1	Impact of exhaust emissions on chemical snowpack composition at Concordia
2	Station, Antarctica
3	Detlev Helmig ^{1,*} , Daniel Liptzin ¹ , Jacques Hueber ¹ , and Joel Savarino ²
4	¹ Institute of Arctic and Alpine Research, University of Colorado, Boulder, CO 80309, USA
5	² Univ. Grenoble Alpes, CNRS, IRD, Grenoble INP, IGE, F-38000 Grenoble, France
6	*corresponding author: detlev.helmig@colorado.edu
7	
8	Manuscript submitted to The Cryosphere - Revised Version
9	September 10, 2019
10	

11 Abstract

12 The chemistry of reactive gases inside the snowpack and in the lower atmosphere was investigated at 13 Concordia Station (Dome C), Antarctica, from December 2012 - January 2014. Measured species included 14 ozone, nitrogen oxides, gaseous elemental mercury (GEM), and formaldehyde, for study of photochemical 15 reactions, surface exchange, and the seasonal cycles and atmospheric chemistry of these gases. The 16 experiment was installed \approx 1 km from the station main infrastructure inside the station clean air sector 17 and within the station electrical power grid boundary. Ambient air was sampled continuously from inlets 18 mounted above the surface on a 10 m meteorological tower. In addition, snowpack air was collected at 19 30 cm intervals to 1.2 m depth from two manifolds that had both above and below surface sampling inlets. 20 Despite being in the clean air sector, over the course of the 1.2-year study, we observed on the order of 21 50 occasions when exhaust plumes from the camp, most notably from the power generation system, were 22 transported to the study site. Continuous monitoring of nitrogen oxides (NO_x) provided a measurement 23 of a chemical tracer for exhaust plumes. Highly elevated levels of NO_x (up to 1000 x background) and 24 lowered ozone (down to ≈50%), most likely from titration with nitric oxide, were measured in air from 25 above and within the snowpack. Within 5-15 minutes from observing elevated pollutant levels above the 26 snow, rapidly increasing and long-lasting concentration enhancements were measured in snowpack air. 27 While pollution events typically lasted only a few minutes to an hour above the snow surface, elevated 28 NO_x levels were observed in the snowpack lasting from a few days to \approx one week. GEM and formaldehyde 29 measurements were less sensitive and covered a shorter measurement period; neither of these species' 30 data showed noticeable concentration changes during these events that were above the normal variability 31 seen in the data. Nonetheless, the clarity of the NO_x and ozone observations add important new insight 32 to the discussion if and how snow-photochemical experiments within reach of the power grid of polar 33 research sites are possibly compromised by the snowpack being chemically influenced (contaminated) by 34 gaseous and particulate emissions from the research camp activities. This question is critical for 35 evaluating if snowpack trace chemical measurements from within the camp boundaries are 36 representative for the vast polar ice sheets.

37 Introduction

Research conducted during the past ≈15 years has revealed an active and remarkable spatial diversity of 38 39 atmospheric oxidation chemistry in the polar lower atmosphere [Grannas et al., 2007]. Ozone plays a 40 fundamental role in controlling the lifetime of many atmospheric trace gases directly and indirectly by 41 modulating atmospheric OH. Unlike the episodic ozone depletion events observed at coastal sites, the 42 opposite effect (i.e. ozone production) has been observed in the Antarctic interior [Crawford et al., 2001; 43 Helmig et al., 2007b; Helmig et al., 2008a; Legrand et al., 2009; Legrand et al., 2016]. The discovery of 44 ozone production chemistry in the remote and pristine Antarctic environment was rather surprising, 45 because hitherto photochemical production in the lower atmosphere had exclusively been associated 46 with polluted urban environments [Molina and Molina, 2004]. Photochemical production and snowpack 47 emissions of nitric oxides (NO_x) have been identified as underlying processes driving this chemistry. NO_x 48 has been shown to be formed from photochemical reactions in the snowpack [Honrath et al., 1999; Jones 49 et al., 2000], with deposited nitrate constituting the reservoir of this chemistry. NOx play a crucial role in 50 snow photochemical reactivity [Murray et al., 2015]. NO_x mixing ratios in interstitial air resulting from 51 photochemical reactions can exceed those in the air above the snowpack by a factor of \approx 50 [Van Dam et 52 al., 2015].

53 This concentration gradient is driving NO_x emission fluxes out of the snowpack into the overlying 54 atmosphere [Jones et al., 2001; Honrath et al., 2002], which, under stable atmospheric conditions, can 55 cause large NO_x enhancements in the atmospheric surface layer [Helmig et al., 2008b; Neff et al., 2008; 56 Frey et al., 2011; Frey et al., 2013], and in the presence of solar irradiance trigger photochemical ozone 57 production, with resulting peak ozone levels that can be double those in the boundary layer [Crawford et 58 al., 2001; Helmig et al., 2008a; Legrand et al., 2016]. Experiments on reactive nitrogen chemistry 59 investigating this rather unexpected ozone production chemistry have built on a variety of atmospheric 60 research strategies, including snowpack air sampling [Dibb et al., 2002; Jacobi et al., 2004; Helmig et al., 61 2007a; Van Dam et al., 2015], snow chambers [Dibb et al., 2002], snow chemical analyses [Dassau et al., 62 2002; Dibb et al., 2007b; France et al., 2011; Erbland et al., 2013], atmospheric monitoring [Frey et al., 63 2011; Kramer et al., 2015; Legrand et al., 2016], surface fluxes [Jones et al., 2001; Honrath et al., 2002; 64 Frey et al., 2011; Frey et al., 2015], and boundary layer vertical profiling [Helmig et al., 2008a; Frey et al., 65 2015].

