

Glacier runoff variations since 1955 in the Maipo River Basin, semiarid Andes of central Chile

Supplementary Information

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Section S1: Geodetic mass balances

This section presents a summary of the methods included in the study by Farías-Barahona et al. (2019a):

- 20 Farías-Barahona, D., Ayala, Á., Bravo, C., Vivero, S., Casassa, G., Seehaus, T., Vijay, S., Loriaux, T., Huenante, J., Buglio, F., Carrasco, J., Lippl, S. and Braun, M.: 60 Years of Glacier Elevation and Mass Changes in Maipo Basin, central Andes of Chile, To be Submitt. to Front. Earth Sci. - Cryospheric Sci., 2019a.

Derivation of the 1955 DEM

- 25 A Digital Elevation Model (DEM) of the Chilean Central Andes was derived from a topographic map at a 1:50,000 scale. The map was based upon aerial photographs taken in 1955 (Hycon flights), as part of an agreement between Chile and the Inter-American Geodetic Survey of the United States. The cartography was produced by the Chilean Military Geographic Institute (IGM) using analogue photogrammetric techniques, which requires an operator that stereoscopically reconstruct the imaging geometry using mechanical or optical devices (Mikhail et al., 2001). The vertical reference used mean sea level corresponding to precise trigonometric levelling with theodolites from tide gauge benchmarks located on the Chilean coast. The maps were
30 georeferenced to the PSAD56 datum, UTM projection zone 19S.

The generation of Digital Elevation Models (DEMs) for 1955 consists of the following steps (Farías-Barahona et al., 2019b):

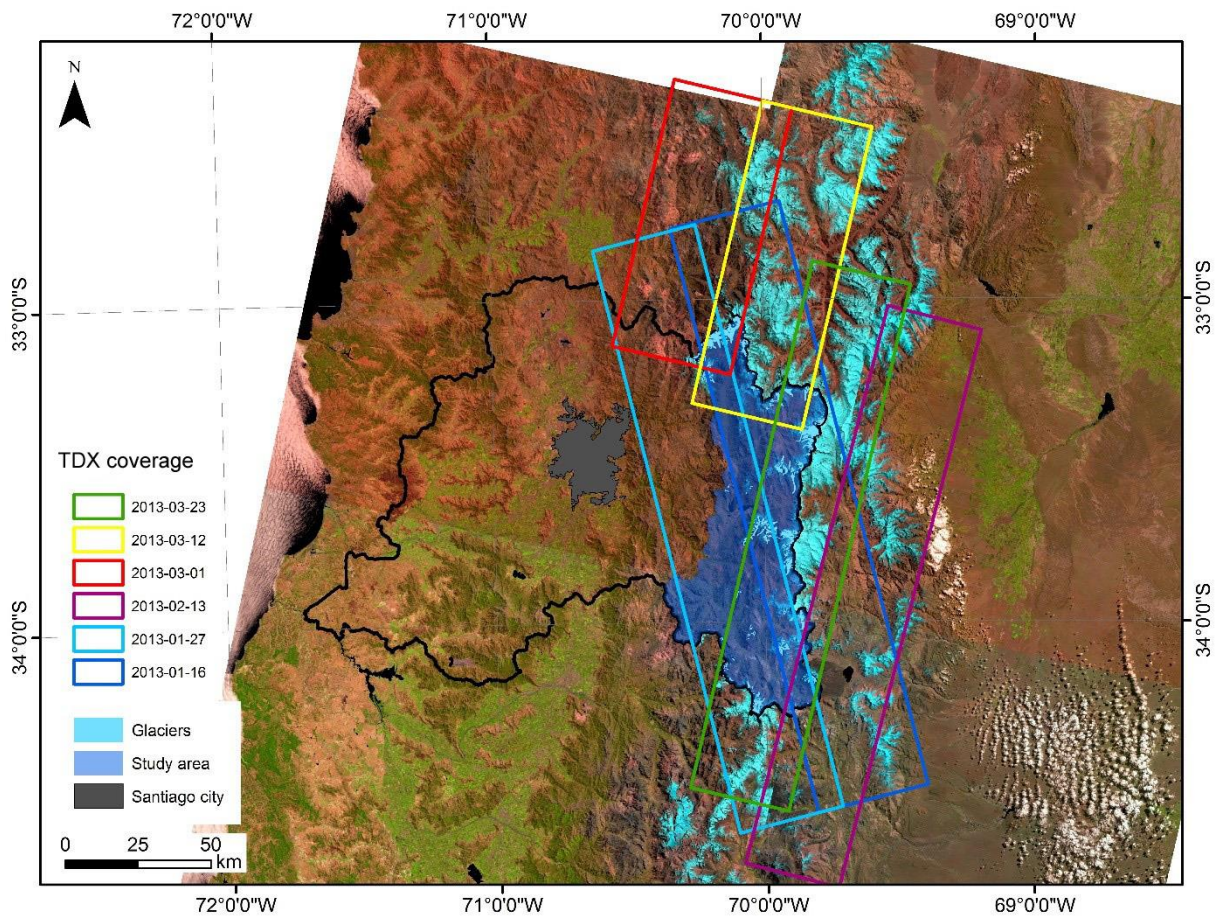
- The 1955 topographic maps containing 50-m contour lines were scanned by the IGM at 1200 DPI.
- The scanned maps were georeferenced to a common base using standard Geographic Information System (GIS) procedures. The original PSAD 56 datum was converted to the WGS84 datum.
- 35 • The 50-m contour lines were digitized.
- The contour lines were interpolated using the kriging method (e.g. Mölg et al., 2017; Farías-Barahona et al., 2019). A DEM with a final resolution of 30 m was then generated from the contours (Farías-Barahona et al., 2019b).

