

# Intrusion, retention, and snowpack chemical effects from exhaust emissions at Concordia Station, Antarctica

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## Abstract

The chemistry of reactive gases inside the snowpack and in the lower atmosphere was investigated at Concordia Station (Dome C), Antarctica, from December 2012 - January 2014. Measured species included ozone, nitrogen oxides, gaseous elemental mercury, and formaldehyde, for study of photochemical reactions, surface exchange, and the seasonal cycles and atmospheric chemistry of these gases. The experiment was installed  $\approx 1$  km from the station main infrastructure inside the station clean air sector and within the station electrical power grid boundary. Ambient air was sampled continuously from inlets mounted above the surface on a 10 m meteorological tower. In addition, snowpack air was collected at 30 cm intervals to 1.2 m depth from two manifolds that had both above and below surface sampling inlets. Despite being in the clean air sector, over the course of the 1.2-year study, we observed on the order of 50 occasions when exhaust plumes from the camp, most notably from the power generation system, were transported to the study site. Continuous monitoring of nitrogen oxides ( $\text{NO}_x$ ) provided a measurement of a chemical tracer for exhaust plumes. Highly elevated levels of  $\text{NO}_x$  (up to 1000 x background) and lowered ozone (down to  $\approx 50\%$ ), most likely from titration with nitric oxide, were measured in air from above and within the snowpack. Within 5-15 minutes from observing elevated pollutant levels above the snow, rapidly increasing and long-lasting concentration enhancements were measured in snowpack air. While pollution events typically lasted only a few minutes to an hour above the snow surface, elevated  $\text{NO}_x$  levels were observed in the snowpack lasting from a few days to  $\approx$  one week. These observations add important new insight to the discussion if and how snow-photochemical experiments within reach of the power grid of polar research sites are possibly compromised by the snowpack being chemically influenced (contaminated) by gaseous and particulate emissions from the research camp activities. This question is critical for evaluating if snowpack trace chemical measurements from within the camp boundaries are representative for the vast polar ice sheets.

## 34 Introduction

35 Research conducted during the past  $\approx 15$  years has revealed an active and remarkable spatial diversity of  
36 atmospheric oxidation chemistry in the polar lower atmosphere [Grannas *et al.*, 2007]. Ozone plays a  
37 fundamental role in controlling the lifetime of many atmospheric trace gases directly and indirectly by  
38 modulating atmospheric OH. Unlike the episodic ozone depletion events observed at coastal sites, the  
39 opposite effect (i.e. ozone production) has been observed in the Antarctic interior [Crawford *et al.*, 2001;  
40 Helmig *et al.*, 2007b; Helmig *et al.*, 2008a; Legrand *et al.*, 2009; Legrand *et al.*, 2016]. The discovery of  
41 ozone production chemistry in the remote and pristine Antarctic environment was rather surprising,  
42 because hitherto photochemical production in the lower atmosphere had exclusively been associated  
43 with polluted urban environments [Molina and Molina, 2004]. Photochemical production and snowpack  
44 emissions of nitric oxides ( $\text{NO}_x$ ) have been identified as underlying processes driving this chemistry.  $\text{NO}_x$   
45 has been shown to be formed from photochemical reactions in the snowpack [Honrath *et al.*, 1999; Jones  
46 *et al.*, 2000], with deposited nitrate constituting the reservoir of this chemistry.  $\text{NO}_x$  play a crucial role in  
47 snow photochemical reactivity [Murray *et al.*, 2015].  $\text{NO}_x$  mixing ratios in interstitial air resulting from  
48 photochemical reactions can exceed those in the air above the snowpack by a factor of  $\approx 50$  [Van Dam *et al.*,  
49 *et al.*, 2015].

50 This concentration gradient is driving  $\text{NO}_x$  emission fluxes out of the snowpack into the overlying  
51 atmosphere [Jones *et al.*, 2001; Honrath *et al.*, 2002], which, under stable atmospheric conditions, can  
52 cause large  $\text{NO}_x$  enhancements in the atmospheric surface layer [Helmig *et al.*, 2008b; Neff *et al.*, 2008;  
53 Frey *et al.*, 2011; Frey *et al.*, 2013], and in the presence of solar irradiance trigger photochemical ozone  
54 production, with resulting peak ozone levels that can be double those in the boundary layer [Crawford *et al.*  
55 *et al.*, 2001; Helmig *et al.*, 2008a; Legrand *et al.*, 2016]. Experiments on reactive nitrogen chemistry  
56 investigating this rather unexpected ozone production chemistry have built on a variety of atmospheric  
57 research strategies, including snowpack air sampling [Dibb *et al.*, 2002; Jacobi *et al.*, 2004; Helmig *et al.*,  
58 2007a; Van Dam *et al.*, 2015], snow chambers [Dibb *et al.*, 2002], snow chemical analyses [Dassau *et al.*,  
59 2002; Dibb *et al.*, 2007b; France *et al.*, 2011; Erbland *et al.*, 2013], atmospheric monitoring [Frey *et al.*,  
60 2011; Kramer *et al.*, 2015; Legrand *et al.*, 2016], surface fluxes [Jones *et al.*, 2001; Honrath *et al.*, 2002;  
61 Frey *et al.*, 2011; Frey *et al.*, 2015], and boundary layer vertical profiling [Helmig *et al.*, 2008a; Frey *et al.*,  
62 2015].

