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**Inconsistency in Precipitation Measurements across Alaska and Yukon  
Border**

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18 **Abstract**

19 This study quantifies the inconsistency in gauge precipitation observations across the border of Alaska  
20 and Yukon. It analyses the precipitation measurements by the national standard gauges (NWS 8-in gauge  
21 and Nipher gauge), and the bias-corrected data to account for wind effect on the gauge catch, wetting loss  
22 and trace events. The bias corrections show a significant amount of errors in the gauge records due to the  
23 windy and cold environment in the northern areas of Alaska and Yukon. Monthly corrections increase  
24 solid precipitation by ~~136~~135% in January, 20% for July at the Barter Island in Alaska, and about 31% for  
25 January and 4% for July at the Yukon stations. Regression analyses of the monthly precipitation data  
26 show a stronger correlation for the warm months (mainly rainfall) than for cold month (mainly snowfall)  
27 between the station pairs, and small changes in the precipitation relationship due to the bias corrections.  
28 Double mass curves also indicate changes in the cumulative precipitation over the study periods. This  
29 change leads to a smaller and inverted precipitation gradient across the border, representing a significant  
30 modification in the precipitation pattern over the northern region. Overall, this study discovers significant  
31 inconsistency in the precipitation measurements across the US and Canada border. This discontinuity is  
32 greater for snowfall than for rainfall, as gauge snowfall observations have large errors in ~~the~~ windy and  
33 cold conditions. This result will certainly impact regional, particularly cross borders, climate and  
34 hydrology investigations.

35

36 **Key words:** snowfall, national precipitation gauge, measurement errors, bias correction,  
37 precipitation gradient and distribution.

38

## 39 1. Introduction

40 It is known that discontinuities in precipitation measurements may exist across the national boundaries  
41 because of the different instruments and observation methods used ([Nitu and Wong, 2010](#); [Sanderson,  
42 1975](#); [Sevruk and Klemm, 1989](#); [Yang et al., 2001](#)). For instance, the National Weather Service (NWS) 8-  
43 inch gauge is used for precipitation measurements in the United States (U.S.), and the Nipher snow gauge  
44 has been used in Canada for decades. Different instruments have also been used in various observational  
45 networks within the same country. In the synoptic network, the Type-B rain gauge and Nipher gauge are  
46 the standard manual instruments for rain and snow observations in Canada ([Mekis and Vincent, 2011](#);  
47 [Metcalf and Goodison, 1993](#)), and recently the Geonor automatic gauges have been installed([Nitu and  
48 Wong, 2010](#); [Sanderson, 1975](#); [Sevruk and Klemm, 1989](#); [Yang et al., 2001](#)). For instance, the National  
49 Weather Service (NWS) 8 inch gauge is used for precipitation measurements in the United States, and the  
50 Nipher snow gauge is the standard instrument over Canada. Different instruments have also been used in  
51 various observational networks within the same country. The Type B rain gauge and Nipher gauge are the  
52 standard instruments for rain and snow observations in Canada, respectively ([Mekis and Vincent, 2011](#);  
53 [Metcalf and Goodison, 1993](#)), and recently the Geonor gauges have been installed at the synoptic  
54 stations across Canada.

55 Instruments also change over time at most operational networks, resulting in significant breaks in data  
56 records. It has been realized that combination of regional precipitation records from different sources may  
57 result in inhomogeneous precipitation time series and can lead to incorrect spatial interpretations ([Yang et  
58 al., 2005](#)). Efforts have been reported to examine the [precipitationP](#) discontinuity within a country  
59 ([Groisman and Easterling, 1994](#); [Sanderson, 1975](#)). [Leeper et al., \(2015\)](#)~~Leeper et al (2014)~~ found that the  
60 US COOP stations reported slightly more precipitation overall (1.5%) with network differences varying  
61 seasonally. The COOP gauges were sensitive to wind biases, particularly over winter when COOP  
62 observed (10%) less precipitation than the [U.S. Climate Reference Network \(USCRN\)](#).~~,-~~ Conversely,  
63 wetting and evaporation losses, which dominate in summer, were sources of bias for USCRN. [Mekis and](#)

64 [Brown, \(2010\) developed adjustment method to link the Nipher gauge and ruler snowfall measurements](#)  
65 [over Canada](#) ~~Yang and Simonenko, (2013)~~~~Yang and Simonenko (2013)~~ compared the measurements  
66 among 6 Russian Tretyakov gauges at the Valdai experimental station, and reported the differences of less  
67 than 5-6% for the study period. These results are useful to determine the homogeneity of precipitation  
68 data collected by a standard gauge within the national and regional networks.

69 Many studies show that the national standard gauges, including the Canadian Nipher, and US 8-inch  
70 gauges, under measure precipitation especially for snowfall ([Goodison, 1981; Goodison et al., 1998;](#)  
71 [Yang et al., 1995, 1998a, 1999](#))~~(Goodison, 1981; Goodison et al., 1998; Yang et al., 1995, 1998a, 1999).~~  
72 Compatibility analysis of precipitation measurements by various national gauges suggests little difference  
73 (less than 5%) for rainfall observations, but a significant discrepancy (up to 110%) for snowfall  
74 measurements (Yang et al., 2001). For instance, the experimental data from Valdai show that the U.S. 8-  
75 inch gauge at Valdai systematically measured 30-50% less snow and mixed precipitation than the  
76 Canadian Nipher gauge (Yang et al., 2001). This difference in national gauge catch has introduced a  
77 significant discontinuity in precipitation records between the U.S. and Canada borders, particularly in  
78 windy and cold regions. Differences in the snow measurements across the US and Canada border has also  
79 been noticed in other studies as a problem to produce gridded products and to develop [precipitation input](#)  
80 [for basin hydrological investigations](#) ([Šeparović et al., 2013; Zhao et al., 2010](#)).

81 ~~P input for basin hydrological study~~ (~~Šeparović et al., 2013; Zhao et al., 2010~~). Although [Yang et al.](#)  
82 ~~(2001)~~~~Yang et al. (2001)~~ compared the relative catch of many national standard gauges, little has been  
83 done to address the inconsistency of precipitation records across the national borders. This is an  
84 important issue, since most regional precipitation data and products have been compiled and derived from  
85 the combination of various data sources, assuming these data and observations were compatible across the  
86 borders and among the national observational networks. [Simpson et al., \(2005\) studied temperature and](#)  
87 [precipitation distributions over the State of Alaska and west Yukon, and documented precipitation](#)  
88 [increase from north to south. They also report differences in mean monthly precipitation across the](#)

89 Alaska-Yukon border, i.e. about 5-15 mm in central-east Alaska and 15-40 mm in central-west Yukon.  
90 (Jones and Fahl, 1994) found a weak gradient in annual precipitation across the AK-YK border, including  
91 the headwaters of the Yukon River. Other studies also discuss precipitation distribution and changes over  
92 the arctic regions (Legates and Willmott, 1990; Serreze and Hurst, 2000; Yang et al., 2005).

93 The objective of this work is to examine the inconsistency in precipitation measurements across the  
94 border between Alaska and Yukon. We analyze both gauge-measured and bias-corrected monthly  
95 precipitation data at several climate stations across the border, and quantify the changes in precipitation  
96 amounts and patterns due to the bias corrections. We also calculate the precipitation gradients across the  
97 border, and discuss precipitation distribution for the warm and cold seasons. The methods and results of  
98 this study are useful for cold region climate and hydrology investigations and applications.

99

## 100 **2. Study Area, Data, data and Methods**

101 The study areas include the northern and central regions of Alaska (AK) and Yukon (YK). We choose 5  
102 climate stations across the Yukon and Alaska border, which use the national standard gauges (NWS 8 in  
103 gauge and the Canadian Nipher gauge) for precipitation observations (Figure 1). These stations can be  
104 classified ~~into~~ 2 groups. The first group, 3 stations about 150 km apart, is the northern region along the  
105 coast of the Beaufort Sea; with the Barter Island station in Alaska and Komakuk and Shingle ~~Point~~  
106 stations in Yukon. The second group is in the central part of the region, ~~i.e.~~ the Eagle station in Alaska  
107 and Dawson station in Yukon, about 130 km apart.

108 The three northern stations selected for this study are located north of the Brooks Range. The approximate  
109 distances to the mountain edge are 100 km for the Barter Island station, 90 km for Shingle Point station,  
110 and 150 km for the Komakuk station. Both stations in Yukon are along the shore line and the station in  
111 Alaska is an island site, very close to the coast line. The altitudes of the stations range from 7 to 49 m  
112 a.s.l. According to Manson and Solomon, (2007), the summer storm tracks are usually from the northwest

113 coming from the open water in the Beaufort Sea and are the greatest contributor to annual precipitation.  
114 The storms are obstructed by the Brooks Range once moving inland. The weather patterns in the  
115 surrounding of the stations might be affected by the mountains, but the stations are not separated by the  
116 Brooks Range. Given this setting, it is expected to see little impact of mountain range on the precipitation  
117 process and distribution along the relatively flat coast line.

118 These stations have been operated by the NWS and Environment Canada (EC) since the early 1970's. The  
119 observations have been done according to the national standards of US and Canada. The detail  
120 information for these stations are given in Table 1, such as the location, period of measurement used for  
121 this work, instrument types for precipitation observations, and a climate summary for yearly temperature,  
122 precipitation ~~(P)~~, and wind speed. —————

123 Yang et al. (2005) have developed a bias corrected daily precipitation dataset for the northern regions  
124 above ~~45°N~~45N. The source data are acquired from the National Centers for Environmental Information  
125 (NCEI), i.e. ~~Climatic Data Center, i.e.~~ a global daily surface data archive for over 8,000 stations around  
126 the world (<https://www.ncdc.noaa.gov/data-access/quick-links#ghcn>). (~~http://www.ncdc.noaa.gov/cgi-~~  
127 ~~bin/res40.pl~~). To focus on the high latitude regions, a subset of the global daily data, about ~~45,000~~  
128 stations located north of ~~45°N~~45N with data records longer-than 20 years during 1973-2003 has been  
129 created. Yang et al. (2005) applied a consistent procedure derived from the WMO Solid Precipitation  
130 Intercomparison (Goodison et al., 1998), using wind speed, temperature, and the precipitation as inputs  
131 (Yang et al., 1998b, 2005), ~~at all the stations over the high latitude regions~~. They quantify the  
132 precipitation gauge measurement biases for the wind-induced undercatch, wetting losses, and trace  
133 amount of precipitation. For the US stations, wind data from the standard height was reduced to the gauge  
134 level of the NWS 8-in gauge (standard height is 1 m). Wind speeds and directions were measured at the  
135 Canadian climatic network; the same approach was applied to estimate the wind speed at the gauge height  
136 (standard height is 2 m) on precipitation days. The corrections were done only for those stations with  
137 wind observations. Unfortunately there are many stations in the US without wind information and this is a

138 ~~challenge to gauge bias corrections. This study uses the updated monthly precipitation, temperature and~~  
139 ~~wind speed data from Yang et al. (2005) for the selected AK and YK stations. The data periods range~~  
140 ~~from 7 to 10 years for the stations, but long enough to examine P patterns in these regions.~~

141 This study uses the updated (until 2013) monthly precipitation, temperature and wind speed data from  
142 Yang et al. (2005) for the selected AK and YK stations (Table 1). The selected data periods range from 7  
143 to 10 years for the stations that are considered long enough to examine precipitation patterns in these  
144 regions. Missing records affect regional climate data analyses. In this study, a threshold of 0°C of  
145 monthly temperature has been used to determine the cold and warm months for snow and rain. Mixed  
146 precipitation has not been classified separately. The frequency of missing values was calculated when the  
147 bias correction was made in Yang et al., (2005). For any month with less than 20 days (~30%) of  
148 measurements, it is excluded from data analysis. Statistical methods to compare the measured and  
149 corrected monthly and yearly precipitation data across the selected border station pairs is used to analyze  
150 these data. It also carries out regression analysis on monthly precipitation records, and calculates the  
151 cumulative precipitation amounts to derive the Double Mass Curves (DMC) over the study period. The  
152 double mass curve (DMC) is a useful tool to evaluate the consistency of observation records over space  
153 and time (Searcy and Hardison, 1960). Some typical issues of observations that DMC can identify,  
154 include changes in the station location, and instruments or sensors. A reference station is needed for DMC  
155 analyses. In this study, the DMC has been applied without a reference station to mainly detect any shifts  
156 between the observed and corrected precipitation. Through the data analyses and comparisons with other  
157 studies, we document the spatial and temporal variations of bias corrections across the border stations.  
158 We also determine the precipitation gradients across the border, and examine the changes, due to the bias-  
159 corrections of the US and Canadian gauge data, in precipitation distributions on both seasonal and yearly  
160 time scales.

