

This document contains two sections.

The first three sections contain the replies. The last section highlights the changes of the revised article in comparison to the originally submitted article.

I would like to thank Mike Wolovik and the two anonymous reviewers for their engagement and open discussion. It has improved the manuscript.

**PDF page counter    Content**

Page 2-5	Replies (minor revision)
Page 6-19	Markup version of revised article

Editor Decision: Publish subject to minor revisions (review by editor) (21 Jul 2019) by Benjamin Smith

Comments to the Author:

Thanks for thorough responses to the referees. There are some places where the language in the MS is a little rough--

**Thanks for taking the time to review the manuscript and for the valuable suggestions. Our comments and replies are in blue.**

Lines 75-78: Not sure what "for instance" is meant to indicate. This phrase needs a prior general phrase so that the reader can tell what general condition the instance is of. I don't see that here.

**We have reduced the unnecessary wordiness.**

134: another "for instance" that isn't quite clear-- the previous list of parameters seems fairly comprehensive, so it's not like those are only a few of the parameters that might be relevant.

**Yes we agree that it is a quite comprehensive list, but some may argue that also the ice properties, such as temperature and concentration of impurities, matter. Hence we have rephrased the entire sentence: "However, it is an open question if this triggers Marine Ice Sheet Instability in the other shelves, because the stability depends **strongly** on the distribution of pinning points, sloping of the bed, the depth and width of submarine troughs, and the softness of the bed, for instance."**

150: "a by expert judgement merged data products" -- this seems to be garbled somehow

**We've shortened this part and keep "data products"**

167: move 'for instance' directly after the word 'wall'

**We have moved as suggested.**

174: "may shifts also" - should be "may also shift"

**Thanks for indicating this mistake, which we have corrected.**

This is likely a partial list. Please look over the MS again to check for similar problems

**Beside the above listed items, we have also changed the following.**

**Line 17 : Delete comma:**

“coastal societies, and economic activities” =>

“coastal societies and economic activities”

**Line 20 : Replace : “ground water” => “groundwater”**

**Line 23 : Replace : “modelled” by “modeled”**

**Line 35 : Add comma:**

“During its transport onto the continental shelf the water” =>

“During its transport onto the continental shelf, the water”

Line 49 : Replace “ice berg” => “iceberg”

Line 53/54 : Add comma :

“In addition to the erection of a submarine wall they” =>

“In addition to the erection of a submarine wall, they”

Line 57 : Add comma :

“In this paper we investigate how“ =>

“In this paper, we investigate how“

Line 59 : Spelling: “sea level raise” => “sea level rise”

Line 71 : Add comma :

“ice shelf geometry and grounding line” =>

“ice shelf geometry, and grounding line”.

Line 81 : Shortened “provide confirmation that the” => “confirm that the”

Line 84 : Delete comma :

“continental shelf break, and“ =>

“continental shelf break and“

Line 88/89 : Rephrase :

“substantially larger than the originally proposed” =>

“substantially longer than the initially proposed”

Line 90 : Word order :

“experiment simulates accurately the ingress” =>

“experiment accurately simulates the ingress”

Line 102 : Delete comma :

“around Siple Island, because the” =>

“around Siple Island because the”

Line 105 : Add word :

“temperature reduces melting of” =>

“temperature reduces **the** melting of”

Line 105 : Shortened : “(Figure 2a and Figure 2c).” => “(Figure 2a and 2c).”

Line 113 : Add word :

“underneath Abbot” =>

“underneath **the** Abbot”

Line 115 : Rephrase :

“with the exception of the Amery Ice Shelf” => “except for the Amery Ice Shelf”

Line 117 : Add word :

“All above reported” =>

“All **the** above reported”

Line 143/144 : Word order :

“such as those seen in in Prydz Bay in front of Amery Ice Shelf or in the George VI Ice Shelf,” =>

“such as those seen in the George VI Ice Shelf or in Prydz Bay in front of Amery Ice Shelf,”

Line 150-152 : Rephrase : “Therefore, we use RTOPO instead of most updated products. Since our

simulations are consistent with former studies using RTOPO, the quality of our simulations could be judged in the light of former studies.” =>

“Therefore, we use RTOPO instead of **the** most updated products. Since our simulations are consistent with **prior** studies using RTOPO, the quality of our simulations could be judged in the light of **previous** studies.”

