

Interactive comment on “Detecting the permafrost carbon feedback: Talik formation and increased cold-season respiration as precursors to sink-to-source transitions” by Nicholas C. Parazoo et al.

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Reviewer 1

Review of “Detecting the permafrost carbon feedback: Talik formation and increased cold-season respiration as precursors to sink-to-source transitions”

The authors ran the Community Land Model (CLM) version 4.5 up to 2300, using RCP 8.5 forcing. They then perform an in depth analysis of permafrost-region dynamics in this simulation, including identifying key events: Talik formation (related the degradation

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of permafrost) and sink-to-source transition, i.e. the point at which the land surface changes from a net sink of carbon from the atmosphere, to a net source. It is interesting to note that this behaviour (starting as a sink and transitioning to a source) is identified across a large fraction of the current permafrost zone. However, the total carbon source is apparently only 11.6 GtC by 2300, which is low compared with previous estimates.

The authors extensively analyse different variables such as thawed volume (a newly defined metric), active layer depth, primary production, respiration and fires, and how these influence talik formation and sink-to-source transition. They find three main drivers of sink-to-source transition: 1. Active layer thickening in cold, carbon-rich high Arctic permafrost 2. Talik formation leading to winter respiration in low Arctic, warmer soils 3. Fire driven carbon source in more productive regions which dry out, and a lot of vegetation is burned. They also showed some indicators of talik formation such as a rapid increase in thawed volume immediately preceding talik formation.

This is a very thorough analysis and a well written paper that will make a great publication in The Cryosphere, after some small revisions. In general I would like to see a bit more analysis about the size of the carbon sources, not just the timing of transition. This could comprise a bit of discussion of the cumulative carbon source (11.6 Gt), and the significance of this - compared to previous estimates, and the time trajectory of the cumulative source (i.e. when does the Arctic as a whole become a net source?). And then if they could break that down to say which of the different types of source (driven by AL, talik or fires) has the bigger contribution to the total source or if these are all comparable magnitude, that would be add some value to the paper. It's fine saying that we should monitor the high Arctic systems as they will become a source soonest, but if this source is likely to be very small, there would not be so much point?

Response:

We thank the reviewer for the astute observation of the small source and excellent recommendation to enhance our discussion. We found an error in our C source budget

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calculation with the new total being 120 Pg C by 2300. We have corrected this error, and included more extensive analysis and discussion of cumulative carbon emissions and sources throughout the paper (see responses below)."

Reviewer: I also suggest considering the soil types at the boreholes. It would hopefully be possible to get an idea of this from a site description or by asking the PI. For example if these are peaty soils that might explain the very slow progression of freeze-thaw compared with CLM (also relatedly, the water content).

Response:

This is a good suggestion. It's difficult to pin down in CLM if mineral soil texture affects thaw rates compared to borehole observations due to the many possible explanation for differences in thaw rates: soil organic content, lateral water flow, surface slope and aspect, ground ice. However, we have looked into soil texture effects at the Alaskan boreholes and find that higher rates of observed soil thawing may be related to 2 factors: (1) relatively dry upper soil at the Gakona and Mould Bay sites, and (2) low surface organic layer and high conductivity of the Barrow2 and Mould Bay soils.

We revised the site descriptions in methods as follows (line 245):

"Mould Bay is a continuous permafrost tundra site with measurements at 63 depths from 0 - 3 m. Mould Bay has almost no organic layer (about 2 cm) and then sandy silt with high thermal conductivity. Barrow is a continuous permafrost tundra site with measurements at 35 depths from 0 - 15 m. The soil at Barrow is represented by silt with a bit of mix with some organics and almost no organic layer on top. Conductivity of the upper layer is ~ 1 W mK⁻¹ for unfrozen and ≥ 2 W mK⁻¹ for frozen soil. Gakona is a continuous permafrost forest tundra site with measurements at 36 depths from 0 - 30 m. Gakona has a thick organic layer of moss (0 to 5 cm), dead moss (from 5 to 13 cm), and peat (from 13 to 50 cm), then silty clay at depth."

We offer an explanation for high observed thaw rates in Section 3.1 (line 359):

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"Overall, we find that simulated patterns of permafrost thermal state change are consistent with available observations, but that the exact thaw rates are uncertain. Although there are many possible explanations for differences in observed and simulated thaw rates, we can attribute high observed thaw rates in part to a combination of (1) relatively dry upper soil at Gakona and Mould Bay, and (2) low surface organic layer and high conductivity of the Barrow and Mould Bay soils. We keep these uncertainties in mind as we examine patterns of change and talik formation simulated into 2300."

We offer some advice for future experiments in the discussion (line 586):

"Controlled experiments demonstrating the sensitivity of talik to parameters that control soil drying such ice impedance or baseflow scalars (e.g. Lawrence et al., 2015), and the effect of organic content and mineral soil texture (Lawrence and Slater, 2008), could provide key insight on soil thermal dynamics in frozen or partially frozen conditions."

We also include an additional column in Table 2 for soil characteristics indicating surface organic layer content and soil type:

Soil Features: Surface organic layer / Soil Type Mould Bay, Canada: Organic layer (~2 cm) / Sandy silt Barrow2, Alaska: Low organic layer / Sandy silt Gakona1, Alaska: Thick organic layer (50 cm) / Silty clay

Reviewer: Finally the paper is rather long and I would suggest reducing in length where possible. I have indicated a couple of points below.

Response:

We removed figures 10 and 11 as suggested by the reviewers, but added a new figure for cumulative carbon sources to address the primary reviewer concerns.