161

162 ~~This study applies statistical methods to compare the measured and corrected monthly and yearly~~  
163 ~~precipitation data across the border station pairs. It also carries out regression analysis on monthly P~~  
164 ~~records, and calculates the cumulative P amounts to derive the double mass curves over the study periods.~~  
165 ~~Through the data analyses and comparisons with other studies, we document the spatial and temporal~~  
166 ~~variations of bias corrections across the border stations. We also determine the precipitation gradients~~  
167 ~~across the border, and examine the changes, due to the bias corrections of the US and Canadian gauge~~  
168 ~~data, in precipitation distributions on both seasonal and yearly time scales.~~

169



### 3. Results

Based on the analyses of the measured precipitation ( $P_m$ ) and corrected precipitation ( $P_c$ ) data, this section presents the results on the bias corrections of monthly and yearly precipitation for each station~~the stations~~, regression and correlation of monthly precipitation~~P~~ data between the stations, and cumulative precipitation via the double mass curves for the warm (monthly temperature  $> 0^\circ\text{C}$ ) and cold seasons (monthly temperature  $< 0^\circ\text{C}$ ).

#### 3.1. Monthly data and corrections

The monthly mean precipitation~~P~~ and bias corrections are illustrated in Figure 2 for the northern group during the corresponding observation period (Table 1). In Figure 2, the missing data percentages are also presented for each month. Barter Island had the lowest percentages of missing data, about 2% as a maximum monthly mean in December. The mean missing percentages for the Komakuk station was about 5% (in May), with the maximum month in July 1984 (16%). For Shingle Point, the mean missing values were 11% for both April and May, with the maximum (26%) in April 1979. Given the small percentages of missing records, its impact is insignificant on monthly mean and yearly precipitation calculations. Figure 2 shows that annual precipitation cycle was centered on August, with an approximate maximum  $P_m$  around 40 to 80 mm between August and September. This maximum was coincident with the monthly mean maximum temperature in the area (around  $10^\circ\text{C}$ ).

For the Barter Island station in AK, the corrections were~~are~~ variable through the months. The monthly corrections increased~~increase~~ the  $P_m$  amount by 3-~~31~~34 mm for snow to 4-9 mm for rain. The relative increases were~~are~~ 59-136% for snow and 20-41% for rain, with a monthly mean of 9 mm (or ~~76%~~78%). The relative changes were~~are~~ usually large for months with low  $P_m$  and small for months with high precipitation. In other words~~word~~, the monthly correction amounts do not always matched~~match~~ with the percentage changes, i.e. a small correction in a dry month can have a large percentage change.

193 It is important to note that gauge measurements at Barter ~~showed~~ show the maximum precipitation~~P~~ in  
194 August ~~and October~~, but the peak shifted to October due to the corrections; i.e. the mean monthly  $P_c$  in  
195 October ~~were~~ 98~~is~~—100% (about 32mm~~70mm~~) more than the  $P_m$  (Figure 2). Closer  
196 examination~~Examination~~ of the monthly precipitation~~P~~ time series for Barter Island (Figure 3) indicated→  
197 indicates that, for most of the years, October was~~is~~ the most significant contributor to the total annual  
198 (232~~1~~% for  $P_m$  and 222~~5~~% for  $P_c$ ). However, there were~~are~~ some years in the study period with the  
199 maximum  $P_m$  in other months; for example, the highest  $P_m$  in 1982 was in September, as documented by  
200 Yang et al., (1998b). Climate data and analyses showed~~Yang et al. 1998. Climate data and analyses show~~  
201 the highest wind speed (4.5 m/s) and cold temperature (about -9°C) for October, indicating higher  
202 undercatch by the US standard gauge for snowfall. On the other hand, the wind speed showed~~shows~~ the  
203 minimum values in July and August (3.3 m/s), coincident with the highest temperatures (4.6 and 4 °C)  
204 (Figure 2). Due to the combination of warm temperatures and low wind speeds, the corrections for  
205 summer months were~~are~~ the lowest at this station (20-27%).

206 For the Komakuk Beach station in Yukon, the corrections increased~~increase~~ the precipitation by 0.7-5.5  
207 mm (or 14%-34%) for snow and 1-2.6 mm (4%-10%) for rain, with a total monthly mean change of 2.64  
208 mm (14~~19~~%) (Figure 2). The monthly maximum precipitation~~P~~ was in August, i.e. 48mm~~47mm~~ and  
209 50mm, respectively, for the  $P_m$  and  $P_c$ . The monthly minimum precipitation~~P~~ was in March, i.e.  $P_m = 4.2$   
210 mm and  $P_c = 5$  mm. For this station, the~~These~~ extremes remained in~~remain~~ the same month after the bias  
211 corrections. The wind speed had~~has~~ the minimum value in Aug. (3.1 m/s) and Sept. (3.2 m/s), and max in  
212 Dec. (4.3m/s) and Jan (4.7m/s). The temperatures were~~are~~ highest in July (6.9°C) and Aug. (5.8°C), and  
213 lowest in Feb and Mar (-25 °C). Given this climate condition, the corrections were~~are~~ lower in the  
214 summer months (mean of 6%) and higher in winter (mean of 23%).

215 The monthly corrections for the Shingle Point station in Yukon ranged~~range~~ from 1-7.63 mm (3%-157%)  
216 for rain to 1-8.24 mm (14%-28%) for snow, with the monthly mean correction of 4.23~~5~~ mm (14~~16~~%).

217 The ~~month of~~ maximum precipitation ~~was in~~ Aug., about 73-76 mm (or 20% of the annual total) (Figure  
218 2). The minimum ~~precipitation was in~~ ~~P~~ was in Feb. with ~~9.2 mm for the measured P; and it shifted to~~  
219 March with ~~9.8 mm for P<sub>m</sub>; and~~ 11 mm for ~~P<sub>c</sub>; the corrected values~~. The monthly wind speeds ~~were~~  
220 generally higher in winter and lower in summer, with the maximum in Feb. (4 m/s) and minimum in May  
221 (2.7 m/s). The temperatures ~~had~~ have a common annual cycle with the maximum in July (11°C) and the  
222 minimum in Feb. (-24.3°C). Because of the higher wind speeds and cold temperatures in the cold months,  
223 the corrections ~~were~~ are greater for the winter season.

224 It ~~was~~ is necessary to compare the correction result across the border in order to quantify the effect of  
225 biases in gauge observations on precipitation analyses, such as ~~precipitation~~ ~~P~~ distribution and seasonal  
226 patterns. The mean snowfall corrections ~~were~~ are about ~~96~~ 100% for Barter Island in ~~Alaska~~ AK and  
227 around 22% for both Shingle Point and Komakuk stations in Yukon; while the rainfall corrections  
228 ~~were~~ are approximately 32% for Barter and ~~7~~ 6% for the two Yukon stations. Bias corrections also  
229 ~~demonstrated~~ demonstrate a clear shift in the ~~maximum precipitation~~ ~~max-P~~ timing for the Barter Island,  
230 but no change for the Yukon stations. This remarkable contrast across the border ~~was~~ is caused mainly by  
231 the difference in gauge types and their catch efficiency. Many experimental studies have shown that the  
232 Canadian Nipher snow gauge catches more snowfall relative to the US gauge (Goodison et al., 1998;  
233 Yang et al., 1998b). For instance, the mean catch ratios for snowfall ~~were~~ are about 40% and 85% for 4  
234 m/s wind speed, respectively, for the NWS 8-in unshielded and Nipher gauges ([Figure 4](#)) ([Yang et al.,](#)  
235 [1998b](#)). (~~Yang et al. 1998, Figure 4~~).

236 For the central group, the maximum and minimum  $P_m$  ~~were~~ is in July and March for the Eagle station  
237 (Figure 5). The corrections did not modify the timings of maximum and minimum ~~amounts; they~~  
238 ~~remained in~~ ~~P~~; July for the maximum ( $P_m=67$  mm and  $P_c=70$  mm), and ~~in~~ March for the minimum ( $P_m=3$   
239 mm and  $P_c=4$  mm) ~~precipitation.~~. The correction ~~increased~~ increases the precipitation by 0.6-1.8 mm  
240 (8%-22%) for snow and 1-3 mm (5%-10%) for rain, with a monthly mean correction of 1.7 mm (12%).

241 The annual temperature cycle for Eagle ~~showed~~shows warmer temperatures ~~relative to~~than in the northern  
242 station, ~~with the maximum of~~around 16.2°C ~~and with temperatures~~ above 0°C ~~during~~from April to mid-  
243 October. Eagle ~~had lower~~has variable wind speeds around 1 m/s (Figure 5).

244 For Dawson station, precipitation ~~was~~is more homogeneous throughout months; varying from 10 mm to  
245 50 mm in October and June, respectively. Another relative maximum occurs in January with  $P_m=38$  mm  
246 (Figure 5). The precipitation correction ~~was~~is small and ~~fluctuated~~fluctuates from 0.3 to 1 mm (or 2%-  
247 4%) for snow and 0.4-1.3 mm (3%-4%) for rain. This small correction ~~was~~is due to the lower undercatch  
248 correction for the Nipher gauge, besides the warmer temperatures and lighter winds. The temperature  
249 annual amplitude ~~was~~is between 16°C in July and -25°C in January, with ~~April to September~~  
250 temperatures above 0°C ~~from April to September.~~ Wind speeds ~~showed~~show a clear annual cycle with  
251 the maximum in May (~~of~~1.6 m/s); and lighter winds in winter months, with ~~the~~ minimum ~~of~~0.4 m/s in  
252 January (~~0.4 m/s).~~

253 The temperature and wind conditions ~~were~~are similar between the Eagle and Dawson ~~stations~~regions,  
254 with ~~the~~mean temperature around 1°C and wind speed of 1m/s. ~~The missing data percentages were also~~  
255 ~~similar for Eagle and Dawson stations; less than 3% for most months, with the maximum of 10% in May~~  
256 ~~2006 for Eagle and 20% in September 2009 for Dawson. The~~ ~~But the~~ bias corrections ~~were~~are quite  
257 different, with the mean corrections of ~~16~~15% for snow and ~~7~~6% for rain at Eagle, and about 2 % ~~and 3%~~  
258 for both rain and snow at Dawson. ~~Overall, the~~The Eagle correction ~~was~~is four times greater ~~at Eagle~~ than  
259 that ~~at~~for Dawson. This discrepancy reflects again the catch difference between the US and Canadian  
260 standard gauges.

261 In order to understand the effect of ~~precipitation~~P bias corrections on regional climate around the AK-YK  
262 border, it ~~was~~is useful to examine and compare the temperature and precipitation features between the  
263 northern and central regions. The monthly mean temperature threshold of 0°C ~~did~~does not occur exactly  
264 at the same time among the 2 groups; the warm months (above 0°C) ~~were~~are between June and

265 September in the north group and between April and September in the central group. Although both  
266 regions ~~had~~have similar mean minimum temperatures, around -24°C and -27°C, the maximum  
267 temperature ~~was considerably lower~~is ~~lower~~s in the north part, ~~with the~~ average of 8°C in the north group  
268 vs. 16°C for the central region. ~~Additionally~~Besides, the monthly mean wind speed ~~was~~is higher for the  
269 northern region, 4 m/s vs. 1 m/s. Therefore, because of the colder temperatures and higher winds in the  
270 northern region, the bias corrections ~~were~~are higher in the north relative to the central region.