Line 153 : Add word :

“Shelf into walled” =>

“Shelf into **the** walled”

Line 154 : Change word :

“ice was grounded” =>

“ice **were** grounded”

Line 155 : Add comma :

“However we would” =>

“However, we would”

Line 157-159 : Rephrase : “These simulations would also uncover if enhanced melting at the western end of the wall may open a backdoor that open a second route” =>

“These simulations **could** also uncover if enhanced melting at the western end of the wall **unlocks** a backdoor that opens a second route”

Line 161/162 : Correct :

“warm water masses flows” =>

“warm water masses flow”

Line 166 : Change word :

“larger than the originally proposed” =>

“larger than the **initially** proposed”

Line 168/169: Rephrase : “Our results suggest that a too small wall blocking only the water flow in the troughs leading to Pine Island, for instance, might” =>

“Our results suggest that a too-small wall, **for instance**, blocking only the water flow in the troughs leading to Pine Island might”

Line 170 : Delete comma :

“hemisphere), if it is not hindered” =>

“hemisphere) if it is not hindered”

Line 172 : Add comma : “However, a small”

Line 175 : Rephrase :

“global warming (Miller et al., 2006) may shifts also the coast easterly” =>

“global warming (Miller et al., 2006) **might also shift** the coast easterly”

Line 181 : Add word :

“The wall blocks in inflow” =>

“The wall blocks in **the** inflow”

Line 203/204 : Rephrase :

“is freely accessible online (for example at

<https://data1.gfdl.noaa.gov/nomads/forms/core/COREv2.html>.” =>

“is freely accessible online at <https://data1.gfdl.noaa.gov/nomads/forms/core/COREv2.html>.”

Caption Figure 1 : Changed text : “Modeled and observed seafloor ocean potential temperatures

( $\theta$ bottom) in the Amundsen Sea Sector of West Antarctica. The plot shows the simulated mean ocean temperatures for the control run (CTRL, years 1948—2007), while circles (ANT XI/3; Miller and Grobe, 1996) or diamonds (ANTXXVI/3; Gohl, 2010) represents observed bottom temperatures taken in 1994 and 2010, respectively. The solid green lines follow shelf ice edges, and the inset shows the location of the area of interest as a red box.”

Caption Figure 2 : Changed text : “The ice shelf edges are highlighted by solid green lines.” => “Solid green lines highlight ice shelf edges.”

Caption Figure 3 : Changed text : “Mean basal melting rates around Antarctica. Longitude-specific changes in modeled basal melting with (WALL) and without (CTRL) the presence of the submarine wall are shown as dashed red and solid blue lines surrounding the center map, respectively. The black dashed line marks the wall’s location in the Amundsen Sea. The map of Antarctica shows the spatial distribution of the melting rate anomaly, where positive numbers (red color) represents increased melting rates if the wall is present (see colorbar). The following abbreviations are used: Pine IG (Pine Island Glacier), Thwaites G (Thwaites Glacier) and Getz IS (Getz Ice Shelf). Green dotted lines depict shelf ice edges of the Ross Ice Shelf, the Filchner-Ronne Ice Shelf and the Amery Ice Shelf.”

# Brief communication: A submarine wall protecting the Amundsen Sea intensifies melting of neighboring ice shelves

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## Abstract

Disintegration of ice shelves in the Amundsen Sea, in front of the West Antarctic Ice Sheet, has the potential to cause sea level rise by inducing an acceleration of ice discharge from upstream grounded ice. Moore et al. (2018) proposed that using a submarine wall to block the penetration of warm water into the sub-surface cavities of these ice shelves could reduce this risk. We use a global sea ice-ocean model to show that a wall shielding the Amundsen Sea below 350 m depth successfully suppresses the inflow of warm water and reduces ice shelf melting. However, these warm water masses get redirected towards neighboring ice shelves, which reduces the net effectiveness of the wall. The ice loss is reduced by 10% integrated over the entire Antarctic continent.

## 15 1 Introduction

One of the consequences of the warming of Earth's climate is sea level rise (Vaughan et al., 2013). Sea level rise will impact coastal societies, and economic activities in these areas. Currently the main contributors to rising global mean sea level are the thermal expansion of the world's ocean, the mass losses emanating from the Greenland Ice Sheet, and the world-wide recession of mountain glaciers and ice caps (Chen et al., 2017; Rietbroek et al., 2016; Shepherd et al., 2012). The remaining smaller sources are continental ground-water depletion (Wada et al., 2012) and the Antarctic Ice Sheet (King et al., 2012; Rietbroek et al., 2016; Shepherd et al., 2012); though Antarctica's sea level contribution has accelerated in recent decades (King et al., 2012; Rietbroek et al., 2016; Rignot et al., 2011).

