



# IcePAC – a Probabilistic Tool to Study Sea Ice Spatiotemporal Dynamic: Application to the Hudson Bay area, Northeastern Canada

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**Abstract.** A reliable knowledge and assessment of the sea ice conditions and their evolution in time is a priority for numerous decision makers in the domains of coastal and offshore management and engineering as well as in commercial navigation. As of today, countless research projects aimed at both modelling and mapping past, actual and future sea ice conditions were completed using sea ice numerical models, statistical models, educated guesses or remote sensing imagery. From these research, reliable information helping to understand sea ice evolution in space and in time are available to stakeholders. However, no research has, as of today, assessed the evolution of the sea ice cover with a frequency modelling approach, by identifying the underlying theoretical distribution describing the sea ice behaviour at a given point in space and time. This project suggests the development of a probabilistic tool, named IcePAC, based on frequency modelling of historical 1978-2015 passive microwave sea ice concentrations maps from EUMETSAT OSI-409 product, to study the sea ice spatiotemporal behaviour in the waters of the Hudson Bay System in Northeastern Canada. Pixel scale models are based on the generalized Beta distribution and generated at a weekly temporal resolution. Results showed coherent with the Canadian Ice Service 1980-2010 Sea Ice Climatic Atlas average freeze-up and meltdown dates for numerous coastal communities in the study area and showed that it is possible to evaluate a range of plausible events, such as the shortest and longest probable ice free season duration, for any given location on the simulation domain. Results obtained in this project open a path towards various analyses on sea ice dynamics that would gain in informative content and value by relying on the kind of probabilistic information and simulation data available from the IcePAC tool.

## 1. Introduction

Numerous scientific projects recognized the link between climate change and changes in the spatiotemporal sea ice dynamics (Andrews et al., 2017; Cavalieri and Parkinson, 2012; Comiso, 2011, 2002; Comiso et al., 2008; Gloersen et al., 1998; Johannessen et al., 2004; Rothrock et al., 1999; Stocker, 2014; Stroeve et al., 2007; Stroeve et al., 2014; Stroeve et al., 2012; Wang and Overland, 2009). The average temperature raise in the Arctic region is said to reach 2.0°C according to Serreze and Barry (2011), which makes it logical to apprehend that changes in the spatiotemporal dynamic of the sea ice cover will tend to amplify.



In this changing environment, an adequate and efficient monitoring of sea ice is of capital importance to better understand the climate and its impacts on marine and coastal areas (Allard and Champagne, 1980; Barnhart et al., 2014; Bernatchez and Dubois, 2004; Bernatchez and Dubois, 2008; Bintanja and Selten, 2014; Davies et al., 2014; Dionne, 1973; Holland et al., 2006; Kowal et al., 2017; Manabe and Stouffer, 1995; Overeem et al., 2011; Peterson et al., 2002; Rahmstorf, 1995; Rahmstorf and Ganopolski, 1999; Vermaire et al., 2013), on the security of economical and logistical activities in northern communities (Aksenov et al., 2017; Andrews et al., 2017; Ho, 2010; Lasserre and Pelletier, 2011; Liu and Kronbak, 2010), on arctic marine fauna protection (Bhatt et al., 2010; Castro de la Guardia et al., 2013; Darnis et al., 2012; Laidre et al., 2015; Post et al., 2013; Wassmann et al., 2011) and on the traditional way of life of Inuit communities (Durkalec et al., 2015; Laidler et al., 2010).

To understand and appreciate the role of the sea ice cover regarding climate, marine and coastal environment management, fauna protection and the cultural traditions of northern communities, access to informative and reliable sea ice dynamics information is fundamental. Engineers, stakeholders, Inuit and northern populations, navigators and scientists must be able to quantify hazards related to the sea ice cover in order to efficiently evaluate, anticipate and minimize the risks of usage, building and exploitation in marine and coastal areas. Given the increase in activity noticed in the Arctic and the North, one can expect the demand in information to also increase.

Despite a large number of Earth observation datasets on the sea ice cover, only a few provide both high temporal and spatial resolutions. National ice services, such as the Canadian Ice Service (CIS), provide the users with sea ice conditions climatology that are a reliable source of descriptive statistics on the sea ice dynamics, such as the average freeze-up or maximum extent date. Nevertheless, these datasets do not carry information on the nature of the underlying statistical distributions of sea ice parameters, such as sea ice concentration (SIC), at any given point and nor do they permit to analyse the probability of occurrence of a specific sea ice condition.

By exploring an innovative probabilistic sea ice concentrations modelling avenue, this study proposes a tool, named IcePAC, to characterize the underlying statistical distributions of the SIC at any point in the Hudson Bay System (Saucier et al., 2004) based on historical passive microwave remote sensing data from 1978 to 2015. These data are then used to analyse the dynamic of SIC in the HBS with a probabilistic perspective and compared to the CIS climatology.