### 271 3.2. Yearly data and corrections

272 ~~The Figure 6 shows the~~ annual  $P_m$  and  $P_c$  time series for 11 years ~~during 1978-1988~~ in the northern group  
273 ~~is presented in Figure 6. There was almost no missing data for the whole period, except 3% for 1978.~~ At  
274 the Barter Island station in Alaska, the yearly  $P_m$  ~~ranged~~ranges from 114 mm to 211 mm, with the long-  
275 term mean of ~~155~~157 mm. The mean annual corrections ~~ranged from~~are ~~about~~ 67 ~~to~~ -138 mm, with a  
276 long-term mean of 101 mm (or 65%).~~%~~. The  ~~$P_c$  corrected  $P$~~  records ~~varied~~vary from 181mm to 343 mm.  
277 The maximum precipitation was in 1985 for both  $P_m$  and  $P_c$  (211 mm and 343 mm, respectively). The  
278 minimum precipitation was in 1983 for the  $P_m$  and  $P_c$  (114 mm and 181 mm, respectively).

279 For Komakuk Beach station in Yukon, the  $P_m$  ~~ranged~~ranges from ~~103 mm~~103mm to 306 mm, ~~with~~ the  
280 ~~missing data between 0 and to 7% among the years. The bias~~ corrections ~~increased~~increase the  
281 precipitation by ~~13 mm~~13mm to ~~45 mm~~45mm (or 8-19%). The long-term mean ~~was~~is about ~~194~~197 mm  
282 for  $P_m$  and ~~220~~223 mm with the corrections. The maximum ~~precipitation occurred  $P$  was~~in 1981, 306  
283 mm and 347 mm,~~respectively~~; for  $P_m$  and  $P_c$ ,~~respectively~~. The minimum ~~precipitation~~ $P$  was in 1988 for  
284 both the  $P_m$  and  $P_c$ , 103 mm and 123 mm, respectively.

285 For Shingle Point station in Yukon, yearly  $P_m$  ~~varied~~varies from ~~126 mm~~127mm to ~~551~~566 mm ~~and~~; the  
286  ~~$P_c$  ranges from 138 to 638~~corrections are 139-88 mm. The mean annual total precipitation ~~was~~is about  
287 ~~302~~306 mm for  ~~$P_m$  the gauge data~~ and ~~341~~345 mm after the corrections (change of ~~13~~12%). The high and

288 low extreme years were 1981 ( $P_m = 551 = 566$  mm,  $P_c = 638654$  mm), and 1988 ( $P_m = 126$  mm,  $P_c = 138$   
289 mm). Shingle station had missing data from 2% in 1983 to 10% in 1979. ( $P_m = 127$  mm,  $P_c = 139$  mm)

290 Figure 7 displays the mean annual precipitation in cold and warm seasons for the northern group. The  
291 According to the gauge measurements showed, the mean annual values  $P$  in this region fluctuates from  
292 155444 mm at Barter Island, 194 mm, 103 mm at Komakuk to 302566 mm at Shingle Point, i.e. The  
293 gauge data suggest a strong precipitation increased  $P$  increase from the west to the east, particularly  
294 between Komakuk Beach and Shingle Point. However, the corrected data ( $P_c$ ) showed show a different  
295 pattern (Figure 7), i.e. higher precipitation  $P$  at Barter than Komakuk, so the gradient across the border  
296 changed the sign and magnitude. This change was caused mainly by the high correction corrections at  
297 the Barter station, particularly for snowfall data during the cold months (Fig. 2).-

298 For the central group, the annual results are shown for 8 years (2006-2013) in Figure 8. The  $P_m$  ranged  
299 from 66 to 391 mm at the Eagle, and the bias corrections were 5-27 mm, correspondingly, which on  
300 average increase the total precipitation by 7%. While at Dawson, the  $P_m$  ranged from 158 to 333 mm, and  
301 the adjustments were from 4 mm to 10 mm, with an average increase in yearly precipitation by 3%. The  
302 gauge data showed a slight increase (12 mm) of mean precipitation from west to the east, i.e. slightly  
303 higher  $P$  in Yukon relative to Alaska. This result is consistent with other studies (Simpson et al., 2002,  
304 2005). The corrected data, on the other hand, suggest a smaller gradient (1 mm) across the border (Figure  
305 9). This change was mainly due to the higher corrections for the US 8-inch gauge at Eagle.

306 For the central group, the results are shown for 8 years (2006-2013) in (Figure 8). The annual  $P_m$  ranges  
307 from 100 to 400 mm at the Eagle, and the corrections are 7-27 mm, or 6-9%, which on average increase  
308 the total precipitation by 7%. While at Dawson, the measured  $P$  ranges from 158 to 353 mm, and  
309 adjustments are 4 mm to 11 mm, with an average increase in yearly precipitation by 3%. The gauge data  
310 show a slight increase (22 mm) of mean  $P$  from west to the east, but the corrected data suggest a smaller

311 ~~gradient (11mm) across the border. This change is mainly due to the higher corrections for the US 8-inch~~  
312 ~~gauge at Eagle (Figure 9).~~

313 Similar to the monthly results, the northern stations ~~exhibited~~~~exhibit~~ higher yearly corrections for  
314 snowfall and rainfall measurements relative to the central group. This ~~was~~ because of higher winds in  
315 the northern stations, i.e. yearly mean wind speeds of 3.8 m/s in the north group and 1 m/s in the central  
316 group. This windy and snowy environment in the north ~~produced~~~~produce~~ higher wind-loss for the  
317 snowfall measurements by the gauges, which ~~was~~ the largest errors in precipitation records in the high  
318 latitudes (~~Benning and Yang, 2005; Yang and Ohata, 2001; Yang et al., 1998b~~)(~~Benning and Yang,~~  
319 ~~2005; Yang and Ohata, 2001; Yang et al., 1998b~~). It is important to note that gauge measured and bias  
320 corrected data ~~showed~~~~show~~ different pattern in seasonal and yearly ~~precipitation~~~~P~~ in the northern region.  
321 In other words, bias corrections of gauge measurements alter the ~~precipitation~~~~P~~ gradient in the northern  
322 areas; this change ~~was~~ mainly due to the difference in the catch efficiency between the US and Canadian  
323 standard gauges. The corrections for the US gauge snow measurements ~~were~~~~are~~ much higher than the  
324 Canadian gauge, particularly in the cold and windy coastal regions.

### 325 **3.3. Regression analysis of monthly data**

326 The scatter plots of corresponding monthly precipitation for the two stations across the border and  
327 between the ~~two~~~~2~~ Yukon stations in Canada are illustrated in Figure 10. For the cold season (Figure  
328 10.A), the gauge data ~~showed~~~~show~~ more snowfall at ~~Barter~~~~Barter~~ for most years. Regression analysis  
329 ~~suggested~~~~suggests~~ a weak relationship, with  $R^2=0.34$ . The corrected data ~~showed~~~~show~~ a similar  
330 relationship, but a shift in the regression line, indicating a greater ~~precipitation~~~~P~~ difference over the cold  
331 season across the border. For the warm season (Figure 10.B), the gauge data ~~showed~~~~show~~ higher  
332 ~~precipitation~~~~P~~ at the Komakuk station, and the regression ~~suggested~~~~suggests~~ a ~~much~~ stronger relationship.  
333 The ~~corrected data revealed~~~~Pe—reveals~~ a closer relationship between these two stations,  
334 ~~proposing~~~~suggesting~~ a smaller gradient for the warm months.

335 The scatter plot between the two stations in the Yukon Territory ~~showed~~show higher precipitationP at  
336 Shingle point for both cold and warm seasons. It also ~~gave~~gives another point of view about the effect of  
337 the correction in this area. Relative to the cold months (Figure 10.C), the corrections ~~were~~are smaller for  
338 the warm months (Figure 10.D), and ~~there is a better~~correlation improved ( $R^2=0.72-0.76.75). However,  
339 the relationship ~~did not~~doesn't change much in both cases between the measured and corrected data. This  
340 wasis because very small amount of corrections ~~for~~due to the lower wind conditionswinds and higher  
341 catch efficiency of the Canadian Nipher gauge.$

342 For the central group, the scatter plot between Eagle and Dawson stations ~~illustrated~~illustrates a clear  
343 difference in precipitation amount for the cold and warm months (Figure 10.E-F). The cold months  
344 ~~showed~~show more precipitationP at Dawson, particularly for the wettest events, while Eagle ~~did~~does not  
345 show any comparable amount. The correlation wasis weak, and insignificant ( $R^2 =0.13$ ). The shift in the  
346 fit line between measured and corrected data wasis also very small. The warm months ~~showed~~show low  
347 precipitation at Dawson; a different pattern from the cold months. The regression wasis better,  $R^2$   
348 ~~=0.5958~~, with a smaller shift due to the corrections.

349 Overall, we ~~obtained~~obtain consistent results among the Alaska and Yukon stations. The correlations  
350 ~~were~~are higher in warm months ( $R^2 = 0.58$  to 0.7675) and lower for the cold season ( $R^2$  between 0.13 and  
351 0.52). This result may suggest that the rainfall wasis more homogeneous over the regions in summer, and  
352 greater difficulty and errors in snowfall measurements during the cold months.

### 353 3.4. Cumulative precipitation via double mass curves (DMC)

354 The DMC plot for Barter Island and Komakuk Beach ~~showed~~shows more  $P_m$  at Komakuk than Barter  
355 (Figure 11.A). The bias corrections ~~led~~lead to a shift of the relationship with a significant increase in the  
356 total precipitationP amount at ~~Barter~~Barterter. Relatively, the total cumulative precipitation for Barter  
357 Island ~~increased~~increases by 65% after the correction and by 1413% at Komakuk. The difference between  
358 the two stations at the last cumulative point (December 1988) is 426440 mm for  $P_m$ , and 393380 mm for



359 | P<sub>c</sub>. This shift ~~represented~~represents a modification in the precipitation difference between these stations,  
360 | i.e. a change in the gradient's direction (Figure 7).

361 | The comparison of cumulative precipitation values between Shingle Point and Komakuk, both in Yukon,  
362 | is illustrated in Figure 11.B. Shingle Point ~~showed~~shows more cumulative precipitation at the end of the  
363 | period (~~P<sub>m</sub>=33223348~~ mm vs. ~~P<sub>m</sub>=21152144~~ mm for Komakuk). Although the relationship ~~was~~is more  
364 | homogeneous between these stations, there ~~was~~is a break in the records around ~~13001000~~ mm for  
365 | Komakuk, maybe associated with changes in instruments or sensors. Examination of the station history  
366 | and information revealed an anemometer issue around the critical time that was fixed by August 1980.  
367 | This may affect wind data and thus the corrected precipitation values. Both stations ~~showed~~have increases  
368 | in total cumulative ~~precipitation by 13%. P by 3%, i.e. a change in precipitation difference from 1204 mm~~  
369 | ~~to 1352 mm between Shingle and Komakuk over the study period (2006-2013).~~

370 | The central stations ~~showed~~show a greater amount of P<sub>m</sub> in Dawson (~~20652202~~ mm) than in Eagle  
371 | (~~19732027~~ mm) over the study period. Bias corrections ~~changed~~change the total ~~precipitation~~P by ~~37%~~  
372 | and ~~73%~~ for ~~Eagle and Dawson~~ and Eagle, respectively, -resulting in a shift in the DMC (Figure 11.C),  
373 | particularly for the last period of time, to ~~21232265~~ mm in Dawson and to ~~21162173~~ mm in Eagle. This  
374 | shift also ~~represented~~represents a slightly smaller precipitation difference between the two stations.  
375 | During Eagle and Dawson. In the 8 years, the cumulative difference ~~decreased~~goes from ~~175 mm to 92~~  
376 | ~~mm to 7.3 mm, over the study period.~~

377 | In summary, the DMC for measured and corrected precipitation ~~showed~~show that the main change ~~was~~is  
378 | due to the difference in their corrections (Figure 11); the north stations ~~showed~~show a greater change  
379 | compared with the central group. The P<sub>c</sub> ~~showed~~shows in all the cases a smaller precipitation difference  
380 | between the two countries. This smaller difference ~~led~~leads to a decrease in the ~~precipitation~~P gradient  
381 | across the border. This result implies that existing precipitation climate maps and information derived

382 | from gauge measurement without bias corrections may over-estimate the precipitation~~P~~ gradient in these  
383 | regions. This overestimation will affect regional climate and hydrology analyses.