In Antarctica, remotely sensed, ~~modelled~~modeled and paleoclimatological-proxy data indicate that the highest potential sea level contribution will come from the West Antarctic Ice Sheet (Bamber et al., 2009; Golledge et al., 2013; Joughin and Alley, 2011; Pollard and DeConto, 2009; Sutter et al., 2016), particularly from the Amundsen Sea Sector, where the progressive thinning of its ice shelves over the last two-and-a-half decades has greatly enhanced rates of ice mass loss emanating from this sector (Pritchard et al., 2012; Rignot et al., 2014; Shepherd et al., 2018). Here ice shelves currently prevent unrestricted flow of ice streams into the ocean. Here, warm high salinity Circumpolar Deep Water (CDW) has been observed to flow onto the continental shelf and flood the cavities underneath the Amundsen Sea Sector's ice shelves, driving

30 high rates of basal melting (Depoorter et al., 2013; Jacobs et al., 2011; Jenkins et al., 2018; Pritchard et al., 2012). Various processes control the flow of warm water masses (a body of ocean water with a common formation history and a defined range of tracers, such as temperature and salinity, is called water mass) predominately via glacially scoured submarine troughs (Bingham et al., 2012; Dutrieux et al., 2014) into the ice shelf cavities. It includes wind-driven Ekman transport, whereby variations in offshore wind stresses, also altered by local sea ice conditions (Kim et al., 2017), lift CDW onto the 35 continental shelf (Dutrieux et al., 2014; Kim et al., 2017; Paolo et al., 2018; Schmidt et al., 2013). During its transport onto the continental shelf, the water mass is transformed into modified CDW (mCDW) by mixing with local, fresher on-shelf water masses (Webber et al., 2017, 2018). In the Amundsen Sea, decadal-scale changes in the draft and intensity of the CDW incursion onto the continental shelf – and ultimately the basal melting of the ice masses fringing this sector of the 40 Antarctica – have also been directly linked to changes in the large-scale oceanic and atmospheric circulation, including the influence of ENSO-induced atmospheric wave trains propagating towards this region from the central tropical Pacific Ocean (Dutrieux et al., 2014; Jenkins et al., 2018; Nakayama et al., 2018; Steig et al., 2012). These processes together drive the detected retreat of ice shelves in the Amundsen Sea through decadal oceanographic variability (Jenkins et al., 2018).

45 Since the West Antarctic Ice Sheet resides on retrograde sloping bedrock topography (Fretwell et al., 2013), it is inherently susceptible to the Marine Ice Sheet Instability (Schoof, 2007; Weertman, 1974), whereby the reduced buttressing effect of thinning ice shelves triggers the retreat of upstream ice, leading to larger ice thickness at the grounding line – the grounding line marks the transition from grounded ice to floating ice. This amplifies the ice flux across the new grounding line, which stretches and thins the ice further and ultimately triggers additional grounding line retreat. This sustained retreat accelerates the transport of grounded inland ice towards the ocean past the grounding line, where it directly contributes to sea level rise 50 (by ice-berg calving and ocean-driven melting).

50 Moore et al. (2018) proposed a targeted geoengineering project that could reduce the risk of this ice sheet instability by protecting the ice shelves from warm Circumpolar Deep Water via the erection of a submarine wall. Wolovik and Moore (2018) tested this idea with a simple flow line model (2-dimensional xz-plane model) of Thwaites Glacier – one of the largest contributors of ice discharge into the Amundsen Sea (Shepherd et al., 2018; Turner et al., 2017). In addition to the erection of a submarine wall, they (cf. Moore et al., 2018) imposed artificial pinning points to enhance the buttressing effect 55 of ice shelves on grounded ice. Both measures successfully reduced ice mass loss emanating from this sector of Antarctica (Wolovick and Moore, 2018).

55 In this paper, we investigate how a submarine wall, shielding the Amundsen Sea Embayment (Figure 2a), reduces the basal melting of rates ice shelves flowing into the Amundsen Sea Embayment. The warm water masses rejected by the wall enhance ice shelves west of the wall. These effects counteract the wall's purpose mitigating sea level rise. In this study, we 60 neglect feedbacks between changes of basal melting rates and advance or retreat, respectively, of impacted ice shelves. We do not analyze how the wall hinders the exchange of nutrients and influences submarine biological processes.

## 2 Model setup

We use a global Finite Element Ocean Model (FESOM; Wang et al., 2014) to test the effects of erecting a wall in front of the Amundsen Sea Sector's ice shelves. The model has a variable horizontal resolution of (minimum 5 km) around Antarctica and its adjacent ice shelf cavities and has 100 vertical levels (z-coordinate). The interaction between the ocean and static ice shelves occurs via the three-equation system that describes the flux of heat and fresh water between the ocean and ice shelf base through an exchange controlling boundary layer (Hellmer and Olbers, 1989; Holland and Jenkins, 1999). FESOM has proven its applicability for oceanographic studies of the Southern Ocean (Hellmer et al., 2012; Nakayama et al., 2014; Timmermann et al., 2012).

While coarse resolution models have been found to underestimate the ocean-induced basal melting of Antarctica's ice shelves (Naughten et al., 2018), our basal melting rates (Figure 1b) are in reasonable agreement with recent observational estimates (Rignot et al., 2013). The model utilizes the ocean bathymetry, ice shelf geometry, and grounding line position data of RTOPO1 (Timmermann et al., 2010). We use the CORE2 forcing for atmospheric conditions (Large and Yeager, 2008) covering the years 1948—2007 to drive the ocean model. This forcing period is run twice. The first full period is considered as spin-up and, hence, we restrict our analysis on the last complete forcing period.