Line-by-line comments

Reviewer: Introduction: L62: 'Shifts in vegetation community' is mentioned in the introduction as being an important factor, but is this considered here? Are you running with

C4

dynamic vegetation? If you are, this should be mentioned and if not, this omission should be discussed later on. L70: Same for soil organic matter export by rivers.

Response:

Since these processes are not considered in this analysis, we added a qualifier statement later on in the discussion (Line 590):

“Other factors affecting soil hydrology and carbon cycling not considered in our CLM4.5 simulations include high spatial resolution in discontinuous permafrost, shifts in vegetation community, lateral flow representation, thermokarst activity and other thaw-related changes to the ground surface, surface slope and aspect, soil heterogeneity, and potentially several other factors (see Jorgenson and Osterkamp (2005) for discussion of some of the many complexities to be considered).”

Reviewer: L86-87: Include more recent references for total permafrost carbon quantity, such as Hugelius et al 2014, *Biogeosciences* (<https://www.biogeosciences.net/11/6573/2014/bg-11-6573-2014.pdf>), and there is also a new paper by Jackson et al coming out in November with revised estimates, this will be in *Annual review of Ecology, Evolution and Systematics* (<http://www.annualreviews.org/doi/abs/10.1146/annurev-ecolsys-112414-054234>).

Response:

References and carbon totals are updated as follows (Line 91):

“Talik as well as longer, deeper active layer thaw stimulate respiration of soil C (Romanovsky and Osterkamp, 2000; Lawrence et al., 2008), making the ~1035 Pg soil organic carbon in near surface permafrost (0-3 m) and ~350 Pg soil organic carbon in deep permafrost (> 3 m) vulnerable to decomposition (Hugelius et al., 2014; Jackson et al., 2017).”

Reviewer: Methods: L138-142 The standard RCP 8.5 only goes until 2100 so presumably some extension is used here? Could you mention what this looks like - for

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example, does it stabilise at some point or does the global temperature and CO₂ just keep increasing? I also think it would be useful for comparing with the permafrost thaw results, to see a plot of the global temperature across the three centuries of future simulation. I suggest adding this at least as a supplementary figure (as there are already a lot of figures in the main manuscript).

Response:

We used the ECP8.5 scenario for the period 2100-2300. We added a time series for air temperature in Figure 1A (shown below) and modified our methodology description as follows (Line 165):

“We use an anomaly forcing method to repeatedly force CLM4.5 with observed meteorological from the CRUNCEP dataset for the period 1996–2005 (data available at dods.ipsl.jussieu.fr/igcmg/IGCM/BC/OOL/OL/CRU-NCEP/) and monthly anomalies added based on a single ensemble member from a CCSM4 Representative Concentration Pathway 8.5 (RCP8.5) simulation for the years 2006-2100 and Extended Concentration Pathway 8.5 (ECP8.5) for the years 2100-2300. Land air temperature for the period 2006-2300, shown in Fig. 1A., is projected to increase steadily in our simulation, with a slight decrease in the rate of warming”

Reviewer: L152-154 “The C source transition represents a shift of ecosystem C balance from a neutral or weak C sink to a long-term source driven by onset of permafrost thaw and respiration of deep SOM” - here you suggest that the deep SOM alone is driving the transition, whereas your analysis suggests that it is driven by different things depending on region. Maybe you can qualify this sentence a bit?

Response:

Great point! We have qualified this sentence as follows (L185):

“The C source transition represents a shift of ecosystem C balance from a neutral or weak C sink to a long-term source as C balance shifts to increasing dominance of

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C source processes including permafrost thaw and fires (Koven, Lawrence and Riley, 2015).”

Reviewer: L 201-204: “For comparison to projected trends in CLM4.5, we recalculate observed trends using the inter-site average from all 9 sites at 3 unique locations: northern Siberia (67°N, 144°E), southwest Siberia (61°N, 115°E), and southeast Siberia (59°N, 131°E).” This is not entirely clear. You were talking about using 6 sites and now it says 9, but then you end up with 3? Can you make it more clear? Did you combine sites into groups based on approximate locations. . .?

Response:

We clarified our analysis as follows (L231):

“To assess observed thaw trends from 1955-1990, we analyze individual sites which report at least 10 months yr-1 of reported monthly mean soil temperature at each layer, and 55 months across the 5 layers (out of 60 possible layer-months per year). Based on these requirements, we find that 6 of 9 sites yield at least 6 years of data over multiple decades, and are well suited for examining historical thaw trends. For comparison of observed trends to historical and projected trends from 1950-2300, we analyze clusters of sites by combining the 9 sites into 3 groups based on approximate locations, and calculate observed trends using the inter-site average at each location. We use 2 sites in northern Siberia (67°N, 144°E), 6 sites in southwest Siberia (61°N, 115°E), and 1 site in southeast Siberia (59°N, 131°E). Site information is shown in more detail in Table 1.”

Reviewer: Results: L228 Do you mean 2300?

Response:

Yes, thank you for identifying this mistake!

Reviewer: L266 “Our simulations show a similar drying pattern in shallow layers (0-1 m depth) in the 4 decades prior to talik onset (Fig. 2D).” The shallow drying does not

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appear to be shown on Figure 2D, only the total column soil moisture?

Response:

We added a line for shallow soil moisture (see attached revision of Fig 2), and revised the text as follows (L07):

“Our simulations show a similar, but very slight, drying pattern in shallow layers in the 4 decades prior to talik onset (1.3% loss of soil moisture over 0-1 m depth; Fig. 2D), accounting for about half of total water storage loss in the column. More significant changes in water balance occur following talik onset, including more rapid drying in shallow layers (~10% over 4 decades) and in the column (~16%), and a substantial increase in sub-surface drainage, as discussed below.”

