Halogen-based reconstruction of Russian Arctic sea ice area from the Akademii Nauk ice core (Severnaya Zemlya)

Andrea Spolaor[1,2], Thomas Opel[3], Joseph R. McConnell[4], Olivia J. Maselli[4], Gunnar Spreen[5], Cristiano Varin[1], Torben Kirchgeorg[1], Diedrich Fritzsche[3], Alfonso Saiz-Lopez[6] and Paul Vallelonga[7]

1. Ca'Foscari University of Venice, Department of Environmental Sciences, Informatics and Statistics, Via Torino 155, 30170 Venice Mestre, Italy

2. Institute for the Dynamics of Environmental Processes, IDPA-CNR, Via Torino 155, 30170 Venice Mestre, Italy

3. Alfred Wegener Institute Helmholtz Centre for Polar and Marine Research, Periglacial Research Section, Telegrafenberg A43, 14473 Potsdam, Germany

4. Desert Research Institute, Department of Hydrologic Sciences, 2215 Raggio Parkway, Reno, NV 89512, USA

5. Norwegian Polar Institute, Fram Centre, Hjalmar Johansens gt. 14, 9296 Tromsø, Norway

6. Atmospheric Chemistry and Climate Group, Institute of Physical Chemistry Rocasolano, CSIC, 28006 Madrid, Spain.

7. Centre for Ice and Climate, Niels Bohr Institute, University of Copenhagen, Juliane Maries Vej 30, Copenhagen Ø 2100 Denmark

ABSTRACT

The role of sea ice in the Earth climate system is still under debate, although it is known to influence albedo, ocean circulation, and atmosphere-ocean heat and gas exchange. Here we present a reconstruction of 1950 to 1998 A.D. sea ice in the Laptev Sea based on the Akademii Nauk ice core (Severnaya Zemlya, Russian Arctic). The chemistry of halogens bromine (Br) and iodine (I) are strongly active and influenced by sea ice dynamics, in terms of physical, chemical and biological process. Bromine reacts on the sea ice surface in auto-catalyzing "Bromine explosion" events causing an enrichment of the Br/Na ratio and hence a bromine excess (Br_{exc}) in snow compared to that in seawater. Iodine is suggested to be emitted from algal communities growing under sea ice. The results suggest a connection between Br_{exc} and spring sea ice area, as well as a connection between iodine concentration and summer sea ice area. The correlation coefficients obtained between Br_{exc} and spring sea ice ($r=0.44$) as well as between iodine and summer sea ice ($r=0.50$) for the Laptev sea suggest that these two halogens could become good candidates for extended reconstructions of past sea ice changes in the Arctic.

1 INTRODUCTION

The rapid and unexpected decrease of Arctic sea ice during recent decades has highlighted the lack of knowledge regarding the mechanisms controlling sea ice growth and decay. Sea ice affects albedo by covering the relatively dark, energy-absorbing ocean waters with highly reflective ice (Francis et al., 2009). Sea ice formation is an important process for driving salinification of surface waters, thereby promoting convection in polar regions (Holland et al., 2001). Sea ice also is an efficient barrier between the ocean and atmosphere, limiting the effectiveness of ocean water to warm the polar atmospheric boundary layer, as well as limiting both the exhalation of CO_2 from CO_2-rich upwelling circum-Antarctic waters and the drawdown of atmospheric CO_2 by downwelling surface waters (Dieckmann and Hellmer, 2010) .

Satellite measurements have documented a rapid decrease in recent Arctic summer sea ice area (Comiso, 2011;Stroeve et al., 2007). During the last 10 years, the four lowest September sea ice minima of the last 35 years have been recorded (Arctic Sea-Ice Monitor – IJIS, www.ijis.iarc.uaf.edu) with the lowest sea ice area of 3.41 million km^2 recorded on 16th September 2012. This is 47% of the average sea ice minimum extent for the 1970-1990 period (7.2 million km^2), but the latest ocean-

atmosphere coupled climate models are unable to replicate the rapid pace of Arctic sea ice retreat(Turner et al., 2012). Such limitations may result from poor parameterization of key physical sea ice processes because the only record available is the relatively short 35-year period of satellite observations.

Accurate reconstruction of sea ice variability before the satellite epoch is important for understanding interactions between sea ice area and both the forcing and effects of climate changes (Wolff et al., 2006). These results also are important for improved model calibration. Many different approaches have been proposed to reconstruct sea ice extent variability. Sediments in marine cores reflect sources and so are used to show pathways of Arctic and sub-Arctic oceanic circulation (Darby, 2003). Ice-rafted debris in marine sediment records indicates the occurrence of floating ice (Lisitzin, 2002), while sea ice-related palaeo-productivity can be inferred from the accumulated remains of microscopic organisms and other biomarkers. Recently, highly branched isoprenoids (IP_{25}) in specific sea ice diatoms in sediment cores have also been proposed and applied as indicators of past sea ice variability (Belt et al., 2007;Müller et al., 2009;Xiao et al., 2013). Coastal records also help to understand the past dynamics of sea ice, producing a clear signal in both coastal sediments and landforms (Polyak et al., 2010). Additionally, integration of various paleoclimate archives and historical observations allows a broad reconstruction of past sea ice variability in the Arctic (Divine and Dick, 2006;Kinnard et al., 2011;Polyakov et al., 2003;Vinje, 2001). A weakness of the available reconstructions of past sea ice area and variability, extending back more than a century, is that they are characterized by poor temporal resolution and/or a limited regional significance.

Many atmospheric conditions are recorded in high temporal resolution in ice cores, such as dust deposition, temperature, solar radiation and atmospheric gas concentration (Wolff et al., 2010) and so they are employed extensively for reconstructing past climate (Petit et al., 1999). The absence of sufficiently reliable and specific proxies has limited their application to reconstructions of sea ice extent and variability. Methanesulphonic acid (MSA), a product of ocean algal emissions, has been used to reconstruct past sea ice changes from both Antarctic and Arctic ice cores (Curran et al., 2003;Isaksson et al., 2005), however MSA is unstable and remobilized in ice cores over centennial to millennial time scales (Röthlisberger et al., 2010). Sodium in ice cores reflects glacial-interglacial sea ice variability but on shorter timescales is strongly influenced by meteorology as well as competing sea ice and open ocean emission sources (Abram et al., 2013;Levine et al., 2014).

Recent studies of the halogen elements Bromine (Br) and Iodine (I) have shown potential for their use as proxies of polar sea ice area extent in both Antarctic and Arctic regions (Spolaor et al., 2014;Sturges

and Barrie, 1988). Bromine is released into the atmosphere as a component of sea salt. An additional source comes from so-called "bromine explosions" that are defined as an autocatalytic sequence of chemical reactions able to generate gaseous bromine compounds such as BrO from bromide trapped in the sea ice (Pratt et al., 2013;Vogt et al., 1996). This is supported by satellite measurements in polar regions that show pronounced springtime increases in atmospheric BrO concentrations (Schonhardt et al., 2012) associated with sea ice presence. These so-called "bromine explosions" (Simpson et al., 2015;Simpson et al., 2007), have the net effect of enriching Br beyond the seawater Br/Na ratio in the snow deposits in the ice caps and subsequently ice cores (Spolaor et al., 2014;Sturges and Harrison, 1986;Sturges and Barrie, 1988). Some studies have found depleted Br compared to sea water abundances in coastal areas (de Gois et al., 2015;Barrie et al., 2001). However all measurements conducted on Arctic and Antarctic ice caps demonstrate a general enrichment of Br (Spolaor et al., 2014) compared to sea water abundances. Other processes can influence the Br excess detected in the snow, however previous findings suggest that Bromine related sea ice chemistry could be the major explanation.

Iodine emissions to the polar atmosphere mainly arise from oceanic biological production and formation of volatile organo-iodine compounds (Atkinson et al., 2012). Laboratory analyses suggest that sea salt iodine contributes less than 2% of total iodine deposition (Spolaor et al., 2013b;Sturges and Barrie, 1988). Although it has been discovered that atmospheric iodine around Antarctica is produced from algae growing on the underside of the relatively thin seasonal sea ice (Saiz-Lopez et al., 2015;Saiz-Lopez et al., 2007), satellite measurements do not show exceptional iodine activity above Arctic sea ice (Spolaor et al., 2014), which may be due to differences in sea ice conditions between both polar environments. In the Arctic, only above summer sea ice satellites are able to determine IO emissions. Arctic boundary layer observations show sporadic enhanced atmospheric IO concentrations related to the presence of ice-free open ocean conditions (Mahajan et al., 2010).

Here we present halogen records of the Akademii Nauk (AN) ice core (Opel et al., 2013) from Severnaya Zemlya to assess their potential for the reconstruction of regional sea ice extend variability in the Russian Arctic and to provide a new regional-scale sea-ice reconstruction, i.e. the easternmost record of the Arctic. Severnaya Zemlya is located in the marine boundary layer and is surrounded by winter Arctic sea ice. The AN ice core features annual resolution, and hence can be used to produce a sensitive climate record for comparison to satellite, ship and land-based observations of sea ice area. Combined with other circumpolar ice caps, this location allows the possibility to produce localized sea ice reconstructions for the whole Arctic region. The bromine excess (Br_{exc}) is expressed in terms of

concentrations (ng g^{-1}) and has been calculated by subtracting the seawater component from the total bromine concentration using sodium as seawater proxy. Iodine concentrations have been used directly without any seawater correction since the sea water contribution (using the average sea water abundance) is less than 1%. These halogen data have been compared with summer and spring sea ice areas from the Laptev and Kara seas, the two Arctic seas east and west of Severnaya Zemlya, respectively. Our results suggest a strong connection between Br_{exc} and spring sea ice changes in the Laptev Sea as well as a positive correlation between iodine and summer sea ice in the Laptev Sea. This work continues investigations already done by our research group on the connection between halogens, (I, Br) and sea ice changes (Spolaor et al., 2013a;Spolaor et al., 2013b). These data are the first investigating the halogen climate signal in the Arctic in the last 50 years and shed new light on the connections between halogen and past sea ice changes.