Considerable oceanic variability has been detected at both seasonal and interannual timescales in front of both Pine Island (Webber et al., 2017) and Dotson Ice Shelf, located between Thwaites Glacier and Getz Ice Shelf, (Jenkins et al., 2018). It is driven by both local and remote forcing. Hence we shall expect some differences between merged hydrographic observations and a simulated long-term mean, while a reliable climatological data set is lacking for our region of interest. Therefore, we use existing observations for comparison with our simulations under the assumption that available observations represent a quasi-mean state. Measured bottom temperatures, predominantly taken in austral summer by the marine cruises *ANT XI/3* (Miller and Grobe, 1996) and *ANT XXVI/3* (Gohl, 2010), provide confirmation that the simulated bottom temperature distribution is reasonable (Figure 1).

We investigate differences in ice shelf basal melting with (WALL) and without (CTRL) the erection of a wall surrounding the Amundsen Sea (Figure 2a). This feature follows the approximate location of the continental shelf break, and blocks any circulation below 350 m depth, such as the CDW inflow from the deep ocean onto the Amundsen Sea Sector's continental shelf (Figure 2). The wall proposed by Moore et al. (2018) blocks only the channelized flow of warm water in troughs leading directly to Pine Island and Thwaites Glaciers, while our wall with a length of about 800 km, is substantially larger than the originally initially proposed wall in size, and shields the entire Amundsen Sea Embayment.

## 3 Results

Consistent with oceanographic observations, our CTRL experiment accurately simulates accurately the ingress and delivery of warm mCDW through submarine troughs towards the ice shelves fringing the Amundsen Sea Sector (Figure 1). Measured bottom temperatures, acquired in austral summer, also strongly agree with the spatial distribution of our simulated temperatures, giving confidence in our abilities to accurately predict basal melting in the present study (Figure 1).

Contrary to our CTRL experiment, our erected wall blocks the ocean below 350 m depth and suppresses the direct inflow of  
95 CDW to the interior of the Amundsen Sea Embayment in front of the western Marie Byrd Land. Consequently, the simulated ocean is generally cooler (Figure 2a) and fresher within the walled region. This colder water column supports enhanced sea ice formation, which releases brine into the underlying ocean across this region. However, the brine-induced salinification is insufficient to compensate the salinity supply of the unobstructed mCDW inflow.

We also detect a slight cooling of the bottom temperatures east of the walled region. The outflow of cooler water masses  
100 from the walled region via the Abbot Ice Shelf's sub-ice shelf cavity (south of Thurston Island) contributes to this cooling (Figure 2a and 2b). The deflected warm water mass flows westward and rises the temperature on the west side of the walled region. This causes the temperature to rise in the westernmost corner of the walled region around Siple Island, because the warm water mass penetrates via the Getz Ice Shelf (between the groundling line of Antarctica and Siple Island) into the walled region.

105 In the walled region, the lower ocean temperature reduces the melting of ice shelves (Figure 2a and Figure 2c). However, the restrained warm water mass advances into the neighboring region, where ice shelves experience intensified melting and amplified ice mass loss (central and western Getz Ice Shelf; Figure 2b). Therefore, the warm water mass that would have otherwise impacted the Amundsen Sea Embayment, shifts to neighboring ice shelves.

Figure 3 depicts the longitudinal distribution of the simulated basal melting rates around Antarctica, with and without  
110 erection of the submarine wall. In the Amundsen Sea Embayment the ice mass loss around Pine Island drops significantly by 85 %. This phenomenon contrasts with increases ice mass loss detected at Getz Ice Shelf (~130° W, eastern Marie Byrd Land), where melting increases by approximately 50%. As discussed above, basal melting is reduced in the western Bellingshausen Sea. In addition to the decreased melting simulated underneath the Abbot Ice Shelf, basal melting at George VI Ice Shelf increases by up to 10%. The wall has little impact on basal melting of ice shelves fed by ice streams from the  
115 East Antarctic Ice Sheet, except for with the exception of the Amery Ice Shelf, where the rate increases by approximately 5 %. The wall in the Amundsen Sea triggers most probably a perturbation that propagates via the Antarctic Coastal Current towards the Prydz Bay in front of the Amery Ice Shelf. All the above reported intensified melting rates are larger than the standard deviation (1-sigma) of the 20 years melting rate.

Beside regional changes of the basal melting rates, we inspect the continental-wide integrated effect. The reduced ice loss in  
120 the Amundsen Sea Embayment is larger than the corresponding enhanced melting at the western end of the wall. The total ice loss by ice shelves around Antarctica is 10% lower for the WALL experiment.