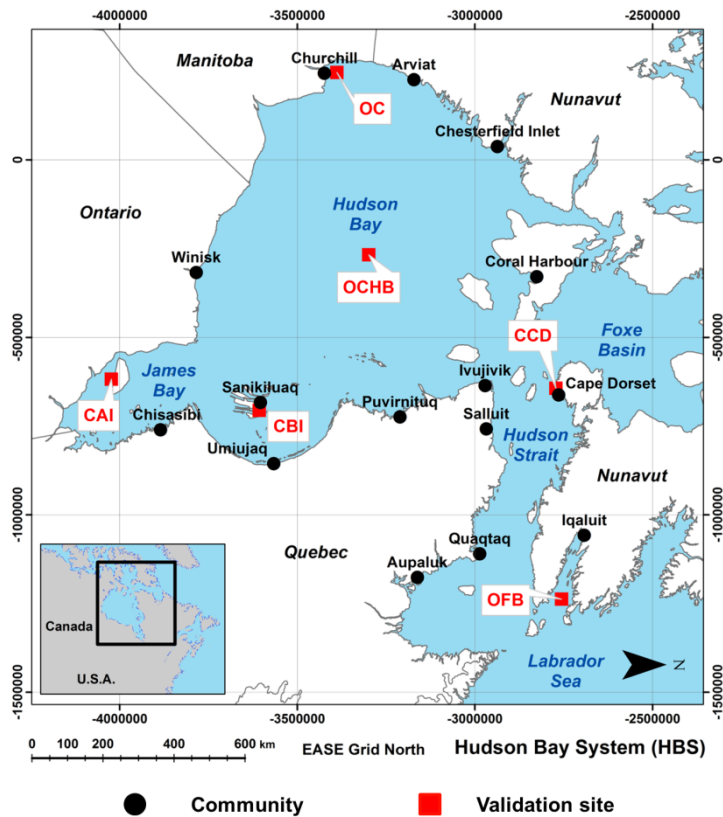
## 2. The Hudson Bay System

The study area is the Hudson Bay System, consisting of the Hudson Bay, the Hudson Strait, James Bay and Foxe Basin (Fig. 1). The HBS is surrounded by the three Canadian provinces of Quebec, Ontario, Manitoba and the territory of Nunavut. It is the largest inland sea on Earth, with a total area of 1 300 000 km<sup>2</sup> (Etkin, 1991; Gagnon and Gough, 2005a; Martini, 1986) and is located in both subarctic and arctic regions. An estimated 20 % of the flux of freshwaters to the Arctic Ocean are thought to come from rivers flowing into the HBS, which represents 900 km<sup>3</sup>/yr<sup>-1</sup> (Déry and Wood, 2005; Déry and Wood, 2004). It is connected to the Labrador Sea by the Hudson Strait and to the Arctic Ocean by the Foxe Basin (Prinsenber, 1986) and is characterized by shallow depths of less than 100 m in the Foxe Basin, of a maximum of 125 m in



65 the Hudson Bay and of more than 200 m in the Hudson Strait (Jones and Anderson, 1994). It has cyclonic currents generated mostly by winds, with a maximum intensity in November (Saucier et al., 2004).

A large number of researches have been done to document the average sea ice behavior in the HBS (CIS, 2013; Gagnon and Gough, 2005a, b; Hochheim and Barber, 2010; Hochheim and Barber, 2014; Kowal et al., 2017; Maxwell, 1986), which goes every year through a complete cryogenic cycle. The sea ice cover in the HBS is essentially constituted of first-year ice, with the exception of possible multi-year ice drifting in the Foxe Basin (CIS, 2013; Etkin and Ramseier, 1993). The sea ice cover initially forms in the northwestern part of the HBS near Southampton Island in late November and progresses towards the southeastern part of the HBS (Hochheim and Barber, 2014; Maxwell, 1986) to finally achieve a complete freeze-up in late December. An annual maximum in extent is usually achieved in April (Gagnon and Gough, 2005b), after which the meltdown begins in May along the northwestern shoreline of the HBS. The meltdown is driven from the shores toward the center of the Bay which usually results into an agglomerate of sea ice in the southern part of the Bay in late July (CIS, 2013). In summary, the HBS is on average entirely frozen in late December and entirely free of ice in mid-August (Mysak et al., 1996; Wang et al., 1994).



80 **Figure 1:** The Hudson Bay System with communities (black dots) and IcePAC validation sites (red squares)



### 3. Data and methods

Sea ice extent has displayed an important decline in the last decades (Cavalieri and Parkinson, 2012; Markus et al., 2009; Meier et al., 2007; Parkinson and Cavalieri, 2002, 2008; Parkinson et al., 1999), as it can be observed with remote sensing images acquired since 1978. Another important source of information on the sea ice cover are model predictions  
85 which come from deterministic models (Hunke et al., 2017; Rousset et al., 2015; Weaver et al., 2001), based on dynamical and thermodynamic equations evolving in synergy inside a modelling framework, or from statistical models, based on statistical tools such as simple and multiple regression analysis (Ahn et al., 2014; Drobot, 2007; Pavlova et al., 2014) to explain an expected sea ice parameter value.

Another statistical approach, focusing on the estimation of the probabilities of occurrence of specific sea ice related  
90 events, has been used in recent research. It is achieved either by using the simple count method (e.g. an event occurred four times in the last ten years, which corresponds to a 40 % probability of occurrence) or by using the frequency modelling method, which consist of adjusting a theoretical distribution to a series of observations, consequently defining the plausible sea ice events for the entire range of probabilities (i.e.  $p = 0$  to 100 %). In this research, the frequency modelling method is used on series of passive microwave historical sea ice concentrations (SIC) remote sensing data to adjust distributions to a  
95 total of 20 738 pixels or sites (i.e. the spatial dimension) for each of the 52 weeks of the year (i.e. the temporal dimension), resulting in a total of 1 078 376 distribution fits.

The datasets used and the protocols followed build the IcePAC tool and to model SIC distributions at every pixel in the HBS are described in the following section.

#### 3.1 Sea ice concentration dataset

100 Sea ice concentration is defined as the proportion of sea ice covering a predefined area, expressed as a percentage (Shokr and Sinha, 2015). In remote sensing, this predefined area is represented by a pixel. The choice of an SIC dataset for frequency modelling in our study is highly influenced by the extent of the HBS (1 300 000 km<sup>2</sup>) and has been made to ensure uniformity and continuity of the series used for analysis.

105 Multiple SIC datasets are generated using either visible/thermal, SAR or passive microwave remote sensing. As the objective of the IcePAC tool is to provide the capacity to evaluate the spatiotemporal evolution of the SIC, the source of data needed to ensure a complete coverage of the HBS (i.e. the spatial dimension) and to ensure continuity in the series (i.e. the temporal dimension). A passive microwave dataset had been chosen as it answers these two needs.

110 The OSI-409 – Reprocessed Global Sea Ice Concentration dataset (Eastwood et al., 2015; Tonboe et al., 2016) was selected as it makes possible the reconstitution of SIC series on more than 30 years with a 12.5 km spatial resolution and is processed with a unique hybrid SIC algorithm. The hybrid algorithm integrates the information taken from the Bootstrap algorithm (Comiso, 1995) when  $SIC < 70\%$ , linearly weights the SIC estimated by the Bootstrap and Bristol (Smith, 1996) algorithms when  $70\% < SIC < 90\%$  and integrates the information from the Bristol algorithm when  $SIC > 90\%$ .