2 DATA AND METHODS

2.1 AKADEMII NAUK ICE CORE

A 724 m ice core from Akademii Nauk (AN) ice cap (80°31'N, 94°49'E, 760 m a.s.l., Figure 1) was drilled from 1999 to 2001(Fritzsche et al., 2002), presenting the easternmost ice core record currently available from the Arctic. Due to the relatively low altitude of the ice cap, the ice core shows evidence of summer melt and infiltration processes (Opel et al., 2009) which may influence some of the atmospheric records preserved in the ice (Fritzsche et al., 2005). Despite a mean annual air temperature of -15.7°C (May 1999 to April 2000), surface melting occurs almost every year when temperatures may rise above 0°C even at the ice cap summit and a considerable amount of Akademii Nauk ice core consists of melt-layers and partly infiltrated firn (Opel et al., 2009). In the literature few studies have reported the effect of meltwater percolation on the ice core climate signal. Pohjola et al. (Pohjola et al., 2002) studied the effect of percolation in the Lomonosovfonna ice core (Svalbard), a site that features melting and infiltration conditions similar to the Akademii Nauk. Their results suggest that, though the original seasonal climate signal could be disturbed especially for the anions associated with strong acids (NO_3^- and SO_4^{2-}), most of the other chemical species and in particular the stable water isotopes are less affected than the strong acids. In the Akademii Nauk ice core, stable isotope records still show seasonal variations, indicating a minor impact of melting and refreezing only. Even though ions and other species may be more affected a deep infiltration and redistribution will be obstructed by the frequent occurrence of melt layers. Therefore, the Akademii Nauk ice core can be considered suitable

for high-resolution (i.e. annual to multi-annual) reconstruction of paleoclimate and atmospheric aerosol loading as already shown for the past century (Weiler et al., 2005) and the past millennium (Opel et al., 2013). Bromine and iodine are photoactive species but recent findings suggest that Br is preserved in the snow after deposition, while iodine can be remobilized in the snow but still maintains the average annual climatic signal (Spolaor et al., 2014). The relatively high average accumulation rate of 0.46 m water equivalent per year, determined for the period 1956-1999, could also help to preserve the halogen signal in the snowpack.

The core chronology is based on counting of annual layers in stable water isotopes, constrained by the identification of reference horizons including the ^{137}Cs nuclear bomb test peak (1963 AD) and volcanic eruptions (Bezymianny 1956 AD). In this pilot study, we focus on the core section 0-29 m, representing the time period 1950-1999. The core chronology for the time period viewed here is well constrained by the detection of the 1956 AD volcanic eruption of Bezymianny (Kamchatka Peninsula) (Opel et al., 2013). Based on comparisons to other dating approaches (linear interpolation, age modeling) we estimate the dating uncertainties to be about ± 1 year, but definitely less than ± 3 years.

2.2 HALOGENS ANALYSIS

Contiguous, longitudinal samples (1.0x0.033x0.033 m) were cut from the Akademii Nauk ice core and shipped frozen to the Ultra-Trace Chemistry Laboratory at the Desert Research Institute for analyses using a unique, continuous ice core measurement system (McConnell et al., 2002). Longitudinal samples are melted consecutively on a carefully cleaned, engraved melter head that splits meltwater from different parts of the sample cross-section into ultra-clean (innermost ~10%), clean (next ~20%) and potentially contaminated (outermost part of the ice core ~ 70%) continuously flowing sample streams. Elemental measurements are made on the ultra-clean sample stream, with ultra-pure nitric acid added immediately after the melter plate to yield an acid concentration of ~1%. The analytical system includes two Thermo-Fisher Element II high-resolution Inductively Coupled Plasma Mass Spectrometers operating in parallel and used to measure simultaneously >30 elements (McConnell et al., 2014;Sigl et al., 2014) including Br, I, and Na, a Picarro L2130 water isotope analyser (Maselli et al., 2013), a Droplet Measurement Technologies SP2 black carbon analyser (McConnell et al., 2007), among other instruments for determination of ammonium, nitrate, hydrogen peroxide and other chemical compounds (Pasteris et al., 2014). Effective depth resolution differs between the instruments in the analytical system and operating parameters but in this study is estimated to be ~0.02 m for Br, I, and Na, with all measurements exactly co-registered in depth. Detection limits are 0.1, 0.003, and 0.06

193 ng/g for Br, I, and Na, respectively.
194
195 **2.3 BACK TRAJECTORY CALCULATIONS**
196 To understand the sources of air masses that influence the bromine and iodine deposition at the drill
197 site of the Akademii Nauk ice core we calculated back trajectories for the period covered by satellite
198 sea ice measurements (1979-2000). Three-day back trajectories for spring (March, April, May, MAM)
199 were chosen to investigate the bromine sources while six-day back trajectories for summer (June, July,
200 August; JJA) were chosen to investigate the iodine sources, where some long-life iodine organic
201 compounds and re-cycling processes can influence total iodine concentrations (Simpson et al., 2015).
202 Back Trajectories were calculated with the model HYSPLIT (*HYbrid Single-Particle Lagrangian*
203 *Integrated Trajectory, Version September 2014)* using NCEP/NCAR Reanalysis data from the National
204 Weather Service's National Centers for Environmental Prediction provided by the NOAA's Air
205 Resources Laboratory (Draxler and Hess, 1998). The arrival height was the drilling site at Akademii
206 Nauk ice cap located at 760 m a.s.l. Trajectories were calculated each 12 hours and vertical motion
207 mode was used. For a better comparison and an indicative estimation of the potential origin of air
208 masses, back trajectories of five-year periods were clustered for each three-month period (MAM and
209 JJA) for the years 1980 to 2000 (Figure 2).
210
211 **2.4 SEA ICE AREA AND ANOMALIES**
212 Time series of the monthly mean sea ice area over the period January 1979 to December 2013 were
213 calculated for three source regions in the Arctic (Figure 1). These correspond to the Kara Sea (496,875
214 km^2), the Laptev Sea (781,875 km^2), and a subset of the Arctic Ocean (536,875 km^2). Observations of
215 sea ice concentrations from passive microwave satellite radiometers were used as input data (Sea Ice
216 Concentrations from Nimbus-7 SMMR and DMSP SSM/I-SSMIS Passive Microwave Data) (Cavalieri
217 et al., 1996, updated yearly). Regional averages were produced for the sea ice concentration datasets
218 which are published at 25 km grid resolution resulting in a single sea ice concentration value for each
219 region every day. The time series were resampled to three-monthly averages for each region and
220 averaged by multiplying the mean sea ice concentration with the area of each region.
221 Typical uncertainties for sea ice concentration observations are between 5% and 10%, depending upon
222 the time of year (greater uncertainty in summer) and the sea ice concentration range (greater
223 uncertainty for lower ice concentrations). Although averaging over larger areas, such as those
224 designated in Figure 1, will reduce the relative uncertainty we estimate the uncertainty of the sea ice

data presented here to be no greater than 10%.

In addition to the satellite measurements we compare our results to the dataset produced by Polyakov et al. (Polyakov et al., 2003) reporting August sea ice anomalies in the Kara, Laptev, East Siberian and Chukchi Seas. The dataset was produced by compiling Russian historical records of fast ice locations in the Arctic seas from ship-based observations, hydrographic surveys and commercial shipping routes and aircraft-based observations. During World War II some missing data (1942-1945) were reconstructed using statistical regression models.

3 RESULTS

3.1 Trajectories and sea ice area: the Laptev Sea basin influence

Air mass back trajectories suggest that the Laptev Sea basin is the predominant source for air masses arriving at Akademii Nauk drill site during springtime (Figure 2). The percentage of the springtime air masses originating in the Laptev basin range from a minimum of 44% (1976-1980) up to 53% (1991-1995), with the percentage of air masses originating from the Kara Sea and Arctic Ocean regions defined here consistently lower. Hence, we consider the Laptev Sea basin to be the most important source region for the spring climate signal present in the Akademii Nauk ice core. For the summer period the sources are more variable with the Laptev Sea basin showing a majority of air mass sources in three of the five 5-year periods investigated. During the decade 1981-1990, the percentage of summer air masses from the Kara Sea exceeds those from the Laptev Sea (in particular between 1986-1990); however no associated changes have been detected in climate proxies such as $\delta^{18}O$ (Opel et al., 2009).

In addition to the back trajectories we calculate sea ice areas for the three assigned basins of the Arctic Ocean and Laptev and Kara Sea regions. The results clearly demonstrate that the greatest variability of sea ice area occurs in the Laptev Sea for both spring and summer sea ice. In particular the Arctic Ocean region shows very small changes in summer minima and the production of first-year sea ice is hence negligible compared to the other two basins. Seasonal changes in Kara Sea ice area are comparable to those calculated for the Laptev Sea (Figure 3). However, considering that the air masses arriving at Akademii Nauk in spring and summer originate primarily from the Laptev Sea and that this region displays the greatest seasonal variability in sea ice area, we consider halogens concentration in the Akademii Nauk ice core to be most likely dominated by sea ice variability in the Laptev Sea. Nonetheless we also evaluate the possibility that the Kara Sea is an important secondary source of halogens.

3.2 Statistical analysis

Given the results of the calculated airmass back trajectory and observed sea ice variability, we compare the yearly average values of Br_{exc} and I concentrations with summer and spring sea ice of Laptev and Kara Seas. Because of the low variability detected in this part of the Arctic Ocean, this region was excluded from statistical evaluation. Apart from the Polyakov anomalies, the parameters are transformed to the logarithmic scale to reduce their asymmetry and thus improve on the adherence to the normal distribution assumption used in the statistical analyses. Inspection of normal probability plots confirms the absence of departures from the normality for the log-transformed parameters. Furthermore, the series of log(I) is de-trended by subtracting a least-squares-fit straight line. Autocorrelation and partial autocorrelation plots confirm the absence of serial correlation in the log-transformed parameters and in the detrended log(I).

Table 1 lists the correlations of the detrended log(I) and log(Br_{exc}) with various calculated parameters of sea ice extent together with the p-values based on the unilateral t test for the correlation coefficient. We consider correlations with a p-value smaller than 0.05 as an indication of a statistically significant positive correlation. The data support a significant positive correlation between log(I) and the logarithm of the Laptev Sea ice area both in Spring (r=0.50, df=20, p=0.009) and Summer (r=0.49, df=20, p=0.011), while there is no evidence of correlation with the Kara Sea ice area in either season (Figure 4a). The results confirm also a positive correlation between log(Br_{exc}) and the logarithm of the Laptev Sea spring sea ice (r=0.44, df=20, p=0.020) but not with Kara Sea spring sea ice (r=0.18, df=20, p=0.205) (Figure 4b).