Most of these studies rely on observations from dedicated campaigns at research stations, including 66 67 photochemistry campaigns at Summit, Greenland [Dibb et al., 2007a], the Antarctic Tropospheric 68 Chemistry Investigation [ANTCI; Eisele and Davis, 2008] at the South Pole, the Chemistry of the Antarctic 69 Boundary Layer and the Interface with Snow (CHABLIS) experiment at Halley [Jones et al., 2008], and the 70 Oxidant Production over Antarctic Land and its Export (OPALE) campaign at Concordia Station [Preunkert 71 et al., 2012]. A common limitation of these studies is that experiments were conducted in proximity to 72 research stations, where use of fuel-powered engines in generators and vehicles cause exhaust emissions 73 with highly elevated concentrations of particulates and gases, particularly of volatile organic compounds 74 (VOCs) and NO_x. A critical question is if and how this pollution, and possibly secondary products formed 75 during the atmospheric transport and deposition, impact the snow chemical position and reactivity, and 76 potentially the findings from this aforementioned literature. This is of particular importance for oxidized 77 nitrogen species. This study yielded for the first time a year-long record of NO_x and O₃ in an Antarctic

- 78 snowpack at Concordia and the atmosphere above it. This experiment also gave us the opportunity to
- real study and evaluate occurrences of pollution episodes, using the NO_x monitoring as a sensitive chemical
- 80 tracer for identification of exhaust plumes.

81 Methods

- 82 Location: This experiment was conducted at the French/Italian Antarctic research station Concordia,
- 83 located at the Dome Circe or Dome Charlie (Dome C, 75.10°S/123.35°E3233 m asl, mean temperature –
- 84 55°C). An experimental site was established at the border of the clean air sector, approximately 1 km to
- the west of the station common buildings (Figure 1). The clean air sector is located in the opposite
- 86 direction of the prevailing wind direction. The site consisted of a 8 m x 2 m x 2.5 m underground laboratory

87



Figure 1: Satellite image of research station Concordia with location of the snow photochemistry experiment indicated by the red text and cross marker. Its location was \approx 1 km west of the station main buildings and power generation plant. The map also indicates the location where the snow pits were collected, the border of the clean air sector (which is west of the red line), and a wind rose for a full year of wind data collected from the met tower that was adjacent to the snow tower manifolds.

- 89 positioned at the border of the clean air area, a 10 m tall meteorological tower, and two snow air sampling
- 90 manifolds for sampling the atmosphere and the snow interstitial air (Figure 2). The installation was in late
- 91 November 2012 with continuous monitoring conducted until January 2014 (14 months).
- 92 <u>Meteorological Tower:</u> A 10 m meteorological tower (Figure 2a) was equipped with two sonic
- 93 anemometers for atmospheric turbulence measurements, and three gas sampling inlets (0.5 m, 2 m, 10
- 94 m) with sampling lines inside a heated conduit running to the laboratory. The upper inlet was attached



Figure 2: (a) Meteorological tower with the station infrastructure in the background. The wooden box to the left is the entry hatch to the underground laboratory. Air sampling inlets were located at 0.5 m, 2 m and 10 m above the surface on the tower. Two snowpack air sampling towers were located approximately 7 m to the left and 10 m to the right of the base of the meteorological tower. (b) One of the two snow pack air sampling manifold (snow tower), with one pair of inlets right on the snow surface, and one inlet pair at ~30 cm height. Four more equivalent sampling inlet pairs are below the snow surface at 30 cm depth intervals extending to a maximum depth of 1.2 m (see Figure 1 in *Van Dam et al.* [2015] for a schematic of a similar installation at Summit, Greenland).

95

to a manual pulley allowing the upper inlet to be lowered for side by side sampling of both inlets fortracking and correcting sampling inlet/lines biases.

98 Snowpack Air Sampling: Two identical multi-inlet snow sampling manifolds ('snow tower') for collection 99 of interstitial and ambient air were constructed, with a similar design to that described by [Seok et al., 100 2009] (Figure 2b). The snow tower consisted of a vertical post of square aluminum alloy (3.8 x 3.8 cm) 101 with eight 60 cm long cross arms spaced vertically every 20 cm. Each of the cross bars supported a pair 102 of sampling inlets. The inlets were fitted with 25 mm Acrodisc hydrophobic polytetrafluoroethylene (PTFE) 103 syringe filters (Pall Life Sciences, Ann Arbor, Michigan, USA) to prevent snow and ice crystals from being 104 pulled into the sampling line. For the installation, a snow pit was dug and the inlets were inserted 105 horizontally into the clean untouched walls of the pit. The hole was then loosely refilled with the 106 excavated snow, re-establishing the stratigraphy as much as possible. Blowing snow then refilled any 107 remaining gaps within the following 2-3 days. The snow tower was kept in place after the campaign, so 108 we have no data on the potential changes in porosity and air flow dynamics that resulted from the 109 installation and subsequent changes in the snow morphology as the snowpack re-equilibrated. Insulated 110 and heated sampling lines connected the sampling inlets to the chemical analyzers in the underground

laboratory. All sampling lines were of 0.64 cm o.d. x 30 m long pre-conditioned PFA tubing, except the 111 112 lines to the gradient inlets on the meteorological tower, which were of 0.78 cm o.d. because they were 113 continuously pumped to maintain a flow of at least 2 l min⁻¹. Air was pulled through the snow tower 114 sampling lines by the combined flow of the gas analyzers (ozone monitor at 1 l min⁻¹, a gaseous elemental 115 mercury (GEM) analyzer at 1 l min⁻¹, NO_x monitor at 1 l min⁻¹). A maximum of two monitors sampled from 116 the snow tower inlets together at a given time to limit the maximum snow air sampling. Since each line 117 connected to a pair of inlets at equal height, the effective flow through each inlet was a maximum of ≈1 l 118 min⁻¹. Each height was sampled for 10 min every 2 hours, resulting in an approximately volume of a sphere 119 with a radius of 25 cm around each inlet every 2 hours. Sampling from the two snow towers was 120 alternated every 24 hours. Each sampling inlet had a thermocouple wire attached for monitoring of the 121 snowpack temperature gradient.

122 <u>Ozone Measurements:</u> Ozone was measured with a Thermo Environmental (TEI) 49i UV absorption 123 monitor that was calibrated against a NOAA Global Monitoring Division reference standard before field 124 shipment.

