19.09	727.73	381.04	no
57	Kaldvatn	62.045	6.395	59	70	0.78	95.7	62.02	1955 (3 m)
58	Nysetervatn	62.352	6.835	55	334	2.36	59.93	29.65	1955 (13 m)
59	Gjevilvatn	62.648	9.490	112	660	21.18	167.83	169.63	1973 (15 m)
60	Engelivatn	63.100	8.545	56	243	1.81	41.51	20.6	1942 (7.5 m)
61	Søvatn	63.226	9.308	70	280	5.17	156.64	101.44	1940 (19.8 m)
62	Rovatn	63.287	9.069	560	13	7.74	352.35	237.87	no
63	Fjergen	63.434	11.910	126	512	13.45	303.99	227.42	1993 (16 m)
64	Funnsjøen	63.480	11.787	119	441	7.99	82.07	60.91	1938 (11.5 m)
65	Lustadvatn	63.991	12.013	91	275	7.11	82.46	68.81	no
66	Follavatn	64.040	11.113	53	182	1.44	420.12	252.29	1923 (9.5 m)
67	Krinsvatn	63.804	10.227	35	87	0.41	413.8	205.67	no
68	Namsvatn	65.019	13.539	98	454	39.44	1009.35	700.8	1951 (14 m)
69	Fustvatn	65.899	13.286	70	39	16.65	970.52	475.8	no
70	Røssvatn	65.858	13.794	91	384	47.78	2513.59	1501.21	no
71	Tustervatn	65.858	13.794	91	384	47.78	2513.59	1501.21	1957 (13 m)
72	Vassvatn	66.397	13.176	35	108	0.81	66.17	16.39	no
73	Storglåmvatn	66.773	14.143	49	529	6.18	72.53	84.79	1964 (12.5 m)
74	Skarsvatn	67.084	14.982	56	162	0.29	164.97	145.08	no
75	Vatnevatn	67.320	14.750	35	4	6.64	196.07	141.18	no
76	Kobbvatn	67.597	15.970	70	8	4.9	782.19	387.22	no
77	Sørfjordvatn	67.549	15.901	70	80	0.31	212.45	116	no
78	Storvatn	67.848	15.503	35	56	6.6	155.58	71.28	no
79	Forsavatn	68.310	16.739	112	29	1.2	250.48	232.54	no
80	Sneisvatn	68.405	15.709	74	17	0.37	86.75	29.45	no
81	Svolværvatn	68.246	14.541	21	4	0.93	21.45	18.5	no
82	Gåslandsvatn	68.723	14.628	140	16	1.54	11.9	7.35	no
83	Skodbergvatn	68.620	17.252	91	101	8.56	128.92	107.41	1953 (6.5 m)
84	Nervatn	68.869	17.867	77	7	1.2	681.76	535.57	no
85	Lysevatn	69.413	17.860	28	22	41.94	281.02	129.46	no
86	Insetvatn	68.677	18.735	126	301	3.72	1267.32	1389.68	no
87	Oksfjordvatn	69.903	21.347	56	9	58.12	256.65	265.83	no
88	Lille Mattisvatn	69.894	23.016	102	64	11.12	267.81	318.95	no
89	Lille Ropelvvann	69.761	30.188	18	51	1.19	20.41	48.87	no
90	Bjørnvatn	69.527	30.139	41	21	3.54	5207.35	18101.09	no
91	Murusjøen	64.46	14.103	168	311	7.19	266.73	346.39	no
92	Limingen	64.693	13.76	140	418	95.7	746.52	673	1955 (9 m)
93	Vekteren	64.894	13.563	119	446	8.8	381.72	310.05	1963 (5.5 m)
94	Saksvatn	64.919	13.482	112	462	1.69	76.14	63.86	no

95	Lenglingen	64.196	13.83	168	354	30.26	467.61	452.54	no
96	Engeren	61.527	12.082	364	472	11.49	231.52	395.05	no
97	Femunden	61.935	11.868	336	664	203.4	807.97	1793.94	no
98	Isteren	61.910	11.779	340	645	80.64	1129.71	2445.91	no
99	Møkeren	60.120	12.318	406	176	12.77	75.24	367.63	1928 (1.2 m)
100	Søndre Øyersjøen	60.209	12.448	417	270	2.06	34.26	66.26	1934 (4 m)
101	Varalden	60.144	12.416	413	203	6.5	103.95	214.11	1929 (4.5 m)
102	Rømsjøen	59.665	11.836	385	138	13.66	65.28	91.89	no

Appendix 2.