384

#### 385 | 4. Summary and Discussion~~diseussion~~

386 | This study documents and quantifies the inconsistency in precipitation measurements in the northern and  
387 | central regions of Alaska/Yukon, with a focus on ~~the~~ station pairs across US-Canada border. ~~–~~The  
388 | monthly bias corrections show large~~a significant amount of~~ errors in the gauge records due to the windy  
389 | and cold environment in the northern areas of Alaska and Yukon. The corrections for gauge undercatch  
390 | increase the snowfall by 136~~135~~% in January for ~~the~~ Barter Island station in Alaska. For the Yukon  
391 | stations, the increase is about 31~~34~~% in January and 4% in July. These represent an annual mean loss of  
392 | 8193 mm (101~~100~~%) in snowfall and 2025~~mm (30%)~~ mm (29%) of rain at Barter, while at Shingle Point  
393 | and Komakuk Beach in Yukon the corrections are, on average, about 2534 mm (21%) for snow and 87.5  
394 | mm (6%) for rain. For Eagle (AK) and Dawson (YK) stations in the central region, the bias corrections  
395 | are small. The monthly~~annual~~ corrections range from 2%~~3%–16% for snow~~, to 22% in winter and from  
396 | 3% to 10% on summer months~~3–7% for rain, much smaller than those for the northern region~~.

397 | On the annual scale, ~~the~~ Barter Island station in AK shows a yearly mean correction around 65%, five  
398 | times greater than the correction at Shingle Point and Komakuk Beach (13~~12~~% and 14%) in Canada. In  
399 | the central region, Eagle station shows an increase by 7%, meanwhile for Dawson the increase is only  
400 | 3%. Thus, the bias correction is twice for Alaska compared to the Yukon stations. Relative to the northern  
401 | region, these corrections are small mainly due to warmer~~warm~~ temperatures and lower~~low~~ winds in the  
402 | central region. These results clearly demonstrate that bias corrections may affect the spatial distribution of  
403 | precipitation across the border.

404 | Regression analyses of the monthly P~~–~~data show small changes in the relationship due to the bias  
405 | corrections. The most evident change in the regression is between Barter Island and Komakuk Beach for

406 both warm and cold seasons. The rest of the scatter plots, for the Komakuk Beach-Shingle Point and  
407 Eagle-Dawson, do not show any appreciable change as the result of the bias corrections. There is a  
408 stronger precipitationP correlation for the warm months (mainly rainfall) than for the cold month (mainly  
409 snowfall) for all the station pairs. The cold months seem to have a greater precipitationP variability across  
410 the regions.

411 The double mass curve analyses demonstrate a significant change in the precipitationP accumulation and  
412 difference between the two stations across the AK-YK border for the northern region, little changes for  
413 the two stations in Yukon, and a smaller change in the central group. These changes, caused by gauge  
414 catch efficiency, alters the precipitationP difference, resulting in a smaller and inverted precipitation  
415 gradient across the border in the northern region. The double mass curve (DMC) is a useful tool to  
416 evaluate the consistency of observation records over space and time (Searcy and Hardison, 1960).~~It is~~  
417 ~~very clear from this study that the significant inconsistency exists in the precipitation measurements~~  
418 ~~across the border. This inconsistency is much greater for snowfall than for rain, as gauge snowfall~~  
419 ~~observation has large errors in the windy and cold conditions. This discrepancy should be taken into~~  
420 ~~account when using the P data across the national borders for regional climate and hydrology~~  
421 ~~investigations.~~

422 ~~The double mass curve (DMC) is a useful tool to evaluate the consistency of observation records over~~  
423 ~~space and time (Searcy and Hardison Clayton, 1960). Some typical issues of observations that DMC can~~  
424 ~~identify include changes in the station locations, and instruments or sensors.~~ Although in this work the  
425 DMC has not been constructed against a reference station, the results clearly show some breaks on the  
426 slope and gaps in the curves, indicating changes in precipitationP relationship across the border that could  
427 be caused by any of the two stations. This information provides the timing when significant changes  
428 occurred in the precipitationP regime. Detail metadata~~Metadata~~ and information for the stations/networks  
429 are necessary to understand the changes in precipitationP observations and to improve the  
430 homogenization of the precipitation records over the high latitudes.

431 This study shows similar monthly  $P_m$  across the north border region and higher  $P_m$  in Yukon than Alaska  
432 over the central region. This result is similar to other studies (Serreze and Hurst, 2000; Simpson et al.,  
433 2005). After the bias corrections, precipitation patterns across the border changed, i.e. higher precipitation  
434 in Barter than Komakuk, in other words, an inverted gradient across the borderline. Over the central  
435 region, the measured mean annual precipitation is slightly higher in Yukon than Alaska, which is also  
436 consistent with Simpson et al., (2002) and (2005). Our results suggest that the gradient between the  
437 central pair of stations becomes smaller after the bias correction. This discrepancy should be taken into  
438 account when using the precipitation data across the national borders for regional climate and hydrology  
439 investigations.

440 Missing data may affect regional precipitation analyses. In this study, we calculated the missing data  
441 percentages for all stations during the corresponding study periods, and set up a threshold of 30% to  
442 exclude those months with higher missing values from monthly precipitation calculations. We compared  
443 the precipitation amounts with and without the application of the threshold. The results do not show any  
444 significant changes in the differences of gauge measured annual mean precipitation across the border,  
445 although this filter affected annual precipitation in certain years. For instance, the northern station pair  
446 (Barter and Komakuk stations) has missing value of 32% on July 1987. Calculations of yearly  
447 precipitation for 1987 with and without this month show 16% and 10% difference at Komakuk and Barter  
448 Island stations, respectively. Over the study period of 11 years, the annual mean bias correction  
449 percentages remain the same (65% in Barter and 13% in Komakuk) with or without the missing months.  
450 The mean annual decrease in bias correction amounts after the consideration of missing data is about 1-  
451 3% in the northern region. This analysis suggests that the effect of missing data for our study is not  
452 significant, particularly with the application of 30% missing threshold. More efforts are needed to further  
453 examine the issues of missing records in climate analyses.

454 Classification of precipitation<sup>P</sup> types is the first step for the bias corrections of gauge records. It is also  
455 important for climate change analyses over the cold regions. Leeper et al., (2015),Leeper et al (2015), in

456 comparison of US CRN with the CO-OP station network precipitation measurements, averaged the  
457 USCRN hourly temperatures data during precipitation<sup>P</sup> periods into an event mean and used it to group  
458 precipitation<sup>P</sup> events into warm (mean temperature > 5C), near-freezing (mean temperature between 0C  
459 and 5C), and freezing (mean temperature < 0C) conditions. Yang et al., (2005)~~Yang et al. (2005)~~ used the  
460 daily mean air temperature to estimate precipitation types (snow, mixed, and rain) when this information  
461 is not available for the northern regions. In this study, monthly mean temperatures have been used to  
462 determine the warm months~~season~~ (mainly for rain) and cold months (mainly for snow). Mixed  
463 precipitation has not been classified separately. This approach is reasonable for our analysis to focus on  
464 the inconsistency in the monthly and yearly P<sub>m</sub><sup>P</sup> records across the border. Data collections and analyses  
465 on shorter timescales, such as daily or hourly steps, are expected to produce better results, since  
466 temperatures vary throughout the days in a month, particularly in the spring and fall seasons. Automatic  
467 sensors will also be important to decide precipitation types at the operational and research networks.

468 The bias-corrected precipitation dataset developed by Yang et al., (2005) has been used for this analysis.  
469 The corrections have been done systematically on a daily time scale that affects the daily P<sub>m</sub> time series.  
470 This analysis focuses on the results of monthly and yearly precipitation data and quantifies the changes in  
471 precipitation pattern across the AK-YK border. Careful analyses of available daily measured P<sub>m</sub> and  
472 corrected P<sub>c</sub> data are necessary, since in the northern regions with low precipitation in winter, the bias  
473 corrections can easily increase the daily P<sub>m</sub> by a factor of up to 4-5 (Benning and Yang, 2005; Kane and  
474 Stuefer, 2015; Yang et al., 1998b, 2005). This means that extreme precipitation events have been very  
475 likely and seriously underestimated by using the gauge records without any bias corrections. The  
476 consequence is certainly significant for climate regime and change investigations. To fill this knowledge  
477 gap, our efforts are underway to examine the daily corrections, particularly on the windy and heavy  
478 precipitation days, and to document the possible underestimation of precipitation extremes over the large  
479 northern regions.

480 ~~Automation~~The bias corrected P dataset developed by Yang et al. (2005) has been used for this analysis.  
481 ~~The corrections have been done systematically on a daily time scale that affects the daily P time series.~~  
482 ~~This analysis focuses on the results of monthly and yearly P data and quantifies the changes in P pattern~~  
483 ~~across the AK-YK border. Careful analyses of available daily measured and corrected P data are~~  
484 ~~necessary, since in the northern regions with low P in winter, the bias corrections can easily increase the~~  
485 ~~daily P by a factor of up to 4.5 (Yang et al., 1998; Benning and Yang, 2005; Yang et al., 2005). This~~  
486 ~~means that extreme P events have been very likely and seriously underestimated by using the gauge~~  
487 ~~records without any bias corrections. The consequence is certainly significant for climate regime and~~  
488 ~~change investigations. To fill this important knowledge gap, our efforts are underway to examine the~~  
489 ~~daily corrections, particularly on the heavy and windy P days, and to document the possible~~  
490 ~~underestimation of P extremes over the large northern regions.~~

491 ~~Finally, automation~~ of the meteorological observation networks and instruments has been a trend over the  
492 past ~~fewseveral~~ decades around the world, including both the developed and developing nations. There is  
493 a large variety of automatic gauges currently used for precipitation measurements at the national networks  
494 (Nitu and Wong, 2010). These gauges differ in the measuring system, orifice area, capacity, sensitivity,  
495 and configuration. The variation in automatic gauges is much greater relative to the manual standard  
496 gauges (~~Goodison et al., 1998; Sevruk and Klemm, 1989~~)(~~Goodison et al., 1998; Sevruk and Klemm,~~  
497 ~~1989~~). As demonstrated by (~~Yang et al., 2001~~)~~Yang et al. (2001)~~ and this study, the use of different  
498 instruments and configurations significantly affect the accuracy and consistency of regional precipitation  
499 data. Fortunately, the Geonor gauge has recently been chosen and used at both the US Climate Reference  
500 Network (USCRN) and the Surface Weather and Climate Network (SWCN) in Canada. This may reduce  
501 the inconsistency in precipitationP measurements across US and Canada borders, although the double and  
502 single Alter wind shields have been installed with the Geonor gauges in US and Canada, respectively.

503 ~~Finally, it is important to emphasize that automatic gauges also significantly under catch snowfall (Wolff~~  
504 ~~et al., 2015) and bias corrections are necessary in order to obtain reliable precipitation data for the cold~~

505 regions and seasons. The WMO SPICE project aims to examine the performance of automatic gauges and  
506 instruments for snowfall observations in various climate conditions. It has tested many different automatic  
507 gauges, including the Geonor gauge, at more than 20 field sites around the globe (Nitu et al., 2012;  
508 Rasmussen et al., 2012; Wolff et al., 2015). The results of this project will be very useful to improve  
509 precipitation data quality and regional climate analyses, including the border regions between US and  
510 Canada.