125 NO_x Measurements: Nitrogen oxides were monitored with a TEI chemiluminescence analyzer (TEI 42C-126 TL). The TEI 42C-TL has two channels. The first channel measures NO via NO + O₃ chemiluminescence. 127 The second channel measures total nitrogen oxides $(NO_x = NO + NO_2)$ by redirecting air through a heated 128 (325°C) molybdenum converter, which causes NO₂—including other oxidized nitrogen compounds—to be 129 converted to NO. NO₂ is then determined by subtracting NO, obtained from the first channel, from the 130 resulting NO_x signal. There are a number of other oxidized nitrogen species that can contribute to the 131 NO₂ measurement [Steinbacher et al., 2007]. The error in the NO₂ measurement increases with rising 132 levels of interfering gases such as nitrous acid (HONO), peroxyacetyl nitrate (PAN), and alkyl nitrates that 133 contribute to the NO₂-mode signal. Consequently, NO₂ concentrations obtained with the TEI 42C-TL 134 represent an estimate for the sum of these oxidized nitrogen species. Field calibrations were conducted 135 with a NIST-traceable 1 ppm NO in N₂ gas standard (Scott-Marin, Inc., Riverside, CA, USA) that was dynamically diluted to low ppb mixing ratios. We did not bring a zero air compressed gas cylinder to 136 137 Concordia. Instead, a low NO_x dilution gas was prepared by pumping ambient air (which had significantly 138 lower NO_x levels than snowpack air) through a cartridge filled with $\approx 1 \text{ dm}^3$ of granular Chemisorbent 139 (Purafill, Doraville, GA). Calibration ranges were from 0.1-25 ppb, and the instrument response was linear 140 within this range. Intercept values of the linear regression, and zero values from sampling of the scrubbed 141 air, were below 0.1 ppb.

142 <u>Formaldehyde Measurements:</u> Formaldehyde was monitored with a commercial analyzer by liquid 143 fluorimetry. Details of the measurement, instrument characterization, and deployment at Concordia 144 Station have been provided by [*Preunkert et al.*, 2015].

145 <u>Gaseous Elemental Mercury Measurements:</u> GEM was measured with a commercial Tekran Model 2537

146 instrument. The analytical protocol, and calibration and instrument characterization details are available

- 147 in [*Angot et al.*, 2016].
- 148 <u>Snow Sampling and NO₃⁻ Determination</u>: The snow pit NO₃⁻ data stem from sampling that was done at and
- 149 near Concordia between January 2009 and December 2010, and at \approx 3 m distance from snow tower 2 in
- 150 January 2014. Snow was collected in pre cleaned 50 ml centrifuge tubes inserted directly on a newly

- 151 scraped wall of a snow pit. Nitrate concentration in snow samples was measured directly in the field, at
- 152 the wet chemistry laboratory of Concordia station. Each sample was melted at room temperatures and
- 153 NO₃⁻ concentrations were determined using a colorimetric method employed routinely at Concordia [*Frey*
- 154 et al., 2009].

155 Results and Discussion

- 156 Results of the year-round snowpack and ambient monitoring, including interpretations on photochemistry
- will be presented elsewhere [*Helmig et al.*, 2019]. Here, we primarily focus on occurrences of pollution
 transport to the sampling site and its penetration into the snowpack.
- 159 Figure 3 shows a photograph of the station main buildings. The power plant is adjacent to the two column 160 structure. Approximately 300 m³ of Special Antarctic Blend (SAB) diesel fuel are burned in the plant for 161 electricity and heat generation per year. The exhaust plume from the 5-m high stack of the power plant 162 can be seen in the picture, blowing towards the west. Due to the typical strong stratification and stability 163 of the atmosphere near the surface, the plume does not rise far above the stack height, but instead gets 164 transported horizontally at a height of \approx 5-10 m above the snow surface. This is a typical exhaust plume 165 dispersion behavior for a cold regions environment, seen at many other polar research stations. The 166 plume typically does not hit the surface within the immediate distance of the stack location. Depending 167 on the actual turbulent mixing conditions, it may take several hundred meters before the stack emissions 168 are encountered right at the surface.
- 169 Of the gases monitored in this experiment, NO_x were the most sensitive tracer for pollution impact We
- 170 chose to concentrate on NO_x as the total of $NO+NO_2$, as this is a more representative indicator for the
- total amount of oxidized nitrogen, whereas NO would only indicate a fraction. Further, the fractionation
- between NO and NO_x is sensitive to other gases (such as ozone) and residence time and snowpack depth.
- 173 Therefore, interpretation of NO data is more ambiguous. NO_x in ambient air at Concordia remained well
- below 1 ppb during background conditions year-round [Helmig et al., 2019], in agreement with
- 175 observations from prior shorter campaign NO_x measurement at Concordia



Figure 3: Photograph of Concordia Station illustrating the dispersion of the exhaust plume from the electrical power generating plant during conditions with a strong surface temperature inversion. The plume dissipated toward the west in the direction of the experimental site. This is a typical situation for a contamination event.

176

[*Frey et al.*, 2011; *Frey et al.*, 2013; *Frey et al.*, 2015]. We did not observe highly elevated NO levels in the surface layer outside of pollution events, as what has been observed at South Pole [*Helmig et al.*, 2008b; *Neff et al.*, 2008]. During pollution events, the 1 ppb threshold was exceeded very quickly in measurements taken from the meteorological tower and from above-surface snow tower inlets, with resulting NO_x mixing ratios rising to as high as close to 200 ppb, representing an up to 1000-fold enhancement over background conditions (Figure 4).

183 During the course of the ≈≈14 month study, a total of ≈50 pollution events were observed, although some 184 events overlapped (Figure 4). Most of these occurrences were relatively short, with elevated mixing ratios 185 above the snowpack lasting from minutes to a few hours. We counted 15 events in total when there was 186 a spike in the above surface NO_x measurements followed by an increase of at least 1 ppb of NO_x in 187 snowpack air. It took up to in excess of 7 days for NO_x in the snowpack air to return to pre-event levels. 188 Integrated over the entire campaign, pollution episodes constituted < 2.0 % of the measurements above 189 the snowpack, and <10% of the measurements within the snowpack. The correlation analyses of pollution 190 occurrences with wind direction clearly defines the direction of the transport. The predominant wind 191 direction sector at Concordia is southeast to northwest (Figure 5a), with southeasterly winds having the overall largest share. NOx levels were consistently well below 1 ppb when winds were from east to 192 193 northeast. The sector with pollution transport is well defined, with wind directions covering 194 approximately 45-120° (Figure 5b). These sectors perfectly line up with the upwind direction of the station 195 power plant (Figure 1), clearly identifying the plant as the source of these pollution occurrences.



