Reply to the comments by Yves Cornet on the manuscript "Snow water equivalent in the Alps as seen by gridded datasets, CMIP5 and CORDEX climate models"

We thank Dr. Yves Cornet for the detailed review of the discussion paper. We have addressed all the points he raised, performing supplementary analyses that in some cases were added in the manuscript or in the new Supplementary material, hopefully improving the overall quality and clarity of the work. We noted that for few comments, the page and lines indicated by the Reviewer do not match exactly the page/lines in the manuscript published on TCD. This was not a problem as we could easily associate the comments to the correct sentences in the text.

Our point-by-point reply (black) to the suggestions and comments of the Reviewer (gray) is reported below.

Research paradigms and hypothesis to be demonstrated

1. “The analysis of snow water equivalent in the Alpine region is thus a very challenging job because of spatial heterogeneity which is not taken into account in the 6 datasets used as reference (2 products derived from satellite observation and the 4 reanalysis), in the GCM and the RCM that have been compared. Nevertheless, regarding “real world” knowledge, figure 2 shows two maps provided by the HISTALP. I tell to the authors why they didn’t use this product as reference. I also tell them why they didn’t qualified HISTALP in a much more detailed way because I think it is a consistent representation of real world than the 6 ones selected as reference.”

The HISTALP gridded dataset provides a limited number of variables, including surface air temperature and total precipitation, and snowfall precipitation estimated from these. It is a good reference dataset but unfortunately it does not provide snow depth or snow water equivalent data, which are the focus of our study. This fact has been better clarified in the text, at P8 L18-20. HISTALP temperature and precipitation climatologies have been shown in the paper and compared to the coarser resolution EOBS dataset to highlight the possible added value that high spatial resolution data can bring (the HISTALP spatial resolution is 0.083° lat/lon against 0.25° lat/lon of EOBS). Moreover, by comparing the two datasets, we highlight that uncertainties do exist also in observational reference datasets, and not only in climate models. Not unexpectedly, the uncertainty in temperature turns out to be lower than that found in precipitation. This is stated at P12 L1-4

2. “Moreover the inter-comparison of the Global and Regional climatic model and the reference datasets without knowing the “real world” evolution and its current situation is somehow disturbing for scientist. As a consequence the use of historical and predicted mean annual cycles from these models seems to me a very critical scientific paradigm which is non-pertinent. To conclude this section, I think that the comparison between models and the so-called references is probably interesting for climatic models developers. The analysis is thus acceptable with the exception of the section dedicated to the future evolution of the snowpack. It’s thus absolutely inappropriate to present it as long as the demonstration of the reliability and the realistic spatial pattern of the SNW output of the models in Mountainous regions is not made. So, the question of major interest to be answered before this predictive operation with dangerous interpretative issues is the enhanced knowledge of the snowpack from finer observations by elaborating spatially representative sampling plan of the phenomenon and developing measurement methods enabling it to be implemented.
We agree that at the present state of knowledge, i.e. “without knowing the real world”, projections of future snow depth are speculations. But the aim of this study is neither to state how snow water equivalent will evolve in the future nor to provide indications of the future state of snow resources. Instead we aim to show (i) how the uncertainty/spread found in the historical period project into the future, to assess the overall agreement on the relative changes with respect to each modelclimatology, and (ii) discuss whether the magnitude of the relative snow changes is similar in coarse and fine scale models. The spatial resolution is one order of magnitude finer in RCMs than in GCMs so that high elevation areas are resolved better in the former. We had two main questions in mind: (1) how does the resolution affect snow depth representation and its future changes in the Alpine environment? And, (2) is there any specific feature emerging in higher resolution projections, or are they indistinguishable from the lower resolution ones? This investigation is corroborated by a recent study comparing “bias corrected” and “non-bias corrected” snowfall projections of EURO-CORDEX RCM models (Frei et al. 2017). In that study bias corrected RCM snowfall was constrained to a snowfall reference dataset derived from 2 km resolution gridded temperature and precipitation data. According to that analysis, the relative change (RCP8.5 vs baseline) of the mean September-May snowfall is comparable whether applying or not the bias correction and the bias adjustment does not seem to have any significant effect on the trend. We added in the introduction (P4 L1–6) a sentence that states in a clearer way what the main purposes of this study are.