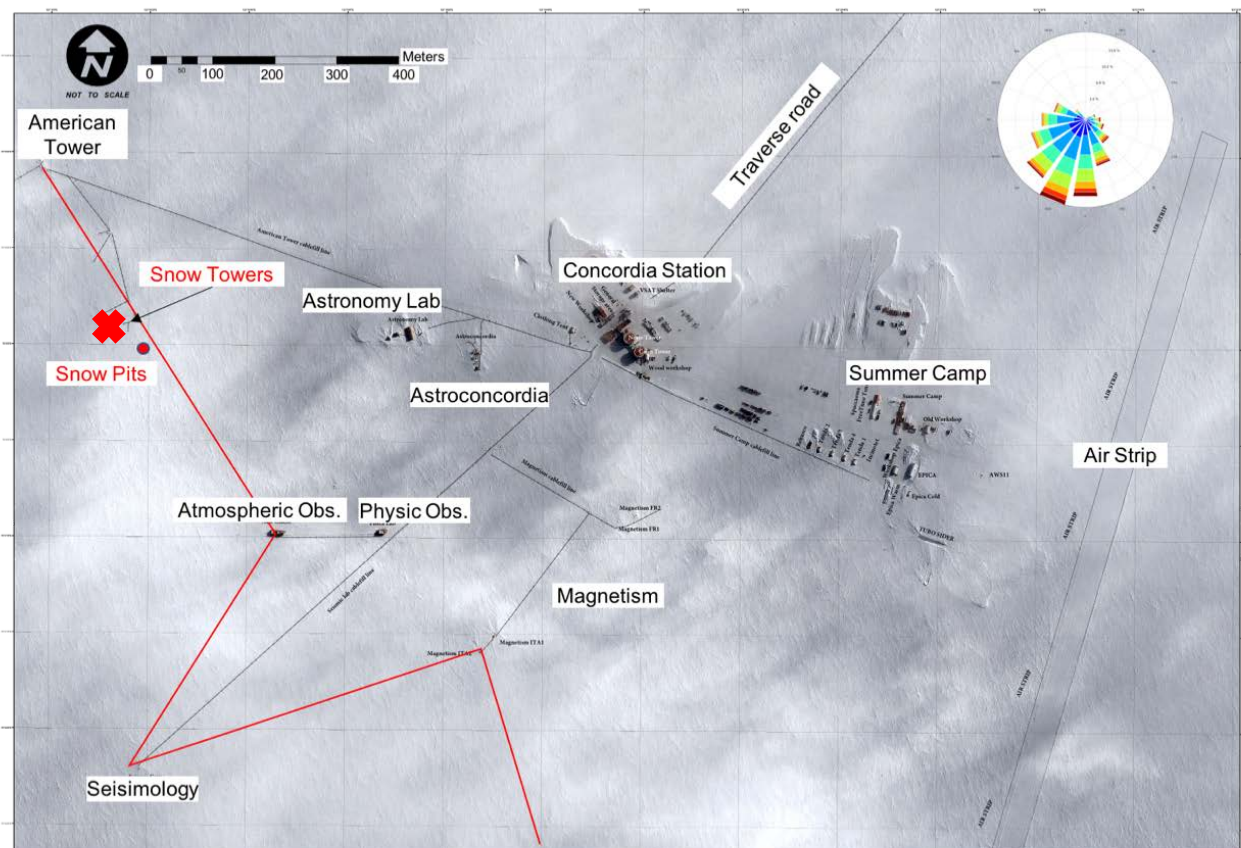
63 Most of these studies rely on observations from dedicated campaigns at research stations, including  
64 photochemistry campaigns at Summit, Greenland [Dibb *et al.*, 2007a], the Antarctic Tropospheric  
65 Chemistry Investigation [ANTCI; Eisele and Davis, 2008] at the South Pole, the Chemistry of the Antarctic  
66 Boundary Layer and the Interface with Snow (CHABLIS) experiment at Halley [Jones *et al.*, 2008], and the  
67 Oxidant Production over Antarctic Land and its Export (OPALE) campaign at Concordia Station [Preunkert  
68 *et al.*, 2012]. A common limitation of these studies is that experiments were conducted in proximity to  
69 research stations, where use of fuel-powered engines in generators and vehicles cause exhaust emissions  
70 with highly elevated concentrations of particulates and gases, particularly of volatile organic compounds  
71 (VOCs) and  $\text{NO}_x$ . A critical question is if and how this pollution, and possibly secondary products formed  
72 during the atmospheric transport and deposition, impact the snow chemical position and reactivity, and  
73 potentially the findings from this aforementioned literature. This is of particular importance for oxidized  
74 nitrogen species. This study yielded for the first time a year-long record of  $\text{NO}_x$  and  $\text{O}_3$  in an Antarctic

75 snowpack at Concordia and the atmosphere above it. This experiment also gave us the opportunity to  
76 study and evaluate occurrences of pollution episodes, using the NO<sub>x</sub> monitoring as a sensitive chemical  
77 tracer for identification of exhaust plumes.

## 78 **Methods**

79 Location: This experiment was conducted at the French/Italian Antarctic research station Concordia,  
80 located at the Dome Circe or Dome Charlie (Dome C, 75.10°S/123.35°E 3233 m asl, mean temperature –  
81 55°C). An experimental site was established at the border of the clean air sector, approximately 1 km to  
82 the west of the station common buildings (Figure 1). The clean air sector is located in the opposite  
83 direction of the prevailing wind direction. The site consisted of a 8 m x 2 m x 2.5 m underground laboratory

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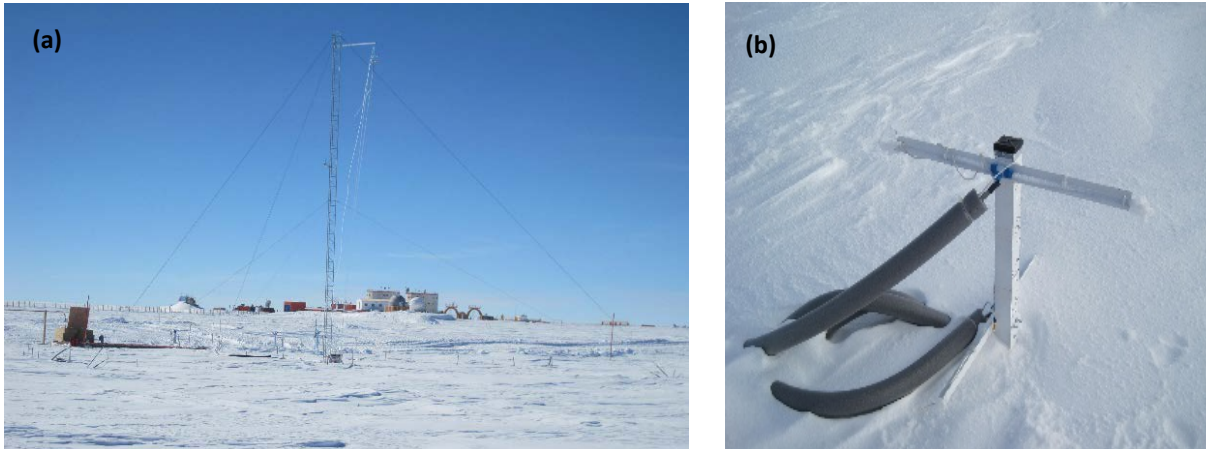


**Figure 1:** Satellite image of research station Concordia with location of the snow photochemistry experiment indicated by the red text and cross marker. Its location was ~1 km west of the station main buildings and power generation plant. The map also indicates the location where the snow pits were collected, the border of the clean air sector (which is west of the red line), and a wind rose for a full year of wind data collected from the met tower that was adjacent to the snow tower manifolds.