Figure 4: NO_x measured from the above surface inlets on the two snow towers, and from all inlets on the meteorological tower. Circles represent 10-min averaged data. Plotted are data that were extracted from occurrences when NO_x was above 1 ppb, well above the background mixing ratio, and any time when the wind direction was from the polluted wind sector.



Figure 5: (a) Concentration wind rose with the relative frequency of NO_x mixing ratio data from the above surface inlets of the two snow towers segregated by 10° sectors for the full year of observation data. This panel shows all data. (b) The same analysis, with wind direction data binned in 20° sectors for events when NO_x in ambient air exceeded 1.2 ppb.





Figure 6. Combined measurements from the two snow towers capturing a pollution event at Concordia during the middle of the winter (Day of Year 191 = July 11). Plotted time series traces correspond to the sampling heights indicated in the legend, with positive numbers giving the height above the snow surface, and negative numbers the depth below the snow surface. The right graph (b) is an enlargement of the data shown in the left (a) with the transitions between the connecting lines between two snow tower measurements removed to show the level of agreement in the data from the two sampling manifolds. The sampling switched between the two snow towers every 24 hours, leading to some abrupt shifts in the snowpack NO_x measurements.

200 One of these elevated NO_x events is further investigated in Figure 6. Here, we show the measurements 201 from six inlets on the snow tower over a one-week period. The sampling of a polluted plume is first 202 observed in the two above surface inlets (orange/red colored data; +10 and +45 cm), by the sudden 203 increase of NO_x from well below 1 ppb to a mixing ratio of \approx 13 ppb. This spike in NO_x lasted for \approx 3 hours. 204 After that time, NO_x in air sampled above the surface dropped very quickly and equilibrated to prior mixing

205 ratios within less than 0.5 hours.

206 A much different behavior was found in the air sampled from within the snowpack, indicated by the data 207 in the blueish colors. The onset of the pollution signal is delayed, by 1-3 hours, with progressively later 208 times towards deeper in the snowpack. Maximum mixing ratios that are reached in the snowpack are 209 lower, i.e. 10-50% of those that were measured above the surface, with mixing ratio maxima becoming 210 progressively smaller with increasing depth. The most remarkable difference between the above and 211 below surface measurements is the longer residence time of the pollution signal in the snowpack. NO_x 212 mixing ratios in air withdrawn from all sampling inlets in the snowpack dropped steadily, but remained 213 elevated in comparison to levels seen before the pollution event for \approx one week. The behavior seen in the measurements from snow tower 1 were in full qualitative, and within $\approx 30\%$ quantitative agreement with 214 215 the concurrent observations from the second snow tower. After the pollution event, NO_x in the snowpack air steadily declined over several days (Figure 6). Fitting of the data to an exponential decay function 216 217 yields similar results for all snowpack depths (Figure 7), with exponential regression fit \mathbb{R}^2 results of ≥ 0.95 . 218



Figure 7. Exponential decay function fits to the NO_x snowpack measurements versus time at four depths for the event starting on Day of Year 191 shown in Figure 6. The start of the event was defined as the time when high NOx was detected above the snowpack. Solutions for the best fit exponential decay functions are given in the legend.

221 Effects of the exhaust transport were also observed in the ozone signal. The ozone record, shown in Figure 222 8, shows a plethora of short positive and negative spikes superimposed on the annual cycle. The up to 223 15-20 ppb sudden ozone increases seen during the austral summer months are attributable to the 224 photochemical ozone production events that occur in the surface layer of the Antarctic Plateau (see 225 discussion in the introduction section). Occurrences of these elevated ozone events at Concordia Station 226 have previously been investigated by Legrand et al. [2009] and Cristofanelli et al. [2018]. Besides these 227 positive ozone spikes, this annual record also shows numerous sudden negative ozone changes that can 228 be attributed to destruction of ozone by titration of NO in the exhaust plume. Up to 50% of the ambient 229 ozone was destroyed in air sampled from the above surface inlets. Similar to NO_x, this signal, albeit weaker 230 and attenuated in time, was also seen in the air sampled from within the snowpack (Figure 8).



Figure 8. Ozone measured from the snow tower inlets throughout the year. Negative spikes in the data coincide with elevated NO_x from exhaust infiltration in the snowpack.

231

232

233 Enhancements in formaldehyde in ambient samples suspected to be influenced by station emissions have 234 previously been noted in measurements taken during the OPALE campaign [Preunkert et al., 2015]. We 235 therefore investigated if pollution signatures were present in formaldehyde measurements taken in the 236 record from our study. Furthermore, we revisited the GEM measurements taken during the campaign 237 [Angot et al., 2016]. Both of these measurements did not quite have the time resolution and sensitivity as the NO_x and ozone monitoring. Formaldehyde measurements overlapped with the ozone and NO_x 238 239 monitoring only for a short period during the 2012/2013 austral summer, and this measurement period 240 suffered from a number of instrument problems. The remaining data did not allow a conclusive 241 evaluation on the behavior of formaldehyde in the snowpack during pollution events. Similarly, we did

- not identify a clear signature of GEM changes in snowpack air that correlated with the NO_x enhancements
- 243 during pollution events.
- Nitrogen oxides undergo reaction with atmospheric oxidants, primarily the OH radical (summer only) and
- ozone, yielding higher oxidized nitrogen species that can partition into the snowpack aqueous and solid
- 246 phase. The frequency, large enhancement, and long duration (in the snowpack) of NO_x pollution events
- 247 constitute an apparent unnatural source of NO_x to the snowpack. One can hypothesize that reaction of
- NO and NO₂ with OH may be a source of HNO_2 and HNO_3 in the snowpack, which would add acidity to the snow. This then also poses the question if and to what degree photochemical processes, building for
- snow. This then also poses the question if and to what degree photochemical processes, building for instance on NO_2^- or NO_3^- as a substrate, may be altered from natural conditions. However, the overall
- quantitative effect is likely relatively minor given the large overall NO₃⁻ reservoir in the solid phase of the
- 252 snowpack.).
- 253 We investigated this question by comparing NO₃⁻ results from snow pit sampling at different locations
- within the camp and at up to 25 km distance of Concordia Station. Nitrate in the snowpack shows a steep
- vertical gradient, with highest levels observed right at the surface, and progressively lower concentrations
- with increasing depth (Figure 9; and Supplement Figure S-1 for a summary graph where data from both
- 257 groups were combined and binned in 15 cm depth intervals). The results from the seven snow pits are
- 258 consistent in the depth profile;
- 259



Figure 9. Nitrate concentration in snow pits in proximity (≈ 20 m) to the snow towers at the border of the clean air sector (warm colors), and sampling locations up and downwind of Concordia (cold colors). Sampling dates are indicated in the figure legend. Horizontal error bars depict the estimated uncertainty of the chemical analysis, i.e. 10% at > 10 ng g⁻¹, 50% between 5-10 ng g⁻¹, and 100% at < 5 ng g⁻¹.