Methodological issues

Weighting procedure in the computation of RMSD, normalized variance and Pearson correlation.

- This procedure is described at p. 7. You assign a weight to each grid value given by the ratio between the area above 1000 m elevation and the area of the grid cell. You should give some arguments to justify that threshold and also to convince me that it is valid whole over the GAR.
- Further in the text (p. 12 l. 3, p. 13 legend caption) you explain the procedure in another way. I think that you should correct that to remain coherent through the whole paper to avoid ambiguity.

The weighting procedure mentioned by the reviewer in the first item above is applied only when snow water equivalent fields are spatially averaged over the Greater Alpine Region, i.e. in the plots shown in Figures 6 and 7, and not in the Taylor diagrams. The spatial averages of Figures 6 and 7 are intended to be representative of the mountains only, so we exclude the areas below 1000 m a.s.l. (we recognize that this threshold is arbitrary but we think that it could be appropriate in the GAR for focusing on high-altitude regions only) using this weighting procedure. The detail of the procedure is as follows: we weigh the snow water equivalent values at each grid cell by the area of the grid cell with mean elevation higher than 1000 m a.s.l. using a Digital Elevation Model at high spatial resolution (1 km), then the weighted values are spatially averaged over the domain of interest, the Greater Alpine Region. This is better explained in the manuscript at P10 L13-19.

For the Taylor diagrams, we calculated the root mean square error (RMSE), normalized standard deviation (NSD) and the correlation coefficient (R) over the full domain, without applying any weighting based on elevation. In this case the multiannual mean snow water equivalent was simply remapped onto the target grid conserving the snow mass from the original and the target grid cells. To this end we used a standard function incorporated in the CDO (Climate Data Operators) software mostly used in the climate community to handle climate model data in netCDF format. The CDO “remapcon” function performs an area-weighted remapping where the interpolation weights are based on the fractional area overlap of the original and target grid cells, following Jones, 1999. This methodology has been explained at P10 L6-12.


- You should also provide a map in figure 1 () for instance with the spatial variation of this weight (area ratio).
Actually the “map of weights” is resolution-dependent so we should provide a map of weights for each dataset considered in this study. These maps would show for each coarse scale grid cell “the fraction of the area of the grid cell with mean elevation higher than 1000 m”, where the topography is taken from the GLOBE Digital Elevation Model at 1 km resolution. This procedure has been better explained at P10 L13-19. We report below (Fig R1) two examples of maps of weights, the former referring to the CFSR reanalysis and the latter to the EC-Earth GCM. We prefer to not provide the maps of weights for each model in the manuscript because we think that this level of detail is too high for the general purpose of the paper.

Fig. R1. “Map of weights” showing the fraction of each grid cell at elevation above 1000 m a.s.l. for the CFSR reanalysis and the global climate model EC-Earth. The reference topography is taken from the 1 km digital elevation model GLOBE (Hastings and Dunbar, 1999).

• - p. 7 l. 9-11 “This procedure allows for a fair comparison between datasets characterized by different spatial resolutions, without introducing uncertainties due to regridding”. I don’t totally agree with you. This procedure will enhance the importance of high area in the computation of “quality” parameters (Taylor diagrams). But high mountain zones are also very heterogeneous as low mountains zones. So the resolution difference effect will persist!!!

This comments reveals a misunderstanding. As explained before the weights are not used to calculate quality parameters (Taylor diagrams) for which we use the full domain. Instead, we used the weights approach to spatially average datasets characterized by very different spatial resolutions over the same domain (GAR above 1000 m a.s.l.), without interpolating the model data. The sentence has been rephrased in the manuscript at P10 L9-11.

Interpolation. In your paper you use several words to describe the mathematical procedure used to change grid resolution (interpolation, reshaping, downscaling, remapping ...). When you reduce the ground sampling distance interpolation is the right terminology. When you degrade the resolution increasing g the ground sampling distance you perform a generalization of your geographical data and I think that this is a spatial aggregation method. You write at p. 10: “To provide a fair comparison of the models and reduce the impact of the horizontal resolution on their performances, in particular on their spatial variance, each GCM is then compared to the MRM after having remapped each individual reference dataset onto the individual GCM grid, so that the reference is reshaped each time according to the model resolution. This approach allows for a fair comparison also for low resolution models.” Despite this statement I think that you should describe in a more detailed way the methods used to “reshape” your grids giving more information about “mass conservation” condition that should be verified.