511 ~~It is however important to emphasize that automatic gauges also significantly under catch snowfall (Wolff~~  
512 ~~et al., 2015) and the bias corrections are necessary in order to obtain reliable P data for the cold regions~~  
513 ~~and seasons. The WMO SPICE project aims to examine the performance of automatic gauges and~~  
514 ~~instruments for snowfall observations in various climate conditions. It has tested many different automatic~~  
515 ~~gauges, including the Geonor gauge, at more than 20 field sites around the globe (Nitu et al., 2012;~~  
516 ~~Rasmussen et al., 2012; Wolff et al., 2015). The results of this project will be very useful to improve P~~  
517 ~~data quality and regional climate analyses, including the border regions between US and Canada~~

518

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522

523        **References**

524        [Benning, J. and Yang, D.: Adjustment of Daily Precipitation Data at Barrow and Nome Alaska for 1995–](#)  
525        [2001, Arctic, Antarct. Alp. Res., 37\(3\), 276–283, doi:10.1657/1523-](#)  
526        [0430\(2005\)037\[0276:AODPDA\]2.0.CO;2, 2005.](#)

527        [Goodison, B. E.: Compatibility of Canadian snowfall and snow cover data, Water Resour. Res., 17\(4\),](#)  
528        [893, doi:10.1029/WR017i004p00893, 1981.](#)

529        [Goodison, B. E., Louie, P. Y. T. and Yang, D.: Report No . 67 WMO SOLID PRECIPITATION](#)  
530        [MEASUREMENT INTERCOMPARISON., 1998.](#)

531        [Groisman, P. Y. and Easterling, D. R.: Variability and Trends of Total Precipitation and Snowfall over](#)  
532        [the United States and Canada, J. Clim., 7, 184–205, doi:10.1175/1520-](#)  
533        [0442\(1994\)007<0184:VATOTP>2.0.CO;2, 1994.](#)

534        [Jones, S. H. and Fahl, C. B.: Magnitude and Frequency of Floods in Alaska and Conterminous Basins of](#)  
535        [Canada, Anchorage, Alaska., 1994.](#)

536        [Kane, D. L. and Stuefer, S. L.: Reflecting on the status of precipitation data collection in Alaska: a case](#)  
537        [study, Hydrol. Res., 46\(4\), 478, doi:10.2166/nh.2014.023, 2015.](#)



538 [Leeper, R. D., Rennie, J. and Palecki, M. A.: Observational Perspectives from U.S. Climate Reference](#)  
539 [Network \(USCRN\) and Cooperative Observer Program \(COOP\) Network: Temperature and Precipitation](#)  
540 [Comparison., J. Atmos. Ocean. Technol., 32\(4\), 703–721, 2015.](#)

541 [Legates, D. R. and Willmott, C. J.: Mean seasonal and spatial variability in gauge-corrected, global](#)  
542 [precipitation, Int. J. Climatol., 10\(2\), 111–127, doi:10.1002/joc.3370100202, 1990.](#)

543 [Manson, G. K. and Solomon, S. M.: Past and future forcing of Beaufort Sea coastal change, Atmosphere-](#)  
544 [Ocean, 45\(2\), 107–122, doi:10.3137/ao.450204, 2007.](#)

545 [Mekis, É. and Vincent, L. a.: An Overview of the Second Generation Adjusted Daily Precipitation](#)  
546 [Dataset for Trend Analysis in Canada, Atmosphere-Ocean, 49\(2\), 163–177,](#)  
547 [doi:10.1080/07055900.2011.583910, 2011.](#)

548 [Metcalf, J. R. and Goodison, B. E.: Correction of Canadian winter precipitation data, in Proc. Eighth](#)  
549 [Symp. on Meteorological Observations and Instrumentation, pp. 338–343, Amer. Meteor. Soc., Anaheim,](#)  
550 [CA., 1993.](#)

551 [Nitu, R. and Wong, K.: CIMO Survey on National Summaries of Methods and Instruments Related to](#)  
552 [Solid Precipitation Measurement at Automatic Weather Stations-Preliminary., 2010.](#)

553 [Nitu, R., Rasmussen, R., Baker, B., Lanzinger, E., Joe, P., Yang, D., Smith, C., Roulet, Y. A., Goodison,](#)  
554 [B., Liang, H., Sabatini, F., Kochendorfer, J., Wolff, M., Hendrikx, J., Vuerich, E., Lanza, L., Aulamo, O.](#)  
555 [and Vuglinsky, V.: WMP intercomparison of instruments and methods for the measurement of solid](#)  
556 [precipitation and snow on the ground : Organization of the Experiment., 2012.](#)

557 [Rasmussen, R., Baker, B., Kochendorfer, J., Meyers, T., Landolt, S., Fischer, A. P., Black, J., Thériault, J.](#)  
558 [M., Kucera, P., Gochis, D., Smith, C., Nitu, R., Hall, M., Ikeda, K. and Gutmann, E.: How Well Are We](#)

559 [Measuring Snow: The NOAA/FAA/NCAR Winter Precipitation Test Bed, Bull. Am. Meteorol. Soc.,](#)  
560 [93\(6\), 811–829, doi:10.1175/BAMS-D-11-00052.1, 2012.](#)

561 [Sanderson, M.: NOTES AND CORRESPONDENCE: A comparison of Canadian and United States](#)  
562 [Standard Methods of Measuring Precipitation, J. Appl. Meteorol., 14, 1197–1199, 1975.](#)

563 [Searcy, J. and Hardison, C.: Double-Mass Curves, United States Department of the Interior, Washington](#)  
564 [DC., 1960.](#)

565 [Šeparović, L., Alexandru, A., Laprise, R., Martynov, A., Sushama, L., Winger, K., Tete, K. and Valin,](#)  
566 [M.: Present climate and climate change over North America as simulated by the fifth-generation](#)  
567 [Canadian regional climate model., 2013.](#)

568 [Serreze, M. C. and Hurst, C. M.: Representation of mean arctic precipitation from NCEP-NCAR and](#)  
569 [ERA reanalyses, J. Clim., 13\(1\), 182–201, doi:10.1175/1520-0442\(2000\)013<0182:ROMAPF>2.0.CO;2,](#)  
570 [2000.](#)

571 [Sevruk, B. and Klemm, S.: Types of standard precipitation gauges, in Proceedings of International](#)  
572 [Workshop on Precipitation Measurement, WMO/IAHS/ETH, vol. 227236, St. Moritz, Switzerland., 1989.](#)

573 [Simpson, J. J., Hufford, G. L., Fleming, M. D., Berg, J. S. and Ashton, J. B.: Long-term climate patterns](#)  
574 [in Alaskan surface temperature and precipitation and their biological consequences, IEEE Trans. Geosci.](#)  
575 [Remote Sens., 40\(5\), 1164–1184, doi:10.1109/TGRS.2002.1010902, 2002.](#)

576 [Simpson, J. J., Hufford, G. L., Daly, C., Berg, J. S. and Fleming, M. D.: Comparing maps of mean](#)  
577 [monthly surface temperature and precipitation for Alaska and adjacent areas of Canada produced by two](#)  
578 [different methods, Arctic, 58\(2\), 137–161, 2005.](#)

579 [Wolff, M. A., Isaksen, K., Petersen-Øverleir, A., Ødemark, K., Reitan, T. and Brækkan, R.: Derivation of](#)  
580 [a new continuous adjustment function for correcting wind-induced loss of solid precipitation: results of a](#)  
581 [Norwegian field study, Hydrol. Earth Syst. Sci., 19\(2\), 951–967, doi:10.5194/hess-19-951-2015, 2015.](#)

582 [Yang, D. and Ohata, T.: A Bias-Corrected Siberian Regional Precipitation Climatology, J.](#)  
583 [Hydrometeorol., 2\(2\), 122–139, doi:10.1175/1525-7541\(2001\)002<0122:ABCSRP>2.0.CO;2, 2001.](#)

584 [Yang, D. and Simonenko, A.: Comparison of Winter Precipitation Measurements by Six Tretyakov](#)  
585 [Gauges at the Valdai Experimental Site, Atmosphere-Ocean, 52\(1\), 39–53,](#)  
586 [doi:10.1080/07055900.2013.865156, 2013.](#)

587 [Yang, D., Goodison, B. E., Metcalfe, J. R., Golubev, V. S., Elomaa, E., Gunther, T., Bates, R., Pangburn,](#)  
588 [T., Hanson, C. L., Emerson, D., Copaciu, V. and Milkovic, J.: Accuracy of tretyakov precipitation gauge:](#)  
589 [Result of wmo intercomparison, Hydrol. Process., 9\(8\), 877–895, doi:10.1002/hyp.3360090805, 1995.](#)

590 [Yang, D., Goodison, B. E., Metcalfe, J. R., Golubev, V. S., Bates, R., Pangburn, T. and Hanson, C. L.:](#)  
591 [Accuracy of NWS 8" standard nonrecording precipitation gauge: Results and application of WMO](#)  
592 [intercomparison, J. Atmos. Ocean. Technol., 15\(1\), 54–68, doi:10.1175/1520-](#)  
593 [0426\(1998\)015<0054:AONSNP>2.0.CO;2, 1998a.](#)

594 [Yang, D., Goodison, B. E., Ishida, S. and Benson, C. S.: Adjustment of daily precipitation data at 10](#)  
595 [climate stations in Alaska: Application of World Meteorological Organization intercomparison results,](#)  
596 [Water Resour. Res., 34\(2\), 241–256, doi:10.1029/97WR02681, 1998b.](#)

597 [Yang, D., Ishida, S., Goodison, B. E. and Gunther, T.: Bias correction of daily precipitation](#)  
598 [measurements for Greenland, J. Geophys. Res., 104\(D6\), 6171, doi:10.1029/1998JD200110, 1999.](#)

599 [Yang, D., Goodison, B., Metcalfe, J., Louie, P., Elomaa, E., Hanson, C., Golubev, V., Gunther, T.,](#)  
600 [Milkovic, J. and Lapin, M.: Compatibility evaluation of national precipitation gage measurements, J.](#)  
601 [Geophys. Res., 106\(D2\), 1481, doi:10.1029/2000JD900612, 2001.](#)

602 [Yang, D., Kane, D., Zhang, Z., Legates, D. and Goodison, B.: Bias corrections of long-term \(1973–2004\)](#)  
603 [daily precipitation data over the northern regions, Geophys. Res. Lett., 32\(19\), L19501,](#)  
604 [doi:10.1029/2005GL024057, 2005.](#)

605 [Zhao, K., Stadnyk, T., Koenig, K. and Crawford, J.: Better Precipitation Product over the Red River](#)  
606 [Basin., 2010.](#)