We also evaluate the correlation between log(Br_{exc}) and I with the Polyakov anomalies dataset. The Polyakov anomalies dataset represents the summer (August) sea ice changes in the last century for several Arctic basins including the Laptev and Kara Seas. The obtained results confirm the finding that log(Br_{exc}) is correlated with Laptev. The correlation between log(Br_{exc}) and the Polyakov anomalies dataset is significant for both Laptev Sea sea ice (r=0.34, df=48, p-value=0.009) and Kara Sea sea ice (r=0.31, df=48, p-value=0.015). As regards to iodine, data indicate a significant correlation between log(I) and the Polyakov anomalies data for Laptev Sea sea ice (r=0.32, df=48, p-value=0.012) but not for Kara Sea ice (r=0.17, df=48, p-value=0.116). These results, together with back trajectory calculation, support the Laptev Sea as the main location of sea ice variability influencing Br_{exc} and I in the Akademii Nauk ice core.

4 DISCUSSION

4.1 Br_{exc} and Laptev Sea spring sea ice

Bromine explosions are defined as the autocatalytic sequence of chemical reactions able to generate gas phase bromine species (such as BrO) and they occur mainly above seasonal sea ice (Figure 5). Bromine explosions lead to enrichment of Br in snow deposition, well beyond the Br/Na ratio observed in seawater. Bromine can be also emitted by other sources, such as biological activity or anthropogenic emissions. In addition heterogeneous reactions on sea-salt particles and sea-salts on snow or sea-ice can contribute to the total emission of Br from the bromine explosion. However satellite images and previous results suggest that these processes are secondary to the sea ice-based bromine emission. There are two calculations that can be used to quantify the influence of bromine explosions in snowpack: bromine excess (Br_{exc}) and bromine enrichment (Br_{enr}). Both calculations use the same Br/Na seawater concentration ratio of 0.006 (Turekian, 1968). Br_{exc} is calculated as Br_{exc} = [Br] – [Na] x 0.006 and indicates how much bromine has been produced by the bromine explosion. Br_{enr} is calculated as Br_{enr} = [Br] / ([Na] x 0.006) and indicates the factor by which bromine has been enriched beyond the seawater ratio. Comparing both Br indicators (Br_{exc} and Br_{enr}) reveals that they have a similar trend over the past 50 years, with a peak in the 1970s and 1980s and a sharp decrease in the last decade of the record (Figure 6). To investigate the connection between Br and sea ice we used Br_{exc} since it allows quantification of the additional bromine fluxes produced by the bromine explosion. Additionally, the Akademii Nauk ice core was subject to summer melt and hence may be susceptible to artefacts based on the different percolation velocities of Br and Na. Air mass back trajectory data indicate the Laptev Sea is the primary springtime source of air masses arriving at Akademii Nauk so we compare Br_{exc} with spring sea ice area in the Laptev Sea. It must be noted that bromine explosions are promoted by first year sea ice (Simpson et al., 2007). Considering the strong sea ice recycling in the Laptev Sea and the small amount of summer sea ice area compared to winter sea ice (Figure 3), we consider the whole spring sea ice area to be an active site for the bromine explosion. A strong influence of the bromine explosion also is expected due to the observation that only 15 to 30 % of sea ice is present by late summer, implying that almost all spring sea ice is seasonal. Furthermore, Laptev Sea ice can be exported by the Transpolar Drift through the Arctic Ocean to the Greenland Sea and North Atlantic, leading to strong recycling of sea ice in the area (Xiao et al., 2013). Comparing the log-transformed Br_{exc} and Laptev Sea spring sea ice we find a positive correlation (r = 0.44). Over the period for which observations are available, noting that both spring sea ice area and Br_{exc} exhibit a steadily decreasing trend (Figure 7a). In 1991 in particular, we observe a sharp decrease of spring sea

ice followed by decreasing production of Br_{exc}. The intensity of bromine explosion and the production of BrO in the Arctic are strongly driven by sea ice presence, in this case spring sea ice. Log-transformed Br_{exc} also was compared with a 50-year record of summer (August) sea ice anomalies for the Laptev basin, calculated by Polyakov et al. (Polyakov et al., 2003) (Figure 7b). We found a positive correlation between the two data series (r=0.31). We emphasize that the degree of correlation may be influenced by the fact that we are not comparing log(Br_{exc}) to the sea ice area at the season of peak emission. Although spring and summer sea ice areas may have different short-term trends, the general trend is similar over the decadal period considered here. Although our results suggest that excess of Br in snow deposition seem mainly driven by the changes in sea ice area, especially first year sea ice, we also consider other factors that could contribute to the total nss-Br in the snow deposition. Bromine explosions are catalyzed by acidity and light, and a change in the acidic condition of the Arctic atmosphere could produce a change in the magnitude of bromine activation and hence bromine deposition (Sander et al., 2006). Other process can interferes on bromine deposition. The recycling of bromine in airborne sea salt aerosols could also impact the bromine concentrations. Quantification of the impacts of acidity changes and other process on Br_{exc} (or Br_{enr}) deposition is beyond the scope of this study and additional studies are necessary to better understand which factors (e.g., sea ice area, acidity, wind pumping, temperature) are the most relevant. However the highly significant correlations and previous research suggests that sea ice is the main driver of Br_{exc}.

4.2 Iodine and Laptev Sea summer sea ice

While the main source of iodine in Antarctica has been attributed to sea ice algae during spring time (Atkinson et al., 2012;Saiz-Lopez et al., 2007;Saiz-Lopez et al., 2015) and recently also to inorganic emissions during winter time (Granfors et al., 2014), in the Arctic there is a lack of knowledge regarding iodine sources. One of the main barriers to identifying Arctic iodine sources is that the atmospheric concentrations of IO are close to satellite detection limits (Figure 5). Iodine could be emitted also by heterogeneous reactions on sea ice, snowpack, or marine aerosol surfaces but the general idea remains that the control of iodine atmospheric concentration is due to biological activities associated with sea ice. The first measurements of iodine in Arctic aerosols (Sturges and Barrie, 1988) found increasing concentrations during late spring and summer. More recently Mahajan et al. (2010) detected iodine emissions in the form of IO, related to the presence of polynyas or areas free of sea ice over Hudson Bay , Canada. These open water areas were identified as hot spots for iodine emission. Arctic sea ice is thicker and less permeable than Antarctic sea ice, constituting a barrier between the

ocean and the atmosphere (Zhou et al., 2013). In this case it is likely that gas phase iodine produced from sea ice algae could escape only from sea ice leads (cracks) or the open ocean surface (Saiz-Lopez et al., 2015). The Akademii Nauk iodine signal is positively correlated with satellite observations of Laptev Sea summer sea ice (r=0.49, Table 1 and Figure 8a) and with spring sea ice (r=0.50). Concentrations of I also have been compared with summer (August) sea ice anomalies for the Laptev basin calculated by Polyakov et al. (Polyakov et al., 2003) (Figure 8b). A significant positive correlation (r=0.32, Table 1) was observed also in this case. No significant correlation was observed with summer or spring sea ice area in the Kara Sea (Table 1), supporting the back-trajectory analysis data. These results are consistent with iodine concentrations from a snow pit sampled at the NEEM ice core site in northwest Greenland. The NEEM record displayed peak iodine concentrations in summer (Spolaor et al., 2014), although only three annual cycles were sampled. In contrast, iodine concentrations from Svalbard were more consistent with spring sea ice area (Spolaor et al., 2013a). The results from Svalbard must be evaluated with caution because seasonal variability is less clear in the Svalbard records and the influence of summer melting and iodine re-emission may be more significant than for the Akademii Nauk ice core. The summer peak of iodine concentrations found in Arctic snow and ice may depend upon both the production of iodine by sea ice algae and processes that control release of iodine to the atmosphere. Arctic sea ice is an effective barrier to gas exchange in the ocean-atmosphere system especially in winter, when there is low ice permeability and minimal convection of Arctic waters (Zhou et al., 2013). Ice permeability follows Arctic temperatures, with maximum permeability in summer and particularly when the sea ice is warmer than -5°C. Water convection is greatest during spring, driven by an unstable brine density profile. As suggested for the Antarctic (Saiz-Lopez et al., ACP, 2015), fracturing is an additional process that may enhance gas exchange, and hence iodine emission from the Arctic Seas, particularly in the highly dynamic Laptev Sea. Phytoplankton blooms occur mainly in late spring and summer (Ardyna et al., 2013;Kevin R. Arrigo et al., 2012), responding to the availability of nutrients and the limited penetration of light through thick winter sea ice. These factors ensure that the most efficient season for production and emission of iodine is the Arctic summer. In summer, the decreased sea ice thickness, increased sea ice fracturing and permeability, and enhanced light penetration and nutrient availability, could create conditions necessary to sustain blooms of sea ice algae responsible for IO production (Kevin R. Arrigo et al., 2012;Allan et al., 2015). In years of decreased Arctic sea ice area, the available substratum for growing algae is reduced and hence IO production is inhibited. This explanation fits well with the data available from Greenland (Spolaor et al., 2014) and presented here for the Akademii Nauk ice core. Satellite

measurements of Arctic IO concentrations exceed detection limits only in the summer, and at that time of maximum IO concentrations above summer sea ice. However, these results can only be used as indications. Improved sensitivity of satellite-borne instruments would greatly enhance the detection and seasonal variability of IO emission sources in the Arctic.

5. CONCLUSIONS

The halogens iodine and bromine reported here from the Akademii Nauk ice core from Severnaya Zemlya offer a new perspective on the variability of sea ice extent in the Arctic. This work follows on previous work to suggest a connection of bromine and iodine chemistry with sea ice changes (Spolaor et al., 2013a;Spolaor et al., 2014;Spolaor et al., 2013b). In particular, Br_{exc} and Br_{enr} have been linked to seasonal sea ice area and here we report a connection between ice-core Br_{exc} and spring sea ice in the Laptev Sea largely because almost all Laptev Sea spring sea ice is seasonal (approximately 80%) and so undergoes continuous renewal. Ice-core iodine appears to be connected with the summer sea ice area; however, its sources are suggested to be biologically-mediated, which makes the interpretation more difficult, and atmospheric concentrations of IO in the Arctic are close to satellite detection limits, limiting the accurate characterization of IO emission sources. Further studies are necessary to better identify the seasonal variability of this element and the impact of acidity on bromine reactivity in Arctic. However, the significant correlation between nss-Br and sea ice during the last 50 years in the Akademii Nauk ice-core record suggests that the nss-Br record primarily reflects changing sea ice conditions in the Laptev sea, bringing important new information for the use of halogens in paleoclimate studies and suggesting that the Akademii Nauk ice core may be a key archive for reconstructing late-Holocene sea ice variations in the Russian Arctic.