261 however, there appears to be a tendency that the four snow pits within the camp have somewhat higher, 262 as well as more variable snow NO_3^- . Deeper into the snowpack the difference between the two groups of 263 data becomes stronger. At 1 m depth, the snow age is approximately ten years, and therefore 264 approximately corresponds to the start of permanent and year-round activities in 2005 (note that summer 265 activities at Concordia Station were established in 1996). Similar vertical NO₃⁻ concentration profiles in 266 the Antarctic snowpack have previously been documented [France et al., 2011]. Further, relatively higher 267 NO_3 concentrations (approximately 5 times) are observed in the upper layers of the snowpack during the 268 Antarctic summer (Supplement Figure S-2) than during the winter. This NO₃- enhancement in surface 269 snow and the seasonal cycle have been linked to the production of HNO₃ in the photoactive summer 270 months from reaction of OH with NO₂, with the NO₃⁻ enhancement being the result of HNO₃ deposition to 271 the snow surface [*Erbland et al.*, 2013]. Elevated NO_3^{-1} concentrations observed at the surface of the 272 snowpack are a common feature at low accumulation regions, with the concentration values depending 273 strongly on the strategy for collection of the first few cm of the snowpack [Erbland et al., 2013; Shi et al., 274 2018]. With the sensitivity of surface snow NO_3^{-1} to the seasonal cycle and, the high variability of NO_3^{-1} 275 observed in surface snow and vertical profiles, the currently available data do not allow a conclusive 276 quantitative assessment of the degree to which the snowpack at Concordia is chemically impaired by 277 ventilation of the snowpack with pollution-NO_x enriched air. Contamination of the snowpack around 278 Concordia station has been noted in previous investigations. Warren et al. [2006] reported an > 3-fold 279 increase in snowpack black carbon concentration after the station was established in 2003, with black 280 carbon levels in pre-2003 snow (from deeper depths) also being in closer agreement with snow collected 281 from further distance to the station. Black carbon is one of the contributing factors for a decrease in light 282 penetration into the snowpack [Warren et al., 2006; France et al., 2011; Libois et al., 2013]. Consequently, 283 the increased presence of black carbon causes a shallower e-folding depth compared to pristine, uncontaminated snow. In addition to the experiments described above, during the 2014 campaign a 284 285 number of dynamic flow-through snow chamber photochemistry experiments were conducted to 286 investigate if there were differences in the reactive chemistry in the snow from near the snow tower site 287 compared to snow sampled 25 km away from camp. These measurements showed on the order of 10-288 20% higher NO_x and less ozone in the outflow of chambers filled with the snow from further away from 289 the camp. We did not conduct a high enough number of repeats for evaluating the repeatability and 290 statistical significance of these results to gauge if and how much of this signal was due to the experimental 291 setup or due to differences in the snowpack chemical composition. Nonetheless, these preliminary 292 findings point towards possible differences in the chemical behavior that potentially are linked to 293 differences in the snow sampling locations and contaminant levels resulting from camp influences that 294 warrant further investigation.

295

296 Summary and Conclusions

With our snowpack sampling manifold we were able to sample snowpack air to a maximum depth of 70 cm below the surface. Up to \approx 2 ppb enhancements in NO_x were observed at that depth from exhaust infiltration. While our experiment was not able to access air deeper (than 70 cm), the observed concentration gradients imply that this transport and contamination extends well beyond the depth that was probed in these measurements. This experiment was a one spot measurement, at \approx 1 km distance from the camp main facilities. We have no data that would allow us to assess to what distance from the camp the snowpack pollution from exhaust infiltration would be noticeable and of importance, but it likely extends well beyond the distance of our site.

306 Several other previous studies have noted challenges in sampling clean air at polar research sites that 307 stem from pollution caused by camp exhaust. For atmospheric sampling, this interference can be 308 mitigated by careful postprocessing/filtering of the monitoring data, or by interrupting the sampling 309 during unfavorable wind conditions, which is particularly critical for integrated aerosol collection [Wolff and Cachier, 1998; Wolff et al., 1998]. Our experiments from Concordia Station emphasize the 310 311 pronounced and longer-lasting influence that station exhaust can have on NO_x levels inside the snowpack 312 air (compared to ambient air). A tendency of potentially enhanced snowpack NO₃⁻ levels in two snow pits collected at the camp, compared to data from three sites at further distance, supports the suspicion that 313 314 the snowpack chemical oxidized nitrogen composition at the station may be compromised (i.e. 315 contaminated) from the re-occurring ventilation of the snowpack with polluted air. A similar conclusion 316 was derived from d¹⁵N nitrate analyses of snow at Summit: Samples collected in the predominant 317 downwind direction of the station generator showed an isotopic signature that had a stronger association with engine exhaust than samples collected at further distance [Fibiger et al., 2016]. Even in the Summit 318 319 clean air sector elemental carbon in snow was 1.8 - 2.4 times higher than in snow collected at 10-20 km 320 distance from the camp [Hagler et al., 2008]. The associations shown in our study argue for further 321 investigation, for instance by a high resolution spatial survey of surface snow composition within and 322 beyond camp boundaries. Given the strong seasonality of NO_3^- , this survey should be done with as close 323 as possible concurrently conducted snow sampling at selected locations to minimize the influence of 324 temporal changes on the NO₃⁻ signature.

These observations emphasize concerns about the representativeness of experimental snow chemistry data collected within a Polar research camp periphery. This raises the question how interpretations from such experiments reflect conditions in the remote Polar environment. Furthermore, our findings should motivate comparison studies with sampling along transects to further distance from the main camp facilities. Comparison of these observations will likely yield new insights for evaluating prior polar research site observations and interpretation of snow photochemistry in the glacial snowpack.