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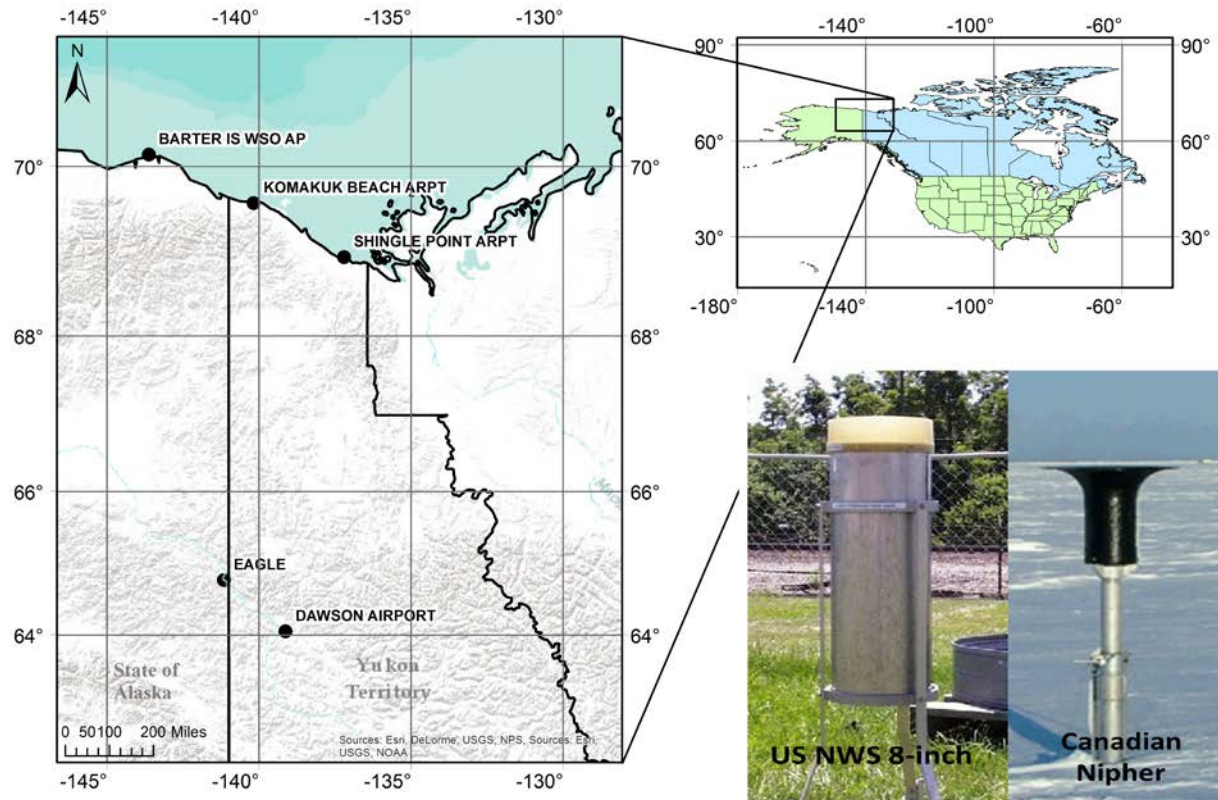
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Table 1: Station information and climate summary

ID	Country	Station Name	Location			Data Period		Measurement Device	Annual Means				
			Latitude (°)	Longitude (°)	Altitude (m)	Start	End	Precipitation gauge	Precipitation (mm)	Missing Precipitation data %	Minimum Temperature (°C)	Maximum Temperature (°C)	Wind Speed (m/s)
700860	US	BARTER IS WSO AP	70.13	-143.63	11	1978	1988	US-8 inch Unshielded	<del>155</del>	<u>0.3</u>	-27.1	4.6	<u>4.0</u>
719690	CA	KOMAKUK BEACH ARPT	69.58	-140.18	7	1978	1988	Nipher Type B gauge	<del>197</del> <u>191.8</u>	<u>2.9</u>	-27.5	7.4	<u>3.94</u>
719680	CA	SHINGLE POINT ARPT	68.95	-137.21	49	1978	1988	Nipher Type B gauge	<del>274</del> <u>302</u>	<u>6</u>	-26.6	10.6	<u>3.4</u>
701975	US	EAGLE	64.78	-141.16	268	2006	2013	US-8 inch Unshielded	<del>253</del> <u>247</u>	<u>0.2</u>	-22.7	15.5	<u>0.94</u>
719660	CA	DAWSON AIRPORT	64.05	-139.13	369	2006	2013	Nipher Type B gauge	<del>275</del> <u>258</u>	<u>0.6</u>	-25.8	15.9	1

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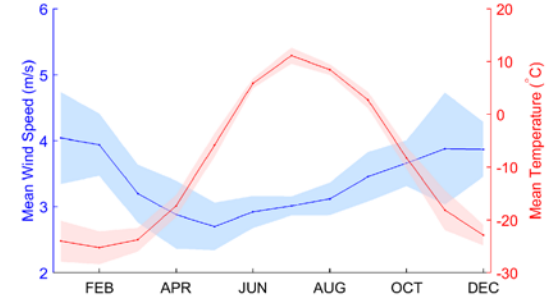
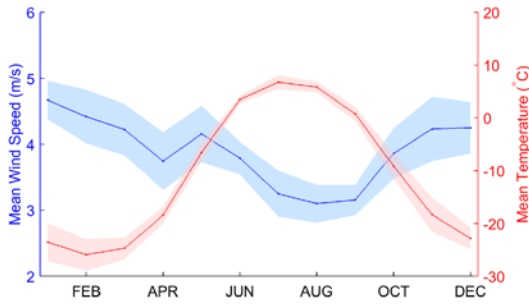
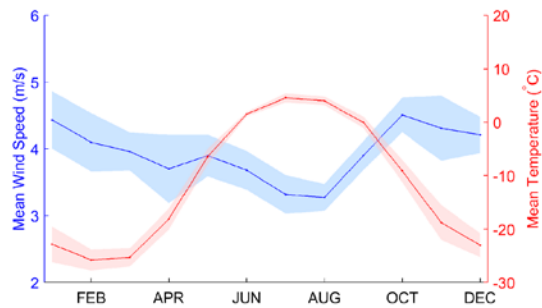
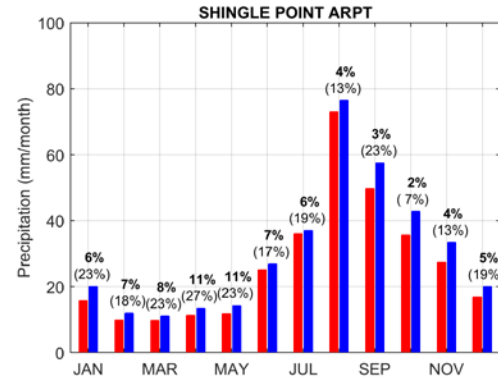
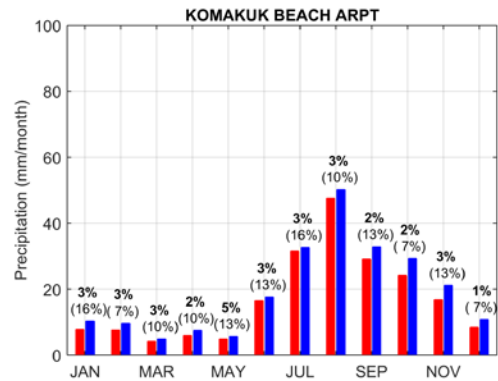
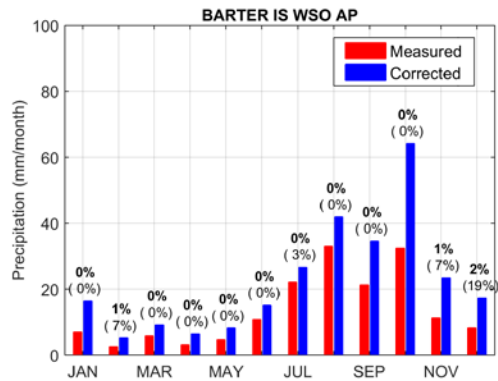


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614 Figure 1: Study areas and locations of selected climate stations, and photosPhotos of the national standard gauges, NWS 8 in gauge (left) and the Nipher  
615 snow gauge (right), respectively, for USA and Canada. ..

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Figure 2: Monthly mean precipitation at 3 stations during 1977 - 1988 (upper panels) and corresponding monthly mean wind speed and air temperature (bottom panels). Shadows represent the 95% confidence interval for the temperature and wind speed. The percentages above the bars represent the missing data for the corresponding time step. The bold percentage is the monthly mean and the one in the parenthesis is the maximum missing value in the study period.

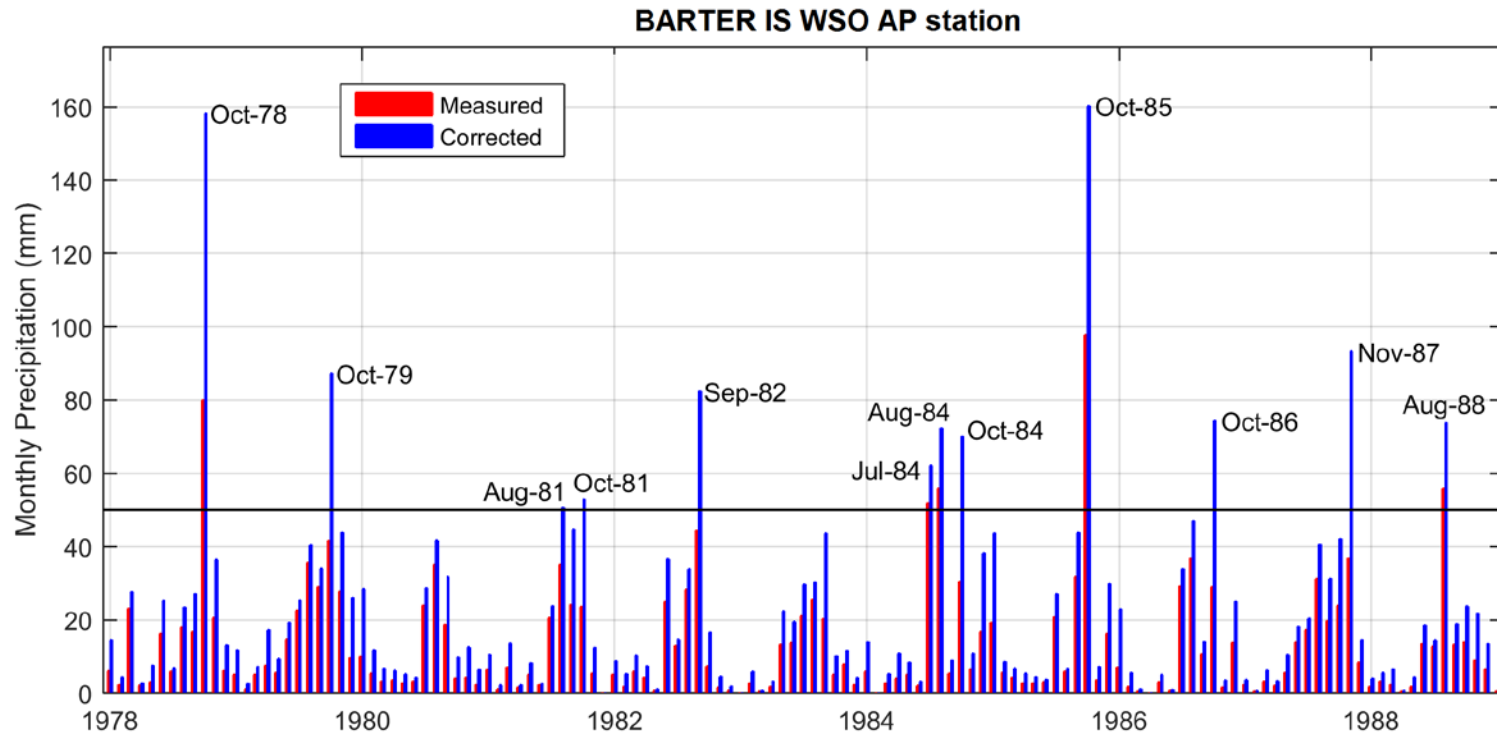
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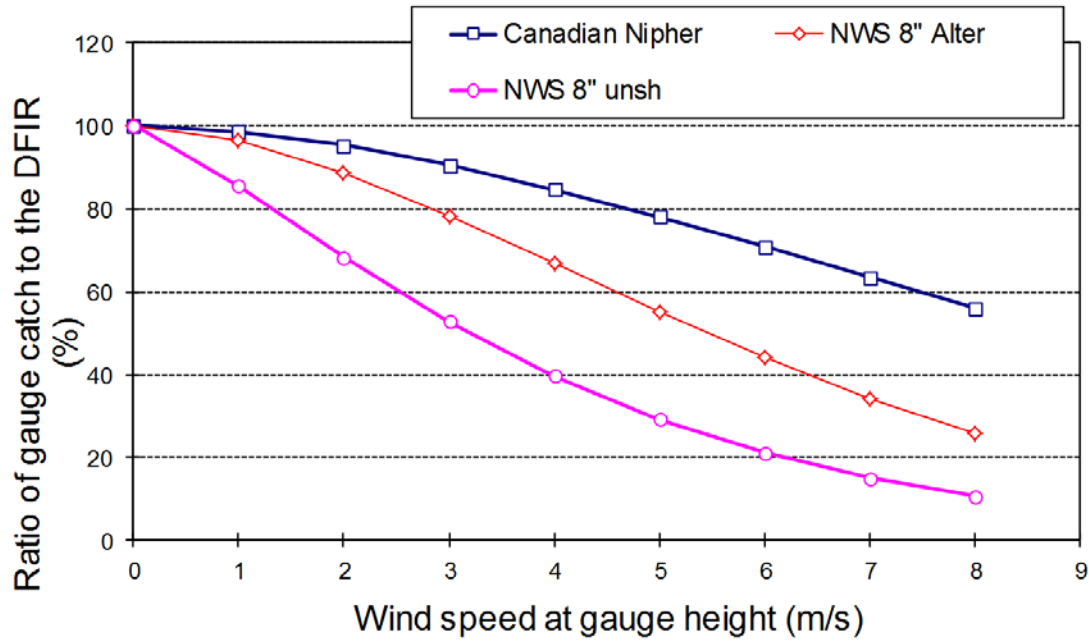
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627 Figure 3: Monthly precipitation records at the Barter station during 1978-1988. The months with more than 50 mm (black line) are labeled.