Summary of ice phenology recordings from 101 Norwegian lakes. Minimum and maximum recordings are given in brackets.

Lake				Break-up		Freeze-up		Frozen lake		Ice free period
no	Lake	Period	n	Median	n	Median	n	Median	n	Median
1	Mjøsa (Hamar)	1910-2001	76	111 (23-139)	74	383 (318-440)	63	392 (350-435)	63	272 (208-401)
2	Storsjø	1910-2011	66	124 (97-140)	48	361 (333-392)	76	390 (349-443)	28	239 (200-276)
3	Lomnessjøen	1919-1997	66	131 (96-147)	69	320 (281-352)	54	327 (302-379)	58	186 (152-248)
4	Osensjøen	1967-2019	49	130 (106-142)	24	360 (336-394)	50	362 (338-406)	22	232 (210-277)
5	Olstappen	1967-2020	53	142 (129-158)			52	309 (285-329)		
6	Aursunden	1902-2020	115	152 (129-175)	58	314 (295-332)	116	324 (295-355)	57	158 (127-186)
7	Atnsjøen	1917-2020	87	145 (122-165)	95	320 (302-347)	98	328 (312-363)	84	176 (144-213)
8	Savalen	1975-2020	45	144 (128-160)			45	323 (306-360)		
9	Narsjø	1930-1961	31	145 (136-164)	29	300 (283-313)	31	311 (293-335)	29	154 (125-175)
10	Gålåvatn	2009-2020	11	145 (124-150)	11	315 (305-326)	11	322 (305-339)	10	175 (162-196)
11	Tesse	1908-2020	74	148 (121-167)			76	330 (311-363)		
12	Aursjø	1967-2020	53	169 (148-181)			53	310 (293-332)		
13	Breidalsvatn	1967-2020	53	168 (147-191)			53	323 (303-347)		
14	Raudalsvatn	1967-2020	53	157 (136-176)			53	329 (313-365)		
15	Gjende	1946-2020	15	149 (137-161)	14	348 (326-377)	19	358 (335-412)	12	194 (175-225)
16	Veslevatn	1971-2018	47	153 (84-182)			47	305 (285-332)		
17	Kaldfjorden	1967-2020	53	159 (136-170)			53	309 (285-332)		
18	Fundin	1970-2020	50	159 (138-174)			48	313 (297-328)		
19	Vinstern	1950-2020	64	163 (147-181)			69	317 (288-339)		
20	Nedre Heimdalsvatn	1975-2020	45	159 (134-171)			45	308 (283-326)		
21	Bygdin	1950-2020	64	170 (153-185)	15	326 (301-382)	65	370 (315-416)	14	157 (130-221)
22	Marsjø	1976-2020	45	160 (135-180)			44	314 (297-328)		
23	Øvre Heimdalsvatn	1969-2020	49	161 (137-188)	12	289 (277-302)	39	294 (279-309)	12	128 (111-151)
24	Elgsjø	1976-2020	45	164 (144-180)			44	306 (291-328)		
25	Leirvatnet	2009-2020	11	182 (157-234)	11	299 (283-312)	11	308 (286-331)	10	120 (55-142)
26	Volbufjorden	1920-1974	55	137 (119-150)	54	320 (305-344)	55	324 (312-353)	54	184 (164-214)
27	Øyangen	1919-1984	65	149 (130-168)	62	318 (299-343)	62	321 (304-344)	61	170 (137-200)
28	Vasetvatnet	2009-2020	11	143 (122-152)	11	307 (294-361)	11	315 (295-363)	10	163 (151-218)
29	Midtre Syndin	2010-2020	10	150 (128-158)	9	309 (280-332)	10	320 (302-334)	8	156 (129-187)
30	Rødungen	1952-2020	41	157 (112-175)	37	312 (301-335)	47	324 (311-366)	31	154 (136-223)
31	Bergsjø	1953-2020	58	160 (146-175)	47	304 (288-343)	56	314 (294-350)	47	144 (127-170)
32	Vangsmjøsa	1940-2020	34	134 (78-149)	33	323 (303-366)	32	375 (315-409)	32	196 (161-276)
33	Krøderen	1900-1964	64	124 (100-161)	7	335 (315-366)	60	338 (306-372)	7	214 (189-255)
34	Fønnebøfjorden	1918-1966	44	131 (104-145)	15	310 (290-321)	47	309 (289-366)	15	174 (152-201)
35	Tunhovdfjorden	1920-2020	73	142 (119-161)	45	329 (275-353)	77	335 (305-362)	41	186 (142-219)