Acknowledgements

This study contributes to the Eurasian Arctic Ice 4k project and was supported by the Deutsche Forschungsgemeinschaft (grant OP 217/2-1 awarded to Thomas Opel). The drilling project on Akademii Nauk ice cap has been funded by German Ministry of Education and Research (BMBF research project 03PL 027A). Analysis and interpretation of the Akademii Nauk ice core at the Desert Research Institute was funded by U.S. National Science Foundation grant 1023672. The research leading to these results has received funding from the European Research Council under the European Community's Seventh Framework Programme (FP7/2007-2013) / ERC grant agreement n° 610055"..

We thank the University of Bremen, supported by the State of Bremen, the German Aerospace, DLR and the European Space Agency for the satellite BrO imagines.

TABLE

Table 1. Correlations (r) of the detrended log(I) and log(Br_{exc}) with the logarithm of the first year sea ice area in the Laptev and Kara seas for the period 1979-1999 and the Polyakov anomalies in the Laptev and Kara seas for the period 1950-1999 (denoted by an asterisk *) (Polyakov et al., 2003). Columns display the r-values and p-values. Bold numbers indicate the statistically significant correlations. No correlation has been calculated between log(Brexc) and summer sea ice since the Bromine explosion occur during spring time.

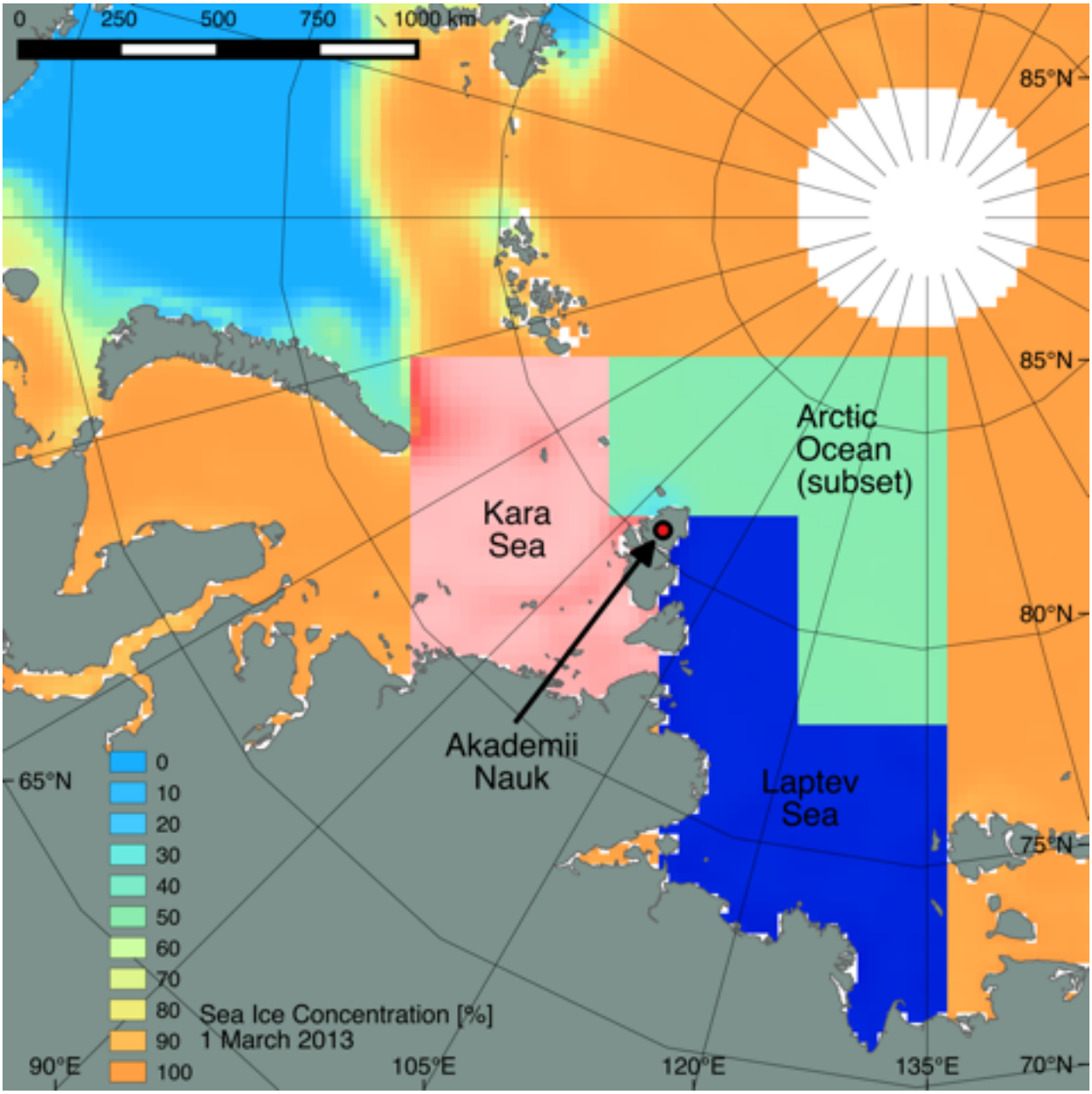
	log (Br_{exc})		log(I)	
	r	p-value	r	p-value
Laptev Sea spring	**0,44**	**0,020**	**0,50**	**0,009**
Kara Sea spring	0,18	0,205	0,04	0,433
Laptev Sea summer	—	—	**0,49**	**0,011**
Kara Sea summer	—	—	-0,03	0,561
Polyakov* - Laptev Sea	**0,31**	**0,009**	**0,32**	**0,012**
Polyakov* - Kara Sea	**0,34**	**0,015**	0,17	0,116

448 **FIGURES**

449 **Figure 1**. Arctic areas considered for sea ice calculations. Kara Sea (pink), Laptev Sea (dark blue) and
450 a subset of the Arctic Ocean (green). The location of the Akademii Nauk ice core drill site on
451 Severnaya Zemlya also is shown.
452
453 **Figure 2**. Air-mass back trajectories calculated for the period 1976 – 2000. Each panel represents the
454 5-year average for spring (MAM, left column) and summer (JJA, right column) seasons.
455
456 **Figure 3**. Sea ice area variation in the period 1979 - 2000 for the three regions defined in Figure 1; (a)
457 Arctic Ocean, (b) Kara Sea and (c) Laptev Sea. Laptev Sea ice shows the greatest seasonal variability.
458
459 **Figure 4**. Scatter plot of log(Br_{exc}) and Spring Laptev sea ice area (a - black squares) and detrended
460 Iodine and summer Laptev sea ice area (b - black circles).
461
462 **Figure 5.** Atmospheric column averages of BrO and IO in the Arctic between 2009 and 2011. In the
463 Arctic, IO concentrations are near the limit of detection by satellite (after Spolaor et al., 2014).
464
465 **Figure 6.** Comparison between bromine excess (Br_{exc}) concentrations, bromine enrichment (Br_{enr})
466 ratios and Sodium (ng g^{-1}) from Akademii Nauk ice core. Raw (grey line) and 3-year smoothed (black
467 line) Br_{exc} concentrations and raw (pink line) and 3-year smoothed (red line) Br_{enr} are shown on
468 logarithmic scales. Raw (light blue) and 3-year smoothed (dark blue) sodium data are also shown.
469
470 **Figure 7**. Br_{exc} compared with sea ice area during spring and summer in the Laptev Sea region. In both
471 panels, raw (grey line) and 3-year smoothed (black line) Br_{exc} data from Akademii Nauk ice core are
472 shown. The Laptev Sea spring sea ice data are based on satellite observations (blue line, panel a).
473 Laptev Sea summer sea ice anomalies (red line, panel b) are from Polyakov et al. (2003).
474
475 **Figure 8**. Iodine compared with Laptev Sea summer sea ice area. In both panels, raw (grey line) and 3-
476 year smoothed (black line) [I] data from Akademii Nauk ice core are shown. The Laptev Sea summer
477 sea ice data are based on satellite observations (blue line, panel a) and the reconstruction of Polyakov et
478 al. (2003) (red line, panel b).
479