85

86 positioned at the border of the clean air area, a 10 m tall meteorological tower, and two snow air sampling  
87 manifolds for sampling the atmosphere and the snow interstitial air (Figure 2). The installation was in late  
88 November 2012 with continuous monitoring conducted until January 2014 (14 months).

89 Meteorological Tower: A 10 m meteorological tower (Figure 2a) was equipped with two sonic  
90 anemometers for atmospheric turbulence measurements, and three gas sampling inlets (0.5 m, 2 m, 10  
91 m) with sampling lines inside a heated conduit running to the laboratory. The upper inlet was attached



**Figure 2:** (a) Meteorological tower with the station infrastructure in the background. The wooden box to the left is the entry hatch to the underground laboratory. Air sampling inlets were located at 0.5 m, 2 m and 10 m above the surface on the tower. Two snowpack air sampling towers were located approximately 7 m to the left and 10 m to the right of the base of the meteorological tower. (b) One of the two snow pack air sampling manifold (snow tower), with one pair of inlets right on the snow surface, and one inlet pair at ~30 cm height. Four more equivalent sampling inlet pairs are below the snow surface at 30 cm depth intervals extending to a maximum depth of 1.2 m (see Figure 1 in *Van Dam et al.* [2015] for a schematic of a similar installation at Summit, Greenland).

92  
93 to a manual pulley allowing the upper inlet to be lowered for side by side sampling of both inlets for  
94 tracking and correcting sampling inlet/lines biases.

95 Snowpack Air Sampling: Two identical multi-inlet snow sampling manifolds ('snow tower') for collection  
96 of interstitial and ambient air were constructed, with a similar design to that described by [*Seok et al.*,  
97 2009] (Figure 2b). The snow tower consisted of a vertical post of square aluminum alloy (3.8 x 3.8 cm)  
98 with eight 60 cm long cross arms spaced vertically every 20 cm. Each of the cross bars supported a pair  
99 of sampling inlets. The inlets were fitted with 25 mm Acrodisc hydrophobic polytetrafluoroethylene (PTFE)  
100 syringe filters (Pall Life Sciences, Ann Arbor, Michigan, USA) to prevent snow and ice crystals from being  
101 pulled into the sampling line. For the installation, a snow pit was dug and the inlets were inserted  
102 horizontally into the clean untouched walls of the pit. The hole was then loosely refilled with the  
103 excavated snow, re-establishing the stratigraphy as much as possible. Blowing snow then refilled any  
104 remaining gaps within the following 2-3 days. The snow tower was kept in place after the campaign, so  
105 we have no data on the potential changes in porosity and air flow dynamics that resulted from the  
106 installation and subsequent changes in the snow morphology as the snowpack re-equilibrated. Insulated  
107 and heated sampling lines connected the sampling inlets to the chemical analyzers in the underground

108 laboratory. All sampling lines were of 0.64 cm o.d. x 30 m long pre-conditioned PFA tubing, except the  
109 lines to the gradient inlets on the meteorological tower, which were of 0.78 cm o.d. because they were  
110 continuously pumped to maintain a flow of at least 2 l min<sup>-1</sup>. Air was pulled through the snow tower  
111 sampling lines by the combined flow of the gas analyzers (ozone monitor at 1 l min<sup>-1</sup>, a gaseous elemental  
112 mercury (GEM) analyzer at 1 l min<sup>-1</sup>, NO<sub>x</sub> monitor at 1 l min<sup>-1</sup>). A maximum of two monitors sampled from  
113 the snow tower inlets together at a given time to limit the maximum snow air sampling. Since each line  
114 connected to a pair of inlets at equal height, the effective flow through each inlet was a maximum of ≈ 1  
115 l min<sup>-1</sup>. Each height was sampled for 10 min every 2 hours, resulting in an approximately volume of a  
116 sphere with a radius of 25 cm around each inlet every 2 hours. Sampling from the two snow towers was  
117 alternated every 24 hours. Each sampling inlet had a thermocouple wire attached for monitoring of the  
118 snowpack temperature gradient.

119 Ozone Measurements: Ozone was measured with a Thermo Environmental (TEI) 49i UV absorption  
120 monitor that was calibrated against a NOAA Global Monitoring Division reference standard before field  
121 shipment.

122 NO<sub>x</sub> Measurements: Nitrogen oxides were monitored with a TEI chemiluminescence analyzer (TEI 42C-  
123 TL). The TEI 42C-TL has two channels. The first channel measures NO via NO + O<sub>3</sub> chemiluminescence.  
124 The second channel measures total nitrogen oxides (NO<sub>x</sub> = NO + NO<sub>2</sub>) by redirecting air through a heated  
125 (325°C) molybdenum converter, which causes NO<sub>2</sub>—including other oxidized nitrogen compounds—to be  
126 converted to NO. NO<sub>2</sub> is then determined by subtracting NO, obtained from the first channel, from the  
127 resulting NO<sub>x</sub> signal. There are a number of other oxidized nitrogen species that can contribute to the  
128 NO<sub>2</sub> measurement [Steinbacher *et al.*, 2007]. The error in the NO<sub>2</sub> measurement increases with rising  
129 levels of interfering gases such as nitrous acid (HONO), peroxyacetyl nitrate (PAN), and alkyl nitrates that  
130 contribute to the NO<sub>2</sub>-mode signal. Consequently, NO<sub>2</sub> concentrations obtained with the TEI 42C-TL  
131 represent an estimate for the sum of these oxidized nitrogen species. Field calibrations were conducted  
132 with a NIST-traceable 1 ppm NO in N<sub>2</sub> gas standard (Scott-Marin, Inc., Riverside, CA, USA) that was  
133 dynamically diluted to low ppb mixing ratios. We did not bring a zero air compressed gas cylinder to  
134 Concordia. Instead, a low NO<sub>x</sub> dilution gas was prepared by pumping ambient air (which had significantly  
135 lower NO<sub>x</sub> levels than snowpack air) through a cartridge filled with ≈ 1 dm<sup>3</sup> of granular Chemisorbent  
136 (Purafill, Doraville, GA). Calibration ranges were from 0.1-25 ppb, and the instrument response was linear  
137 within this range. Intercept values of the linear regression, and zero values from sampling of the scrubbed  
138 air, were below 0.1 ppb.