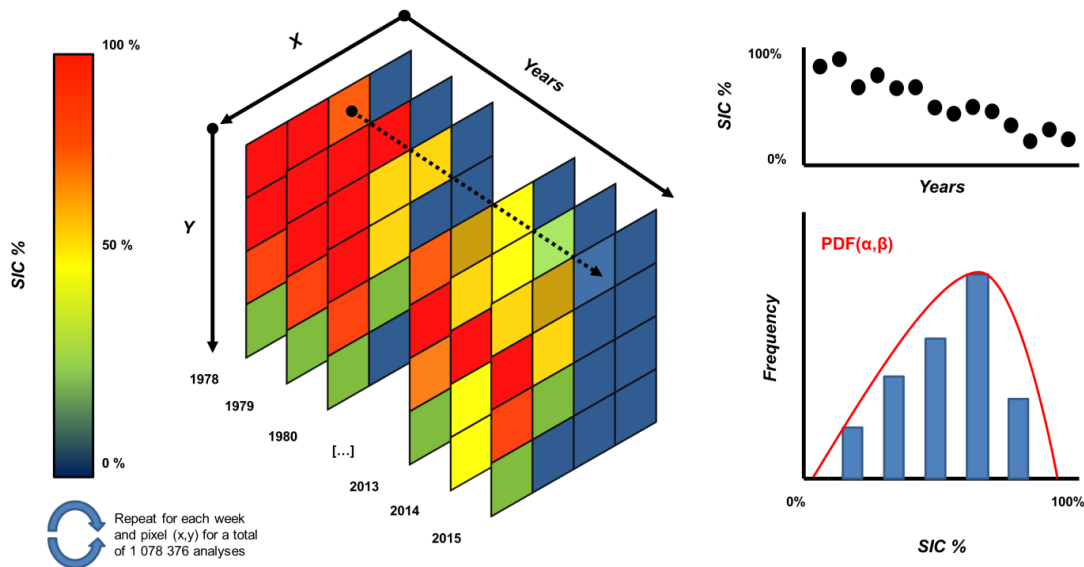


The data has been clipped to the HBS extent using Natural Earth (NaturalEarth, 2014) vector dataset with an estimated spatial resolution of 500 m (Wessel and Smith, 1996), well beyond the resolution of the OSI-409 data.

### 115 3.2 Frequency modelling

The core of the IcePAC tool relies on the use of frequency modelling which consist on describing the probable conditions at a given site with a simplified model fitted on historical SIC data, describing the underlying frequency distribution of the phenomenon. The first step in this approach is to build the time series of historical data, then to ensure their quality using preliminary tests and finally to identify and fit the model on the data (Fig. 2).

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**Figure 2:** The process of frequency modelling from series building to model fit

#### 3.2.1 Building the sea ice concentration (SIC) series

125 The SIC series are built to represent the SIC state for a specific week, for all years between 1978 and 2015. It is on these series that we adjust a theoretical distribution to estimate the probabilities of SIC related event. First, the daily OSI-409 data have to be averaged every 7 days to create weekly datasets. This operation is made following a 365-days “no-leap” calendar convention (i.e. every year has 365 days), separated in 52 weeks (December 31<sup>st</sup> is included in week 52). Second, the data for each week number are stacked in chronological order, from 1978 to 2015. Three types of series are observed in  
 130 the HBS (Fig. 3): the ICE series, where we observe constant high SIC, the MIZ (Marginal Ice Zone) series, where we observe a transition from sea ice to open water, and the OW (Open Water) series, where we observe constant low SIC.

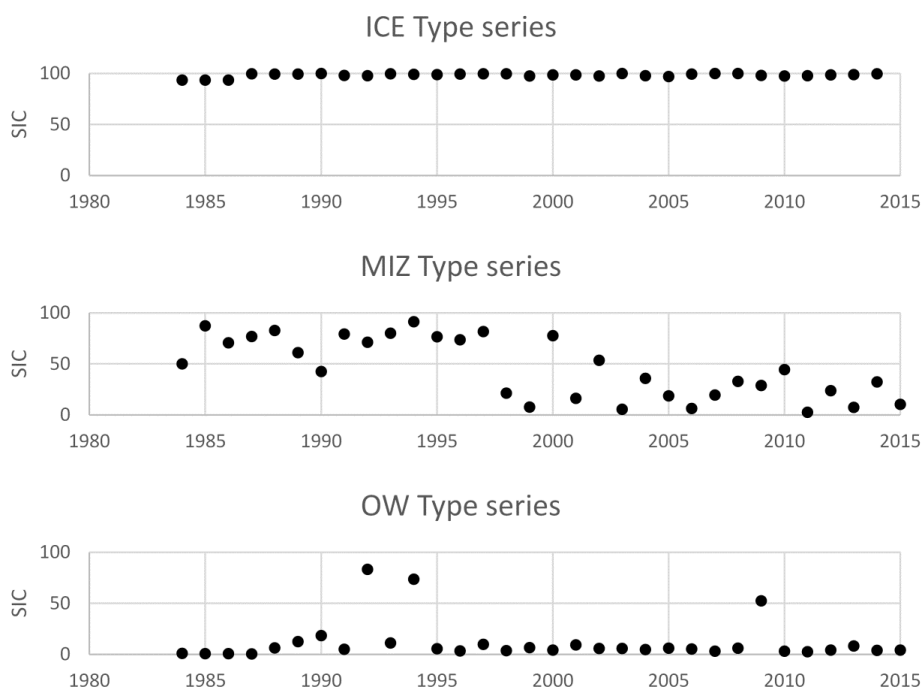


Figure 3: The three types of SIC series encountered in the HBS

### 135 3.2.2 Preliminary tests

Series have to go through a set of preliminary tests to assess their stationarity, homogeneity and independence, assuring they are suitable for frequency modelling. The tests used in IcePAC are the Mann-Kendall test for stationarity (Mann, 1945), the Wald-Wolfowitz test for independence (Wald and Wolfowitz, 1940) and the Wilcoxon test for homogeneity (Wilcoxon, 1945).

140 Series are considered independent if the subsequent unique SIC observations have no incidence on one another; they are considered homogeneous if they are reputed to be from the same distribution and; they are stationary if they are not affected by a trend. In the case of a detected non-stationarity, the trend must be modeled and subtracted from the series (Cave and Pearson, 1914).

145 All the time series, either original or detrended in regards to the Mann-Kendall test, satisfied the preliminary tests for signification level of  $\alpha = 0.05$ .

### 3.2.3 Trend estimation and removal

150 The estimation of a trend on a percentage data series has the particularity that the trend must, in order to be coherent with the physics of the studied phenomenon, be bounded in a  $[0,1]$  domain (Baum, 2008). In other words, we must ensure we do not measure a trend that generates SIC values larger than 100 % or smaller than 0 %. To guarantee the respect of this criterion, a generalized linear model with a logit link function has been used to estimate the trend. The logit link function



linearizes the SIC values using the logit transformation (Eq. 1) and it is on these transformed values that a linear regression of the form  $\alpha x + \beta + \varepsilon$  is measured.

$$\text{logit}(SIC) = \ln(SIC/1 - SIC) \text{ where } SIC \text{ is defined on } ]0,1[ \quad (1)$$

155 For the trend to be removed from the series, an inverse transformation (Eq. 2) must be applied to the estimated trend to bring back its logit values into SIC values.

$$SIC = \exp\left(\frac{\exp(\text{logit}(SIC))}{(1 + \exp(\text{logit}(SIC)))}\right) \quad (2)$$