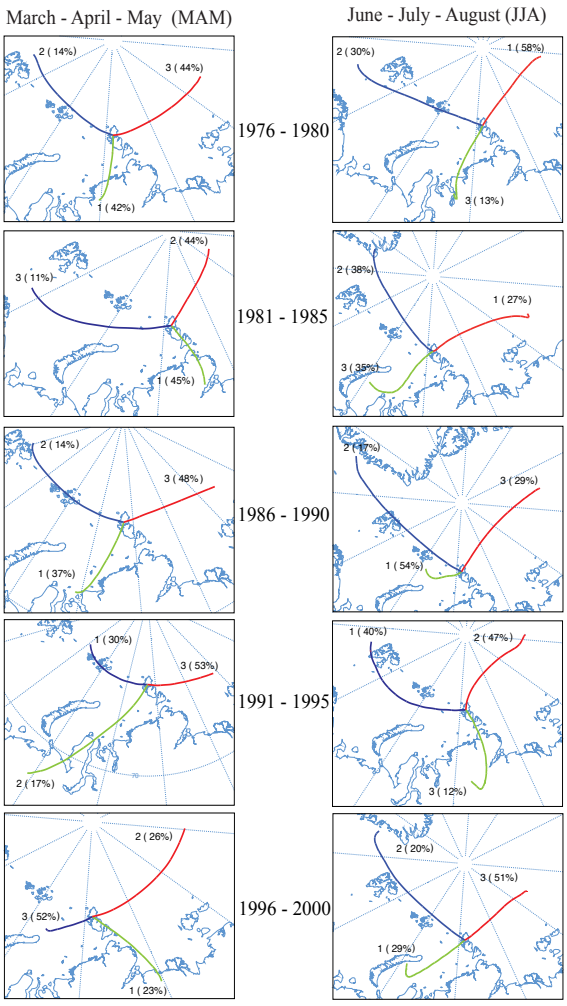
480 **FIGURE 1**

481

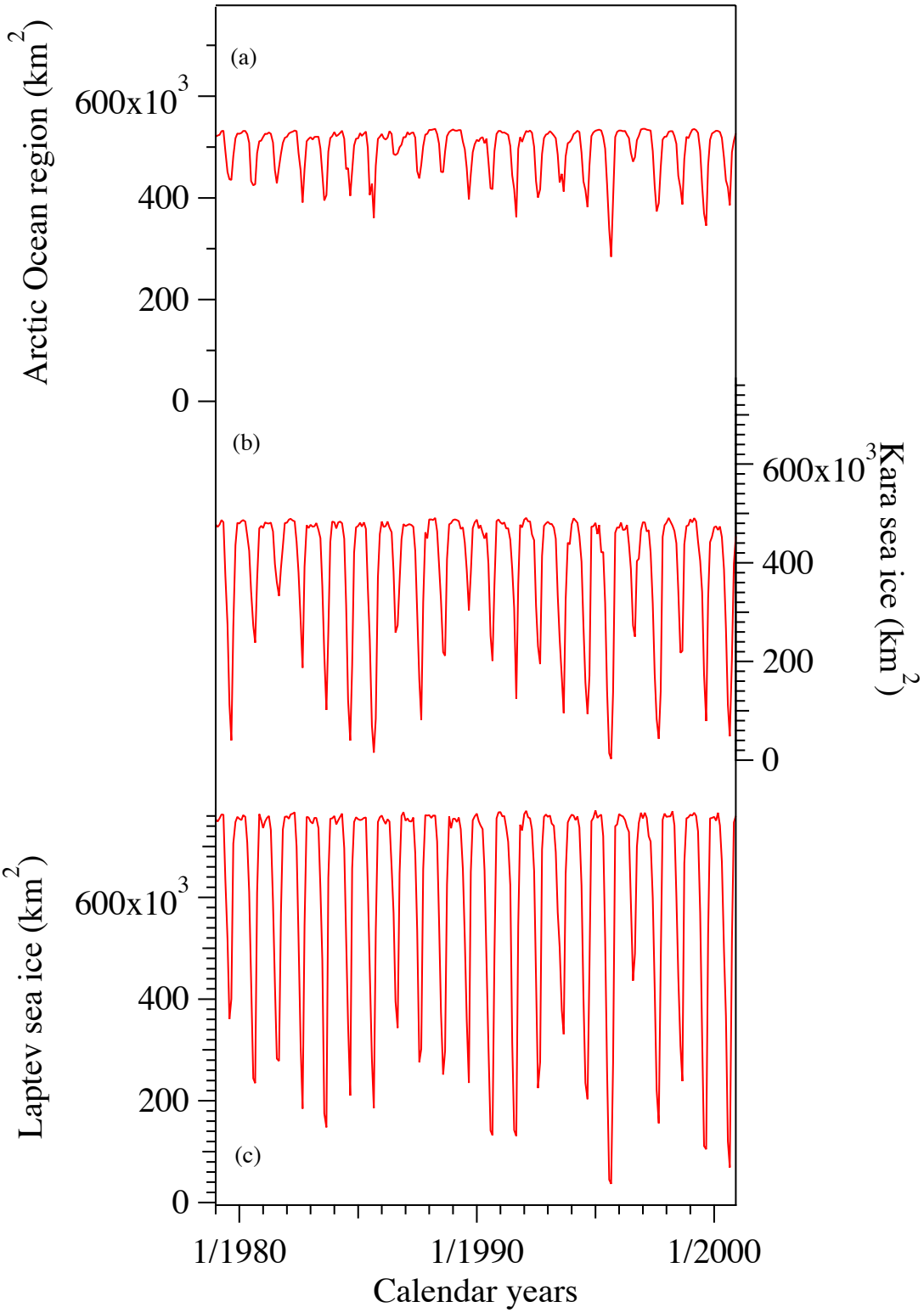


482

FIGURE 2


March - April - May (MAM)
June - July - August (JJA)
1976 - 1980
2 (14%)
3 (44%)
1 (42%)
2 (30%)
1 (58%)
3 (13%)
1981 - 1985
2 (44%)
3 (11%)
1 (45%)
2 (38%)
1 (27%)
3 (35%)
1986 - 1990
2 (14%)
3 (48%)
1 (37%)
2 (17%)
3 (29%)
1 (54%)
1991 - 1995
1 (30%)
3 (53%)
2 (17%)
1 (40%)
2 (47%)
3 (12%)
1996 - 2000
2 (26%)
3 (52%)
1 (23%)
2 (20%)
3 (51%)
1 (29%)

488 **FIGURE 3**
489
490



491
492

FIGURE 4

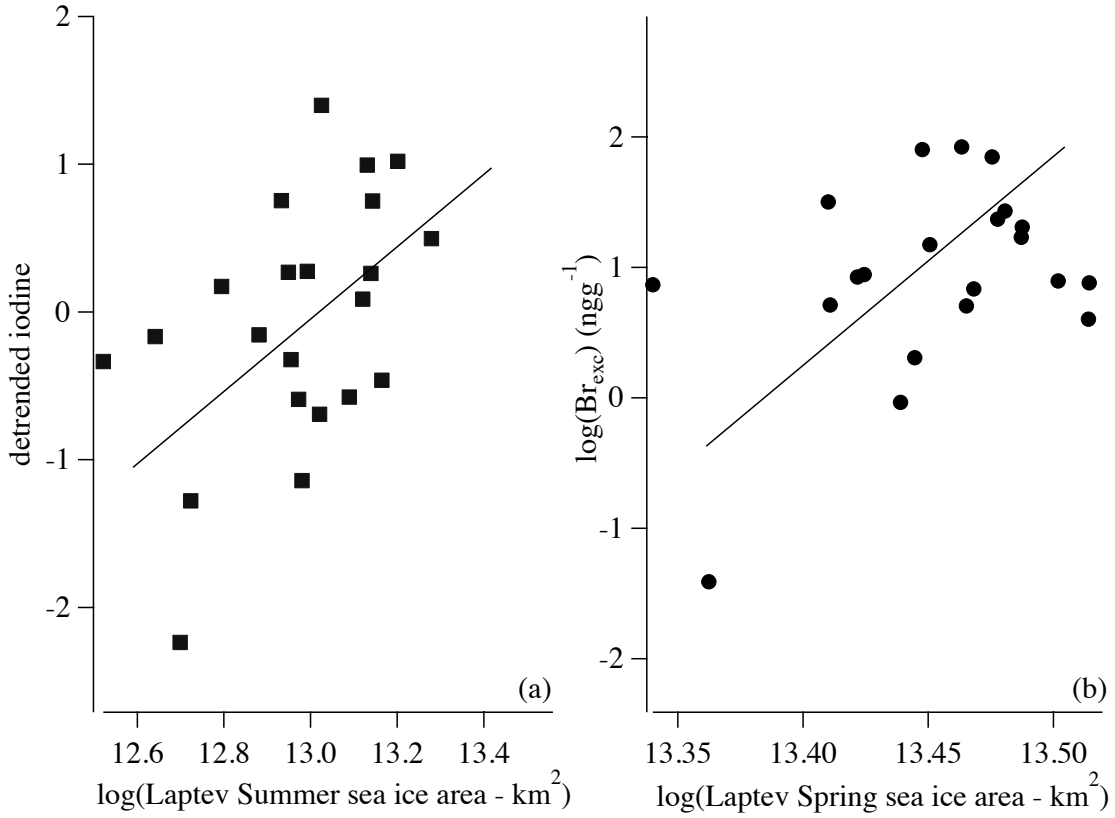
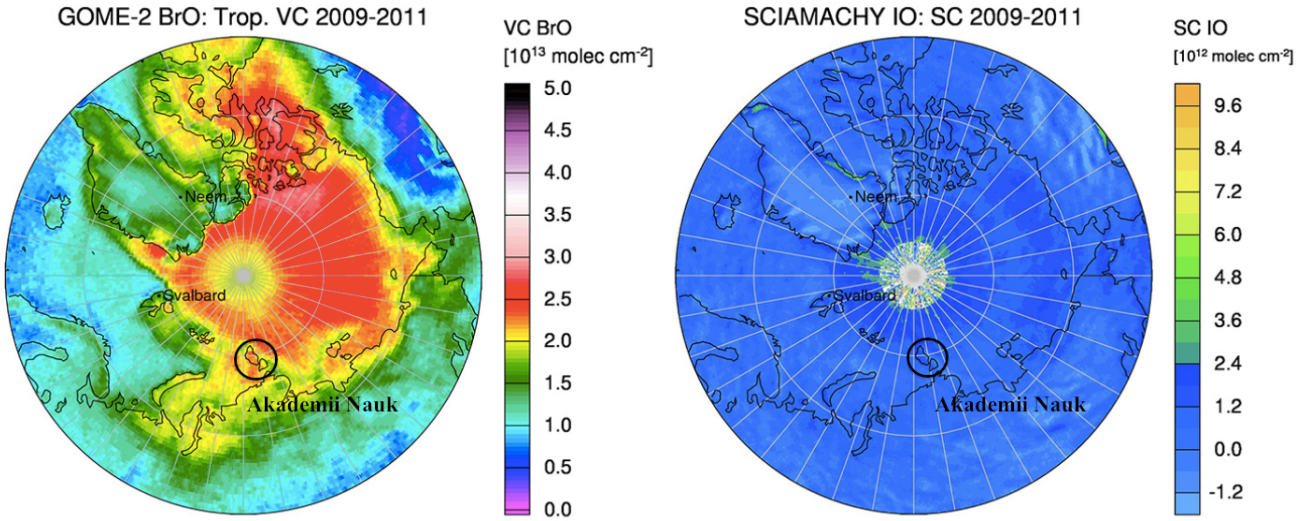
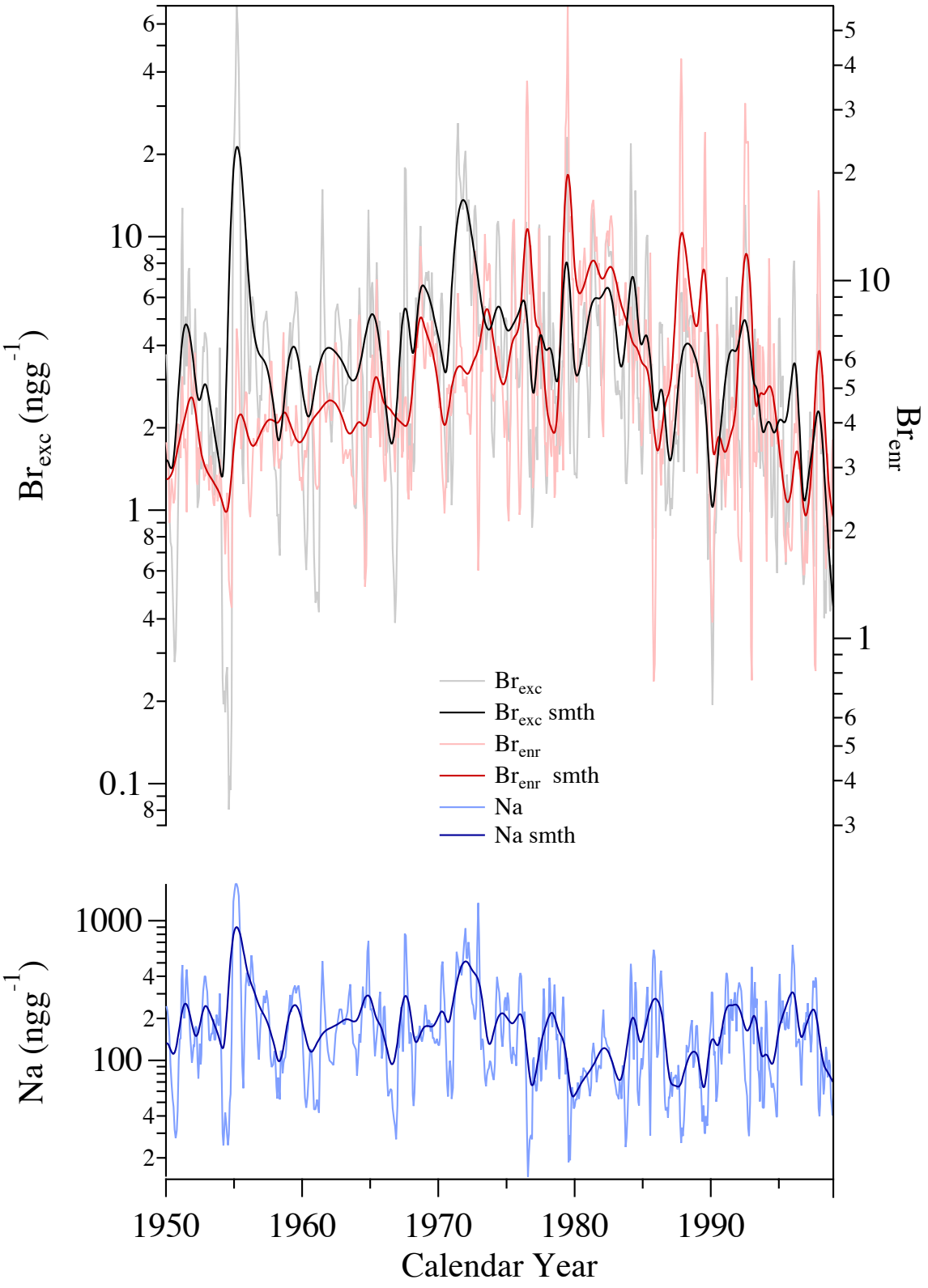