331

332 Author contribution:

333 DH oversaw the study, participated in field work, conducted data analyses and quality control, and 334 prepared the manuscript.

- 335 DL conducted data analyses, quality control, prepared figures, and contributed to the manuscript 336 preparation.
- 337 JH fabricated the instrumental equipment and data acquisition system, and participated in the field work.
- JS contributed to the study design, participated in field work, conducted data analyses and interpretation,
- and contributed to the manuscript preparation.
- 340
- 341

342 Acknowledgements:

- 343 We thank all staff and scientists who helped with the setup and maintenance of the experiment at
- Concordia Station. A. Mass, W. Vicars, and A. Barbero helped with some of the field experimental work
- and data processing. A. Barbero provided the photograph in Figure 3. The French partners acknowledge
- technical support from the C2FN (French National Center for Coring and Drilling, handled by INSU), a grant
- 347 from Labex OSUG@2020 (Investissements d'avenir ANR10 LABX56), and the program 1011 SUNITEDC
- of the Institut Polaire Paul Emile Victor (IPEV). The U.S. scientists were supported through a grant from
- 349 the National Science Foundation (NSF), PLR#1142145.

350 References

- Angot, H., Magand, O., Helmig, D., Ricaud, P., Quennehen, B., Gallee, H., Del Guasta, M., Sprovieri, F.,
- Pirrone, N., Savarino, J., and Dommergue, A.: New insights into the atmospheric mercury cycling in
- 353 central Antarctica and implications on a continental scale, *Atmospheric Chemistry and Physics*, *16*, 8249-
- 354 8264, doi:10.5194/acp-16-8249-2016, 2016.
- 355 Crawford, J. H., Davis, D. D., Chen, G., Buhr, M., Oltmans, S., Weller, R., Mauldin, L., Eisele, F., Shetter, R.,
- Lefer, B., Arimoto, R., and Hogan, A.: Evidence for photochemical production of ozone at the South Pole
- 357 surface, *Geophysical Research Letters*, *28*, 3641-3644, doi:10.1029/2001gl013055, 2001.
- 358 Cristofanelli, P., Putero, D., Bonasoni, P., Busetto, M., Calzolari, F., Camporeale, G., Grigioni, P., Lupi, A.,
- Petkov, B., Traversi, R., Udisti, R., and Vitale, V.: Analysis of multi-year near-surface ozone observations at the WMO/GAW "Concordia" station (75 degrees 06 ' S, 123 degrees 20 ' E, 3280 m a.s.l. - Antarctica),
- 361 *Atmospheric Environment*, 177, 54-63, doi:10.1016/j.atmosenv.2018.01.007, 2018.
- 362 Dassau, T. M., Sumner, A. L., Koeniger, S. L., Shepson, P. B., Yang, J., Honrath, R. E., Cullen, N. J., Steffen,
- 363 K., Jacobi, H. W., Frey, M., and Bales, R. C.: Investigation of the role of the snowpack on atmospheric
- 364 formaldehyde chemistry at Summit, Greenland, *Journal of Geophysical Research*, *107*, 2002.
- Dibb, J. E., Albert, M., Courville, Z., Anastasio, C., Galbavy, E. S., Atlas, E., Beyersdorf, A. J., Blake, D. R.,
- 366 Meinardi, S., Rowland, F. S., Swanson, A. L., Blake, N. J., Bocquet, F., Cohen, L., Helmig, D., Burkhart, J. F.,
- 367 Frey, M. M., Friel, D. K., Hutterli, M. A., Chen, G., Conway, T. J., and Oltrnans, S. J.: An overview of air-
- snow exchange at Summit, Greenland: Recent experiments and findings, *Atmospheric Environment*, *41*,
 4995-5006, 2007a.
- Dibb, J. E., Arsenault, M., Peterson, M. C., and Honrath, R. E.: Fast nitrogen oxide photochemistry in
 Summit, Greenland snow, *Atmospheric Environment*, *36*, 2501-2511, 2002.
- 372 Dibb, J. E., Whitlow, S. I., and Arsenault, M.: Seasonal variations in the soluble ion content of snow at
- 373 Summit. Greenland: Constraints from three years of daily surface snow samples, *Atmospheric*
- 374 *Environment*, *41*, 5007-5019, doi:10.1016/j.atmosenv.2006.12.010, 2007b.
- Eisele, F. L., and Davis, D. D.: Antarctic tropospheric chemistry investigation (ANTCI) 2003, *Atmospheric Environment*, 42, 2747-2748, doi:10.1016/j.atmosenv.2007.09.074, 2008.
- 377 Erbland, J., Vicars, W. C., Savarino, J., Morin, S., Frey, M. M., Frosini, D., Vince, E., and Martins, J. M. F.:
- 378 Air-snow transfer of nitrate on the East Antarctic Plateau Part 1: Isotopic evidence for a photolytically
- driven dynamic equilibrium in summer, *Atmospheric Chemistry and Physics*, *13*, 6403-6419,
- doi:10.5194/acp-13-6403-2013, 2013.
- 381 Fibiger, D. L., Dibb, J. E., Chen, D., Thomas, J. L., Burkhart, J. F., Huey, L. G., and Hastings, M. G.: Analysis
- of nitrate in the snow and atmosphere at Summit, Greenland: Chemistry and transport, *Journal of Geophysical Research-Atmospheres*, *121*, 5010-5030, doi:10.1002/2015jd024187, 2016.
- France, J. L., King, M. D., Frey, M. M., Erbland, J., Picard, G., Preunkert, S., MacArthur, A., and Savarino,
 J.: Snow optical properties at Dome C (Concordia), Antarctica; implications for snow emissions and snow
 chemistry of reactive nitrogen, *Atmospheric Chemistry and Physics*, *11*, 9787-9801, doi:10.5194/acp-119787-2011, 2011.
- 388 Frey, M. M., Brough, N., France, J., King, M. D., Savarino, J., Anderson, P., Jones, A., and E., W.:
- 389 Atmospheric nitrogen oxides (NO and NO2) at Dome C: first observations & implications for reactive
- 390 nitrogen cycling above the East Antarctic Ice Sheet, , Presented at Air-Ice Chemistry Interactions
- 391 workshop, June 6-7, Columbia University, New York, USA, 2011.