### 3.2.4 Distribution selection and fit

160 The selection of adequate candidate distributions to fit on the series is largely limited by the bounded nature of the data. The selected distribution has to be bounded to  $[0,1]$  and to be available in a generalized form in order to adapt to detrended series which are bounded to  $[-1,1]$ . Their generalized forms are to be used with the position parameter  $a$  fixed to  $-1$  and the scale parameter  $b$  fixed at  $2$ , in coherence with the phenomenon. It must also present enough flexibility in shape to adapt to the different type of SIC series in the HBS domain. Two different distributions, the generalized Beta distribution  
 165 and Johnson's SB distribution (Johnson, 1949), both bounded and displaying flexibility in shape, have been fitted on the series using the maximum likelihood estimator (MLE) (NIST, 2013) and compared by measuring the root mean square error between observations and adjusted curves as well as the Akaike Information Criterion (Akaike, 1998) and the Bayesian Information Criterion (Schwarz, 1978).

170 The generalized Beta distribution (Eq. 3) has four parameters which are the  $p$  ( $p > 0$ ) and  $q$  ( $q > 0$ ) shape parameters, the position parameter  $a = -1$  and the scale parameter  $b = 2$ . In Eq. 3,  $B$  is the Beta function. This distribution has been used before in climatology (Dirkson, 2017; Falls, 1974; Henderson-Sellers, 1978; Sulaiman et al., 1999), in seismology (Lallemant and Kiremidjian, 2015), to study air pollution (Nadarajah, 2008) and in hydrology (Chen and Singh, 2017; Yao, 1974).

$$f(x) = \frac{(x - a)^{p-1}(b - x)^{q-1}}{B(p, q)(b - a)^{p+q-1}}, \text{ where } a \leq x \leq b; p, q > 0 \quad (3)$$

175

The Johnson's SB (Eq. 4) is, under its generalized form, a four parameters distribution, for which the  $\gamma$  and  $\delta$  ( $\delta > 0$ ) are the shape parameters, the parameter  $a = -1$  is position and the parameter  $b = 2$  is the scale. This distribution has been used



before in meteorology (Cugerone and Michele, 2015; Wakazuki, 2013), in forestry (Rennolls and Wang, 2005) and in hydrology (D'Adderio et al., 2016).

180

$$f(x) = \frac{\delta}{\sqrt{2\pi}} \frac{(b-a)}{(x-a)(b-x)} \left[ -\frac{1}{2} \left\{ \gamma + \delta \ln \left( \frac{x-a}{b-x} \right) \right\}^2 \right], \text{ where } b, \delta > 0 \quad (4)$$

Two approaches were tested for distribution selection. The first approach considered fitting the distributions to the series according to their Mann-Kendall test results. The second approach considered fitting the distributions to systematically detrended series in order to preserve spatial coherency in the IcePAC outputs.

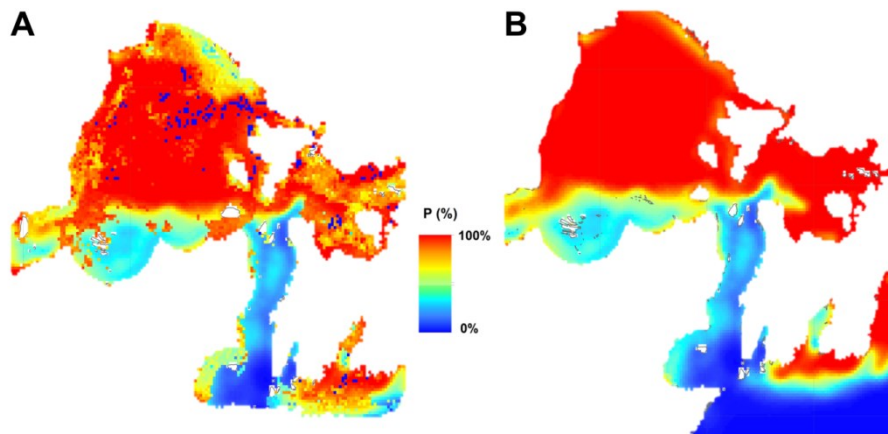
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The use of a systematic trend removal approach is justifiable by the fact that natural phenomena are considered by nature non stationary (Lins, 2012; Rao et al., 2012) and by the relative shortness of the series which can have an effect on the conclusions of the Mann-Kendall test (Hirsch et al., 1982).

190

The two approaches, tested on 958 randomly chosen series, brought similar conclusion, the Beta distribution outperforms the Johnson's SB distribution in 71.4 % of the time. For the remaining 28.6 % adjustments where the Johnson's SB distribution did perform better, both the information criterion and the RMSE showed a non-significant difference with the Beta distribution. As expected, using an approach which systematically corrects for trends in series before distribution fit improved the spatial coherency of the resulting probability maps generated by IcePAC (Fig. 4). In scope of these results and to preserve parsimony in IcePAC, it was decided to use only the Beta distribution for all series, for which a systematic trend removal was applied.

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**Figure 4:** Comparison of IcePAC results for non-systematic (i.e. based on Mann-Kendall test result) trend removal (A) and systematic trend removal (B), on the resulting map for the probability of observing a SIC > 95% on week #1





### 3.2.5 Model queries

200 Queries with the IcePAC tool are possible via three important functions resulting from the distribution fit, the probability density function (PDF), the cumulative distribution function (CDF) and the percent point function (PPF). These functions are obtained using the fitted parameters from the Beta distribution independently for each of the 20 738 points.

The PDF is obtained by fitting a theoretical distribution on the frequency histogram of the SIC observations. The selected distribution in IcePAC is the Beta distribution for which the parameters of Eq. (3) were estimated using the  
205 Maximum Likelihood Estimator (MLE). The PDF allows visualizing how the probability is distributed between values and how it evolves between different values.

Derived from the PDF, the CDF (Eq. 5) allows estimating the probability of a given range of SIC values. It corresponds to the area under the PDF curve for a specific range of SIC values, usually from 0 % to  $SIC_{MAX}$ . As a CDF  
210 example, one could query to know the probability of non-exceedance ( $p$ ) for a sea ice concentration of 25 % ( $SIC_{MAX}$ ) for week #1 for the year to come.

$$F_x(SIC) = \int_0^{SIC_{MAX}} f_x(SIC) dt \quad (5)$$

The inverse function of the CDF, the PPF (Eq. 6) allows to estimate the SIC value for a given probability of non-exceedance. As a PPF example, one could query to know the  $SIC_{MAX}$  for a probability of non-exceedance of 55 % for week  
215 number 1 for the year to come.

$$Q = F^{-1}(p) \quad (6)$$

Since the IcePAC fits are made on detrended series (e.g. residuals), the trend has to be taken into account when processing queries, meaning that it has to be either removed from the  $SIC_{MAX}$  value if the CDF is used or added to the result  
220 of the PPF query in order to render a physically valid result. The query flowchart is presented in Fig. (5).