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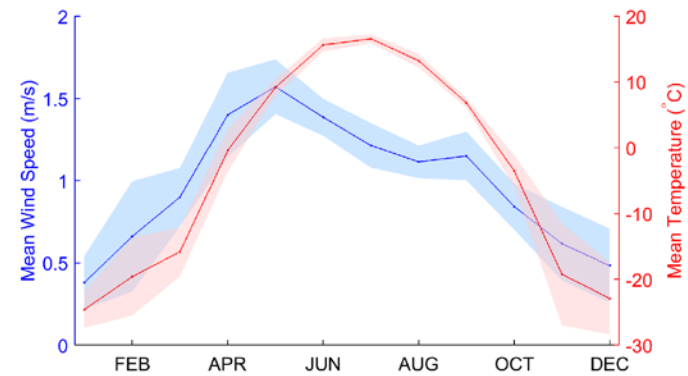
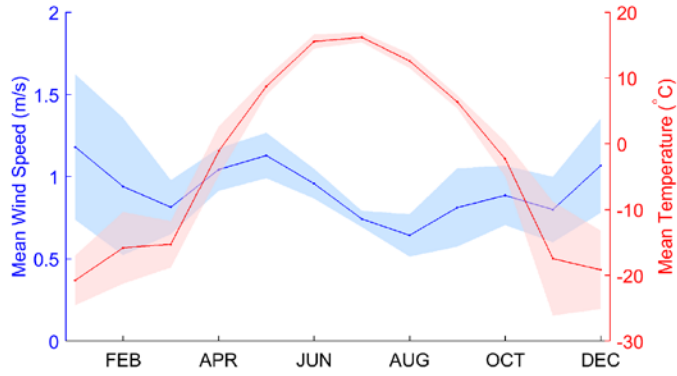
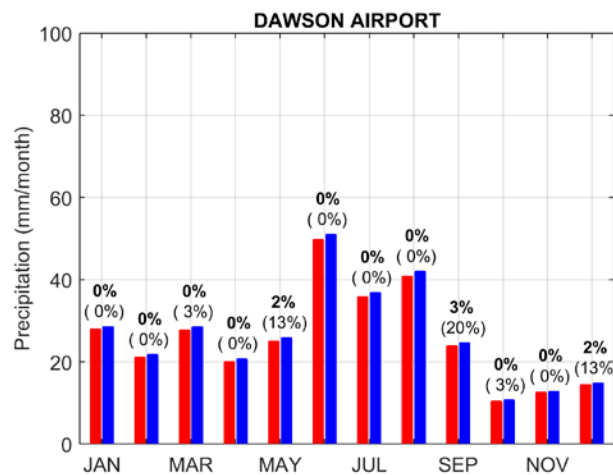
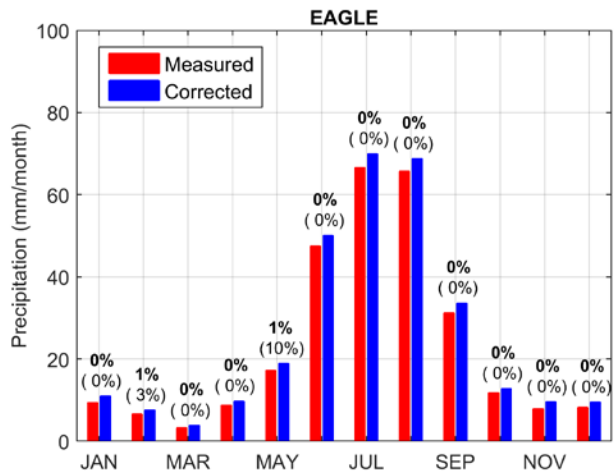
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632 Figure 4: Comparison of the catch ratio of snowfall as a function of wind speed at gauge height for the Alter-shielded or unshielded NWS 8-inc standard  
 633 gauge and the Canadian Nipher snow gauge. DFIR is the Double Fence Intercomparison Reference (Yang et al. 1998)

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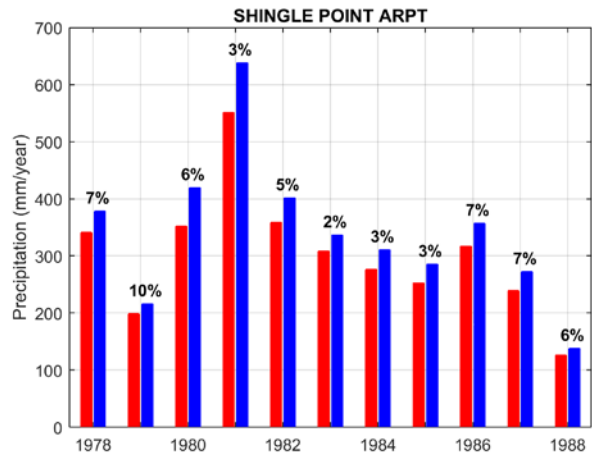
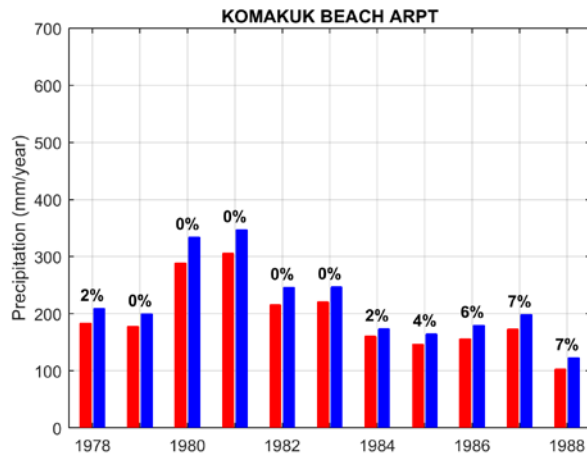
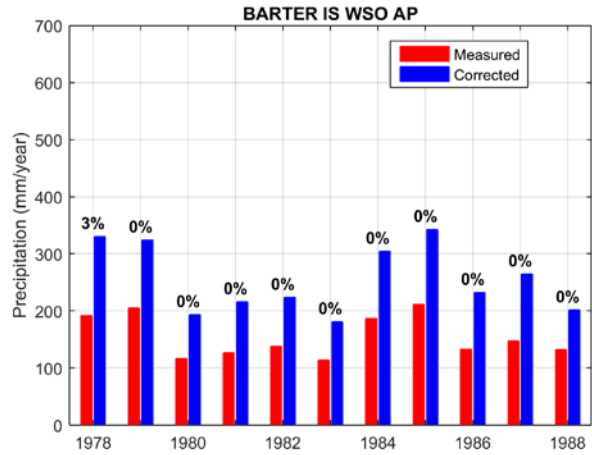
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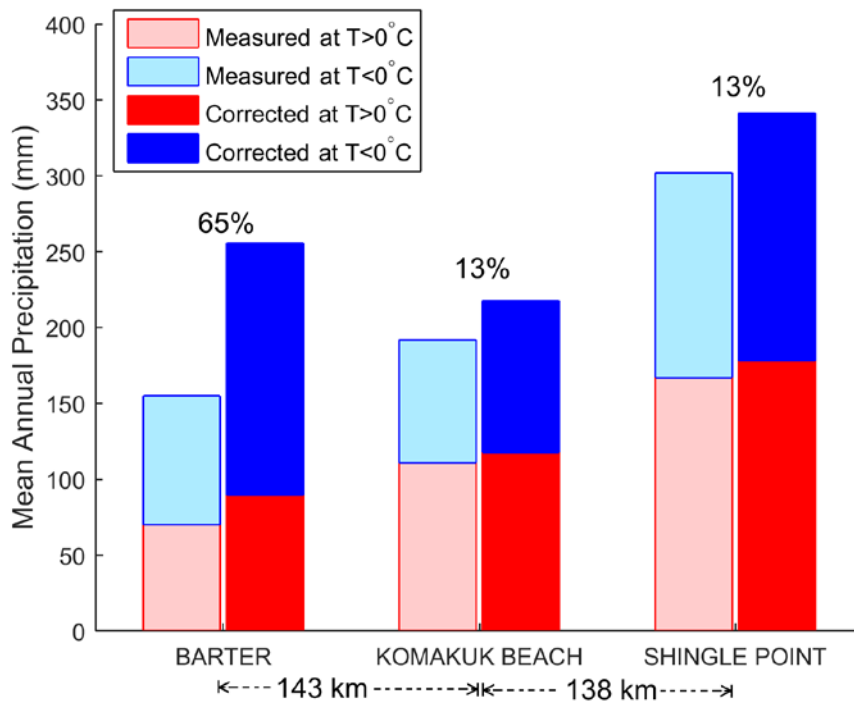
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Figure 5: Monthly mean precipitation at 2 stations during 2006 - 2013 (upper panels) and corresponding monthly mean wind speed and air temperature (bottom panels). Shadows represent the 95% confidence interval for the temperature and wind speed. The percentages above the bars represent the missing data for the corresponding time step. The bold percentage is the monthly mean and the one in the parenthesis is the maximum missing value in the study period.



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641 Figure 6: Annual precipitations during 1978-1988 for the 3 stations in the northern group across the  
 642 border. The percentages above the bars represent the missing data for the corresponding year.