Reviewer: L281-2 “we find more pronounced tilting of the thawed layer with time and depth” This is not obvious to me from the plot. I might just suggest deleting this.

Response:

“Pronounced tilting” is a misleading description of the thaw pattern. We revised this description as follows (L323):

“In the 3 decades leading up to talik onset, we find gradual deepening of the thawed layer to 3-4 m and penetration of thaw period into Jan-Feb.”

Reviewer: L290-2 “the rate of thawing and drainage in response to permafrost thaw may be under- estimated in deeper CLM4.5 layers near bedrock due to reduced heat capacity.” Sorry if I am missing something here but it doesn’t seem to me like reduced heat capacity would reduce the rate of thawing, but rather than it would thaw more quickly because less heat is needed to thaw? Can you check this? Thanks.

Response:

Yes, thanks for catching this inaccurate statement. We clarified as follows (L332).

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“We note that bedrock soil is not hydrologically active in CLM4.5, and thus the rate of thawing and drainage in response to permafrost thaw may be overestimated in deeper CLM4.5 layers near bedrock due to reduced heat capacity.”

Reviewer: L313-315 “however, our comparison to observations suggests that simulated thaw rates in this region and for similar permafrost temperatures are underestimated”. This is not totally clear. Which comparison with obs? Are you referring to the comparison against thaw rates in Siberia which comes in the following section? Or are you inferring this from your comparison against borehole temperatures?

Response:

We clarified the discussion as follows (L355):

“Thaw onset in CLM4.5 is variable in the region containing Gakona (southeast Alaska) with earliest onset by mid-century (~2050s, Fig. 1A); however, our comparison to borehole temperature data at Gakona suggests that simulated thaw rates in southwest Alaska and across pan-Arctic regions with similar permafrost temperatures are underestimated, and that earliest onset may occur sooner than predicted.”

Reviewer: L335-336 “5 Siberian borehole sites which recorded at least 5 years of data spanning multiple decades.” Earlier you were talking about having 6 sites (or 9, or 3) and here it is 5. Please just clarify this a bit!

Response:

Should be 6, thank you! We corrected as follows (L383):

“We focus first on site-specific long-term historical trends by analyzing the 6 Siberian borehole sites which recorded at least 5 years of temperature data spanning multiple decades: Drughina, Lensk, Macha, Oimyakon, Uchur, and Chaingda.”

Reviewer: L337 “Records at these locations show a decrease in thaw volume” Do you mean an increase? On the next line it also refers to ‘negative trends’, which doesn’t

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seem to fit with the plots/results (or the expectations!). Maybe these things should read ‘increase in thaw volume’ and ‘positive trends’?

Response:

Yes, this should read “increase in thaw volume” and “positive trend”. We corrected as follows (385):

“Records at these locations show an increase in thaw volume with an average positive trend of 0.19 m months yr⁻¹ from 1955 – 1990 (Table 1, Fig 5). All sites except Drughina show positive trends”

Reviewer: L345-346 “(layer thickness increases exponentially with depth along the Siberian tran- C4 sect)” This is not totally clear. Do you mean that active layer thickness increases exponentially with latitude. . .? Is that data shown somewhere? (It doesn’t necessarily need to be, maybe just write ‘data not shown’ if it isn’t)

Response:

It’s not clear why we added this statement, and it doesn’t appear to affect the analysis, so we removed it and modified the text as follows (L390):

“Further examination indicates that active layer thickness at Drughina actually decreased to 0.8 meters from 1989-1990 compared to 1.2 meters in the 1970s (data not shown). Drughina also shows smaller average thaw volume magnitude compared to other sites, consistent with shallower thaw. Together, these findings indicate that active layer thickness is decreasing at Drughina.”

Reviewer: L348-355. I’m not sure how much this is adding overall. It gets a bit confusing. Where you say “(vertical dashed line)”, I would change to ‘(vertical dashed lines on Figure 5)’, assuming this is what you’re referring to? Anyway, it gets confusing when talking about groups of sites and it being hard to identify those groups. I would maybe condense these lines to something along the lines of “There is considerable spatial variability in thaw trends, for example site X is this far from site Y [relatively close] but

C10

with Z difference in trends [relatively large]. Talik formation occurs at several sites, at different times between 1957 and 1990 (shown by vertical dashed lines on Figure 5). We acknowledge the difficulty. . .”

Response:

This section sets up empirical evidence for increasing permafrost degradation from west to east in Siberia. We therefore clarify and condensed this section as follows (L395):

“There is considerable spatial variability in thaw volume and trends, but in general thaw trends increase from west (0.18 m months yr⁻¹) to east (0.51 m months yr⁻¹). Talik forms at several sites, at different times between 1957 and 1990 (shown by vertical dashed lines on Fig. 5), with earlier talik to the west consistent with higher mean initial thaw volumes.”