36	Pålsbufjorden	1924-1966	37	145 (121-153)	32	310 (294-355)	39	321 (305-424)	31	166 (143-279)
37	Møsvatn	1926-2018	86	152 (134-176)			30	341 (319-360)		
38	Seljordvatn	1940-1972	30	115 (89-132)	26	359 (322-386)	23	367 (349-413)	24	244 (279-211)
39	Hjartsjå	1919-1998	74	121 (91-139)	43	328 (311-354)	70	334 (313-388)	42	207 (184-261)
40	Vinjevatn	1909-1963	46	133 (103-146)	16	313 (296-344)	46	317 (297-375)	16	182 (150-220)
41	Totak	1818-1966	79	146 (124-169)	25	348 (332-371)	20	373 (349-408)	22	207 (186-230)
42	Eptevatn	1951-2000	45	114 (22-136)	36	340 (315-382)	49	346 (327-386)	32	224 (182-318)
43	Lygne	1924-2009	72	112 (22-137)	71	362 (441-313)			60	253 (212-363)
44	Sandvinvatn	1908-1998	59	106 (33-131)	61	383 (224-437)	64	398 (359-453)	46	276 (225-342)
45	Vangsvatn	1898-1989	69	113 (38-138)	46	347 (316-402)	78	354 (327-420)	61	236 (197-333)
46	Vassbygdvatn	1915-1987	69	116 (56-139)	56	356 (277-401)	65	371 (330-435)	54	242 (158-305)
47	Tyin	1931-2019	26	170 (148-198)	29	335 (314-372)	30	338 (318-373)	24	166 (128-208)
48	Veitastrondvatn	1918-1991	65	137 (76-152)	52	353 (311-416)	61	356 (326-428)	50	217 (171-284)
49	Rørvikvatn	1950-1997	47	137 (91-166)	47	335 (374-310)	48	342 (322-397)	46	199 (159-236)
50	Hersvikvatn	1934-1988	45	83 (18-115)	47	370 (335-413)	42	372 (337-412)	47	292 (245-395)
51	Nautsundvatn	1908-1983	55	106 (33-130)	75	353 (314-426)	75	353 (314-426)	54	248 (215-348)
52	Hestadfjorden	1914-1995	70	117 (17-140)	75	358 (320-423)	77	371 (323-446)	65	242 (192-382)
53	Jølstervatn	1920-1968	22	112 (67-137)	24	384 (340-434)	12	392 (352-430)	24	310 (235-406)
54	Blåmannsvatn	1930-1981	15	95 (39-122)	40	348 (323-407)	39	380 (332-436)	40	347 (221-407)
55	Lovatn	1899-1979	72	108 (18-132)	44	388 (347-436)	51	388 (355-440)	42	281 (227-395)
56	Hornindalsvatn	1950-1970	20	105 (58-128)	19	371 (359-414)	8	406 (378-422)	19	275 (232-363)
57	Kaldvatn	1950-1983	33	113 (82-135)	32	342 (314-385)	28	373 (340-423)	32	228 (179-341)
58	Nysetervatn	1950-1972	16	145 (120-180)	13	324 (309-376)	17	331 (312-381)	13	190 (163-329)
59	Gjevilvatn	1950-2020	13	151 (133-163)	15	347 (321-377)	18	356 (323-387)	8	194 (171-226)
60	Engelivatn	1951-1983	24	144 (118-158)	25	343 (298-344)	27	343 (321-368)	25	186 (147-344)
61	Søvatn	1940-1984	44	146 (118-250)	42	325 (308-347)	42	332 (313-362)	41	180 (62-229)
62	Rovatn	1950-1981	28	126 875-135)	31	361 (325-413)	31	374 (341-416)	27	235 (200-302)
63	Fjergen	1952-1968	28	152 (122-160)	27	318 (294-335)	34	325 (309-366)	21	166 (141-191)
64	Funnsjøen	1951-1972	18	151 (131-169)	21	322 (297-341)	18	335 (310-362)	17	177 (141-204)
65	Lustadvatn	1950-1969	13	140 (127-147)	12	327 (306-341)	17	338 (314-353)	10	194 (164-210)
66	Follavatn	1950-1971	20	138 (118-155)	19	321 (303-343)	20	333 (312-367)	18	178 (163-222)
67	Krinsvatn	1996-2019	16	119 (92-134)	15	340 (246-384)	16	367 (327-4379	11	224 (181-262)
68	Namsvatn	1908-1968	57	163 (137-184)	19	319 (301-341)	58	323 (291-351)	17	164 (126-183)
69	Fustvatn	1930-1972	34	135 (84-162)	37	315 (280-347)	39	329 (288-372)	30	182 (151-249)
70	Røssvatn	1912-1957	46	160 (141-182)	44	354 (310-406)	45	370 (337-417)	44	198 (144-248)
71	Tustervatn	1907-1968	54	156 (137-178)	44	328 (304-366)	50	343 (308-391)	41	174 (127-216)
72	Vassvatn	1950-1979	29	137 (113-175)	29	340 (314-363)	29	361 (330-412)	28	199 (158-232)
73	Storglåmvatn	1948-1973	17	178 (162-210)	20	329 (288-361)	20	342 (294-391)	12	151 (89-187)
74	Skarsvatn	1950-1980	30	144 (124-165)	29	300 (278-320)	30	310 (286-355)	28	154 (129-183)
75	Vatnevatn	1951-1972	21	141 (127-247)	21	325 (300-351)	20	345 (325-373)	20	186 (85-214)
76	Kobbvatn	1916-1978	61	149 (128-167)	58	330 (304-386)	60	339 (310-392)	56	185 (140-245)
77	Sørfjordvatn	1950-1967	7	145 (133-159)	10	310 (294-331)	15	322 (306-384)	5	171 (152-191)