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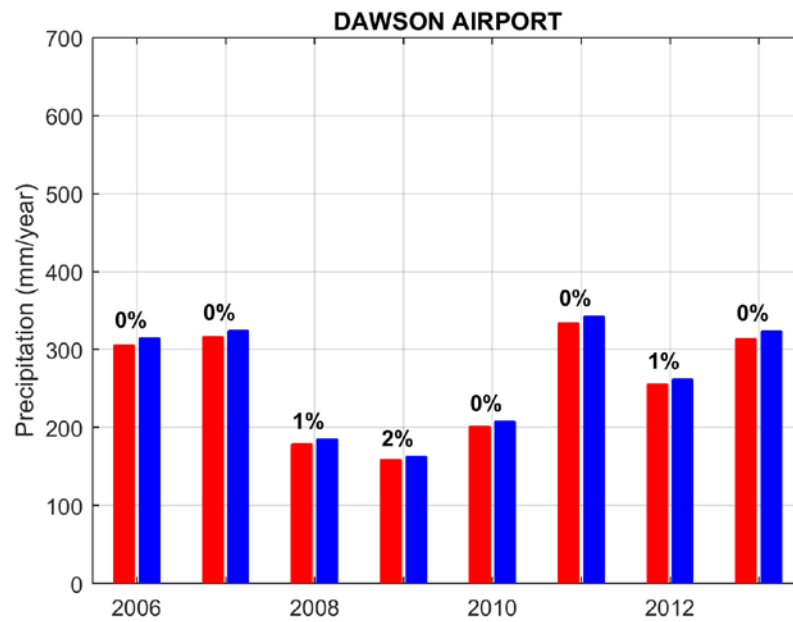
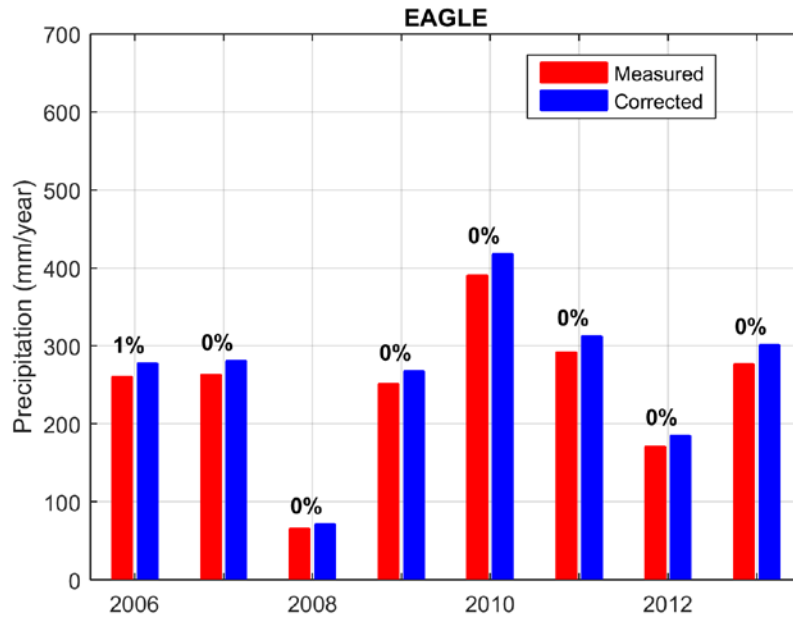
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647 Figure 7: Mean Annual (1978-1988) measured and corrected precipitation for cold ( $T < 0^{\circ}\text{C}$ ) and warm

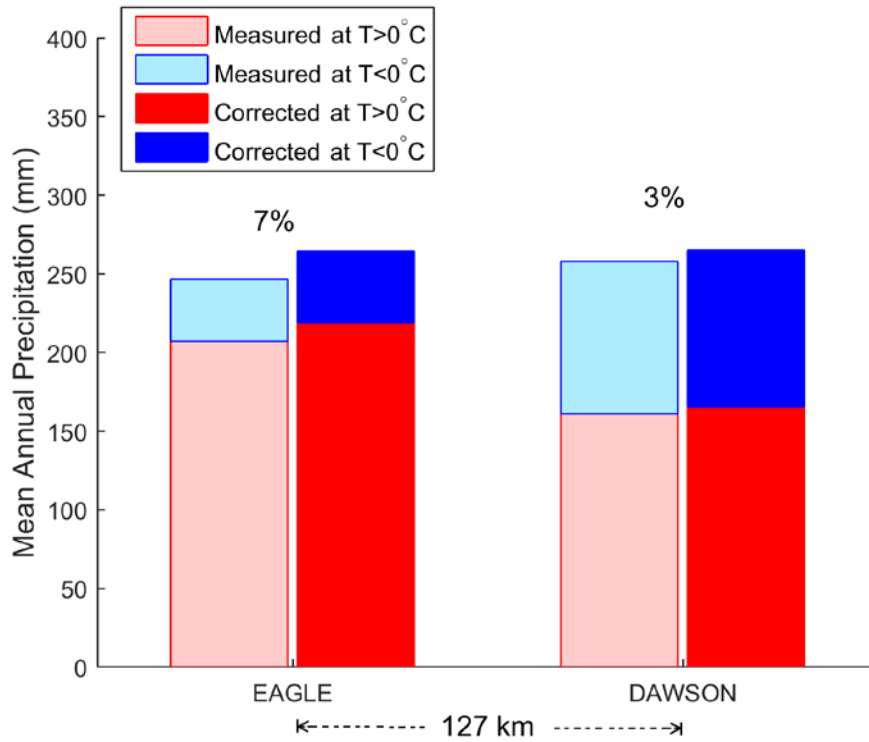
648 ( $T > 0^{\circ}\text{C}$ ) months. The percentages are the changes from measured to corrected precipitation.- The

649 approximate horizontal distance between the stations is displayed at the bottom.

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 652 Figure 8: Annual precipitations during 2006-2013 for two stations in the central part of the AK/YK  
 653 border. The percentages above the bars represent the missing data for the corresponding year.  
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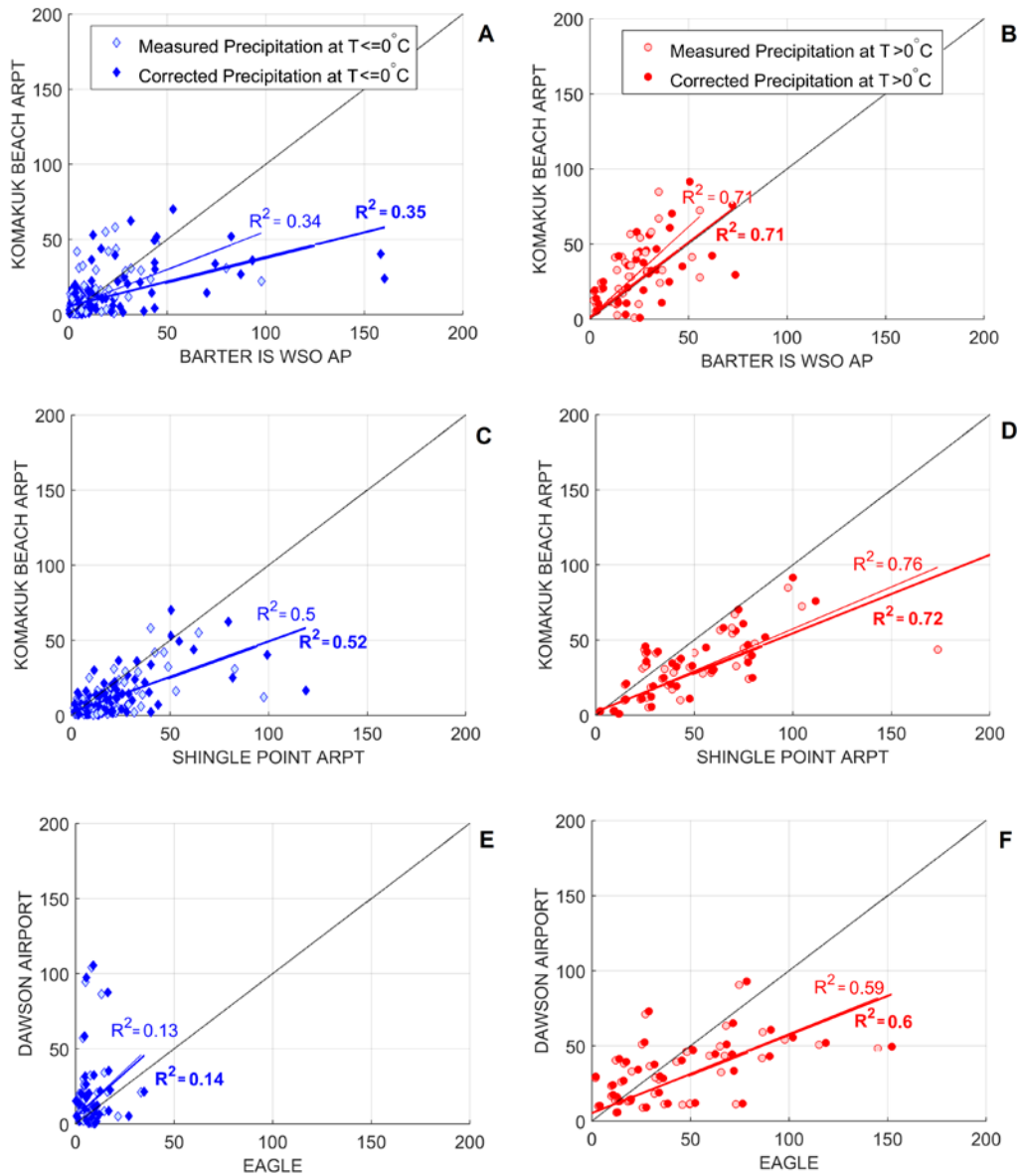
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658 Figure 9: Mean Annual (2006-2013) measured and corrected precipitation for cold ( $T < 0^{\circ}\text{C}$ ) and warm

659 ( $T > 0^{\circ}\text{C}$ ) months. The percentages are the change from measured to corrected precipitation. The

660 approximate horizontal distance between the stations is displayed at the bottom.

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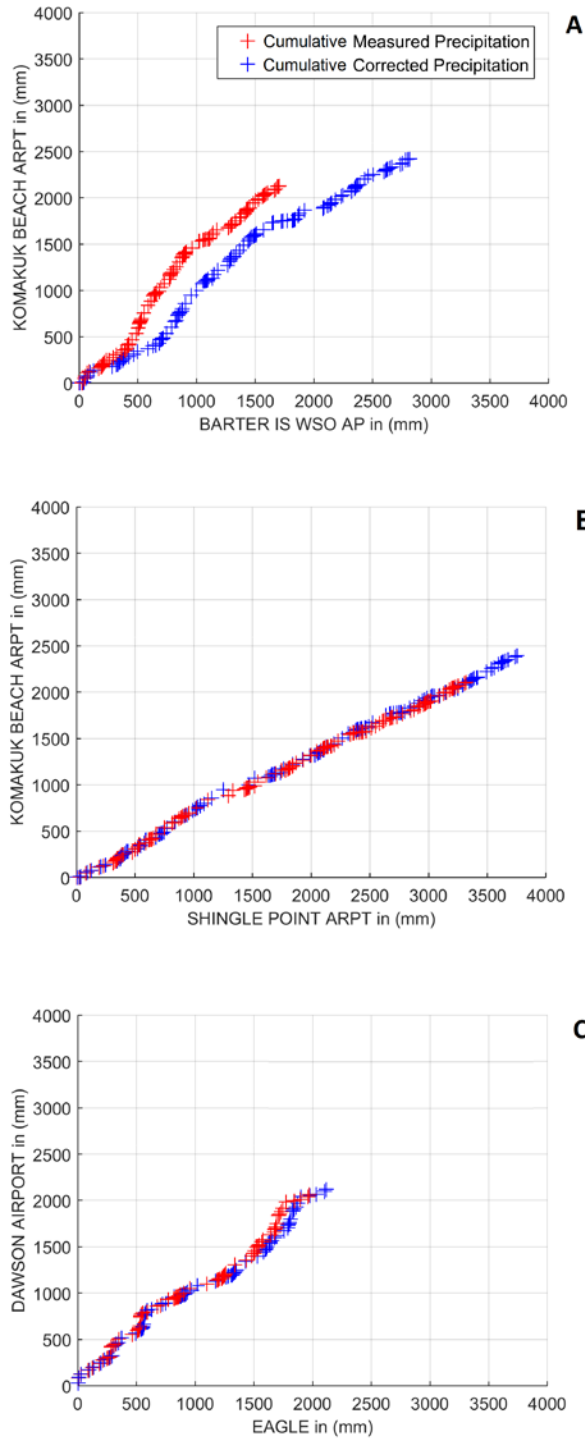
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Figure 10: Scatter plots between station pairs for the measured and corrected precipitation (mm). The red color shows warm months and the blue represents the cold months. A and B - Barter and Komakuk comparison across the border, the highest corrected values for Barter (AK) are labeled with the date to compare with Figure 4. C and D - Komakuk and Shingle Point comparison within Canada. E and F- Eagle vs-. Dawson across the border for the central group.



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Figure 11: Double mass curves between station pairs. The red color shows the warm months and blue represents the cold months. The top and the central plots compare the stations for the northern group and the bottom one is the central station comparison across the border.