We remapped the six reference datasets onto each climate model grid using a conservative remapping, in detail the CDO remapcon function (CDO 2015). This function performs an area-weighted remapping, where
the interpolation weights are based on the fractional area overlap of the source and the destination grid cells, following Jones, 1999. Such interpolation weights applied to the source field allow to conserve the fluxes or water budgets from the source to the destination grid. This procedure has been better explained adding some details at P10 L18-19.

CDO 2015: Climate Data Operators. Available at: http://www.mpimet.mpg.de/cdo

Mean (central position statistics) computation. You compute the average of the references (satellite and Reanalysis products).

- The number of observations (6) is very small and one of these observations is obviously an outlier (20CR). Why did you this outlier?

Thank you for this comment. We agree that after the assessment of its poor performance in representing SNW climatology, the 20CR reanalysis should not be considered as a “reference”. We repeated all the analysis excluding the 20CR from the Multi-Reference-Mean, and consequently figures 5,6 and 7 have been updated in the main text. The new procedure is explained at P12 L31-34;

- This computation is performed on non-independent observation. For instance, it is quite clear for Global SWE Climatology and CFSR. So you give an exaggerated weight to those to datasets!!!

The interdependency of the Global SWE Climatology and the CFSR snow outputs has been better clarified in the text (P6 L2-3; P12 L10-17 and following). Both products integrate, but to different extents, the Special Sensor Microwave Imager (SSM/I) data. The Global SWE Climatology is specifically derived from Special Sensor Microwave Imager (SSM/I) data. The CFSR snow output is mainly based on the Noah land surface model first guess, and a daily snow analysis based on several inputs - among others Special Sensor Microwave Imager (SSM/I) data – is used to constrain the model first guess (Meng et al., 2012). In detail, CFSR snow depth/SWE are limited in the upper and lower boundaries by the snow analysis (it cannot be larger than twice and lower than half the snow analysis) but the temporal evolution of snow depth and SWE is determined by the Noah model. In conclusion, as the similarity between the two datasets is in the similar range of variability, we decided to include both.

Same comments about the computation of the MMM (35 GCM or 7 HiRes GCM with ground sampling distance smaller than 1.25°).

- GCM are probably not independent (see table 1) and some are probably highly correlated and have exaggerated weight!

We totally agree that the climate models are not independent from one another, and several previous studies (e.g. Knutti et al., 2013; Sanderson et al., 2015) focus on this issue. For example in Figure R2 below, now included in the Supplementary Material as Fig. S02, we report the spatial distribution of the DJFMA SNW in the 8 GCMs with horizontal resolution not coarser than 1.25°, referred to as “high-resolution” HiRes GCMs in the manuscript. These high-resolution models are CMCC-CM, EC-Earth, MRI-CGCM3, BCC-CSM1-1-M and four models from the CESM-family. Out of the four CESM-family models, one, CESM1-CAM5, shows a distinct behaviour. The other three (CESM1-BGC, CESM1-FASTCHEM and CCSM4) present very similar SNW patterns (Figure R2, last row), and similar RMSE, NSD and correlations values (Figure 5b in the main text). In order to have a model ensemble including models as independent as possible, we consider in the ensemble mean (MMM-HiRes) only one out of the three (CESM1-BGC). The analyses of figures 6 and 7 are then based on 6 high-resolution models, namely CMCC-CM, EC-Earth, MRI-CGCM3, BCC-CSM1-1-M, CESM1-CAM5 and CESM1-BGC.
In Figure 5a MMM-HiRes refers to the multi-model-mean of these 6 models. This choice has been explained in text at P13 L5-8 and later on at P18 L14-22.

Concerning GCMs with spatial resolution coarser than 1.25° it is difficult to evaluate their degree of interdependence from the Taylor diagrams. However, owing to their overall poor performances in the representation of SNW, and not being the focus of the paper, the aspect of their interdependency is not investigated further (explained in the main text at P18 L20-22).

![Maps of GCMs with spatial resolution coarser than 1.25°](image)

Fig R2: Multiannual mean (1980-2005) snow water equivalent in the GCMs with spatial resolution finer or equal to 1.25°.