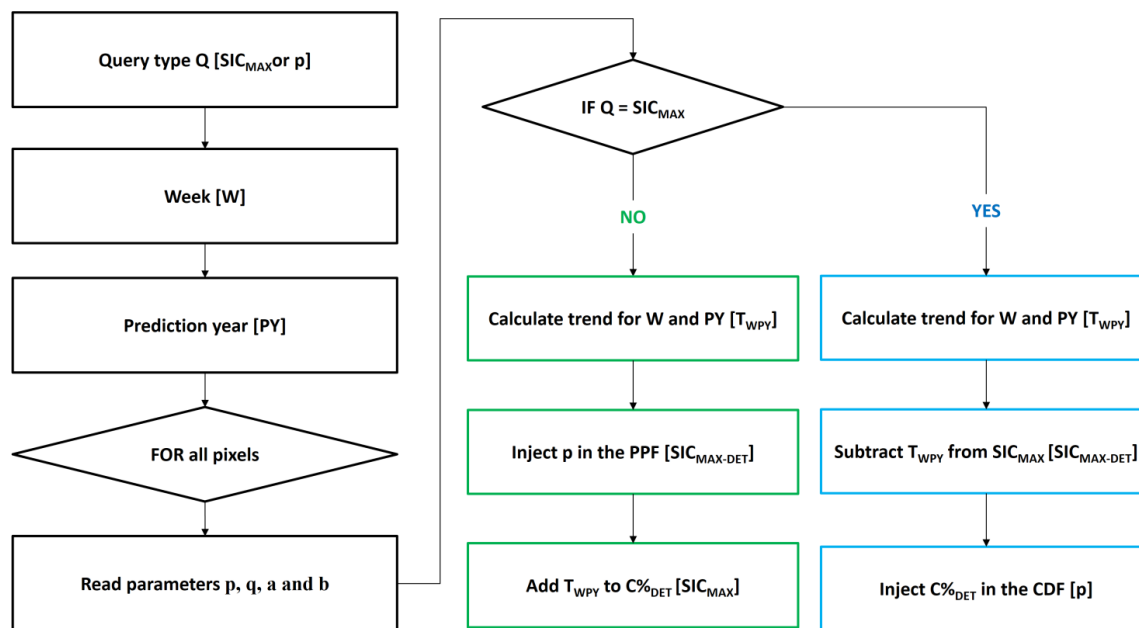


Figure 5: IcePAC queries flowchart

### 3.2.6 IcePAC versus observations in 2016

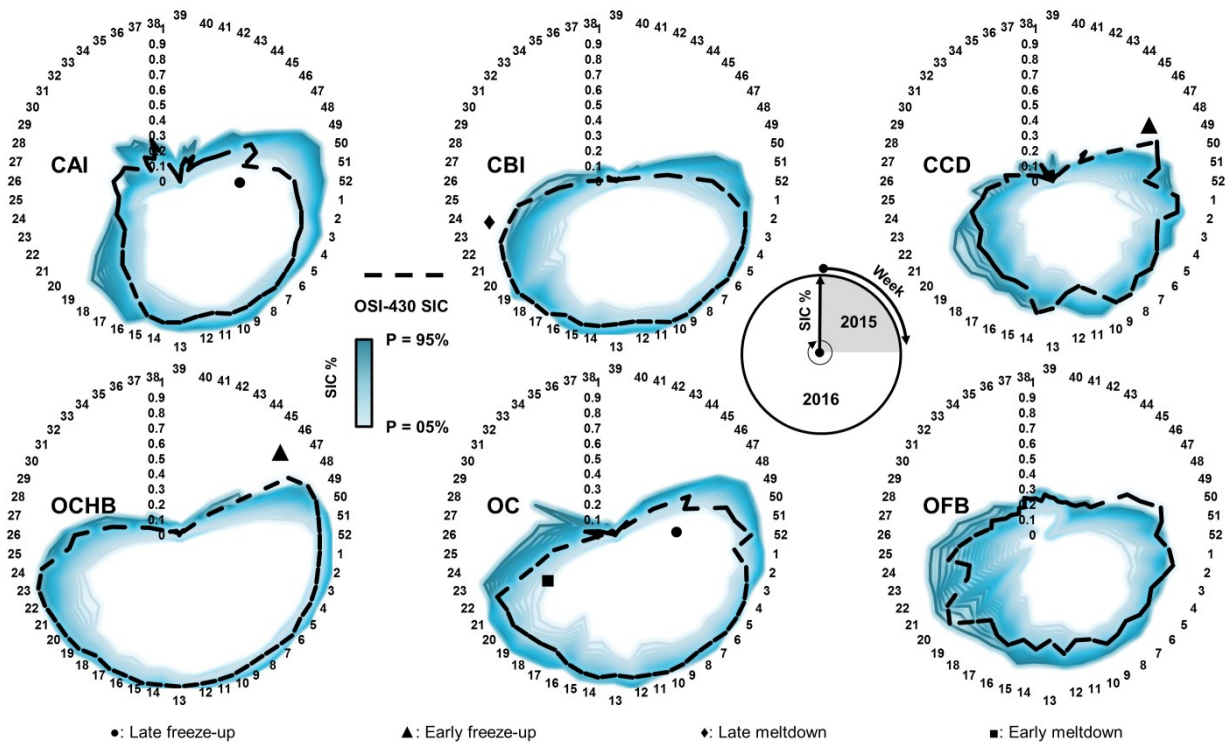
225 The assessment of the validity of IcePAC predictions was done by comparing IcePAC weekly outputs time series for the entire year with the OSI-430 product, an independent data source not used in IcePAC development but based on the same SIC retrieval algorithm (Tonboe et al., 2016). The comparison was made on the 2015-2016 sea ice season, not included in the input data used for fitting all underlying IcePAC Beta distributions.

230 Six different comparison sites were selected to represent different sea ice dynamics (see Fig. 1). Three coastal sites, the Akismi Island (CAI), Cape Dorset (CCD) and Belcher Islands (CBI), were sampled to assess the behavior of IcePAC predictions along the coastline, at different latitudes. Also, three offshore sites, Frobisher Bay (OFB), Central Hudson Bay (OCHB) and Churchill (OC) were sampled to assess the behavior of IcePAC predictions offshore, at different latitudes.