139 Snow Sampling and NO<sub>3</sub><sup>-</sup> Determination:

140 The snow pit NO<sub>3</sub><sup>-</sup> data stem from sampling that was done at and near Concordia between January 2009  
141 and December 2010, and at ≈3 m distance from snow tower 2 in January 2014. Snow was collected in pre  
142 cleaned 50 ml centrifuge tubes inserted directly on a newly scraped wall of a snow pit. Nitrate  
143 concentration in snow samples was measured directly in the field, at the wet chemistry laboratory of  
144 Concordia station. Each sample was melted at room temperatures and NO<sub>3</sub><sup>-</sup> concentrations were  
145 determined using a colorimetric method employed routinely at Concordia [Frey *et al.*, 2009].

## 146 **Results and Discussion**

147 Results of the year-round snowpack and ambient monitoring, including interpretations on photochemistry  
148 will be presented elsewhere [Helmig *et al.*, 2019]. Here, we primarily focus on occurrences of pollution  
149 transport to the sampling site and its penetration into the snowpack.

150 **Figure 3** shows a photograph of the station main buildings. The power plant is adjacent to the two column  
151 structure. Approximately 300 m<sup>3</sup> of Special Antarctic Blend (SAB) diesel fuel are burned in the plant for  
152 electricity and heat generation per year. The exhaust plume from the 5-m high stack of the power plant  
153 can be seen in the picture, blowing towards the west. Due to the typical strong stratification and stability  
154 of the atmosphere near the surface, the plume does not rise far above the stack height, but instead gets  
155 transported horizontally at a height of  $\approx$  5-10 m above the snow surface. This is a typical exhaust plume  
156 dispersion behavior for a cold regions environment, seen at many other polar research stations. The  
157 plume typically does not hit the surface within the immediate distance of the stack location. Depending  
158 on the actual turbulent mixing conditions, it may take several hundred meters before the stack emissions  
159 are encountered right at the surface.

160 Of the gases monitored in this experiment, NO<sub>x</sub> were the most sensitive tracer for pollution impact We  
161 chose to concentrate on NO<sub>x</sub> as the total of NO+NO<sub>2</sub>, as this is a more representative indicator for the  
162 total amount of oxidized nitrogen, whereas NO would only indicate a fraction. Further, the fractionation  
163 between NO and NO<sub>x</sub> is sensitive to other gases (such as ozone) and residence time and snowpack depth.  
164 Therefore, interpretation of NO data is more ambiguous. NO<sub>x</sub> in ambient air at Concordia remained well  
165 below 1 ppb during background conditions year-round [Helmig *et al.*, 2019], in agreement with  
166 observations from prior shorter campaign NO<sub>x</sub> measurement at Concordia



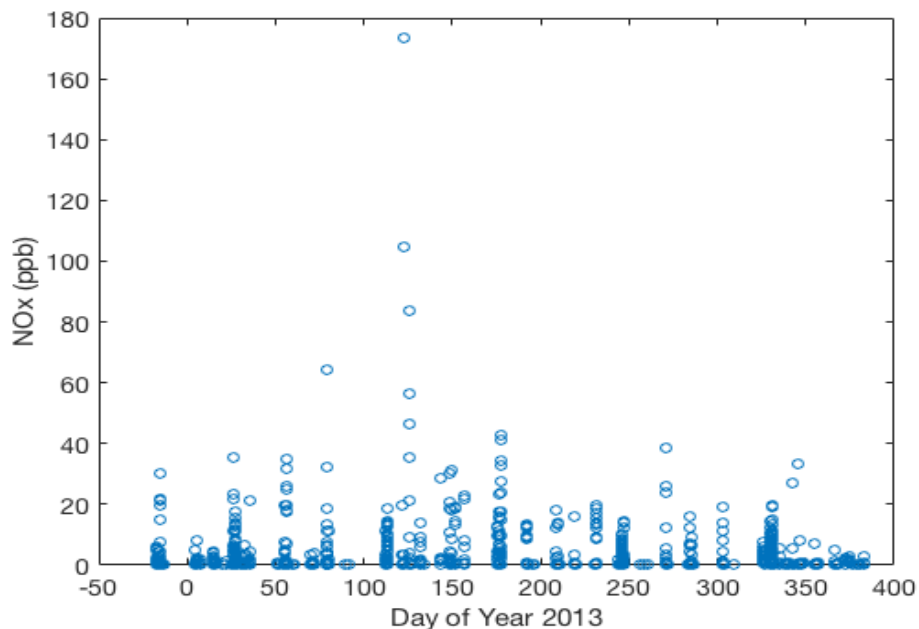
**Figure 3:** Photograph of Concordia Station illustrating the dispersion of the exhaust plume from the electrical power generating plant during conditions with a strong surface temperature inversion. The plume dissipated toward the west in the direction of the experimental site. This is a typical situation for a contamination event.

167

168 [Frey et al., 2011; Frey et al., 2013; Frey et al., 2015]. We did not observe highly elevated NO levels in the  
169 surface layer outside of pollution events, as what has been observed at South Pole [Helmig et al., 2008b;  
170 Neff et al., 2008]. During pollution events, the 1 ppb threshold was exceeded very quickly in  
171 measurements taken from the meteorological tower and from above-surface snow tower inlets, with  
172 resulting NO<sub>x</sub> mixing ratios rising to as high as close to 200 ppb, representing an up to 1000-fold  
173 enhancement over background conditions (Figure 4).