- 392 Frey, M. M., Brough, N., France, J. L., Anderson, P. S., Traulle, O., King, M. D., Jones, A. E., Wolff, E. W.,
- and Savarino, J.: The diurnal variability of atmospheric nitrogen oxides (NO and NO2) above the
- Antarctic Plateau driven by atmospheric stability and snow emissions, *Atmospheric Chemistry and*
- 395 *Physics*, *13*, 3045-3062, doi:10.5194/acp-13-3045-2013, 2013.
- 396 Frey, M. M., Roscoe, H. K., Kukui, A., Savarino, J., France, J. L., King, M. D., Legrand, M., and Preunkert,
- 397 S.: Atmospheric nitrogen oxides (NO and NO2) at Dome C, East Antarctica, during the OPALE campaign,
- 398 *Atmospheric Chemistry and Physics*, *15*, 7859-7875, doi:10.5194/acp-15-7859-2015, 2015.
- 399 Frey, M. M., Savarino, J., Morin, S., Erbland, J., and Martins, J. M. F.: Photolysis imprint in the nitrate
- stable isotope signal in snow and atmosphere of East Antarctica and implications for reactive nitrogen
 cycling, *Atmospheric Chemistry and Physics*, *9*, 8681-8696, doi:10.5194/acp-9-8681-2009, 2009.
- 402 Grannas, A. M., Jones, A. E., Dibb, J., Ammann, M., Anastasio, C., Beine, H. J., Bergin, M., Bottenheim, J.,
- 403 Boxe, C. S., Carver, G., Chen, G., Crawford, J. H., Domine, F., Frey, M. M., Guzman, M. I., Heard, D. E.,
- Helmig, D., Hoffmann, M. R., Honrath, R. E., Huey, L. G., Hutterli, M., Jacobi, H. W., Klan, P., Lefer, B.,
- 405 McConnell, J., Plane, J., Sander, R., Savarino, J., Shepson, P. B., Simpson, W. R., Sodeau, J. R., von Glasow,
- 406 R., Weller, R., Wolff, E. W., and Zhu, T.: An overview of snow photochemistry: evidence, mechanisms
- 407 and impacts, *Atmospheric Chemistry and Physics*, 7, 4329-4373, 2007.
- 408 Hagler, G. S. W., Bergin, M. H., Smith, E. A., Town, M., and Dibb, J. E.: Local anthropogenic impact on
- 409 particulate elemental carbon concentrations at Summit, Greenland, *Atmospheric Chemistry and Physics*,
- 410 *8*, 2485-2491, doi:10.5194/acp-8-2485-2008, 2008.
- Helmig, D., Bocquet, F., Cohen, L., and Oltmans, S. J.: Ozone uptake to the polar snowpack at Summit,
 Greenland, *Atmospheric Environment*, *41*, 5061-5076, doi:10.1016/j.atmosenv.2006.06.064, 2007a.
- Helmig, D., Hueber, J., Liptzin, D., and Savarino, J.: Photo-reactive gases in the snowpack, atmosphere,
 and their surface exchanges at Dome C, Antarctica, J. Geophys. Res., Manuscript in preparation., 2019.
- 415 Helmig, D., Johnson, B., Oltmans, S. J., Neff, W., Eisele, F., and Davis, D. D.: Elevated ozone in the
- 416 boundary layer at South Pole, *Atmospheric Environment*, *42*, 2788-2803,
- 417 doi:10.1016/j.atmosenv.2006.12.032, 2008a.
- Helmig, D., Johnson, B. J., Warshawsky, M., Morse, T., Neff, W. D., Eisele, F., and Davis, D. D.: Nitric oxide
- 419 in the boundary-layer at South Pole during the Antarctic Tropospheric Chemistry Investigation (ANTCI),
 420 *Atmospheric Environment*, 42, 2817-2830, 2008b.
- 421 Helmig, D., Oltmans, S. J., Carlson, D., Lamarque, J. F., Jones, A., Labuschagne, C., Anlauf, K., and Hayden,
- 422 K.: A review of surface ozone in the polar regions, *Atmospheric Environment*, *41*, 5138-5161,
- 423 doi:10.1016/j.atmosenv.2006.09.053, 2007b.
- Honrath, R. E., Lu, Y., Peterson, M. C., Dibb, J. E., Arsenault, M. A., Cullen, N. J., and Steffen, K.: Vertical
 fluxes of NOx, HONO, and HNO3 above the snowpack at Summit, Greenland, *Atmospheric Environment*,
 36, 2629-2640, 2002.
- Honrath, R. E., Peterson, M. C., Guo, S., Dibb, J. E., Shepson, P. B., and Campbell, B.: Evidence of NOx
 production within or upon ice particles in the Greenland snowpack, *Geophysical Research Letters*, *26*,
 695-698, 1999.
- 430 Jacobi, H. W., Bales, R. C., Honrath, R. E., Peterson, M. C., Dibb, J. E., Swanson, A. L., and Albert, M. R.:
- 431 Reactive trace gases measured in the interstitial air of surface snow at Summit, Greenland, Atmospheric
- 432 Environment, 38, 1687-1697, doi:10.1016/j.atmosenv.2004.01.004, 2004.