235 Figure 6 displays weekly SIC predictions outputs from IcePAC, in the  $p = 5$  to 95 % range with a 5 % pace. Each comparison site shows a modeled dynamic which is coherent with reality and with the OSI-430 2015-2016 observation data. As it could be expected, it is during the freeze-up and meltdown periods that the range of probable SIC values is the largest (i.e. the spacing between light and dark blue lines in Fig. 6), compared to stable cover periods where the range of probable SIC values narrows down. Variations in and out of the predicted range of SIC values can be attributed to the natural variation of ice conditions, the occurrence of an highly improbable event, the OSI-409 algorithm bias or a combination of these factors. Some anomalies of various intensities can be detected in the series such as: early 2015 freeze-up events at points



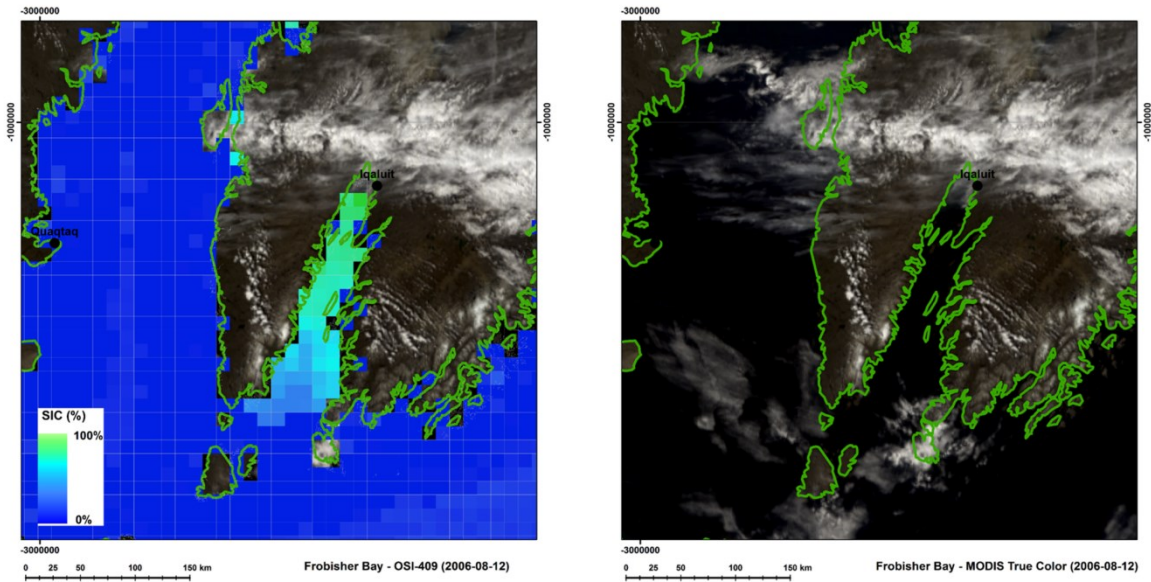
240 CCD and OCHB (▲ in Fig. 6); late 2015 freeze-up events at points CAI and OC (● in Fig. 6); late 2016 meltdown event at point CBI (◆ in Fig. 6); and an early 2016 meltdown event at point OC (■ in Fig. 6), which all are in agreement with the anomaly maps of the NSIDC’s Arctic Sea Ice News and Analysis (<https://nsidc.org/arcticseaicenews/>) as well as anomaly maps based on OSI-409.



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**Figure 6:** IcePAC weekly outputs versus OSI-409 SIC observations for the 2015-2016 sea ice season

The OFB point, near Iqaluit, has a behavior that indicates an underlying error. In fact, neither the predictions, nor the observations reach a SIC value of 0 %, which is improbable in our study area. After an inquiry, the reason behind this  
 250 discrepancies with the validation data is that the OSI-409 product does not adequately estimate the SIC in this area, as it can be seen in Fig. (7). The Frobisher Bay is usually ice free around mid-July, which is never the case in the OSI-409 dataset. Possible sources of error include possible “land spill-over effect” (i.e. land contamination) on estimated SIC and the use of the NSIDC sea ice monthly maximum extent climatology as a mask for restricting areas where sea ice is likely to occur. Both of these bias sources are highly plausible explanations of the error detected in Frobisher Bay as the sensor resolution of  
 255 the passive microwave bands used in the OSI hybrid algorithm goes up to 56 km, which is larger than some parts of Frobisher Bay and as the NSIDC climatology maps indicate Frobisher Bay as potentially frozen during the summer month, which is erroneous as a comparison with MODIS data confirms.



260

**Figure 7:** Frobisher Bay OSI-409 error and time concordant MODIS True Color composite

Such errors make it important to emphasize that the results obtained from the model are to be used with care and ideally in combination with other sources of information such as local knowledge, other remote sensing imagery and historical sea ice maps from national sea ice services.

#### 265 4. Analysis of Hudson Bay sea ice spatiotemporal dynamic

The major asset of the IcePAC tool is that its output data gives a probabilistic perspective on relevant sea ice event, comparatively to the usual static descriptive statistics. Therefore, IcePAC gives not only the capacity to determine the mean event, but also to estimate the range of plausible events for a given site and date.

270 Here, the IcePAC outputs are used to assess the sea ice spatiotemporal dynamic given different probability scenarios and in terms of three cover indicators which are the length of the ice free season (or its corollary the ice covered season), the probable complete melt week and the probable complete freeze-up week.