detrended iodine
2
1
0
-1
-2
(a)
12.6
12.8
13.0
13.2
13.4
log(Laptev Summer sea ice area - km^2)
log(Br$_{exc}$) (ngg^{-1})
2
1
0
-1
-2
(b)
13.35
13.40
13.45
13.50
log(Laptev Spring sea ice area - km^2)

FIGURE 5

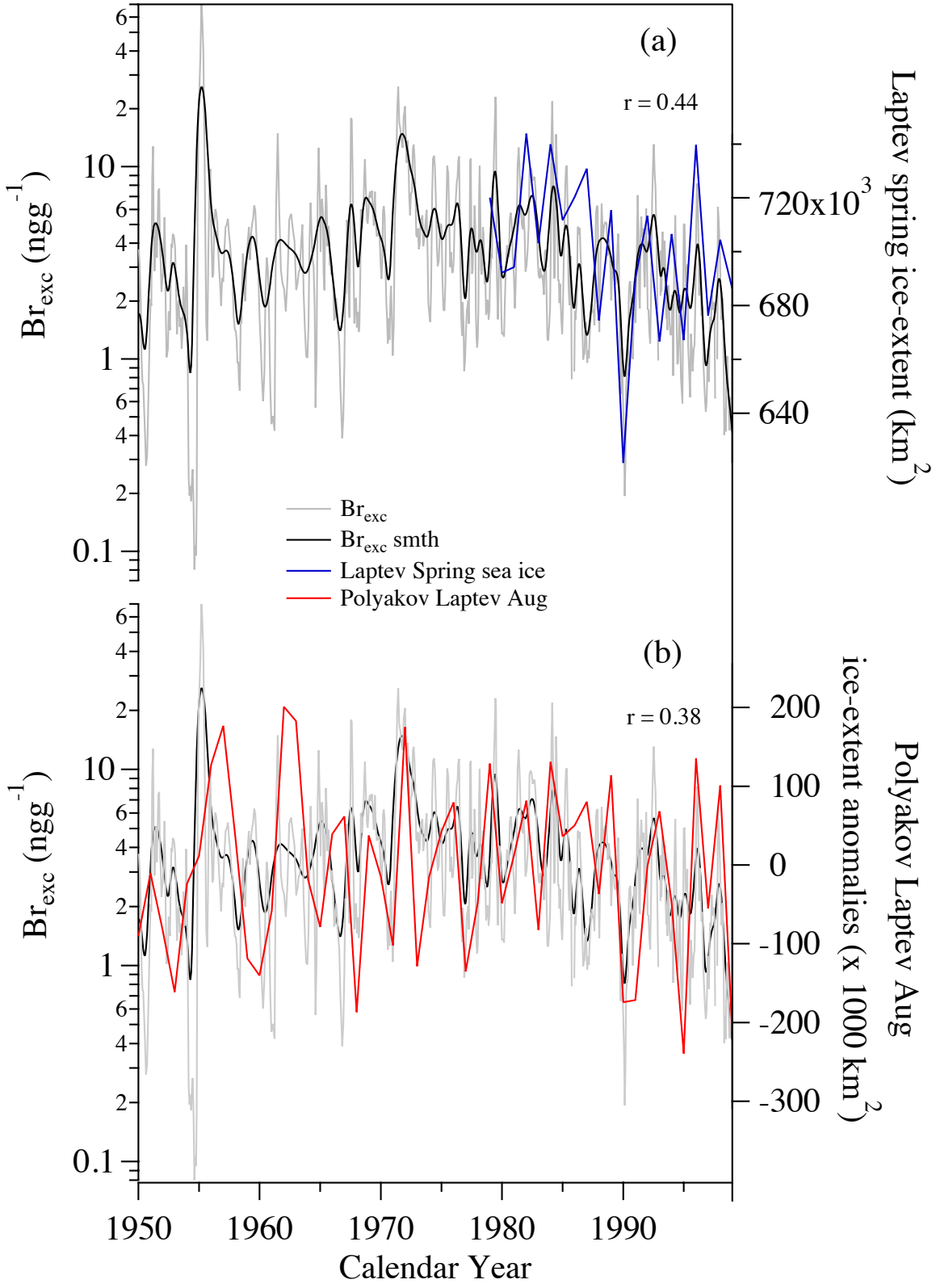

GOME-2 BrO: Trop. VC 2009-2011
VC BrO
[10^13 molec cm^-2]
5.0
4.5
4.0
3.5
3.0
2.5
2.0
1.5
1.0
0.5
0.0
Ny-Ålesund
Svalbard
Akademii Nauk
SCIAMACHY IO: SC 2009-2011
SC IO
[10^12 molec cm^-2]
9.6
8.4
7.2
6.0
4.8
3.6
2.4
1.2
0.0
-1.2
Ny-Ålesund
Svalbard
Akademii Nauk