Reviewer: L364-366 “The simulated trend in thaw volume shows a change in sign at northern locations (blue), acceleration of thaw at southwest sites (orange), and reduction of thaw at the southeast sites (brown)”. This sentence suggests that the thaw volume reduces at the southeast sites, I guess because the thaw volume in CLM is less than the observations, but I would be inclined to interpret this instead as: the CLM simulation always had a too-small thaw volume, and there was never any ‘decrease in thaw volume’ in the simulation. But it is not possible to tell from the plot - Why did you not include the historical CLM simulation on the plot so it would overlap with the observed period? I also wouldn’t say there is a “change of sign” at the northern locations. I guess you refer to the fact that the thaw volume is slightly decreasing at the northern sites historically, and increasing in the future? But as you say this is a very small trend so I would probably instead interpret this as a relatively stable site that shifts to degradation towards the end of the century? I also find figure 6 a bit confusing with the symbols and what they represent. So apparently the circle represents ‘thaw onset in January’ but this happens considerably sooner in the simulation than ‘thaw onset in

C11

March’, which doesn’t make sense to me? Thaw onset should get earlier each year? In the main text it implies this is actually referring to deep thaw lasting throughout the winter (but this is not implied in the figure caption!). Maybe it would just be best not to include the symbols at all. This would make the plot simpler to interpret. Lines 369 “Our simulations show a shift to accelerated soil thaw beginning in the early 2080s”. This sounds like you are referring to the whole simulation and not just this one particular point - can you make this more clear?

Response:

We apologize for the confusion this section created. We added historical thaw simulations from CLM4.5 so the simulated record is continuous from 1950-2300. We also removed the thaw onset symbols, which were meant to represent the progression of thaw later into the cold season (from early winter (Oct-Dec) to deep winter (Jan-Apr)) prior to talik onset. The revised figure 6 is attached, The revised text is shown below (405):

“We recompute observed thaw trends at regional clusters using combined records at the 2 sites in northern Siberia (blue), 6 sites in southwest Siberia (orange), and 1 site in southeast Siberia (brown, Table 1) and compare to historical and projected thaw volume trends in CLM4.5 (Fig. 6). Northern locations show a consistent pattern of low thaw volume (< 10 m month yr⁻¹) and negligible thaw trend (~0 m month yr⁻¹) in the historical simulations and observed record from 1950-2000. Thaw projections in northern Siberia indicate continued stability of permafrost through the early 22st century, followed by a shift to accelerated soil thaw in the early 2120, marked by onset of deep soil thaw late in the cold season.”

“Southern locations show a systematic underestimate of mean thaw volume (< 20 m month yr⁻¹) compared to observations (~40 m month yr⁻¹) from 1950-2000. Simulated thaw trends are negligible prior to 2000, but these likely represent an underestimate given low simulated thaw volumes and significant positive observed trends in

C12

both southeast and southwest Siberia beginning in the 1960s following talik onset (Fig. 5). Thaw projections show more abrupt shifts in thaw volume in the early 21st century in the southwest (~2025) and in the mid 21st century (~2050) in the southeast. The strong discrepancy between observed and simulated thaw and talik onset in southern Siberia warrants close monitoring and continued investigation of this region through sustained borehole measurements and additional model realizations of potential future warming.”

Reviewer: L382 “A total of 6.8 million km² of land is projected to transition”. This is not clear in the abstract which reads like it’s only around 3 million km².

Response:

6.8 million km² refers to all NHL regions, within and outside the permafrost zone. This is clarified in Section 3.3 (L425).

“A total of 6.8 million km² of land is projected to transition, peaking in the late 21st century, with most regions transitioning prior to 2150 (4.8 million km² or 70%, Fig. 7B, solid black). C source transitions which occur in the permafrost zone, accounting for 6.2 million km² of land (91% of all C source transitions), also form talik at some time from 2006-2300 (Fig. 7C). The remaining C source transitions (0.6 million km², or 9%) occur outside the permafrost zone, primarily in eastern Europe.”

We also rephrase the abstract to be more consistent with this section (L29):

“Widespread talik at depth is projected across most of the NHL permafrost region (14 million km²) by 2300, 6.2 million km² of which is projected to become a long term C source, emitting 10 Pg C by 2100, 50 Pg C by 2200, and 120 Pg C by 2300, with few signs of slowing. Roughly half of the projected C source region occurs in predominantly warm sub-Arctic permafrost following talik onset. This region emits only 20 Pg C by 2300, but the CLM4.5 estimate may be biased low by not accounting for deep C in yedoma. Accelerated decomposition of deep soil C following talik onset shifts the

C13

ecosystem C balance away from surface dominant processes (photosynthesis and litter respiration), but sink-to-source transition dates are delayed by 20-200 years by high ecosystem productivity, such that talik peaks early (~2050s, borehole data suggests sooner) and C source transition peaks late (~2150-2200). The remaining C source region is in cold northern Arctic permafrost, which shifts to a net source early (late 21st century), emits 5 times more C (95 Pg C) by 2300, and prior to talik formation due to the high decomposition rates of shallow, young C in organic rich soils coupled with low productivity.”

Reviewer: L388 “followed by gradual decline to 0.5 Pg C by 2300”. Does this suggest the temperature has stabilised and things are moving back towards equilibrium or is it more complicated than that? Could you comment? Including the supplementary plots of temperature trajectories that I suggested earlier might clear this one up.

Response:

Temperatures continue to rise based on ECP8.5 which suggests it is more complicated. We clarify this section, including more detailed analysis of C source magnitudes and categories, as follows (L431):

“Net C emissions from C source transition regions are a substantial fraction of the total NHL C budget over the next 3 centuries (Fig. 8). The cumulative pan-Arctic C source increases slowly over the 21st century, reaching 10 Pg C by 2100 with RCP8.5 warming, then increases more rapidly to 70 Pg C by 2200 and 120 Pg C by 2300 with sustained ECP8.5 warming (Fig. 8, solid black). This pan-Arctic source represents 86% of cumulative emissions in 2300 from the larger NHL talik region (crosses), despite the 2 fold smaller land area, and exceeds the talik region through 2200 due to mitigating widespread vegetation C gains (Koven et al., 2015). Cumulative emissions over all NHL land regions (diamonds, > 55N) increase in similar fashion to the talik region, reaching 120 Pg C by 2200 and 220 by 2300, with no sign of slowing.”