78	Storvatn	1950-1967	9	144 (134-158)	10	335 (305-373)	14	353 (334-376)	7	190 (149-239)
79	Forsavatn	1920-1965	44	141 (117-159)	46	316 (288-350)	46	325 (294-403)	44	176 (134-206)
80	Sneisvatn	1951-1972	19	147 (129-180)	20	303 (274-340)	21	311 (284-395)	17	156 (121-194)
81	Svolværvatn	1951-1971	19	149 (116-168)	20	318 (297-318)	20	334 (310-393)	19	173 (135-206)
82	Gåslandsvatn	1950-1972	22	141 (111-172)	21	316 (283-359)	22	327 (309-367)	21	179 (128-229)
83	Skodbergvatn	1950-1971	11	151 (146-167)	9	331 (317-361)	9	336 (329-362)	8	177 (152-212)
84	Nervatn	1950-1987	37	144 (104-160)	36	305 (289-324)	36	319 (301-362)	34	164 (142-192)
85	Lysevatn	1950-1972	19	158 (127-177)	20	307 (296-332)	20	323 (305-359)	17	152 (125-170)
86	Insetvatn	1914-1958	43	152 (133-181)	45	296 (279-322)	45	314 (291-367)	43	145 (109-181)
87	Oksfjordvatn	1955-1970	15	156 (135-170)	5	339 (301-329)	15	339 (319-352)	5	171 (140-301)
88	Lille Mattisvatn	1950-1972	16	139 (129-151)	11	298 (280-314)	16	319 (296-246)	11	168 (143-312)
89	Lille Ropelvvann	1957-1984	27	161 (128-172)	11	309 (299-322)	27	311 (294-332)	11	152 (141-181)
90	Bjørnvatn	1912-1967	55	151 (130-182)	53	306 (286-327)	55	311 (289-366)	52	157 (117-189)
91	Murusjøen	1926-2001	66	142 (121-155)	74	327 (305-354)	66	336 (311-366)	65	184 (157-223)
92	Limingen	1930-1960	27	152 (131-176)	11	348 (289-385)	30	369 (316-424)	11	200 (120-240)
93	Vekteren	1959-1986	23	157 (144-168)	46	321 (296-341)	15	339 (305-413)	20	164 (145-190)
94	Saksvatn	1960-1990	31	164 (141-180)	29	306 (279-331)	28	309 (298-334)	29	147 (174-119)
95	Lenglingen	1925-2003	76	144 (118-158)	76	329 (307-383)	77	339 (312-385)	74	187 (157-235)
96	Engeren	1911-1983	72	139 (119-157)	72	347 (299-396)	71	350 (311-386)	71	204 (156-244)
97	Femunden	1900-1995	82	148 (128-173)	83	328 (305-353)	83	343 (313-386)	79	177 (152-214)
98	Isteren	1950-1986	34	148 (113-157)	35	309 (283-335)	35	319 (291-385)	34	162 (134-206)
99	Møkeren	1911-2007	65	121 (91-141)	47	332 (261-363)	65	341 (303-446)	37	212 (128-244)
100	Søndre Øyersjøen	1960-2005	42	126 (99-138)	18	334 (305-363)	39	336 (308-367)	18	210 (179-240)
101	Varalden	1950-1983	26	123 (91-135)	22	335 (312-367)	27	350 (315-378)	19	210 (181-245
102	Rømsjøen	1945-1995	46	119 (33-138)	44	339 (305-376)	48	359 (333-398)	43	224 (171-293)