551 **FIGURE 6**
552
553



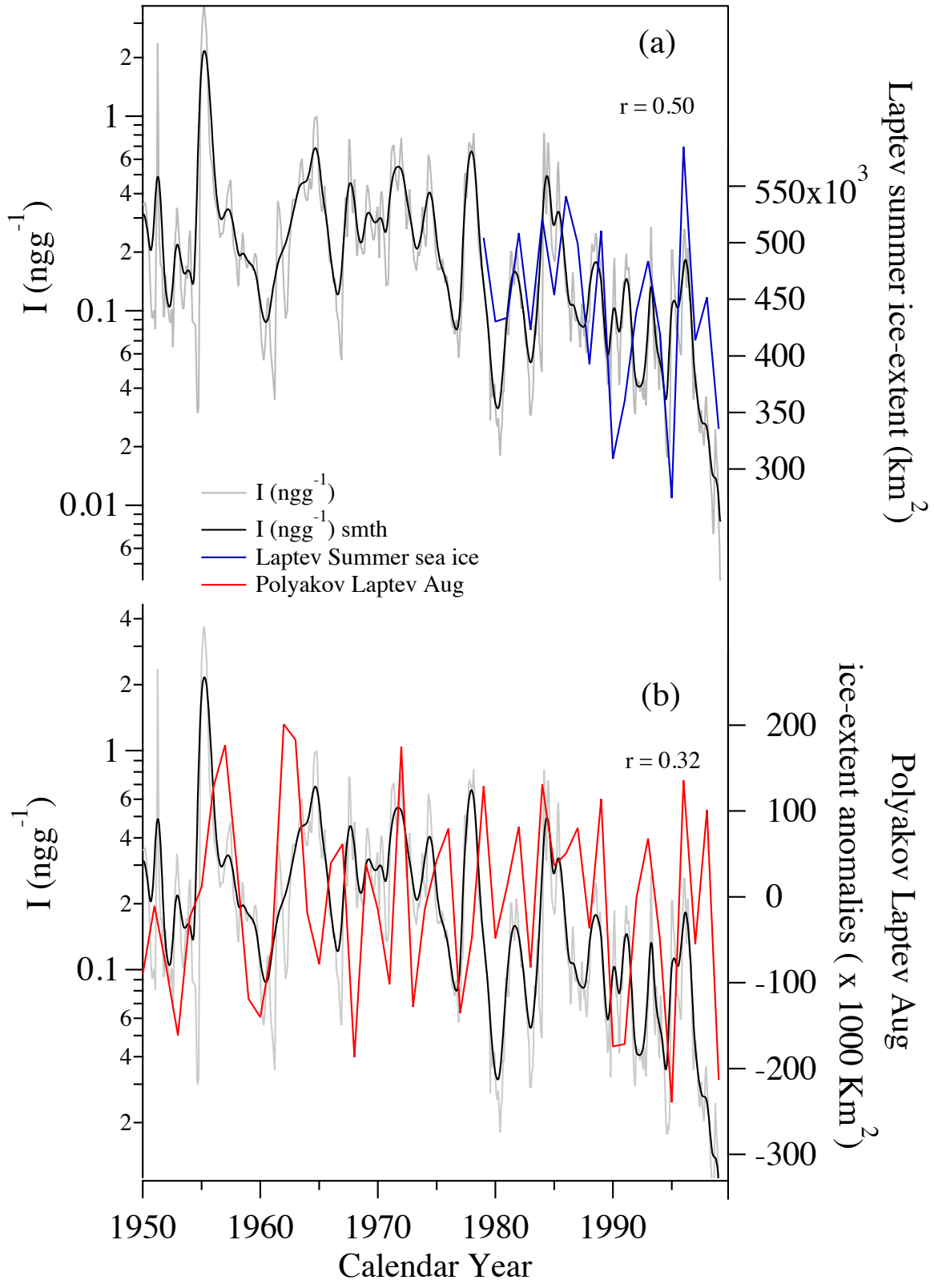
554

555 **FIGURE 7**
556



557
558

559 **FIGURE 8**



560
561

562 **REFERENCES**

563 Abram, N. J., Wolff, E. W., and Curran, M. A. J.: A review of sea ice proxy information from polar ice
564 cores, Quaternary Science Reviews, 79, 168-183, http://dx.doi.org/10.1016/j.quascirev.2013.01.011,
565 2013.
566 Allan, J. D., Williams, P. I., Najera, J., Whitehead, J. D., Flynn, M. J., Taylor, J. W., Liu, D.,
567 Darbyshire, E., Carpenter, L. J., Chance, R., Andrews, S. J., Hackenberg, S. C., and McFiggans, G.:
568 Iodine observed in new particle formation events in the Arctic atmosphere during ACCACIA, Atmos.
569 Chem. Phys., 15, 5599-5609, 10.5194/acp-15-5599-2015, 2015.
570 Ardyna, M., Babin, M., Gosselin, M., Devred, E., Bélanger, S., Matsuoka, A., and Tremblay, J.-E.:
571 Parameterization of vertical chlorophyll a in the Arctic Ocean: impact of the subsurface chlorophyll
572 maximum on regional, seasonal, and annual primary production estimates, Biogeosciences, 10, 4383–
573 4404, doi:10.5194/bg-10-4383-2013, 2013.
574 Atkinson, H. M., Huang, R. J., Chance, R., Roscoe, H. K., Hughes, C., Davison, B., Sch‚àö‚àÇnhardt,
575 A., Mahajan, A. S., Saiz-Lopez, A., Hoffmann, T., and Liss, P. S.: Iodine emissions from the sea ice of
576 the Weddell Sea, Atmos. Chem. Phys. Discuss., 12, 11595-11639, 2012.
577 Barrie, L. A., Yi, Y., Leaitch, W. R., Lohmann, U., Kasibhatla, P., Roelofs, G. J., Wilson, J.,
578 McGovern, F., Benkovitz, C., Melieres, M. A., Law, K., Prospero, J., Kritz, M., Bergmann, D.,
579 Bridgeman, C., Chin, M., Christensen, J., Easter, R., Feichter, J., Land, C., Jeuken, A., Kjellstrom, E.,
580 Koch, D., and Rasch, P.: A comparison of large-scale atmospheric sulphate aerosol models (COSAM):
581 overview and highlights, Tellus B, 53, 615-645, 2001.
582 Belt, S. T., Massé, G., Rowland, S. J., Poulin, M., Michel, C., and LeBlanc, B.: A novel chemical fossil
583 of palaeo sea ice: IP25, Organic Geochemistry, 38, 16-27, 2007.
584 Comiso, J. C.: Large Decadal Decline of the Arctic Multiyear Ice Cover, J Climate, 25, 1176-1193,
585 10.1175/jcli-d-11-00113.1, 2011.
586 Curran, M. A., Van Ommen, T. D., Morgan, V. I., Phillips, K. L., and Palmer., M. R.: Ice core
587 evidence for Antarctic sea ice decline since the 1950s., Science, 302, 1203-1206, 2003.
588 Darby, D. A.: Sources of sediment found in sea ice from the western Arctic Ocean, new insights into
589 processes of en- trainment and drift patterns, J. Geophys. Res., 108, 3257, doi:10.1029/2002JC001350,
590 2003.
591 de Gois, J., Vallelonga, P., Spolaor, A., Devulder, V., Borges, D. G., and Vanhaecke, F.: Bromine
592 isotope ratio measurements in seawater by multi-collector inductively coupled plasma-mass
593 spectrometry with a conventional sample introduction system, Anal Bioanal Chem, 1-8,
594 10.1007/s00216-015-8820-1, 2015.
595 Dieckmann, G. S., and Hellmer, H. H.: The importance of sea ice: An overview in Sea ice, edited by:
596 Dieckmann, D. N. T. a. G. S., Blackwell Publishing Ltd, UK, 2010.
597 Divine, D. V., and Dick, C.: Historical variability of sea ice edge position in the Nordic Seas, Journal
598 of Geophysical Research: Oceans, 111, C01001, 10.1029/2004jc002851, 2006.
599 Draxler, R. R., and Hess, G. D.: An overview of the HYSPLIT_4 modeling system of trajectories,
600 dispersion, and deposition, Aust. Meteor. Mag., 47, 295-308, 1998.
601 Francis, J. A., Chan, W., Leathers, D. J., Miller, J. R., and Veron, D. E.: Winter Northern Hemisphere
602 weather patterns remember summer Arctic sea-ice extent, Geophys. Res. Lett., 36, L07503,
603 doi:10.1029/2009GL037274, 2009.
604 Fritzsche, D., Wilhelms, F., Savatyugin, L. M., Pinglot, J. F., Meyer, H., Hubberten, H. W., and Miller,
605 H.: A new deep ice core from Akademii Nauk ice cap, Severnaya Zemlya, Eurasian Arctic: first results,
606 Annals of Glaciology, 35, 25-28, 2002.

607 Fritzsche, D., Schütt, R., Meyer, H., Miller, H., Wilhelms, F., Opel, T., and Savatyugin, L. M.: A 275
608 year ice-core record from Akademii Nauk ice cap, Severnaya Zemlya, Russian Arctic, Annals of
609 Glaciology, 42, 361-366, 2005.
610 Granfors, A., Ahnoff, M. M., Mills, M. M., and Abrahamsson, K.: Organic iodine in Antarctic sea ice:
611 A comparison between winter in the Weddell Sea and summer in the Amundsen Sea, JOURNAL OF
612 GEOPHYSICAL RESEARCH, doi: 10.1002/2014JG002727, 2014.
613 Holland, M. M., Bitz, C. M., Eby, M., and Weaver, A. J.: The Role of Ice-Ocean Interactions in the
614 Variability of the North Atlantic Thermohaline Circulation, J. Climate, 14, 656–675, 2001.
615 Isaksson, E., Kekonen, T., Moore, J., and Mulvaney, R.: The methanesulfonic acid (MSA) record in a
616 Svalbard ice core,, Ann. Glaciol., 42, 345–351, 2005.
617 Kevin R. Arrigo, Donald K. Perovich, Robert S. Pickart, Zachary W. Brown, Gert L. van Dijken, Kate
618 E. Lowry, Matthew M. Mills, Molly A. Palmer, William M. Balch, Frank Bahr, Nicholas R. Bates,
619 Claudia Benitez-Nelson, Bruce Bowler, Emily Brownlee, Jens K. Ehn, Karen E. Frey, Rebecca Garley,
620 Samuel R. Laney, Laura Lubelczyk, Jeremy Mathis, Atsushi Matsuoka, B. Greg Mitchell, G. W. K.
621 Moore, Eva Ortega-Retuerta, Sharmila Pal, Chris M. Polashenski, Rick A. Reynolds, Brian Schieber,
622 Heidi M. Sosik, Michael Stephens, and Swift, J. H.: Massive Phytoplankton Blooms Under Arctic Sea
623 Ice, Science, 336, 1408, 10.1126/science.1215065, 2012.
624 Kinnard, C., Zdanowicz, C. M., Fisher, D. A., Isaksson, E., De Vernal, A., and Thompson, L. G.:
625 Reconstructed changes in Arctic sea ice over the past 1,450 years, Nature, 479, 509–512, 2011.
626 Levine, J. G., Yang, X., Jones, A. E., and Wolff, E. W.: Sea salt as an ice core proxy for past sea ice
627 extent: A process-based model study, Journal of Geophysical Research: Atmospheres, 119,
628 2013JD020925, 10.1002/2013jd020925, 2014.
629 Lisitzin, A. P.: Sea-ice and Iceberg Sedimentation in the Ocean: Recent and Past, edited by: Springer-
630 Verlag, Berlin, Heidelberg, 2002.
631 Mahajan, A. S., Shaw, M., Oetjen, H., Hornsby, K. E., Carpenter, L. J., Kaleschke, L., Tian-Kunze, X.,
632 Lee, J. D., Moller, S. J., Edwards, P., Commane, R., Ingham, T., Heard, D. E., and Plane, J. M. C.:
633 Evidence of reactive iodine chemistry in the Arctic boundary layer, Journal of Geophysical Research:
634 Atmospheres, 115, D20303, 10.1029/2009jd013665, 2010.
635 Maselli, O., Fritzche D, L., L., J., M., and Meyer, H.: Comparison of water isotope ratio determinations
636 using two cavity ring-down instruments and classical mass spectrometry in continuous ice-core
637 analysis, Isotopes in Environmental and Health Studies, doi:10.1080/10256016.2013.781598, 2013.
638 McConnell, J. R., Lamorey, G. W., Lambert, S. W., and Taylor, K. C.: Continuous Ice-Core Chemical
639 Analyses Using Inductively Coupled Plasma Mass Spectrometry, Environ Sci Technol, 36, 7-11,
640 10.1021/es011088z, 2002.
641 McConnell, J. R., Edwards, R., Kok, G. L., Flanner, M. G., Zender, C. S., Saltzman, E. S., Banta, J. R.,
642 Pasteris, D. R., Carter, M. M., and Kahl, J. D. W.: 20th-Century Industrial Black Carbon Emissions
643 Altered Arctic Climate Forcing, Science, 317, 1381-1384, 2007.
644 McConnell, J. R., Maselli, O. J., Sigl, M., Vallelonga, P., Neumann, T., Ansch√°tz, H., Bales, R. C.,
645 Curran, M. A. J., Das, S. B., Edwards, R., Kipfstuhl, S., Layman, L., and Thomas, E. R.: Antarctic-
646 wide array of high-resolution ice core records reveals pervasive lead pollution began in 1889 and
647 persists today, Sci. Rep., 4, 10.1038/srep05848
648 http://www.nature.com/srep/2014/140728/srep05848/abs/srep05848.html - supplementary-information,
649 2014.
650 Müller, J., Masse, G., Stein, R., and Belt, S. T.: Variability of sea-ice conditions in the Fram Strait over
651 the past 30,000 years, Nature Geosci, 2, 772-776, 2009.