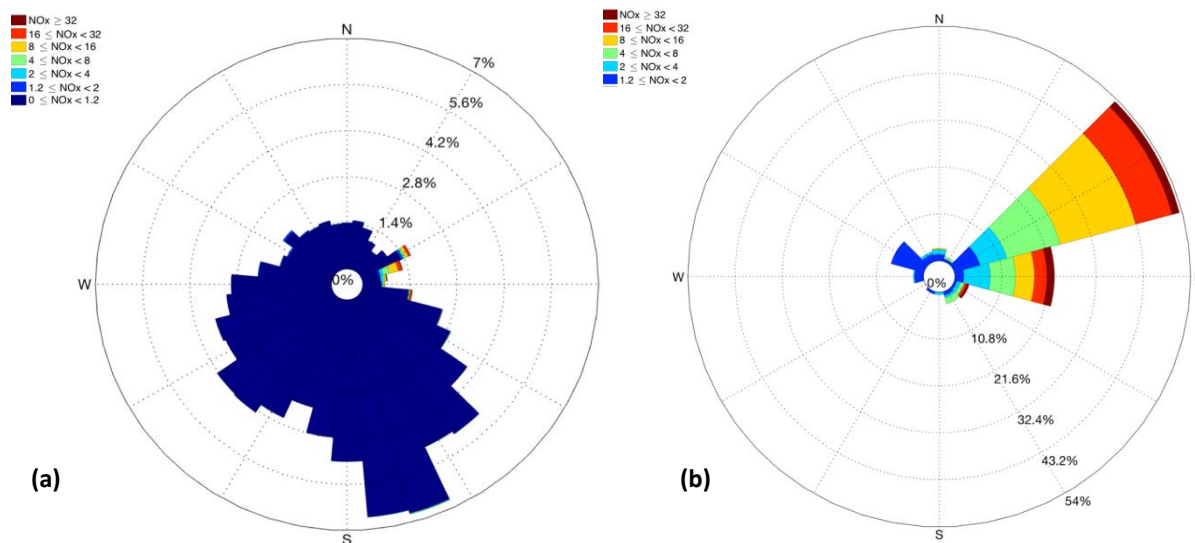
174 During the course of the ≈14 month study, a total of ≈50 pollution events were observed, although  
175 some events overlapped (Figure 4). Most of these occurrences were relatively short, with elevated mixing  
176 ratios above the snowpack lasting from minutes to a few hours. We counted 15 events in total when  
177 there was a spike in the above surface NO<sub>x</sub> measurements followed by an increase of at least 1 ppb of NO<sub>x</sub>  
178 in snowpack air. It took up to in excess of 7 days for NO<sub>x</sub> in the snowpack air to return to pre-event levels.  
179 Integrated over the entire campaign, pollution episodes constituted < 2.0 % of the measurements above  
180 the snowpack, and <10% of the measurements within the snowpack. The correlation analyses of pollution  
181 occurrences with wind direction clearly defines the direction of the transport. The predominant wind  
182 direction sector at Concordia is southeast to northwest (Figure 5a), with southeasterly winds having the  
183 overall largest share. NO<sub>x</sub> levels were consistently well below 1 ppb when winds were from east to  
184 northeast. The sector with pollution transport is well defined, with wind directions covering  
185 approximately 45-120° (Figure 5b). These sectors perfectly line up with the upwind direction of the station  
186 power plant (Figure 1), clearly identifying the plant as the source of these pollution occurrences.

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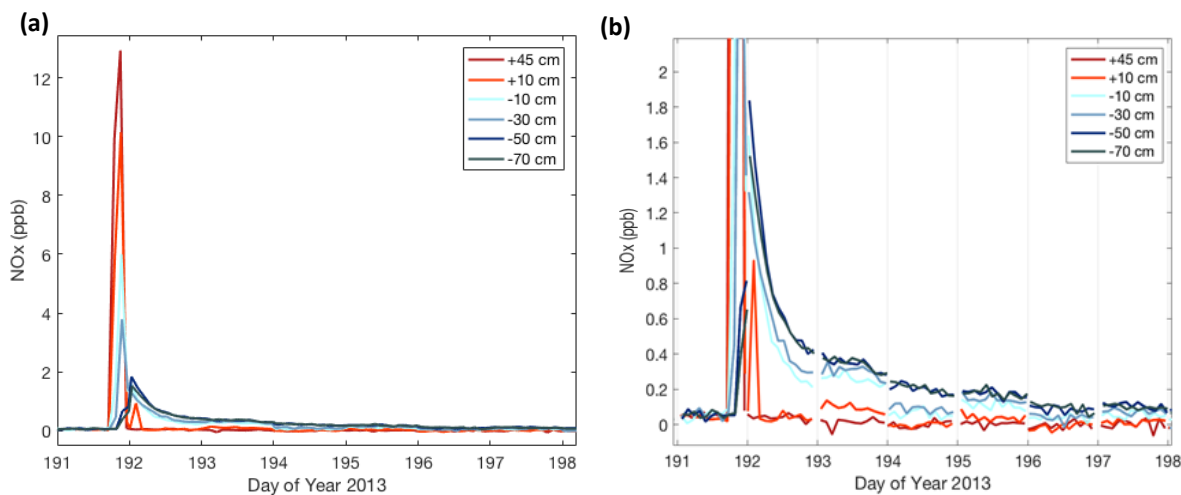


**Figure 4:** NO<sub>x</sub> measured from the above surface inlets on the two snow towers, and from all inlets on the meteorological tower. Circles represent 10-min averaged data. Plotted are data that were extracted from occurrences when NO<sub>x</sub> was above 1 ppb, well above the background mixing ratio, and any time when the wind direction was from the polluted wind sector.

188



**Figure 5:** (a) Concentration wind rose with the relative frequency of NO<sub>x</sub> mixing ratio data from the above surface inlets of the two snow towers segregated by 10° sectors for the full year of observation data. This panel shows all data. (b) The same analysis, with wind direction data binned in 20° sectors for events when NO<sub>x</sub> in ambient air exceeded 1.2 ppb.