What is the statistical pertinence of this aggregation method to determine central position statistics? Why a mean computation to “compensating extreme behaviours” (p. 10 l. 25) ? Why don’t you compute the median (for instance) in this particular case (few and not independent observations, outliers).

We agree that it is interesting to explore the case in which the median is used as metrics, instead of the mean. In order to address this comment we explored two different approaches:

1) We considered as “reference” the median of the 5 datasets (NSIDC, CFSR, MERRA, ERAI/Land and 20CR), we calculated the median of CMIP5 models (full ensemble and HiRES ensemble) and we repeated the analysis of the Taylor diagram. The results are shown in Figure R3. The median is shifted towards very small SNW values as NSIDC, CFSR and 20CR provide low snow. Consistently the normalized standard deviation of the MERRA reanalysis exceeds 2.5 and that of ERA-Interim lies outside the range of the plot, as well as for many climate models.

2) we consider as “reference” the average of 4 datasets (NSIDC, CFSR, MERRA, ERAI/Land), hence excluding the 20CR reanalysis from the “reference” statistics, as suggested by both reviewers. In this case we have a well-balanced ensemble, with NSIDC, CFSR showing low snow and MERRA, ERAI/Land showing large snow amounts. The results are shown in the new Figure 5a of the revised manuscript.
Considering the results of the two approaches we think that this second metric is the most appropriate to describe the “ensemble behaviour” of the reference datasets and, therefore, it has been reported in the manuscript.

Absence of spatial (geographical) analysis of the differences between the various spatial grids. To compare the grids you use 3 “quality” parameters reported in the Taylor diagrams. Even if Pearson-r is a measure of the association between variables and allows a global comparison of spatial patterns, I think that a spatial (or geographical) analysis of residuals (or the differences) is recommended to understand the effect of spatial localization. Doing so you should be able to improve the discussion of some climatological factors that are not integrated in the same way in the models and related to air mass circulations for instance: North-South of the Alps – humid and cold air-mass flow from the North or East-West - continentality and humid air-mass flux coming from the Adriatic towards South Eastern Alpine and Pre-alpine domain) on eventual systematic and variable biases.. This spatial analysis should thus be done for some specific and well-chosen models.

Here the referee states that “To compare the grids, the spatial (or geographical) analysis of residuals (or the differences) is recommended”, but actually this information is already contained in Figures 2,3,4 which show the differences between the various datasets (reanalysis, RCMs, GCMs) and the reference climatologies (EOBS for precipitation and temperature, NSIDC for snow water equivalent). Particularly important are, in our opinion, temperature and precipitation biases, that clearly show, for a given model, possible weaknesses related to the representation of the air mass circulation. These plots are commented in the corresponding section 4.1.1-4.1.3

Some specific comments

p. 1 l. 2. I’m very surprised about your conception of high spatial resolution in this abstract and in the whole text. For remote sensors (you use satellite data) hectometric and higher ground sampling distance corresponds to low and very low spatial resolution which don’t allow any description of bio-geo-physical
processes on the Earth surface characterised by very high spatial frequency that are typical of mountainous area and especially the spatial variability of snow cover characteristics.

We agree that the definition of high resolution depends on the context. In the abstract P1 L2 (“high resolution, regional, observation-based gridded datasets”), “high resolution” refers to the typical spatial scales at which snow processes occur, i.e. less than 1 km. This has been clarified in the text (P1 L2) Later on, when speaking about the resolutions of global climate models, the concepts of “high” and “low” resolutions refer to the typical horizontal grid size of the state-of-the-art numerical climate models (CMIP5), ranging from 70 to 400 km. In this case “high resolution GCMs” are those with grid size equal or finer than 1.25° (about 125 km). We added in the introduction (P3 L3-5) a sentence to clarify these definitions.

p. 1 l. 20 “The shift of the 0_C isotherm to higher elevations ...” Is it demonstrated overall on the GAR?

Yes, because of an overall increase of surface temperatures (see i.e. Gobiet et al., 2014; Hantel et al., 2012; Serquet et al 2011; Beniston, 2003). We added the references in the text at P2 L7

p. 1 l. 22 “...decrease in the solid-to-total precipitation ratio in low- and mid-altitude mountain areas.”