652 Opel, T., Fritzsche, D., Meyer, H., Schütt, R., Weiler, K., Ruth, U., Wilhelms, F., and Fischer, H.: 115
653 year ice-core data from Akademii Nauk ice cap, Severnaya Zemlya: high-resolution record of Eurasian
654 Arctic climate change, 2009.
655 Opel, T., Fritzsche, D., and Meyer, H.: Eurasian Arctic climate over the past millennium as recorded in
656 the Akademii Nauk ice core (Severnaya Zemlya), Clim. Past, 9, 2379-2389, 10.5194/cp-9-2379-2013,
657 2013.
658 Pasteris, D. R., McConnell, J. R., Das, S. B., Criscitiello, A. S., Evans, M. J., Maselli, O. J., Sigl, M.,
659 and Layman, L.: Seasonally resolved ice core records from West Antarctica indicate a sea ice source of
660 sea-salt aerosol and a biomass burning source of ammonium, Journal of Geophysical Research:
661 Atmospheres, 119, 2013JD020720, 10.1002/2013jd020720, 2014.
662 Petit, J. R., Jouzel, J., Raynaud, D., Barkov, N. I., Barnola, J. M., Basile, I., Bender, M., Chappellaz, J.,
663 Davis, M., Delaygue, G., Delmotte, M., Kotlyakov, V. M., Legrand, M., Lipenkov, V. Y., Lorius, C.,
664 Pepin, L., Ritz, C., Saltzman, E., and Stievenard, M.: Climate and atmospheric history of the past
665 420,000 years from the Vostok ice core, Antarctica, Nature, 399, 429-436, 1999.
666 Pohjola, V. A., Moore, J. C., Isaksson, E., Jauhiainen, T., Van de Wal, R. S. W., Martma, T., Meijer, H.
667 A. J., and Vaikmäe, R.: Effect of periodic melting on geochemical and isotopic signals in an ice core
668 from Lomonosovfonna, Svalbard, J. Geophys. Res. Atmos., 107, 1–14, 2002.
669 Polyak, L., Alley, R. B., Andrews, J. T., Brigham-Grette, J., Cronin, T. M., Darby, D. A., Dyke, A. S.,
670 Fitzpatrick, J. J., Funder, S., Holland, M., Jennings, A. E., Miller, G. H., O'Regan, M., Savelle, J.,
671 Serreze, M., St. John, K., White, J. W. C., and Wolff, E.: History of sea ice in the Arctic, Quaternary
672 Sci. Rev., 29, 1757–1778, 2010.
673 Polyakov, I. V., Alekseev, G. V., Bekryaev, R. V., Bhatt, U. S., Colony, R., Johnson, M. A., Karklin,
674 V. P., Walsh, D., and Yulin, A. V.: Long-term ice variability in Arctic marginal seas, J Climate, 16,
675 2078-2085, 2003.
676 Pratt, K. A., Custard, K. D., Shepson, P. B., Douglas, T. A., Pohler, D., General, S., Zielcke, J.,
677 Simpson, W. R., Platt, U., Tanner, D. J., Gregory Huey, L., Carlsen, M., and Stirm, B. H.:
678 Photochemical production of molecular bromine in Arctic surface snowpacks, Nature Geosci., 6, 351-
679 356, doi:10.1038/ngeo1779, 2013., 2013.
680 Röthlisberger, R., Crosta, X., Abram, N. J., Armand, L., and Wolff, E. W.: Potential and limitations of
681 marine and ice core sea ice proxies: an example from the Indian Ocean sector, Quaternary Science
682 Reviews, 29, 296-302, 2010.
683 Saiz-Lopez, A., Mahajan, A. S., Salmon, R. A., Bauguitte, S. J. B., Jones, A. E., Roscoe, H. K., and
684 Plane, J. M. C.: Boundary Layer Halogens in Coastal Antarctica, Science, 317, 348-351, 2007.
685 Saiz-Lopez, A., Blaszczak-Boxe, C. S., and Carpenter, L. J.: A mechanism for biologically induced
686 iodine emissions from sea ice, Atmos. Chem. Phys., 15, 9731-9746, 10.5194/acp-15-9731-2015, 2015.
687 Sander, R., Burrows, J., and Kaleschke, L.: Carbonate precipitation in brine – a potential trigger
688 for tropospheric ozone depletion events, Atmos. Chem. Phys., 6, 4653-4658, 10.5194/acp-6-4653-
689 2006, 2006.
690 Schonhardt, A., Begoin, M., Richter, A., Wittrock, F., Kaleschke, L., Gomez Martin, J. C., and
691 Burrows, J. P.: Simultaneous satellite observations of IO and BrO over Antarctica, Atmos Chem Phys,
692 12, 6565--6580, 2012.
693 Sigl, M., McConnell, J. R., Toohey, M., Curran, M., Das, S. B., Edwards, R., Isaksson, E., Kawamura,
694 K., Kipfstuhl, S., Kruger, K., Layman, L., Maselli, O. J., Motizuki, Y., Motoyama, H., Pasteris, D. R.,
695 and Severi, M.: Insights from Antarctica on volcanic forcing during the Common Era, Nature Clim.
696 Change, 4, 693-697, 10.1038/nclimate2293

http://www.nature.com/nclimate/journal/v4/n8/abs/nclimate2293.html - supplementary-information, 2014.
Simpson, W. R., Carlson, D., Hoenninger, G. n. G., Douglas, T. A., Sturm, M., Perovich, D., and Platt, U.: First-year sea-ice contact predicts bromine monoxide (BrO) levels at Barrow, Alaska better than potential frost flower contact, Atmos. Chem. Phys., 7, 621-627, 2007.
Simpson, W. R., Brown, S. S., Saiz-Lopez, A., Thornton, J. A., and Glasow, R. v.: Tropospheric Halogen Chemistry: Sources, Cycling, and Impacts, Chemical Reviews, 115, 4035-4062, 10.1021/cr5006638, 2015.
Spolaor, A., Gabrieli, J., Martma, T., Kohler, J., Björkman, M. B., Isaksson, E., Varin, C., Vallelonga, P., Plane, J. M. C., and Barbante, C.: Sea ice dynamics influence halogen deposition to Svalbard, The Cryosphere, 7, 1645-1658, 10.5194/tc-7-1645-2013, 2013a.
Spolaor, A., Vallelonga, P., Plane, J. M. C., Kehrwald, N., Gabrieli, J., Varin, C., Turetta, C., Cozzi, G., Kumar, R., Boutron, C., and Barbante, C.: Halogen species record Antarctic sea ice extent over glacial-interglacial periods, Atmos. Chem. Phys., 13, 6623-6635, 10.5194/acp-13-6623-2013, 2013b.
Spolaor, A., Vallelonga, P., Gabrieli, J., Martma, T., Björkman, M. P., Isaksson, E., Cozzi, G., Turetta, C., Kjær, H. A., Curran, M. A. J., Moy, A. D., Schönhardt, A., Blechschmidt, A. M., Burrows, J. P., Plane, J. M. C., and Barbante, C.: Seasonality of halogen deposition in polar snow and ice, Atmos. Chem. Phys., 14, 9613-9622, 10.5194/acp-14-9613-2014, 2014.
Stroeve, J., Holland, M. M., Meier, W., Scambos, T., and Serreze, M.: Arctic sea ice decline: Faster than forecast, Geophys. Res. Lett., 34, L09501, 10.1029/2007GL029703, 2007.
Sturges, W. T., and Harrison, R. M.: Bromine in marine aerosols and the origin, nature and quantity of natural atmospheric bromine, Atmospheric Environment (1967), 20, 1485-1496, 1986.
Sturges, W. T., and Barrie, L. A.: Chlorine, Bromine and Iodine in arctic aerosols, Atmos Environ, 22, 1179-1194, 10.1016/0004-6981(88)90349-6, 1988.
Turekian, K. K.: Oceans, edited by: Englewood, C., N.J., Prentice-Hall, 1968.
Turner, J., Bracegirdle, T., Phillips, T., Marshall, G. J., and Hosking, J. S.: An Initial Assessment of Antarctic Sea Ice Extent in the CMIP5 Models, J Climate, In press, doi: 10.1175/jcli-d-12- 00068.1, 2012.
Vinje, T.: Anomalies and Trends of Sea-Ice Extent and Atmospheric Circulation in the Nordic Seas during the Period 1864,Äì1998, J Climate, 14, 255-267, 10.1175/1520-0442(2001)014<0255:aatosi>2.0.co;2, 2001.
Vogt, R., Crutzen, P. J., and Sander, R.: A mechanism for halogen release from sea-salt aerosol in the remote marine boundary layer, Nature, 383, 327-330, 1996.
Weiler, K., Fischer, H., Fritzsche, D., Ruth, U., Wilhelms, F., and Miller, H.: Glaciochemical reconnaissance of a new ice core from Severnaya Zemlya, Eurasian Arctic, Annals of Glaciology, 51, 64-74, 2005.
Wolff, E. W., Fischer, H., Fundel, F., Ruth, U., Twarloh, B., Littot, G. C., Mulvaney, R., Rothlisberger, R., de Angelis, M., Boutron, C. F., Hansson, M., Jonsell, U., Hutterli, M. A., Lambert, F., Kaufmann, P., Stauffer, B., Stocker, T. F., Steffensen, J. P., Bigler, M., Siggaard-Andersen, M. L., Udisti, R., Becagli, S., Castellano, E., Severi, M., Wagenbach, D., Barbante, C., Gabrielli, P., and Gaspari, V.: Southern Ocean sea-ice extent, productivity and iron flux over the past eight glacial cycles, Nature, 440, 491-496, 2006.
Wolff, E. W., Barbante, C., Becagli, S., Bigler, M., Boutron, C. F., Castellano, E., De Angelis, M., Federer, U., Fischer, H., and Fundel, F.: Changes in environment over the last 800,000 years from chemical analysis of the EPICA Dome C ice core, Quaternary Sci Rev, 29, 285-295, 2010.

742 Xiao, X., Fahl, K., and Stein, R.: Biomarker distributions in surface sediments from the Kara and
743 Laptev seas (Arctic Ocean): indicators for organic-carbon sources and sea-ice coverage, Qua Sci Rev,
744 79 40-52, 2013.
745 Zhou, J., Delille, B., Eicken, H., Vancoppenolle, M., Brabant, F., Carnat, G., Geilfus, N.-X.,
746 Papakyriakou, T., Heinesch, B., and Tison, J.-L.: Physical and biogeochemical properties in land- fast
747 sea ice (Barrow, Alaska): Insights on brine and gas dynam- ics across seasons, J. Geophys. Res.-
748 Oceans, 118, 3172-3189, 10.1002/jgrc.20232, 2013.

749