Reviewer: L405-412. Here you are talking about NBP as positive, increasing, but in

C14

the plot (and according to your stated sign convention), a source is represented by negative NBP, decreasing. Please make this paragraph consistent.

Response:

Corrected as follows (L458):

“In these regions, thaw volume is low (< 50 m months yr⁻¹) and shows a weak relationship to NBP (NBP decreases much faster than thaw volume) prior to C source onset (indicated by large green circle in Fig. 8A).”

Reviewer: L437 - Again wrong sign convention for NBP?

Response:

Corrected as follows (489):

“and talik formation occurs when these regions are weak sinks (NBP > 0 g C m⁻² yr⁻¹).”

We have also revised Figure 8 (now Figure 9, attached) with arrows indicating C source or sink for clarification:

Reviewer: L437-438 “In general, C sources in these regions are more sensitive to C emissions from deep soil thaw” Have you actually quantified how much of the C is coming from deep soil. . . ?

Response:

Unfortunately, we can't quantify C from deep soils since vertical resolved C flux output is not available. We have revised the statement to reflect an inferred contribution from deep soils (L490):

“In general, C source onset under high thaw volume indicates these regions are more sensitive to C emissions from deep soil thaw.”

Reviewer: Figure 10: I think this could also be a supplementary figure or removed altogether, maybe giving slightly more detail on the numbers where it's mentioned in

C15

the text.

Response:

We have removed Figure 10 and revised the text as follows (L470):

“A broader analysis of soil thaw statistics over all regions and periods indicates that most C source transitions (~ 2.3 million km², or 77% of land where C source leads talik) occur at active layer depths below 3 m and thaw season penetration into November.”

Reviewer: Figure 11 / Lines 453-461. I am struggling to interpret the upper plot on this Figure, and I am wondering whether this part adds much to the analysis. Since the analysis is already long I might suggest removing this paragraph and figure.

Response:

Agreed. The main point is the difference in GPP between warm and cold permafrost regions, which is discussed in previous sections. We have removed this paragraph and figure.

Reviewer: I suggest doing a bit more quantification of the contribution to total carbon sources. If there is a total of 11PgC emitted by 2300, what fraction of that comes from the three different 'categories' of points in the trimodal distribution? I think it would be really useful to know which are the important carbon sources - or whether they are all similar. (I have made this comment again above)

Response:

We have added a new figure (Figure 8, attached) quantifying the different carbon sources (below). We also found that our cumulative C emission estimate was off by a factor of 10 (didn't convert from year to decade), bringing our new C emission to 120 PgC by 2300.

We also revised the text as follows:

Abstract, to include total and regional contributions:

“Widespread talik at depth is projected across most of the NHL permafrost region (14 million km²) by 2300, 6.2 million km² of which is projected to become a long term C source, emitting 10 Pg C by 2100, 50 Pg C by 2200, and 120 Pg C by 2300, with few signs of slowing. Roughly half of the projected C source region occurs in predominantly warm sub-Arctic permafrost following talik onset. This region emits only 20 Pg C by 2300, but the CLM4.5 estimate may be biased low by not accounting for deep C in yedoma. Accelerated decomposition of deep soil C following talik onset shifts the ecosystem C balance away from surface dominant processes (photosynthesis and litter respiration), but sink-to-source transition dates are delayed by 20-200 years by high ecosystem productivity, such that talik peaks early (~2050s, borehole data suggests sooner) and C source transition peaks late (~2150-2200). The remaining C source region is in cold northern Arctic permafrost, which shifts to a net source early (late 19th century), emitting 80 Pg C by 2300, and prior to talik formation due to the high decomposition rates of shallow, young C in organic rich soils coupled with low productivity.”

The opening of section 3.3 to include total emissions and different categories:

“Fig. 7A plots the decade in which NHL ecosystems are projected to transition to long-term C sources over the next 3 centuries (2010-2300). A total of 6.8 million km² of land is projected to transition, peaking in the late 21st century, with most regions transitioning prior to 2150 (4.8 million km² or 70%, Fig. 7B, solid black). C source transitions which occur in the permafrost zone, accounting for 6.2 million km² of land (91% of all C source transitions), also form talik at some time from 2006-2300 (Fig. 7C). The remaining C source transitions (0.6 million km², or 9%) occur outside the permafrost zone, primarily in eastern Europe.”

“Net C emissions from C source transition regions are a substantial fraction of the total NHL C budget over the next 3 centuries (Fig. 8). The cumulative pan-Arctic C source increases slowly over the 21st century, reaching 10 Pg C by 2100 with RCP8.5

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warming, then increases more rapidly to 70 Pg C by 2200 and 120 Pg by 2300 with sustained ECP8.5 warming (Fig. 8, solid black). This pan-Arctic source represents 86% of cumulative emissions in 2300 from the larger NHL talik region (crosses), despite the 2 fold smaller land area, and exceeds the talik region through 2200 due to mitigating widespread vegetation C gains (Koven et al., 2015). Cumulative emissions over all NHL land regions (diamonds, > 55N) increase in similar fashion to the talik region, reaching 120 Pg C by 2200 and 220 by 2300, with no sign of slowing.”