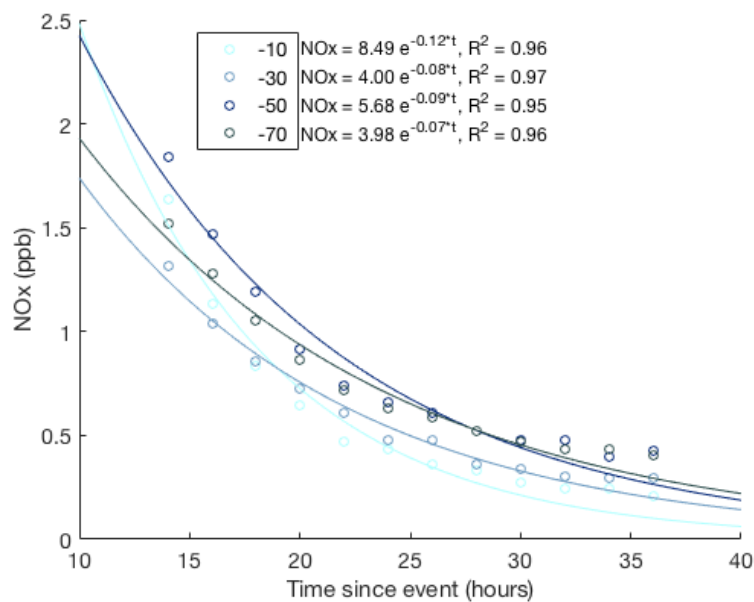


**Figure 6.** Combined measurements from the two snow towers capturing a pollution event at Concordia during the middle of the winter (Day of Year 191 = July 11). Plotted time series traces correspond to the sampling heights indicated in the legend, with positive numbers giving the height above the snow surface, and negative numbers the depth below the snow surface. The right graph (b) is an enlargement of the data shown in the left (a) with the transitions between the connecting lines between two snow tower measurements removed to show the level of agreement in the data from the two sampling manifolds. The sampling switched between the two snow towers every 24 hours, leading to some abrupt shifts in the snowpack NO<sub>x</sub> measurements.



191 One of these elevated NO<sub>x</sub> events is further investigated in **Figure 6**. Here, we show the  
 192 measurements from six inlets on the snow tower over a one-week period. The sampling of a polluted  
 193 plume is first observed in the two above surface inlets (orange/red colored data; +10 and +45 cm), by the  
 194 sudden increase of NO<sub>x</sub> from well below 1 ppb to a mixing ratio of ≈ 13 ppb. This spike in NO<sub>x</sub> lasted for  
 195 ≈3 hours. After that time, NO<sub>x</sub> in air sampled above the surface dropped very quickly and equilibrated to  
 196 prior mixing ratios within less than 0.5 hours.

197 A much different behavior was found in the air sampled from within the snowpack, indicated by the data  
 198 in the blueish colors. The onset of the pollution signal is delayed, by 1-3 hours, with progressively later  
 199 times towards deeper in the snowpack. Maximum mixing ratios that are reached in the snowpack are  
 200 lower, i.e. 10-50% of those that were measured above the surface, with mixing ratio maxima becoming  
 201 progressively smaller with increasing depth. The most remarkable difference between the above and  
 202 below surface measurements is the longer residence time of the pollution signal in the snowpack. NO<sub>x</sub>  
 203 mixing ratios in air withdrawn from all sampling inlets in the snowpack dropped steadily, but remained  
 204 elevated in comparison to levels seen before the pollution event for ≈ one week. The behavior seen in the  
 205 measurements from snow tower 1 were in full qualitative, and within ≈ 30% quantitative agreement with  
 206 the concurrent observations from the second snow tower. After the pollution event, NO<sub>x</sub> in the snowpack  
 207 air steadily declined over several days (**Figure 6**). Fitting of the data to an exponential decay function  
 208 yields similar results for all snowpack depths (**Figure 7**), with exponential regression fit R<sup>2</sup> results of ≥ 0.95.  
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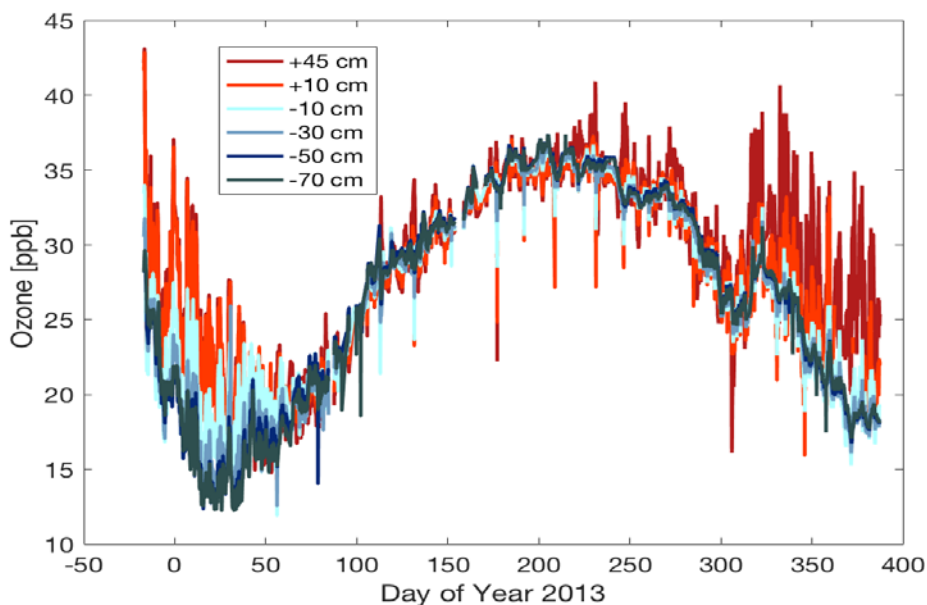


**Figure 7.** Exponential decay function fits to the NO<sub>x</sub> snowpack measurements versus time at four depths for the event starting on Day of Year 191 shown in Figure 6. The start of the event was defined as the time when high NO<sub>x</sub> was detected above the snowpack. Solutions for the best fit exponential decay functions are given in the legend.

210

211

212 Effects of the exhaust transport were also observed in the ozone signal. The ozone record, shown in **Figure**  
213 **8**, shows a plethora of short positive and negative spikes superimposed on the annual cycle. The up to  
214 15-20 ppb sudden ozone increases seen during the austral summer months are attributable to the  
215 photochemical ozone production events that occur in the surface layer of the Antarctic Plateau (see  
216 discussion in the introduction section). Occurrences of these elevated ozone events at Concordia Station  
217 have previously been investigated by *Legrand et al.* [2009] and *Cristofanelli et al.* [2018]. Besides these  
218 positive ozone spikes, this annual record also shows numerous sudden negative ozone changes that can  
219 be attributed to destruction of ozone by titration of NO in the exhaust plume. Up to 50% of the ambient  
220 ozone was destroyed in air sampled from the above surface inlets. Similar to NO<sub>x</sub>, this signal, albeit weaker  
221 and attenuated in time, was also seen in the air sampled from within the snowpack (**Figure 8**).