- 433 Jones, A. E., Weller, R., Anderson, P. S., Jacobi, H. W., Wolff, E. W., Schrems, O., and Miller, H.:
- 434 Measurements of NOx emissions from the Antarctic snowpack, *Geophysical Research Letters*, 28, 1499-435 1502, 2001.
- 436 Jones, A. E., Weller, R., Wolff, E. W., and Jacobi, H. W.: Speciation and rate of photochemical NO and
- 437 NO2 production in Antarctic snow, *Geophysical Research Letters*, 27, 345-348, 2000.
- 438 Jones, A. E., Wolff, E. W., Salmon, R. A., Bauguitte, S. J. B., Roscoe, H. K., Anderson, P. S., Ames, D.,
- 439 Clemitshaw, K. C., Fleming, Z. L., Bloss, W. J., Heard, D. E., Lee, J. D., Read, K. A., Hamer, P., Shallcross, D.
- 440 E., Jackson, A. V., Walker, S. L., Lewis, A. C., Mills, G. P., Plane, J. M. C., Saiz-Lopez, A., Sturges, W. T., and
- Worton, D. R.: Chemistry of the Antarctic Boundary Layer and the Interface with Snow: an overview of
 the CHABLIS campaign, *Atmospheric Chemistry and Physics*, *8*, 3789-3803, doi:10.5194/acp-8-3789-
- 443 2008, 2008.
- Kramer, L. J., Helmig, D., Burkhart, J. F., Stohl, A., Oltmans, S., and Honrath, R. E.: Seasonal variability of
- 445 atmospheric nitrogen oxides and non-methane hydrocarbons at the GEOSummit station, Greenland,
 446 *Atmos. Chem. Phys.*, *15*, 6827-6849, doi:10.5194/acp-15-6827-2015, 2015.
- Legrand, M., Preunkert, S., Jourdain, B., Gallee, H., Goutail, F., Weller, R., and Savarino, J.: Year-round
 record of surface ozone at coastal (Dumont d'Urville) and inland (Concordia) sites in East Antarctica, *Journal of Geophysical Research*, *114*, 1-12, doi:10.1029/2008jd011667, 2009.
- 450 Legrand, M., Preunkert, S., Savarino, J., Frey, M. M., Kukui, A., Helmig, D., Jourdain, B., Jones, A. E.,
- 451 Weller, R., Brough, N., and Gallée, H.: Inter-annual variability of surface ozone at coastal (Dumont
- d'Urville, 2004–2014) and inland (Concordia, 2007–2014) sites in East Antarctica, Atmos. Chem. Phys.,
- 453 *16*, 8053-8069, doi:10.5194/acp-16-8053-2016, 2016.
- Libois, Q., Picard, G., France, J. L., Arnaud, L., Dumont, M., Carmagnola, C. M., and King, M. D.: Influence of grain shape on light penetration in snow, *Cryosphere*, *7*, 1803-1818, doi:10.5194/tc-7-1803-2013, 2013.
- 457 Molina, M. J., and Molina, L. T.: Megacities and atmospheric pollution, *Journal of the Air & Waste* 458 *Management Association, 54*, 644-680, 2004.
- 459 Murray, K. A., Kramer, L. J., Doskey, P. V., Ganzeveld, L., Seok, B., Van Dam, B., and Helmig, D.: Dynamics
- of ozone and nitrogen oxides at Summit, Greenland. II. Simulating snowpack chemistry during a spring
- 461 high ozone event with a 1-D process-scale model, *Atmospheric Environment*, *117*, 110-123,
- 462 doi:10.1016/j.atmosenv.2015.07.004, 2015.
- 463 Neff, W., Helmig, D., Grachev, A., and Davis, D.: A study of boundary layer behavior associated with high
- 464 NO concentrations at the South Pole using a minisodar, tethered balloon, and sonic anemometer,
 465 Atmospheric Environment, 42, 2762-2779, 2008.
- Preunkert, S., Ancellet, G., Legrand, M., Kukui, A., Kerbrat, M., Sarda-Esteve, R., Gros, V., and Jourdain,
 B.: Oxidant Production over Antarctic Land and its Export (OPALE) project: An overview of the 2010-2011
 auromatic comparison for the computing Research, 117, doi:10.1020/2011id017145_2012
- summer campaign, *Journal of Geophysical Research*, *117*, doi:10.1029/2011jd017145, 2012.
- 469 Preunkert, S., Legrand, M., Frey, M. M., Kukui, A., Savarino, J., Gallee, H., King, M., Jourdain, B., Vicars,
- 470 W., and Helmig, D.: Formaldehyde (HCHO) in air, snow, and interstitial air at Concordia (East Antarctic
- 471 Plateau) in summer, *Atmospheric Chemistry and Physics*, *15*, 6689-6705, doi:10.5194/acp-15-6689-2015,
- 472 2015.
- 473 Seok, B., Helmig, D., Williams, M. W., Liptzin, D., Chowanski, K., and Hueber, J.: An automated system for
- 474 continuous measurements of trace gas fluxes through snow: an evaluation of the gas diffusion method

- 475 at a subalpine forest site, Niwot Ridge, Colorado, *Biogeochemistry*, *95*, 95-113, doi:10.1007/s10533-009476 9302-3, 2009.
- 477 Shi, G. T., Hastings, M. G., Yu, J. H., Ma, T. M., Hu, Z. Y., An, C. L., Li, C. J., Ma, H. M., Jiang, S., and Li, Y. S.:
- 478 Nitrate deposition and preservation in the snowpack along a traverse from coast to the ice sheet summit
- 479 (Dome A) in East Antarctica, *Cryosphere*, *12*, 1177-1194, doi:10.5194/tc-12-1177-2018, 2018.
- 480 Steinbacher, M., Zellweger, C., Schwarzenbach, B., Bugmann, S., Buchmann, B., Ordonez, C., Prevot, A. S.
- 481 H., and Hueglin, C.: Nitrogen oxide measurements at rural sites in Switzerland: Bias of conventional
- 482 measurement techniques, *Journal of Geophysical Research*, *112*, doi:10.1029/2006jd007971, 2007.
- 483 Van Dam, B., Helmig, D., Toro, C., Doskey, P., Kramer, L., Murray, K., Ganzeveld, L., and Seok, B.:
- 484 Dynamics of ozone and nitrogen oxides at Summit, Greenland: I. Multi-year observations in the 485 snowpack, *Atmospheric Environment*, *123, Part A*, 268-284,
- 486 doi:http://doi.org/10.1016/j.atmosenv.2015.09.060, 2015.
- 487 Warren, S. G., Brandt, R. E., and Grenfell, T. C.: Visible and near-ultraviolet absorption spectrum of ice
- 488 from transmission of solar radiation into snow, *Applied Optics*, 45, 5320-5334,
- 489 doi:10.1364/ao.45.005320, 2006.
- 490 Wolff, E. W., and Cachier, H.: Concentrations and seasonal cycle of black carbon in aerosol at a coastal
- 491 Antarctic station, Journal of Geophysical Research-Atmospheres, 103, 11033-11041,
- 492 doi:10.1029/97jd01363, 1998.
- 493 Wolff, E. W., Legrand, M. R., and Wagenbach, D.: Coastal Antarctic aerosol and snowfall chemistry,
- 494 *Journal of Geophysical Research-Atmospheres, 103,* 10927-10934, doi:10.1029/97jd03454, 1998.