#### 4 Conclusions

In this study, a submarine wall erected along the continental shelf of the Amundsen Sea is found to suppress the inflow of circumpolar deep water onto the continental shelf. This freshens and cools water masses residing shoreward of the wall, resulting in significantly reduced basal melting rates of the ice-shelves located there. However, inflowing warm Circumpolar  
125

Deep Water (CDW) seaward of this wall is found to be redirected westward towards Getz Ice Shelf, where it enhances basal melting by up to 50%. In particular, the ice shelves to the west (central and west Getz Ice Shelf) show steeply increased melting rates. Hence the wall reduces the ice loss of the most vulnerable ice shelves along the margin of the Western Antarctic Ice Sheet, which is, however, not compensated by enhanced melting in the west. Integrated over Antarctica, the 130 wall decreases ice loss by 10 %. Our results indicate that suppressing the flow of warm water masses into a restricted group of ice shelves results in redirecting it towards a different location. There it enhances basal melting and, ultimately, amplifies ice mass loss. However, it is an open question if this triggers Marine Ice Sheet Instability in the other shelves, because the stability depends strongly on the distribution of pinning points, sloping of the bed, the depth and width of submarine troughs, and the softness of the bed, ~~for instance~~. The onshore bed properties of the eastern Marie Byrd Land, where Pine Island and 135 Thwaites Glaciers are located, are most likely vulnerable to the Marine Ice Sheet Instability. Numerous modelling studies show a relic ice cap in the western Marie Byrd Land on the elevated bed-rock topography even after part of the West Antarctic Ice Sheet (WAIS) has collapsed (e.g. DeConto and Pollard, 2016; Feldmann and Levermann, 2015; Golledge et al., 2015; Winkelmann et al., 2015). Hence the western Marie Byrd Land is favorable for a stable situation. Though, the bed properties under the ice are still known insufficiently.

140 Our fully coupled sea ice-ocean model, which includes ice shelves and ocean-ice shelf interaction, is driven by a prescribed atmospheric forcing. Hence any feedbacks, such as changing ocean surface conditions that impact the atmosphere and change the atmospheric forcing on the ocean, are not included. Therefore, small anomalies between both simulations (CTRL vs. WALL), such as those seen in the George VI Ice Shelf or in Prydz Bay in front of Amery Ice Shelf ~~or in the George VI Ice Shelf~~, could vanish if we would include atmosphere-ocean feedbacks. Here only simulations coupled to the atmosphere 145 would allow confirming the robustness of these features.

The used bedrock topography and ice shelf geometry data set influences the melting rate of individual ice shelf caverns, as it has been shown for the smaller Crosson and Dotson Ice Shelves draining also into the Amundsen Sea (Goldberg et al., 2019). Hence most updated data sets would be preferred; however existing inconsistencies between the most updated data 150 sets, which are seen in differences of the reported grounding positions, free-board heights and bedrock elevation, require the use of ~~a by expert judgement merged~~ data products, such as RTOPO. Therefore, we use RTOPO instead of the most updated products. Since our simulations are consistent with former prior studies using RTOPO, the quality of our simulations could be judged in the light of former previous studies.

Would we detect the penetration of warm water masses via the Getz Ice Shelf into the walled region if we use other bathymetry or bedrock topography data sets? If all ice ~~was were~~ grounded between the western end of the wall and the coast 155 line, we would not see any flow of warm water into the walled region. However, we would detect enhanced melting, which may open up a route into the protected region. Hence fully coupled ice sheet-ocean model simulations, where the geometry of ice shelves are changed by melting and refreezing, would reveal the vulnerability of the Getz Ice Shelf. These simulations ~~would could~~ also uncover if enhanced melting at the western end of the wall ~~may open unlocks~~ a backdoor that opens a second route to the ice shelves prone to Marine Ice Sheet Instability.

160 Regardless of the used bathymetry data set, we are confident that the main findings of this study are robust: a wall shielding the Amundsen Sea Embayment reduces basal melting rates within the protected region, the rejected warm water masses flows along the wall westward, west of the wall warmer water masses drive enhanced basal melting. The wall proposed by Moore et al. (2018), which blocks only the circulation in troughs leading directly to Pine Island and Thwaites Glaciers, would have a length of about 50–100 km and would need 10–50 km<sup>3</sup> of material. By comparison, the construction of the 165 Suez Channel required the excavation of about 1 km<sup>2</sup> of material (Moore et al., 2018). The simulated wall (length of about 800 km) is substantially larger than the originally initially proposed wall in size, and it shields the entire Amundsen Sea Embayment.