Appendix 3.

Variation in average time of ice break-up, time of lake freeze-up, time when lake is completely frozen, and length of ice-free period. The full model is formulated as (see description of parameters in the main text):

$Y = \mu + \alpha_1$Ele $+ \alpha_2$Lat $+ \alpha_3$Long $+ \alpha_4$Ele*Lat $+ \alpha_5$Ele*Long $+ \alpha_6$Long*Lat $+ \alpha_7$Ele*Long*Lat $+ \alpha_8$Distance $+ \alpha_9$Area $+ \alpha_{10}$Catch $+ \alpha_{11}$Flow $+ \varepsilon$

Selection of the best model was based on AIC. The full model and the three best models are presented, with the best model given in bold. AIC and ΔAIC is given.

Time of ice break-up:

No.	**Model formulation (n = 101)**	AIC	ΔAIC
0	Full model	695.8	5.0
1	**$Y = \mu + \alpha_1$Ele $+ \alpha_2$Lat $+ \alpha_3$Long $+ \alpha_4$Ele*Lat $+ \alpha_5$Ele*Long $+ \alpha_6$Long*Lat $+ \alpha_7$Ele*Long*Lat $+ \alpha_{11}$Flow**	**690.8**	**0**
2	$Y = \mu + \alpha_1$Ele $+ \alpha_2$Lat $+ \alpha_3$Long $+ \alpha_4$Ele*Lat $+ \alpha_5$Ele*Long $+ \alpha_6$Long*Lat $+ \alpha_7$Ele*Long*Lat $+ \alpha_8$Distance $+ \alpha_{11}$Flow	691.2	0.4
5	$Y = \mu + \alpha_1$Ele $+ \alpha_2$Lat $+ \alpha_3$Long $+ \alpha_4$Ele*Lat $+ \alpha_5$Ele*Long $+ \alpha_6$Long*Lat $+ \alpha_7$Ele*Long*Lat	691.7	0.9

Time of lake freeze-up:

No.	**Model formulation (n = 86)**	AIC	ΔAIC
0	Full model	719.5	11.8
1	**$Y = \mu + \alpha_1$Ele $+ \alpha_2$Lat $+ \alpha_3$Long $+ \alpha_4$Ele*Lat $+ \alpha_6$Long*Lat $+ \alpha_8$Area**	**707.7**	**0**
2	$Y = \mu + \alpha_1$Ele $+ \alpha_2$Lat $+ \alpha_3$Long $+ \alpha_6$Long*Lat $+ \alpha_8$Area	708.1	0.4
3	$Y = \mu + \alpha_1$Ele $+ \alpha_2$Lat $+ \alpha_3$Long $+\alpha_4$Ele*Lat $+ \alpha_5$Ele*Long $+\alpha_6$Long*Lat $+ \alpha_8$Area	709.6	1.9

Time when lake is completely frozen:

No.	**Model formulation (n = 97)**	AIC	ΔAIC
0	Full model	838.0	12.5
1	**$Y = \mu + \alpha_1$Ele $+ \alpha_2$Lat $+ \alpha_3$Long $+ \alpha_6$Long*Lat $+ \alpha_8$Area**	**825.5**	**0.0**

2	Y = $\mu + \alpha_1$Ele + α_2Lat + α_3Long + α_8Area	827.6	2.1
3	Y = $\mu + \alpha_1$Ele- + α_2Lat + α_3Long + α_5Ele*Long + α_8Area	828.3	2.8

Length of ice-free period:

No.	**Model formulation (n = 86)**	AIC	ΔAIC
0	Full model	808.4	13.8
1	**Y = $\mu + \alpha_1$Ele + α_2Lat + α_3Long + α_5Ele*Long + α_6Long*Lat + α_8Area**	**794.6**	**0**
2	Y = $\mu + \alpha_1$Ele + α_2Lat + α_3Long + α_4Ele*Lat + α_6Long*Lat + α_8Area + α_{10}Flow	795.2	0.6
3	Y = $\mu + \alpha_1$Ele + α_2Lat + α_3Long + α_4Ele*Lat + α_6Long*Lat + α_8Area + α_9Catch	796.2	1.6

Appendix 4.

Test for temporal variation in time of lake freeze-up, time when lake is completely frozen, and length of ice-free period for 98 lakes in Norway. Lake identity is modelled as a random factor, and year is always included in the model as a fixed effect. The full model is formulated as (see description of parameters in the main text):

$Y = \mu + \alpha_1 Ele + \alpha_2 Lat + \alpha_3 Long + \alpha_4 Ele{*}Lat + \alpha_5 Ele{*}Long + \alpha_6 Long{*}Lat + \alpha_7 Ele{*}Long{*}Lat + \alpha_8 Distance + \alpha_9 Area + \alpha_{10} Catch + \alpha_{11} Flow + \alpha_{12} Year + \alpha_{13} Regulated + \alpha_{14} Amplitude + \varepsilon$.

Selection of the best model was based on AIC. The full model and the three best models are presented, with the best model given in bold. AIC and ΔAIC is given.

Time of lake freeze-up:

No.	**Model formulation**	AIC	ΔAIC
0	Full model	25 776.8	63.7
1	**$Y = \mu + \alpha_1 Ele + \alpha_2 Lat + \alpha_3 Long + \alpha_8 Area + \alpha_{12} Year + \alpha_{13} Regulated + \alpha_{14} Amplitude$**	**25 713.1**	**0**
2	$Y = \mu + \alpha_1 Ele + \alpha_2 Lat + \alpha_3 Long + \alpha_6 Long{*}Lat + \alpha_9 Area + \alpha_{12} Year + \alpha_{13} Regulated + \alpha_{14} Amplitude$	25 713.8	0.7
3	$Y = \mu + \alpha_1 Ele + \alpha_3 Long + \alpha_9 Area + \alpha_{12} Year + \alpha_{13} Regulated + \alpha_{14} Amplitude$	25 716.6	3.5

Time when lake is completely frozen:

No.	**Model formulation**	AIC	ΔAIC
0	Full model	35 781.8	67.1
1	**$Y = \mu + \alpha_1 Ele + \alpha_2 Lat + \alpha_3 Long + \alpha_6 Long{*}Lat + \alpha_9 Area + \alpha_{12} Year + \alpha_{13} Regulated + \alpha_{14} Amplitude$**	**35 714.7**	**0**
2	$Y = \mu + \alpha_1 Ele + \alpha_2 Lat + \alpha_3 Long + \alpha_8 Area + \alpha_{11} Year + \alpha_{12} Regulated + \alpha_{13} Amplitude$	35 715.0	0.3
3	$Y = \mu + \alpha_1 Ele + \alpha_2 Lat + \alpha_3 Long + \alpha_6 Long{*}Lat + \alpha_8 Area + \alpha_{11} Year$	35 716.6	1.9

Length of ice-free period:

No.	**Model formulation**	AIC	ΔAIC
0	Full model	27 547.9	55.9
1	**$Y = \mu + \alpha_1$Ele + α_2Lat + α_3Long + α_6Long*Lat + α_9Area+ α_{12}Year + α_{13}Regulated + α_{14}Amplitude**	**27 492.0**	**0**
2	$Y = \mu + \alpha_1$Ele + α_2Lat + α_3Long + α_8Area + α_{11}Year + α_{12}Regulated + α_{13}Amplitude	27 494.1	2.1
3	$Y = \mu + \alpha_1$Ele + α_2Lat + α_3Long + α_6Long*Lat + α_8Area + α_{11}Year	27 496.7	4.7

Appendix 5.

Test for temporal variation in time of ice break-up. Lake identity is modelled as a random factor, and year is always included in the model as a fixed effect. NAO is included in the model as both a linear and a non-linear effect. The full model is formulated as (see description of parameters in the main text):

$Y = \mu + \alpha_1 Ele + \alpha_2 Lat + \alpha_3 Long + \alpha_4 Ele*Lat + \alpha_5 Ele*Long + \alpha_6 Long*Lat + \alpha_7 Ele*Long*Lat + \alpha_8 Distance + \alpha_9 Area + \alpha_{10} Catch + \alpha_{11} Flow + \alpha_{12} Year + \alpha_{13} Regulated + \alpha_{14} Amplitude + \alpha_{15} NAO + \alpha_{16} NAO^2 + \varepsilon$.

Selection of the best model was based on AIC. The full model and the three best models are presented, with the best model given in bold. AIC and ΔAIC is given.

No.	**Model formulation**	AIC	ΔAIC
0	Full model	33 367.0	56.0
1	$Y = \mu + \alpha_1 Ele + \alpha_2 Lat + \alpha_3 Long + \alpha_5 Ele*Long + \alpha_6 Long*Lat + \alpha_{12} Year + \alpha_{15} NAO$	33 311.0	**0**
2	$Y = \mu + \alpha_1 Ele + \alpha_2 Lat + \alpha_3 Long + \alpha_6 Long*Lat + \alpha_{12} Year + \alpha_{15} NAO$	33 311.2	0.2
3	$Y = \mu + \alpha_1 Ele + \alpha_2 Lat + \alpha_3 Long + \alpha_5 Ele*Long + \alpha_{12} Year + \alpha_{15} NAO$	33 315.5	4.5

Appendix 6.

Test for non-linear temporal trends in ice phenology in 30-years periods. Lakes with >50 years of records of both date of break-up and date of frozen lake.