What do you mean by low and mid-altitude? Does that definition depends on the climatological sub-domain within the GAR?

At this point we are presenting a general picture, not focused on the Alps, and with “low and mid-altitude” actually we intend “areas with temperatures closer to the melting point”. We have better specified this in the text at P2 L5-8, thank you.

p.3 l. 14 What do you mean by large scale? The notion of scale in your document is somehow perturbing for cartographers and geographers that are specifically doing multiscale spatial analysis (see also p. 3 l. 14, p. 16 l. 1 for instance)! A map with a scale of 1:100000 is a large scale map that allows the representation and analysis of local physical phenomenon with small autocorrelation distance (high spatial variability). At the contrary a map with a scale of 1:1000000 is a small scale map that allows the representation and analysis of global phenomenon.

This sentence was present in a preliminary version of the paper but it has been removed in the version published on the online TCD http://www.the-cryosphere-discuss.net/tc-2016-280/tc-2016-280.pdf. Interestingly, "large-scale" and "small-scale" have opposite meaning in cartography (as pointed out by the reviewer) and in climate/geophysical fluid dynamics, where the large scales are those with the largest spatial extent and the small scales are those with smaller spatial extent. Curious discrepancy (in fact, opposite meaning) of terms in two neighbouring fields of research.

p. 6 l. 31-32 “Global climate models, also the most spatially resolved ones, do not take into proper account elevations above 1500 m a.s.l. over the GAR.”

It’s really a critical issue because it seems that “a very weak increasing trend towards heavier snowfalls has persisted since the 1960s” until 1999 in the Swiss Alps above the altitude of 1300 m as demonstrated by LATERNSER and SCHNEEBELI (2003, DOI: 10.1002/joc.912), for instance. But this research emphases the snow cover extent using low spatial resolution AVHRR images and you correctly state that satellite products provide a reliable picture of snow cover extent which is not the case for snow depth or snow water equivalent (p. 2 l. 20 and 21).

Yes, moreover, the period over which those trends are calculated (1931-1999) does not consider the last 17 years, generally characterized by low snow.
We chose the ERA-Interim Land grid as it has intermediate resolution between RCM and GCM grids. This explanation has been added in the methodology.

It has been corrected in the text, thank you.

An alternative approach has been devised to provide a fair comparison of the GCMs. Each GCM is compared to the MRM after having conservatively remapped each reference dataset onto the individual GCM grid, so that the reference is reshaped each time according to the model resolution. This approach allows for a fair evaluation of the GCM at the model's grid, regardless of its resolution.

In the document the concept of resolution is confused with that of Ground Sampling Distance!

Thank you. We have corrected it in the manuscript. In climate models, “resolution” refers to the physical distance (meters or degrees) between two consecutive gridpoints, in latitudinal, longitudinal or vertical direction, on the grid used to compute the equations (IPCC, 2013)

Yes, we changed into “non-mountainous midlatitude regions”, thank you

Comments on the document’s form (text, units, figures ...)

Units must be controlled:

- p. 3 l. 16 “~80° km spatial resolution”??
- p. 8, l. 20 “105 kg/m3 is not consistent with the unit used to describe the SNW in the reference datasets and the GCM (figure 2 for instance→kg/m2). It seems that you did this unit conversion to compute the mean annual cycle (figure 5 p. 13)

p. 9 figure caption 2 “... with horizontal resolution higher than 1.25°.” → “... with horizontal resolution finer than 1.25°”

p. 9 figure 2 “Panels (j,k) report the multiannual mean of the DJFMA accumulated snowfall derived from the HISTALP dataset. You should give a precision about the unit. I guess the unit in mm refers to the water equivalent volume per area unit! This value could be expressed using the same unit than the reference datasets and the GCM (kg/m2) assuming that 1 mm corresponds to 1 l/m2 ~ 1 kg/m2 !

p11. figure 3

- Labels are not readable even for points corresponding to cases with large NSD !
- At that scale the large amount of points near the origin must be grouped in one class with a legend identifying all the point (dates or model) in the cluster.
- Colour is not the best graphical variable to identify the signification of the point reported in the legend and it is probably not necessary. If points are grouped combine in one class and write the composition of the class. If points are dispersed then then label is sufficient!!!

All the above technical comments have been accepted and modified in the manuscript accordingly. Thank you.