170 Our results suggest that a ~~too too~~-small wall, for instance, blocking only the water flow in the troughs leading to Pine Island, for instance, might be bypassed by warm water masses. For dynamical reasons, the (geostrophic) flow of water masses turns to the left (on the Southern hemisphere), if it is not hindered by a topographic obstacle. Therefore, warm water masses might even recirculate into the ostensibly protected area if the wall is too small, as the inflow of warm water masses through the 175 Getz Ice Shelf into the walled region suggests. However, a small wall that only protects Pine Island successfully, may redirect the warm water to neighboring ice shelves with a retrograde bed (for example, Thwaites Glacier). There it increases basal melting and may trigger Marine Ice Sheet Instability. The detected poleward shift of westerly winds in the Southern Ocean under global warming (Miller et al., 2006) ~~may shifts might~~ also ~~shift the coast~~ easterly winds along Antarctica's coast poleward, which lifts further the interface of warm water masses (isothermal) along the continental slope (Spence et al., 2014). Ultimately warm water masses could enter the continental shelf directly beside the contemporary path following topographic depressions (troughs). Under these circumstances, the bypassing of a short wall seems to be inevitable, if the wall does not block the entire Amundsen Sea Embayment from coast to coast.

180 Iron is a micronutrient essential for algal production in the Amundsen Sea (St-Laurent et al., 2017) and the erected wall affects its availability. The wall blocks in ~~the~~ inflow of warm and iron-rich CDW and influences the outflow of iron-rich glacial melt-water coming from melting ice shelves. How the changed nutrient supply impacts the marine biological web or the uptake and sequestration of carbon dioxide by the ocean is unclear and goes beyond this study.

185 Geoengineering aims to attenuate the impact of the ongoing anthropogenic climate change, such as sea level rise, but the results of this study suggest that such proposals could have adverse side effects. To evaluate these effects of using submarine walls to protect Antarctica's ice shelves in greater detail, the use of fully coupled ice-sheet-shelf-ocean-atmosphere models should be utilized in future analyses. These models of sufficiently high spatial resolution could simulate accurately changes in sub-ice shelf cavity geometry (including track grounding line migration and ice-shelf thinning) as well as the influx of warm water masses (mCDW) to these locations.

## 190 Acknowledgments

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his comments, which greatly improved this manuscript. We thank both anonymous reviewers and Mike Wolovick for their comments and engagement in the review process, which improved the manuscript.

## 195 **Author contributions**

CR designed the study and wrote the manuscript. ÖG and QW developed and configured the model. ÖG ran the simulations. ÖG and VK performed the analysis. VK and CR prepared the figures. All authors contributed to the interpretation of the results and proofreading of the manuscript.

## **Competing interests**

200 The authors declare that they have no conflict of interest.

## **Code and Data availability**

The FESOM1.4 model code is available at <https://swrepo1.awi.de/projects/fesom/> after registration. The here used atmospheric forcing data set named “CORE-II” (Large and Yeager, 2008) is freely accessible online (for example at <https://data1.gfdl.noaa.gov/nomads/forms/core/COREv2.html>). The topography data set RTOPO could be obtained from <https://doi.pangaea.de/10.1594/PANGAEA.741917>. The temporal average of the fractional basal melting changes between the CTRL and the WALL simulations is obtainable from Zenodo via <https://dx.doi.org/10.5281/zenodo.3240250>. The remaining data is available from the first author ÖG upon reasonable request.