Lake			Break-up		Freeze-up		Frozen lake		Ice free period	
no	Lake	Period	n	Median	n	Median	n	Median	n	Median
1	Mjøsa (Hamar)	1910-2001	76	111 (23-139)	74	383 (318-440)	63	392 (350-435)	63	272 (208-401)
2	Storsjø	1910-2011	66	124 (97-140)	48	361 (333-392)	76	390 (349-443)	28	239 (200-276)
3	Lomnessjøen	1919-1997	66	131 (96-147)	69	320 (281-352)	54	327 (302-379)	58	186 (152-248)
5	Olstappen	1967-2020	53	142 (129-158)			52	309 (285-329)		
6	Aursunden	1902-2020	115	152 (129-175)	58	314 (295-332)	116	324 (295-355)	57	158 (127-186)
7	Atnsjøen	1917-2020	87	145 (122-165)	95	320 (302-347)	98	328 (312-363)	84	176 (144-213)
11	Tesse	1908-2020	74	148 (121-167)			76	330 (311-363)		
12	Aursjø	1967-2020	53	169 (148-181)			53	310 (293-332)		
13	Breidalsvatn	1967-2020	53	168 (147-191)			53	323 (303-347)		
14	Raudalsvatn	1967-2020	53	157 (136-176)			53	329 (313-365)		
17	Kaldfjorden	1967-2020	53	159 (136-170)			53	309 (285-332)		
19	Vinstern	1950-2020	64	163 (147-181)			69	317 (288-339)		
21	Bygdin	1950-2020	64	170 (153-185)	15	326 (301-382)	65	370 (315-416)	14	157 (130-221)
26	Volbufjorden	1920-1974	55	137 (119-150)	54	320 (305-344)	55	324 (312-353)	54	184 (164-214)
27	Øyangen	1919-1984	65	149 (130-168)	62	318 (299-343)	62	321 (304-344)	61	170 (137-200)
31	Bergsjø	1953-2020	58	160 (146-175)	47	304 (288-343)	56	314 (294-350)	47	144 (127-170)
33	Krøderen	1900-1964	64	124 (100-161)	7	335 (315-366)	60	338 (306-372)	7	214 (189-255)
35	Tunhovdfjorden	1920-2020	73	142 (119-161)	45	329 (275-353)	77	335 (305-362)	41	186 (142-219)
39	Hjartsjå	1919-1998	74	121 (91-139)	43	328 (311-354)	70	334 (313-388)	42	207 (184-261)
44	Sandvinvatn	1908-1998	59	106 (33-131)	61	383 (224-437)	64	398 (359-453)	46	276 (225-342)
45	Vangsvatn	1898-1989	69	113 (38-138)	46	347 (316-402)	78	354 (327-420)	61	236 (197-333)
46	Vassbygdvatn	1915-1987	69	116 (56-139)	56	356 (277-401)	65	371 (330-435)	54	242 (158-305)
48	Veitastrondvatn	1918-1991	65	137 (76-152)	52	353 (311-416)	61	356 (326-428)	50	217 (171-284)
51	Nautsundvatn	1908-1983	55	106 (33-130)	75	353 (314-426)	75	353 (314-426)	54	248 (215-348)
52	Hestadfjorden	1914-1995	70	117 (17-140)	75	358 (320-423)	77	371 (323-446)	65	242 (192-382)
55	Lovatn	1899-1979	72	108 (18-132)	44	388 (347-436)	51	388 (355-440)	42	281 (227-395)
68	Namsvatn	1908-1968	57	163 (137-184)	19	319 (301-341)	58	323 (291-351)	17	164 (126-183)
71	Tustervatn	1907-1968	54	156 (137-178)	44	328 (304-366)	50	343 (308-391)	41	174 (127-216)
76	Kobbvatn	1916-1978	61	149 (128-167)	58	330 (304-386)	60	339 (310-392)	56	185 (140-245)
90	Bjørnvatn	1912-1967	55	151 (130-182)	53	306 (286-327)	55	311 (289-366)	52	157 (117-189)
91	Murusjøen	1926-2001	66	142 (121-155)	74	327 (305-354)	66	336 (311-366)	65	184 (157-223)
95	Lenglingen	1925-2003	76	144 (118-158)	76	329 (307-383)	77	339 (312-385)	74	187 (157-235)
96	Engeren	1911-1983	72	139 (119-157)	72	347 (299-396)	71	350 (311-386)	71	204 (156-244)
97	Femunden	1900-1995	82	148 (128-173)	83	328 (305-353)	83	343 (313-386)	79	177 (152-214)
99	Møkeren	1911-2007	65	121 (91-141)	47	332 (261-363)	65	341 (303-446)	37	212 (128-244)