##### 4.1 Analysis with the IcePAC tool

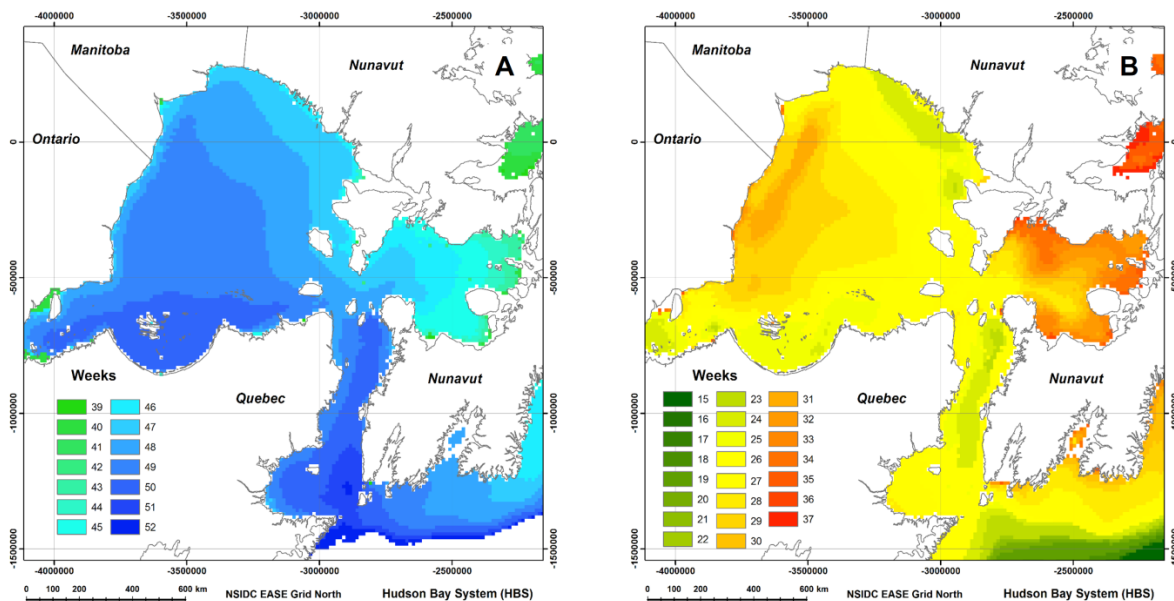
275 Before presenting any results, the ice indicators analyzed in the next paragraphs must be clearly defined. First, the probable complete melt week corresponds to, for varying probability scenarios, to the first week for which the SIC is below 15 % at a given pixel. Second, the probable complete freeze-up week corresponds to, for varying probability scenarios, to the



first week for which the SIC is above 15 % at a given pixel. Finally, the probable duration of the ice free season corresponds to the gap, in weeks, between the different probable meltdown and freeze-up weeks.

The use of the 15 % limit to define presence or absence of sea ice is a convention used by many authors (Andersen et al., 2006; Cavalieri et al., 1997; Cavalieri et al., 1999; Divine and Dick, 2006; Gloersen et al., 1993; Pang et al., 2018) for  
 280 SIC derived with passive microwave data and was therefore used in this analysis.

To estimate the aforementioned ice indicators, the probable SIC value for a given non-exceedance probability ( $p$ ) was extracted from IcePAC for every location and week. For this analysis, a range of non-exceedance probabilities going from 5 % to 95 % were evaluated, with a step of 5 % between each analysis. Time series of the results were compiled and it is on these series that the complete freeze-up and meltdown events were identified. Figure 8 shows the estimated freeze-up and  
 285 meltdown events weeks for  $p = 50\%$ .



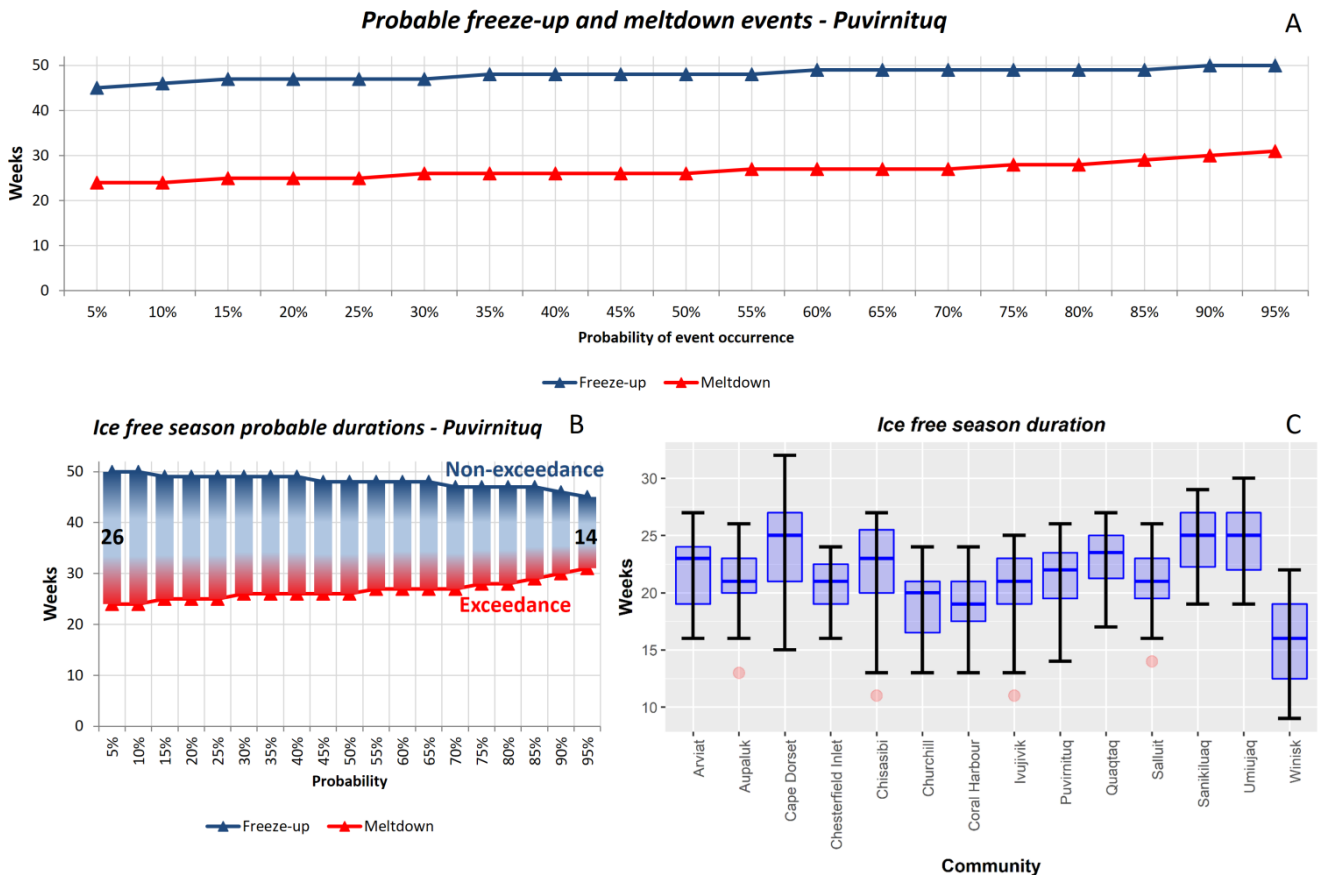
**Figure 8:** Estimated freeze-up (A) and meltdown (B) weeks for  $p=50\%$  in the HBS

This detection process was repeated for every of the 20 738 pixels of the domain and for each probability step. It is  
 290 worth to note that while the meltdown is described as a non-exceedance event ( $SIC < 15\%$ ), directly deducted from  $p$ , the freeze-up is actually defined as an exceedance event ( $SIC > 15\%$ ), deducted from  $1-p$ .