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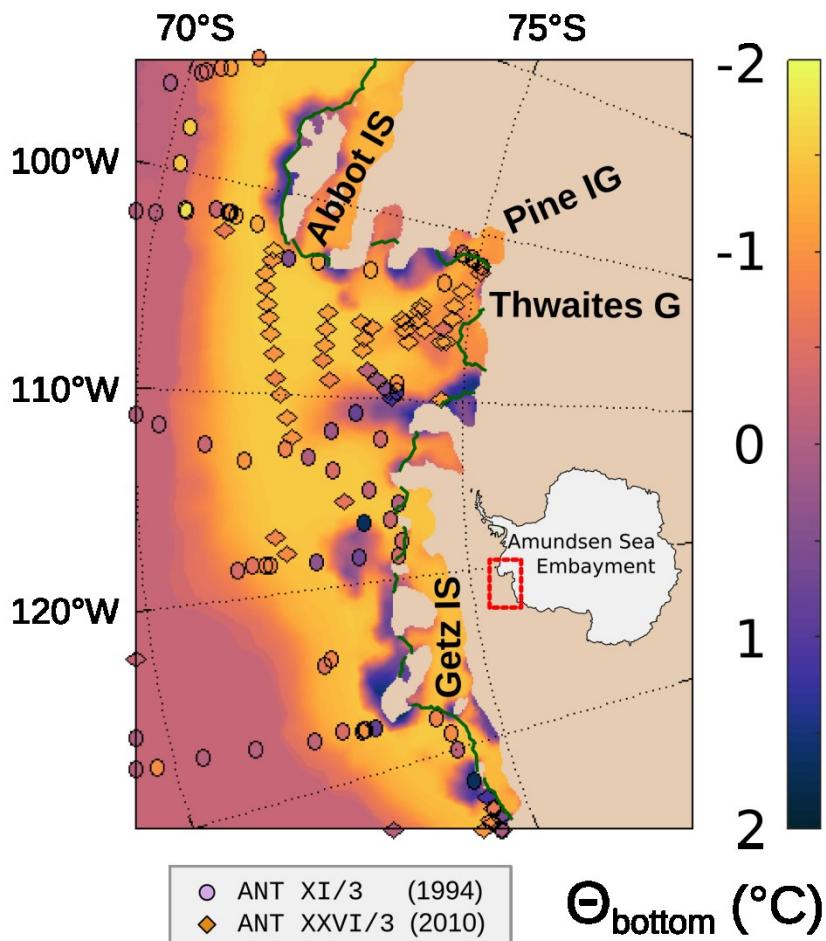
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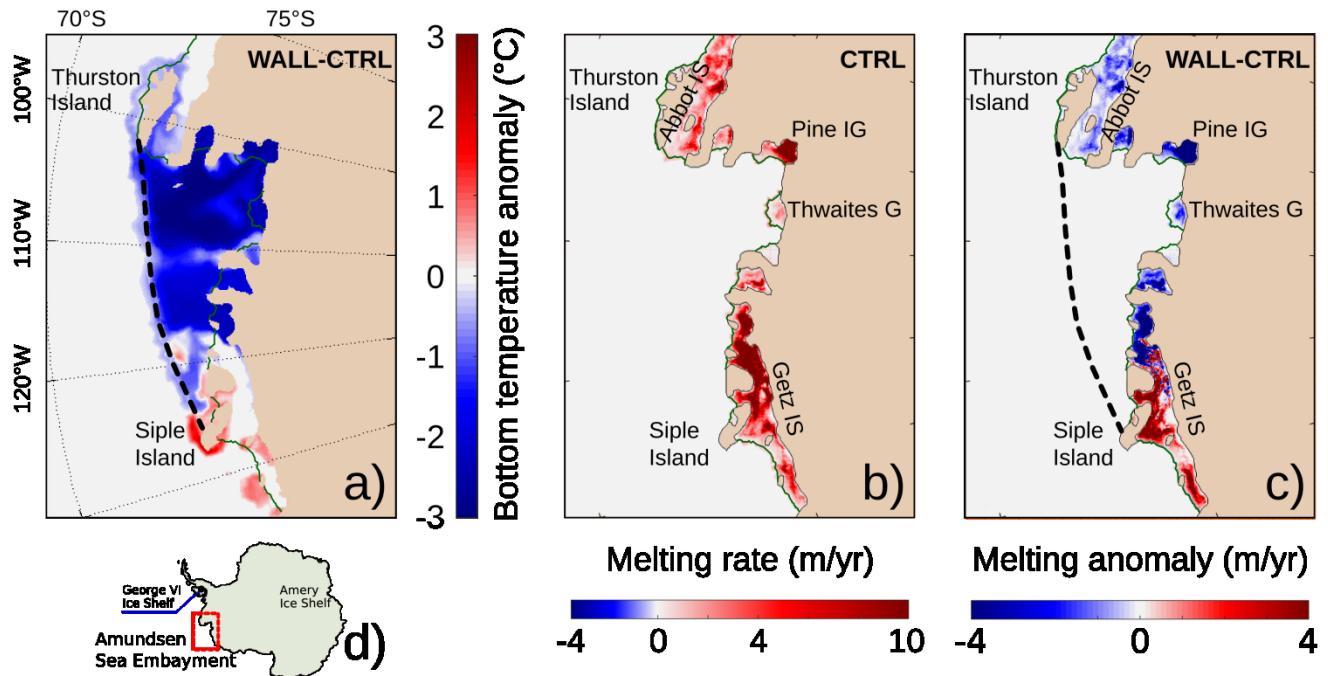
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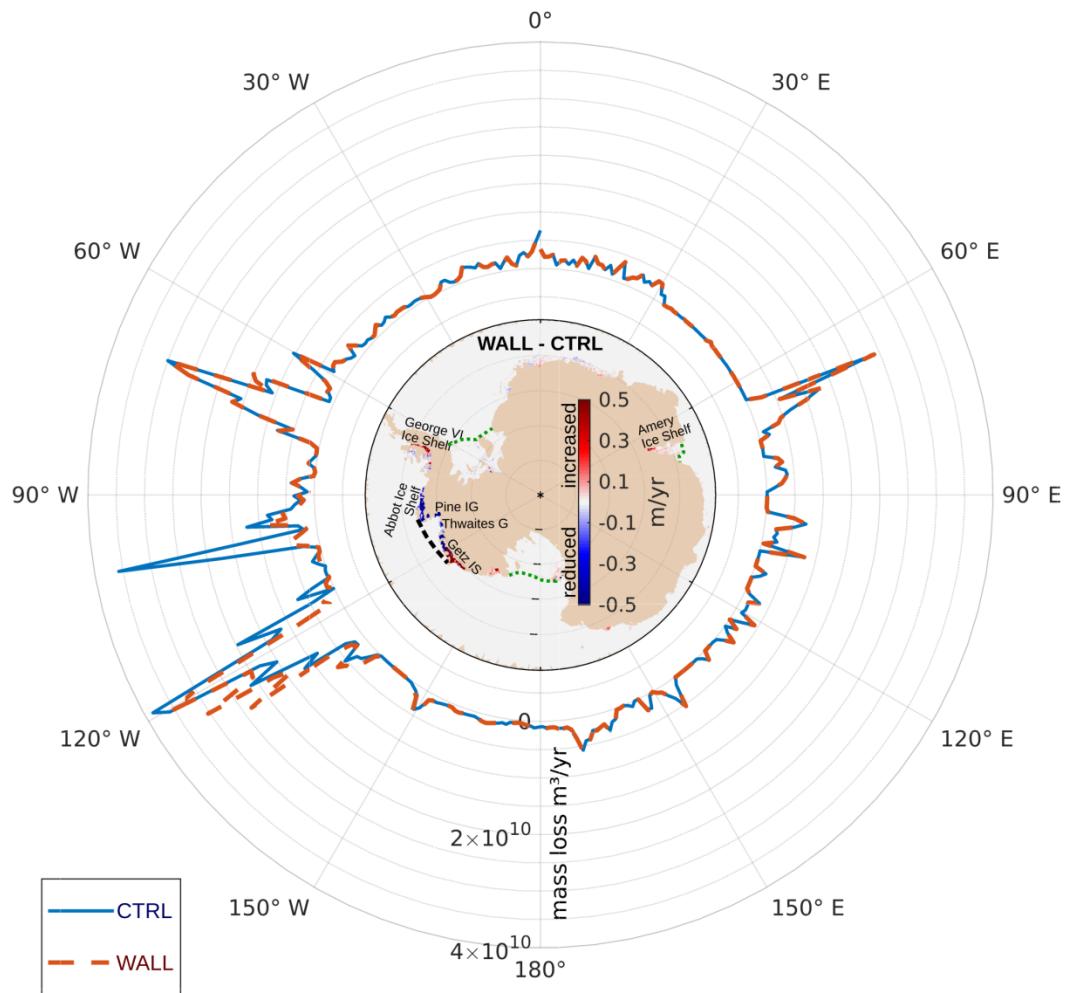
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360 | **Figure 1** ~~Modeled~~ and observed seafloor ocean potential temperatures ( $\Theta_{\text{bottom}}$ ) in the Amundsen Sea Sector of West Antarctica. ~~Inset shows study location.~~ The plot shows the simulated mean ocean temperatures for the control run (CTRL, years 1948–2007), while ~~individual observed bottom temperatures are represented by~~ circles (ANT XI/3; Miller and Grobe, 1996) or diamonds (ANTXXVI/3; Gohl, 2010) ~~represents observed bottom temperatures~~ taken in 1994 and 2010, respectively. ~~The solid green lines follow the shelf ice edge~~ is drawn as solid green line, and the inset shows the location of ~~the~~ area of interest as a red box.