Derivation of the 2013 DEM

A DEM of the study area for year 2013 was derived from TanDEM-X as follows:

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- A set of TanDEM-X scenes from March 2013 was selected (see Figure S1).
 - The TanDEM-X scenes were processed based on differential SAR interferometry.
 - The interferograms were filtered and a phase unwrapping procedure was performed using the minimum cost flow algorithm. The data were then converted to differential heights.

A detailed description of the derivation of the TanDEM-X DEM processing can be found in Braun et al. (2019).



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Figure S1. TanDEM-X (TDX) coverage for the Maipo River Basin in 2013. Background image from USGS Landsat 8 satellite data from February 26, 2014.

DEM differencing and uncertainties

- 50 The geodetic mass balances for the periods 1955-2000 and 2000-2013 were derived using DEM differencing with the SRTM DEM. This product was obtained by bi-static radar interferometry that were acquired simultaneously in the C-band and X-band frequencies between the 11th and 22nd of February 2000 (Farr et al., 2007). This date corresponds to the ablation period in the Southern Hemisphere. Using optical satellite imagery, it was checked that melt-favourable conditions dominated over the Central Andes at the time of the SRTM campaign. The DEMs were co-registered (Nuth and Kääb, 2011) using a mask of stable areas with the Shuttle Radar Topography Mission (SRTM-C) as reference.
- 55 The uncertainty in the geodetic mass balances was explicitly quantified based on standard error propagation methods (Braun et al., 2019). Four sources of uncertainty were recognized: i) volume-to-mass conversion, ii) glacier outlines, iii) radar signal penetration of SRTM, and iv) errors in the DEM differencing.

Table S1: Parameters in TOPKAPI-ETH's sub-surface flux module for the 1955-2016 time period

Parameter	Calibrated value		Units
	Areas of the Maipo River Basin (*)		
	Vegetated areas	Non-vegetated areas	
Depth	6.5	2.1	°C
Depth (lower layer)	27.6	12.2	m
Vertical hydraulic conductivity at saturation	0.074	0.045	m s ⁻¹
Vertical hydraulic conductivity at saturation (lower layer)	0.026	0.056	m s ⁻¹
Horizontal hydraulic conductivity at saturation	0.003	0.015	m s ⁻¹
Horizontal hydraulic conductivity at saturation (lower layer)	0.438	0.605	m s ⁻¹
Water content ratio at saturation	0.65	0.66	-
Water content ratio at saturation (lower layer)	0.66	0.31	-
Residual water content ratio	0.1	0.1	-
Residual water content ratio (lower layer)	0.1	0.1	-
Brooks-Corey exponent for the permeability-saturation curve (horizontal)	3.5	3.5	-
Brooks-Corey exponent for the permeability-saturation curve (horizontal) (lower layer)	3.5	3.5	-
Clapp-Hornberger exponent for the percolation equation (vertical)	13.0	13.0	-
Clapp-Hornberger exponent for the percolation equation (vertical) (lower layer)	13.0	13.0	-
Soil drainage coefficient. Used for upper and lower soil layers	1	1	-
Minimum values of soil moisture to have ETA greater zero (ratio of maximum soil moisture between 0-1)	0.1	0.1	-
Minimum values of soil moisture for ETA to be equal to ETP (ratio of maximum soil moisture between 0-1)	0.35	0.35	-

(*): The classification was made based on the land use maps developed by the National Forest Corporation (CONAF) database (<https://sit.conaf.cl/>).

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Table S2: Summary of the performed simulations

Name of simulation	Method	Domain (*)	Period	Scenario (**)	Comments
SIM-1A	TOPKAPI-ETH	Maipo River Basin	1955-2016	Observed climate	The subperiod 2003-2013 is used for model calibration. The subperiod 1984-2002 is used for model validation. Snow parameters are chosen based on standard values in the literature. Soil parameters are automatically calibrated.
SIM-1B	TOPKAPI-ETH	Selected glacier x	1955-2000	Observed climate	Parameters SRF and TF are used to fit the geodetic mass balance of each glacier. The rest of the parameters are chosen based on standard values in the literature.
	TOPKAPI-ETH	Selected glacier x	2000-2016	Observed climate	Parameters SRF and TF are used to fit the geodetic mass balance of each glacier. The rest of the parameters are chosen based on standard values in the literature.
SIM-1C	Extrapolation	All glaciers	1955-2016	Observed climate	Extrapolation calculated from SIM-1B
SIM-2A	TOPKAPI-ETH	Maipo River Basin	2000-2099	Synthetic scenario y	As in SIM-1A
SIM-2B	TOPKAPI-ETH	Selected glacier x	2000-2099	Synthetic scenario y	As in SIM-1B
SIM-2C	Extrapolation	All glaciers	2000-2099	Synthetic scenario y	Extrapolation calculated from SIM-2B

(*): Selected glaciers x go from 1 to 26, (**): Future scenarios y go from 1 to 10.

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