“The geographic pattern of C sink-to-source transition date is reversed compared to that of talik formation, with earlier transitions at higher latitudes (the processes driving these patterns are discussed in detail below). Overall, the lag relationship between talik onset and C source transition exhibits a tri-modal distribution (Fig. 7D), with peaks at negative time lag (C source leads talik onset; Median Lag = -5 to -6 decades), neutral time lag (C source synchronized with talik onset; Median Lag = -2 to 1 decade), and positive time lag (C source lags talik; Median Lag = 12 decades; red shading in Fig. 7C). Roughly half of these regions (3.2 million km²) show neutral or positive time lag (lag \geq 0). This pattern, characteristic of the sub-Arctic (< 65°N), represents the vast majority of C source transitions after 2150 (Fig. 7B, dotted), but only accounts for 17% of cumulative emissions (20 Pg C by 2300, Fig. 8, dotted). The remaining regions (3.0 million km²) in the Arctic and high Arctic (> 65°N) show negative time lag and account for most of late 21st century sources (Fig. 7B, dashed) and cumulative emissions (95 Pg C by 2300, or 79%; Fig. 8, dashed). C sources in regions not identified as talik (0.63 million km²) either show talik presence at the start of our simulation, or are projected to transition in the absence of permafrost or in regions of severely degraded permafrost (Fig. 7C, dash dotted). This region contributes only 5 Pg C (4%) of cumulative C emissions in 2300.”

The following statements on line 479 referring to high SOM emissions:

“The total area of land in which SOM exceeds 100 kg C m⁻² represents 2/3 of all land where C sources lead talik onset (2.0 million km²), and peaks at a negative time lag of

C18

-5 to -6 decades (Fig. 7D, green bars), which perfectly aligns with the peak distribution of negative time lags. Cumulative C emissions from regions of SOM > 100 kg C m⁻² are also 2/3 of total C emissions (80 Pg C; Fig 8, green).”

Line 495 for fire emissions:

“The regions where fire C emissions exceed 25 g C m⁻² yr⁻¹, representing our threshold for C source transition, are exclusively boreal ecosystems, account for 1/3 of all land with negative lags (~1.1 million km²), and align perfectly with the peak distribution of positive time lags (Fig. 7D, red bars) and cumulative C emissions (20 Pg C in 2300, Fig. 8, red).”

Lines 608 and 635 in the discussion:

“About half of this region (3.2 million km²) shows a pattern of accelerated soil C respiration following talik onset, which shifts the surface C balance of photosynthetic uptake and litter respiration from net C sinks to long term net sources totaling 20 Pg C across 3.2 million km² of NHL land by 2300.”

“We identify an equally large region of land in the high Arctic, representing ~3.0 million km², which is projected to transition to a long term C source much sooner than the sub-Arctic in the absence of talik, and emit 5 times as much carbon by 2300 (~95 Pg C).”

And line 688 in the conclusions

Reviewer: “6.8 million km² of land impacted in Siberia and North America will produce an integrated C source of 90 Pg C by 2100 and 120 Pg C by 2200.” L472 “GPP and total respiration show nearly linear increases (~15% per decade)” Minor point but an increase of 15% per decade would be exponential, not linear.

Response:

Revised as follows (L527):

C19

“GPP and combined respiration increase by ~15% per decade for each permafrost regime surrounding the decade of C source transition with peak fluxes in the growing season (Fig. 11 A - D).”

Reviewer: L485-487 “The trend in soil vs litter respiration explains almost the entire trend in net ecosystem C balance from neutral to net source (Fig. 12 G – H).” I’m not sure what you mean by this. I would have thought the trend in NBP is determined by the difference between GPP and total respiration, rather the difference between soil and litter respiration? And how do you draw these conclusions from the plots? Please could you make this clearer! Thanks.

Response:

We have revised as follows (L541):

“The trend in the respiration difference in warm and cold permafrost, which increase by similar amounts (~100 g C m⁻² yr⁻¹), thus reflects an increasing dominance of respiration from younger and older soil C pools, respectively. These trends are identical to the corresponding NBP trends, which decrease by 100 g C m⁻² yr⁻¹ over the same period from neutral to net source (Fig. 12 G – H), such that the differences between GPP and respiration driving the NBP trends are explained almost entirely by the increasing fraction of soil vs litter respiration.”

Reviewer: L494 NF was only defined at the beginning of the paper and never used again until this point, which means that I had forgotten what it meant by now. You don’t use this abbreviation much so I suggest you don’t need it.

Response:

NF has been changed to non-frozen.

Reviewer: Discussion: L526-529 “Experiments demonstrating the sensitivity of talik to soil drying within the active layer across soil hydrology schemes in previous (CLM4), current (CLM4.5), and newly available (CLM5) versions of CLM could provide key in-

C20

sight on soil thermal dynamics in frozen or partially frozen conditions.” This comes out of the blue a bit, and I am also not clear about why this would be useful? Comparing different hydrology schemes could be useful if one is more realistic than the others or includes different processes? But it is not clear that this would be the case in different CLM versions? It might be better to look at a purpose-built permafrost model, for example, or a model that resolves discontinuous permafrost.