**Figure 8.** Ozone measured from the snow tower inlets throughout the year. Negative spikes in the data coincide with elevated NO<sub>x</sub> from exhaust infiltration in the snowpack.

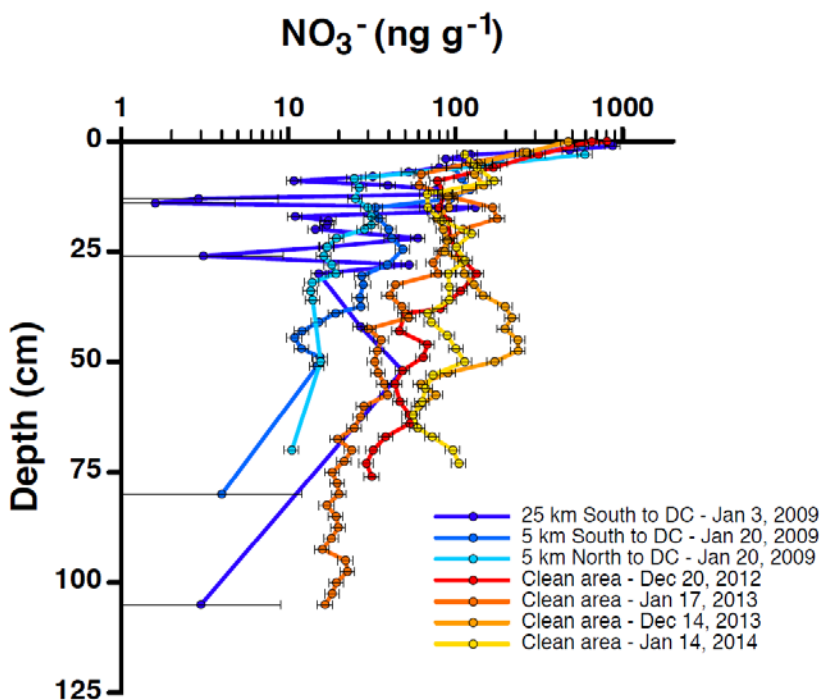
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224 Enhancements in formaldehyde in ambient samples suspected to be influenced by station emissions have  
225 previously been noted in measurements taken during the OPALE campaign [*Preunkert et al., 2015*]. We  
226 therefore investigated if pollution signatures were present in formaldehyde measurements taken in the  
227 record from our study. Furthermore, we revisited the gaseous elemental mercury (GEM) measurements  
228 taken during the campaign [*Angot et al., 2016*]. Both of these measurements did not quite have the time  
229 resolution and sensitivity as the NO<sub>x</sub> and ozone monitoring. Formaldehyde measurements overlapped  
230 with the ozone and NO<sub>x</sub> monitoring only for a short period during the 2012/2013 austral summer, and  
231 this measurement period suffered from a number of instrument problems. The remaining data did not  
232 allow a conclusive evaluation on the behavior of formaldehyde in the snowpack during pollution events.

233 Similarly, we did not identify a clear signature of GEM changes in snowpack air that correlated with the  
234 NO<sub>x</sub> enhancements during pollution events.

235 Nitrogen oxides undergo reaction with atmospheric oxidants, primarily the OH radical (summer only) and  
236 ozone, yielding higher oxidized nitrogen species that can partition into the snowpack aqueous and solid  
237 phase. The frequency, large enhancement, and long duration (in the snowpack) of NO<sub>x</sub> pollution events  
238 constitute an apparent unnatural source of NO<sub>x</sub> to the snowpack. One can hypothesize that reaction of  
239 NO and NO<sub>2</sub> with OH may be a source of HNO<sub>2</sub> and HNO<sub>3</sub> in the snowpack, which would add acidity to the  
240 snow. This then also poses the question if and to what degree photochemical processes, building for  
241 instance on NO<sub>2</sub><sup>-</sup> or NO<sub>3</sub><sup>-</sup> as a substrate, may be altered from natural conditions. Further, the  
242 transformations of NO<sub>x</sub> into these higher oxidized species may potentially leave a long-term chemical  
243 signature in the snowpack (such as of NO<sub>3</sub><sup>-</sup>).

244 We investigated this question by comparing NO<sub>3</sub><sup>-</sup> results from snow pit sampling at different locations  
245 within the camp and at up to 25 km distance of Concordia Station. Nitrate in the snowpack shows a steep  
246 vertical gradient, with highest levels observed right at the surface, and progressively lower concentrations  
247 with increasing depth (Figure 9; and Supplement Figure S-1 for a summary graph where data from both  
248 groups were combined and binned in 15 cm depth intervals). The results from the seven snow pits are  
249 consistent in the depth profile;

250



**Figure 9.** Nitrate concentration in snow pits in proximity ( $\approx 20$  m) to the snow towers at the border of the clean air sector (warm colors), and sampling locations up and downwind of Concordia (cold colors). Sampling dates are indicated in the figure legend. Horizontal error bars depict the estimated uncertainty of the chemical analysis, i.e. 10% at  $> 10$  ng g<sup>-1</sup>, 50% between 5-10 ng g<sup>-1</sup>, and 100% at  $< 5$  ng g<sup>-1</sup>.