365 | Figure 2 Simulated potential ocean temperature anomaly (WALL – CTRL) a), simulated basal ice shelf melting rates (CTRL) in  
 b) and its anomaly c). The subplot 2a) shows the simulated potential ocean temperature anomaly (WALL – CTRL) on the seafloor  
 370 | of the Amundsen Sea Embayment and its adjacent ice shelf cavities. The location of the wall is marked as a dashed line, and the  
 | middle subplot b) show the simulated melting rates for the control run (CTRL) and the right subplot c) shows basal melting anomaly (WALL - CTRL). **Solid green lines highlight The ice shelf edges are highlighted by solid green lines.** The following abbreviations are used: Abbot IS (Abbot Ice Shelf), Pine IG (Pine Island Glacier),  
 Thwaites G (Thwaites Glacier) and Getz IS (Getz Ice Shelf).



375 | **Figure 3 Mean basal melting rates around Antarctica.** Longitude-specific changes in ~~modelled~~ modeled basal melting with (WALL) and without (CTRL) the presence of the submarine wall are shown as dashed red and solid blue lines surrounding the center map, respectively. ~~The black dashed line marks~~ The the wall's location in the Amundsen Sea ~~is marked by the black dashed line~~. The map of Antarctica shows the spatial distribution of the melting rate anomaly, where positive numbers (red color) represents increased melting rates if the wall is present (see colorbar). The following abbreviations are used: Pine IG (Pine Island Glacier), Thwaites G (Thwaites Glacier) and Getz IS (Getz Ice Shelf). ~~Green dotted lines depict~~ The shelf ice edges of the Ross Ice Shelf, the Filchner-Ronne Ice Shelf and the Amery Ice Shelf ~~are depicted by green dotted lines~~.

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