Response:

We agree it's not clear how exactly this would be useful since the model differences in soil hydrology and snow are not systematically differences. We don't think a purpose-built permafrost model is likely to be any better since the processes represented are typically the same as in CLM. We agree that resolving discontinuous permafrost would be good, but it's unclear how to do that other than with increased resolution, but this would miss other important processes such as lateral flow. Something more controlled like a parameter sensitivity study with variable ice impedance or baseflow scalar, such as examined in Lawrence et al., 2015, is likely to provide the best insight. We therefore replace our comment with a more general comment about uncertainties, and suggest more controlled experiments similar to Lawrence (L586):

“Controlled experiments demonstrating the sensitivity of talik to parameters that control soil drying such ice impedance or baseflow scalars (e.g. Lawrence et al., 2015), and the effect of organic content and mineral soil texture (Lawrence and Slater, 2008), could provide key insight on soil thermal dynamics in frozen or partially frozen conditions. Other factors affecting soil hydrology and carbon cycling not considered in our CLM4.5 simulations include high spatial resolution in discontinuous permafrost, shifts in vegetation community, lateral flow representation, thermokarst activity and other thaw-related changes to the ground surface, surface slope and aspect, soil heterogeneity, and potentially several other factors (see Jorgenson and Osterkamp (2005) for discussion of some of the many complexities to be considered).”

C21

Reviewer: L538-540 “thus appears to be driven by combination of warming and increased nitrogen availability resulting from permafrost thaw”. I would suggest changing “appears to be” to “may be”. . . you haven't looked at nitrogen here, and it is not totally clear/agreed that permafrost thaw will increase nitrogen availability.

Response:

We have revised as suggested

Reviewer: L605 “Our main results emphasize an emergence of cold season processes” Not sure what you mean by ‘an emergence of cold season processes’. . .? Rephrase?

Response:

Rephrased as follows (L692):

“Our main results emphasize an increasingly important impact of NHL cold season warming on earlier spring thaw, longer non-frozen seasons, and increased depth and seasonal duration of soil thaw.”

Reviewer: L629-630 “Active-layer deepening leads to C sink-to-source transitions in some re- gions, talik-driven permafrost loss in others,. . .” This should be rephrased,

Response:

Reviewer: “C sink-to- source transitions are caused by active layer deepening in some regions, talik-driven permafrost loss in others,…” Otherwise the sentence doesn't quite make sense, it sounds like active-layer deepening is causing all of the other things!

Rephrased as suggested.

Reviewer: Hope you find these comments helpful. Best wishes!

Response:

Thank you, much improved now!

C22

C23

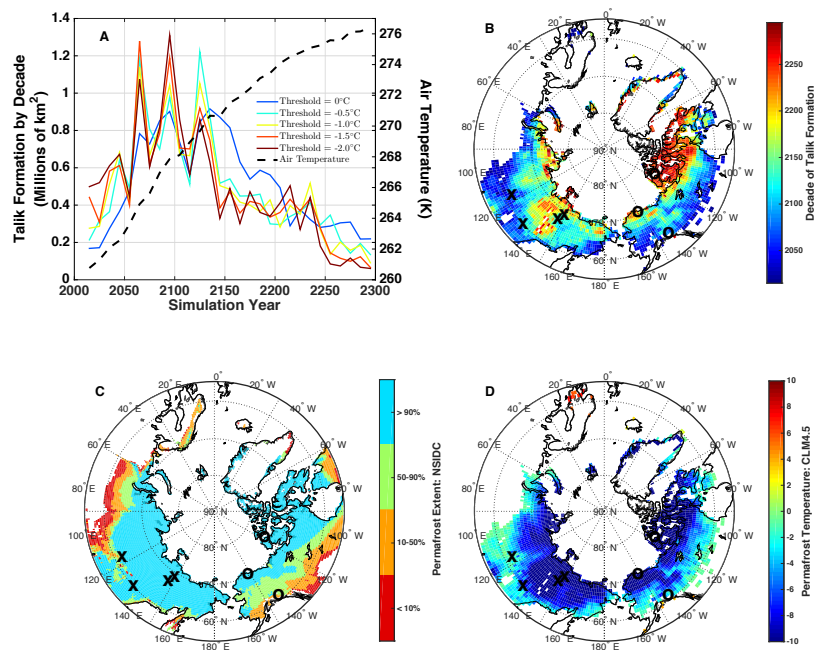


Fig. 1. Figure 1. Decade of projected talik formation and correlation to initial state of simulated permafrost temperature and observed permafrost extent. (A) Time series and (B) map of the simulated decade o

C24

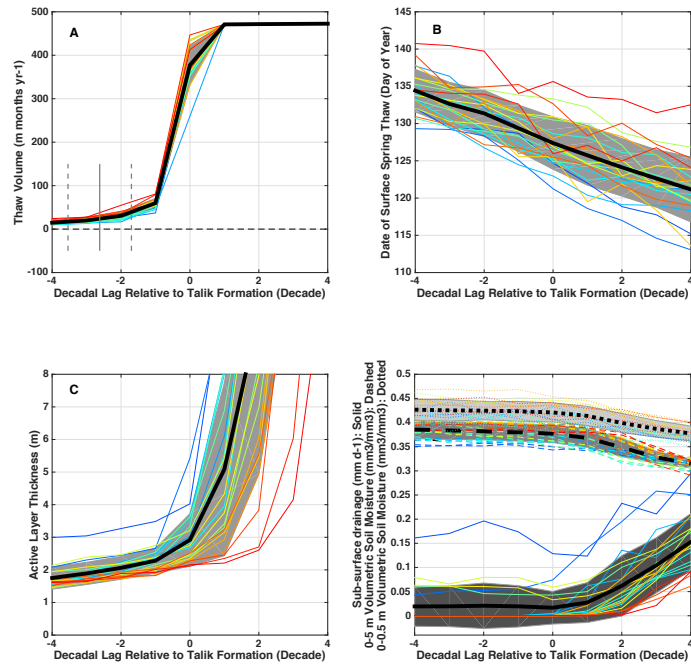


Fig. 2. Figure 2. Patterns showing the progression of soil thaw in the decades surrounding talik onset. Individual lines represent averages across the subset of talik forming regions for each decade from the