251

252 however, there appears to be a tendency that the four snow pits within the camp have somewhat higher,  
253 as well as more variable snow  $\text{NO}_3^-$ . Deeper into the snowpack the difference between the two groups of  
254 data becomes stronger. At 1 m depth, the snow age is approximately ten years, and therefore  
255 approximately corresponds to the start of permanent and year-round activities in 2005 (note that summer  
256 activities at Concordia Station were established in 1996). Similar vertical  $\text{NO}_3^-$  concentration profiles in  
257 the Antarctic snowpack have previously been documented [France *et al.*, 2011]. Further, relatively higher  
258  $\text{NO}_3^-$  concentrations (approximately 5 times) are observed in the upper layers of the snowpack during the  
259 Antarctic summer (Supplement Figure S-2) than during the winter. This  $\text{NO}_3^-$  enhancement in surface  
260 snow and the seasonal cycle have been linked to the production of  $\text{HNO}_3$  in the photoactive summer  
261 months from reaction of OH with  $\text{NO}_2$ , with the  $\text{NO}_3^-$  enhancement being the result of  $\text{HNO}_3$  deposition to  
262 the snow surface [Erbland *et al.*, 2013]. Elevated  $\text{NO}_3^-$  concentrations observed at the surface of the  
263 snowpack are a common feature at low accumulation regions, with the concentration values depending  
264 strongly on the strategy for collection of the first few cm of the snowpack [Erbland *et al.*, 2013; Shi *et al.*,  
265 2018]. With the sensitivity of surface snow  $\text{NO}_3^-$  to the seasonal cycle and, the high variability of  $\text{NO}_3^-$   
266 observed in surface snow and vertical profiles, the currently available data do not allow a conclusive  
267 quantitative assessment of the degree to which the snowpack at Concordia is chemically impaired by  
268 ventilation of the snowpack with pollution- $\text{NO}_x$  enriched air. Contamination of the snowpack around  
269 Concordia station has been noted in previous investigations. Warren *et al.* [2006] reported an > 3-fold  
270 increase in snowpack black carbon concentration after the station was established in 2003, with black  
271 carbon levels in pre-2003 snow (from deeper depths) also being in closer agreement with snow collected  
272 from further distance to the station. Black carbon is one of the contributing factors for a decrease in light  
273 penetration into the snowpack [Warren *et al.*, 2006; France *et al.*, 2011; Libois *et al.*, 2013]. Consequently,  
274 the increased presence of black carbon causes a shallower e-folding depth compared to pristine,  
275 uncontaminated snow. In addition to the experiments described above, during the 2014 campaign a  
276 number of dynamic flow-through snow chamber photochemistry experiments were conducted to  
277 investigate if there were differences in the reactive chemistry in the snow from near the snow tower site  
278 compared to snow sampled 25 km away from camp. These measurements showed on the order of 10-  
279 20% higher  $\text{NO}_x$  and less ozone in the outflow of chambers filled with the snow from further away from  
280 the camp. We did not conduct a high enough number of repeats for evaluating the repeatability and  
281 statistical significance of these results to gauge if and how much of this signal was due to the experimental  
282 setup or due to differences in the snowpack chemical composition. Nonetheless, these preliminary  
283 findings point towards possible differences in the chemical behavior that potentially are linked to  
284 differences in the snow sampling locations and contaminant levels resulting from camp influences that  
285 warrant further investigation.

286

## 287 **Summary and Conclusions**

288 With our snowpack sampling manifold we were able to sample snowpack air to a maximum depth of 70  
289 cm below the surface. Up to  $\approx 2$  ppb enhancements in  $\text{NO}_x$  were observed at that depth from exhaust  
290 infiltration. While our experiment was not able to access air deeper (than 70 cm), the observed  
291 concentration gradients imply that this transport and contamination extends well beyond the depth that  
292 was probed in these measurements.

293 This experiment was a one spot measurement, at  $\approx 1$  km distance from the camp main facilities. We have  
294 no data that would allow us to assess to what distance from the camp the snowpack pollution from  
295 exhaust infiltration would be noticeable and of importance, but it likely extends well beyond the distance  
296 of our site.

297 Several other previous studies have noted challenges in sampling clean air at polar research sites that  
298 stem from pollution caused by camp exhaust. For atmospheric sampling, this interference can be  
299 mitigated by careful postprocessing/filtering of the monitoring data, or by interrupting the sampling  
300 during unfavorable wind conditions, which is particularly critical for integrated aerosol collection [Wolff  
301 and Cachier, 1998; Wolff *et al.*, 1998]. Our experiments from Concordia Station emphasize the  
302 pronounced and longer-lasting influence that station exhaust can have on  $\text{NO}_x$  levels inside the snowpack  
303 air (compared to ambient air). A tendency of potentially enhanced snowpack  $\text{NO}_3^-$  levels in two snow pits  
304 collected at the camp, compared to data from three sites at further distance, supports the suspicion that  
305 the snowpack chemical oxidized nitrogen composition at the station may be compromised (i.e.  
306 contaminated) from the re-occurring ventilation of the snowpack with polluted,  $\text{NO}_x$ -enriched air. A  
307 similar conclusion was derived from  $\text{d}^{15}\text{N}$  nitrate analyses of snow at Summit. Samples collected in the  
308 predominant downwind direction of the station generator showed an isotopic signature that had a  
309 stronger association with engine exhaust than samples collected at further distance [Fibiger *et al.*, 2016].  
310 Even in the Summit clean air sector elemental carbon in snow was 1.8 – 2.4 times higher than in snow  
311 collected at 10-20 km distance from the camp [Hagler *et al.*, 2008]. The associations shown in our study  
312 argue for further investigation, for instance by a high resolution spatial survey of surface snow  
313 composition within and beyond camp boundaries. Given the strong seasonality of  $\text{NO}_3^-$ , this survey should  
314 be done with as close as possible concurrently conducted snow sampling at selected locations to minimize  
315 the influence of temporal changes on the  $\text{NO}_3^-$  signature.

316 These observations emphasize concerns about the representativeness of experimental snow chemistry  
317 data collected within a Polar research camp periphery. This raises the question how interpretations from  
318 such experiments reflect conditions in the remote Polar environment. Furthermore, our findings should  
319 motivate comparison studies with sampling along transects to further distance from the main camp  
320 facilities. Comparison of these observations will likely yield new insights for evaluating prior polar  
321 research site observations and interpretation of snow photochemistry in the glacial snowpack.

322

#### 323 **Author contribution:**

324 DH oversaw the study, participated in field work, conducted data analyses and quality control, and  
325 prepared the manuscript.

326 DL conducted data analyses, quality control, prepared figures, and contributed to the manuscript  
327 preparation.

328 JH fabricated the instrumental equipment and data acquisition system, and participated in the field work.

329 JS contributed to the study design, participated in field work, conducted data analyses and interpretation,  
330 and contributed to the manuscript preparation.

331

332

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