Figure 9A presents the probable freeze-up and meltdown events for the coastal community of Puvirnituk, located in  
 295 the northeastern part of the Hudson Bay. In this figure we notice that meltdown has a 25% probability to be completed by week 25 (Jun. 18<sup>th</sup> to Jun. 24<sup>th</sup>), only a 10% probability of being completed for week 24 (Jun. 11<sup>th</sup> to Jun. 17<sup>th</sup>) and is certain to be completed by week 31 (Jul. 30<sup>th</sup> to Aug. 05<sup>th</sup>). According to the two curves plotted in Fig. (9A), we can state that a



complete freeze-up and complete meltdown at Puvirnituk can be respectively expected, with a very high probability of occurrence ( $p > 95\%$ ) on weeks 50 (Dec. 10<sup>th</sup> to Dec. 16<sup>th</sup>) and 31 (Jul. 30<sup>th</sup> to Aug. 05<sup>th</sup>). Figure 9B shows an assessment of the range of probable ice free season duration made for the same coastal community. Here, the freeze-up curve is inverted (1-P) as the shortest possible ice free season duration is a combination of the latest possible meltdown (high exceedance probability) and the earliest possible freeze-up (high non-exceedance probability). By comparing the space between the two curves for the 5% to 95% probability range, we observe that the shortest possible ice free season at Puvirnituk is of 14 weeks and that the longest is of 26 weeks. Figure 9C shows the ice free season duration estimated for numerous coastal communities located in the study area using the method described for Fig. (9B). Particularly remarkable cases can be noticed such as Cape Dorset, which displays a large variability in possible ice free season duration given the polynya located along the coast in front of the community and Winisk, located in the southcentral part of the Hudson Bay, where ice tends to form early and melt late, which explains the shorter ice free seasons when compared to other communities.



310 **Figure 9:** Probable freeze-up and meltdown events for the coastal community of Puvirnituk (A), range of probable ice free season duration for the same community (B) and probable ice free season lengths for all coastal communities of the HBS (C)



## 4.2 Comparison with the Canadian Ice Service Atlas

The HBS has been under scrutiny for many decades by the Canadian Ice Service (CIS) and they have gathered numerous sea ice conditions maps in the area from which they built a 30 years sea ice climatological atlas (CIS, 2013), portraying the average sea ice conditions, freeze-up and melting dates. As the information provided by the CIS atlas is given based on average values, a comparison with the  $p = 50\%$  IcePAC output gives an outlook on the coherency of the model when compared with another source of data, built around a different methodology. The comparison, as displayed in Table 1, confirms that the freeze-up and melting dates identified by IcePAC are realistic when compared to the CIS historical data. Small differences in weeks are present and may be linked to a multitude of factors, the most important being the different methodologies used to generate the data. The CIS atlas is based on ice maps made by trained sea ice analysts and calculated with descriptive statistics while IcePAC is based on algorithmically generated ice maps and calculated from adjusted frequency models. Also, there is a difference in spatial resolution between the two products; CIS uses mostly Synthetic Aperture Radar (SAR) images at 50 m resolution while IcePAC uses passive microwave images resampled at 12.5 km resolution. Finally, there is a timeframe difference as the CIS atlas is built with data from 1980-2010 (30 years) while IcePAC is built with data from 1978-2015 (37 years).

**Table 1:** Comparison between the CIS atlas and IcePAC  $p = 50\%$  modeled occurrence weeks for the freeze-up and meltdown events at selected sites in the HBS. (F = Freeze-up, M = Melting)

COMMUNITY	CIS (F)	CIS (M)	P=50% (F)	P=50% (M)	DIFF. (F)	DIFF. (M)
Arviat	47	25	47	24	0	1
Aupaluk	48	28	48	27	0	1
Cape Dorset	49	24	49	24	0	0
Chesterfield Inlet	47	25	46	25	1	0
Chisasibi	NA*	27	49	26	NA*	1
Churchill	47	27	47	27	0	0
Coral Harbour	45	26	46	27	-1	-1
Ivujivik	48	25	48	27	0	2
Puvirnituq	48	25	48	26	0	-1
Quaqtaq	48	25	49	25	-1	0
Salluit	49	24	48	26	-1	-2
Sanikiluaq	48	25	50	25	-2	0
Umiujaq	49	25	50	25	-1	0
Winisk	46	30	47	31	-1	-1

\* NA = Not covered by Canadian Ice Service 30 years sea ice atlas



## 330 5. Conclusion

The IcePAC tool permits an assessment of plausible sea ice related events for the entire range of probabilities for 20 738 sites (pixels) in the Hudson Bay System, for all 52 weeks of the year. It is based on local (pixel) models that use the generalized Beta distribution to describe the sea ice behaviour with four parameters at each site (with position and scale being fixed), based on historical 1978-2015 information from passive microwave imagery (OSI-409). From these  
335 parameters, IcePAC generates spatialized sea ice probabilistic information that can be used in any geographic information system or a web based map interface for further analyses.

An analysis has been made to define, for each pixel in the simulation domain, the plausible scenarios for each probability. A subsequent comparison with 1980-2010 Canadian Ice Service sea ice climatology atlas (CIS, 2013) showed that the information generated with the IcePAC tool, for the  $p = 50\%$  case, renders coherent probabilities for freeze-up and  
340 meltdown events all over the HBS. Another analysis, focused the community of Puvirnituq, showed that it is possible to evaluate locally the “worst” and “best” case scenarios in term of ice free season length.

The model outputs generated with the IcePAC tool are bringing a novel probabilistic perspective regarding important events related to sea ice dynamics that was not available before. With its capacity to be transposed to other areas and to be easily updated, the IcePAC tool brings information on probable freeze-up and meltdown weeks as well as on ice presence  
345 and ice free season lengths. This relevant information will help decision makers apprehend the possible scenarios of sea ice dynamics and to prepare for an effective mitigation of climate change impacts on the coastal and offshore environments.

### Author Contribution

CG defined the objectives of this research project, processed the data, coded the IcePAC tool, analysed the results and wrote this research paper. MB and KC participated in defining the objectives, guiding the development of the tool, analysing the  
350 results and writing this paper.

### Competing Interests

The authors declare no conflicts of interest.

### Acknowledgements

The authors would like to acknowledge the Adaptation Platform of Natural Resources Canada (NRCAN) for the funding of  
355 this research (AP060). They would also like to recognize and thank Professor Salaheddine El Adlouni from the University of Moncton, Canada, for his comments and suggestions on this work.





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