C25

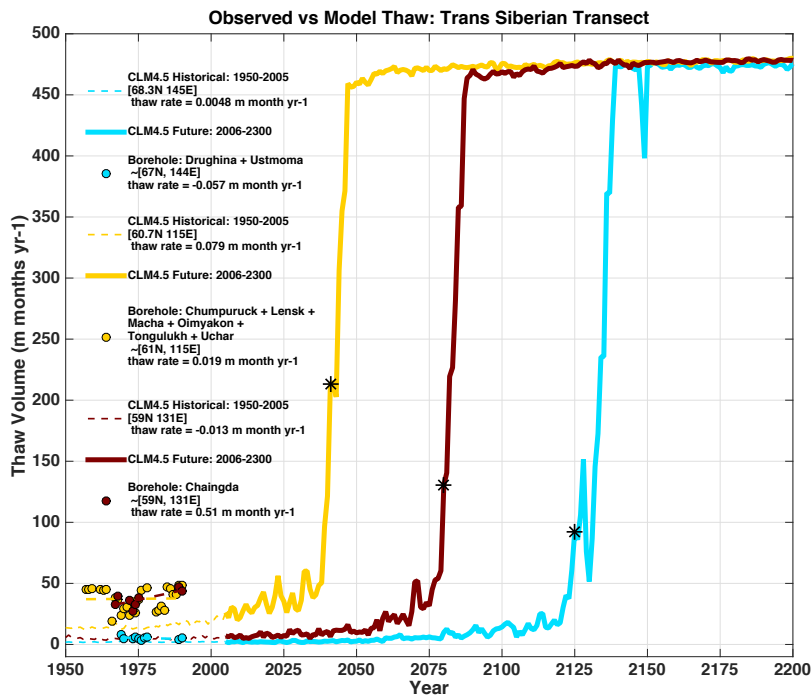


Fig. 3. Figure 6. Comparison of observed soil thaw to historical and future simulations at sites along the East Siberian Transect (crosses in Fig. 1). Observed thaw (filled circles) from 1955-1990 is based on

C26

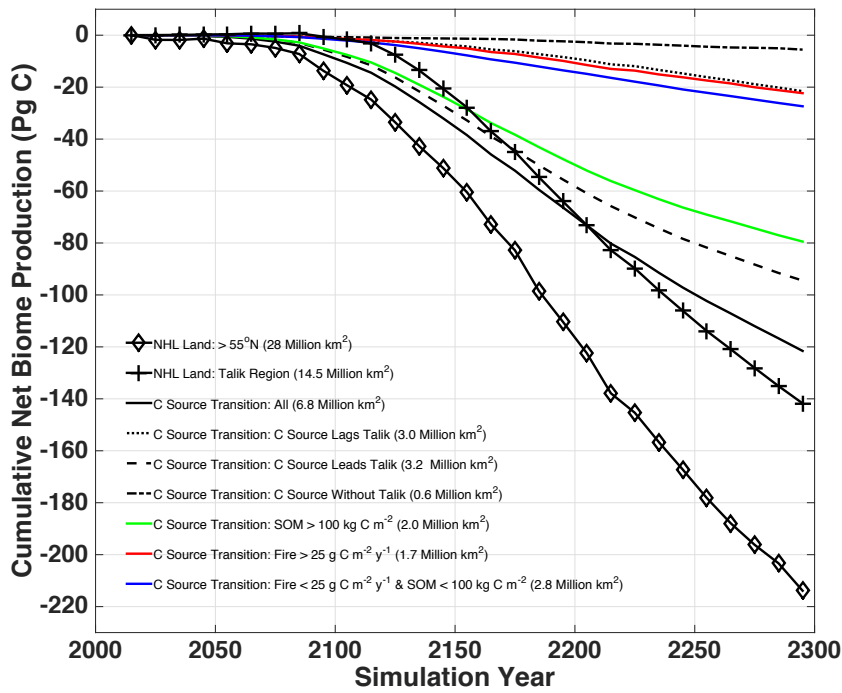


Fig. 4. Figure 8. Cumulative net biome production (NBP) over northern high latitude (NHL) regions ($> 55^{\circ}\text{N}$) from 2010 to 2300, where $\text{NBP} < 0$ is a net source. NHL regions are divided into the following categories

C27

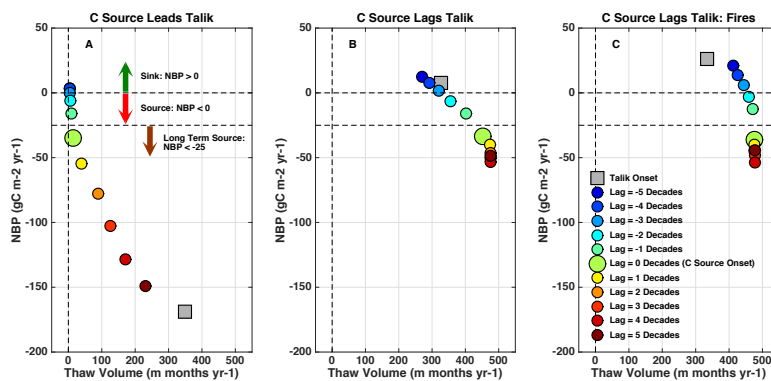


Fig. 5. Figure 9. Net biome production (NBP) as a function of thaw volume. Symbols represent NBP and thaw volume values averaged over regions which transition to long term C source from 2060-2140, binned into

C28