



Supplement of

Contrasted geomorphological and limnological properties of thermokarst lakes formed in buried glacier ice and ice-wedge polygon terrain

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S1 Bathymetric mapping through lake ice with ground-penetrating radar (GPR)

The difference between the dielectric properties of ice, water, and the sediments under lake bottoms makes GPR an effective method for determining lake depth profiles in the field. In this study, we used a pulseEKKO PRO controller manufactured by Sensors and Software coupled with 100 megahertz (MHz) antennas to obtain depth profiles from lake K and lake L. GPR profiles were calibrated and correlated with water depth measurements. Profiles were post-processed using Sensor and Software EKKO Project Version 5 proprietary software. Post-processing included time-zero correction and integration of GPS data, topography, and horizontal filtering to improve visualization of horizontal reflectors.

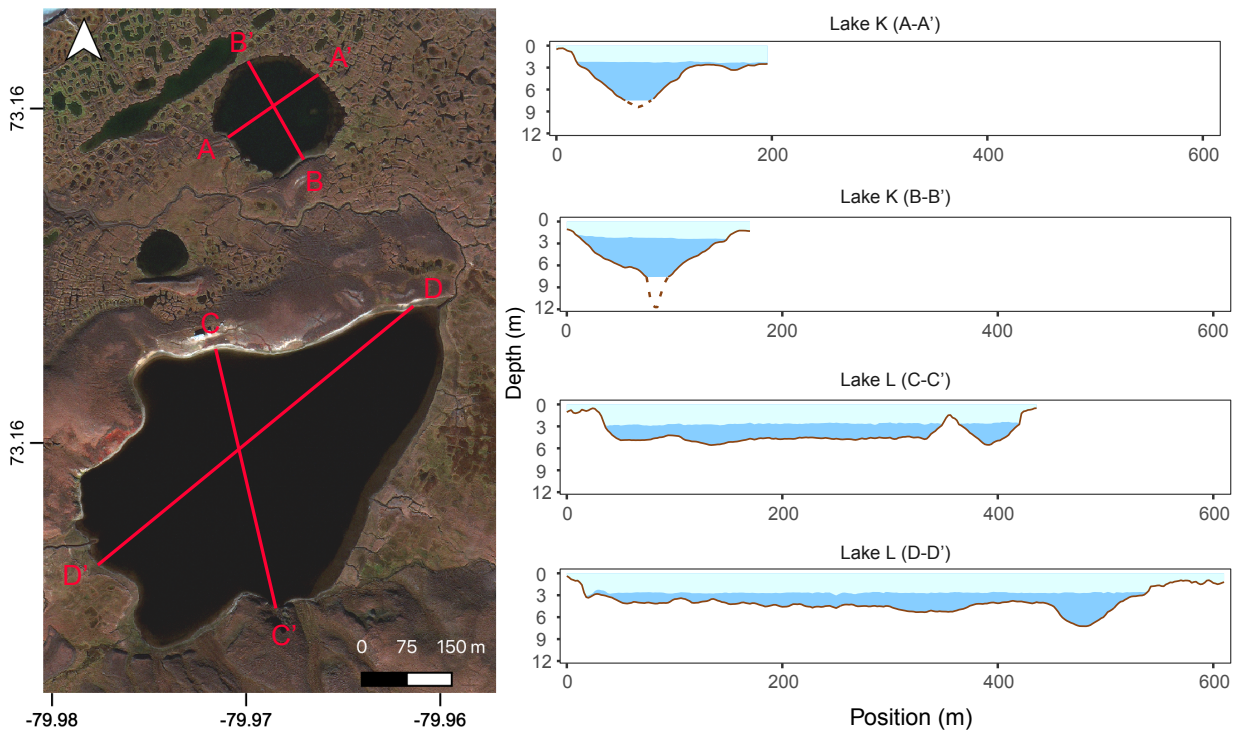
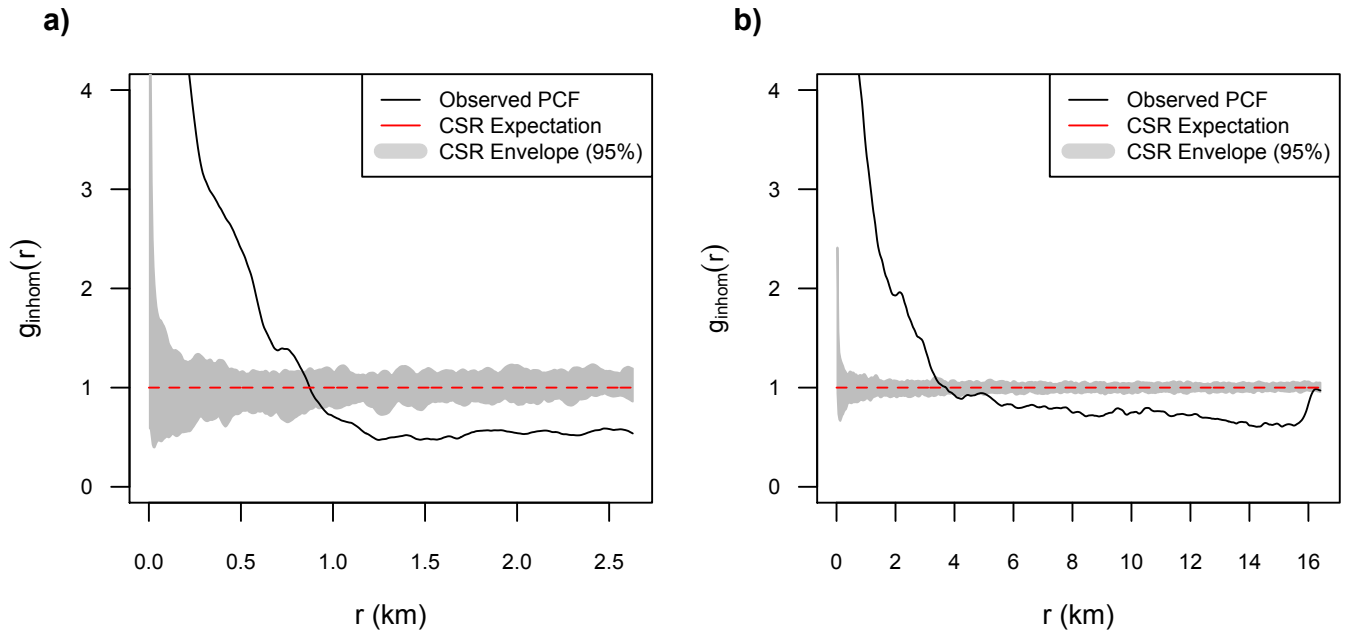


Figure S1: Ground penetrating radar profile of lakes K and L (background: GeoEye, 2010)

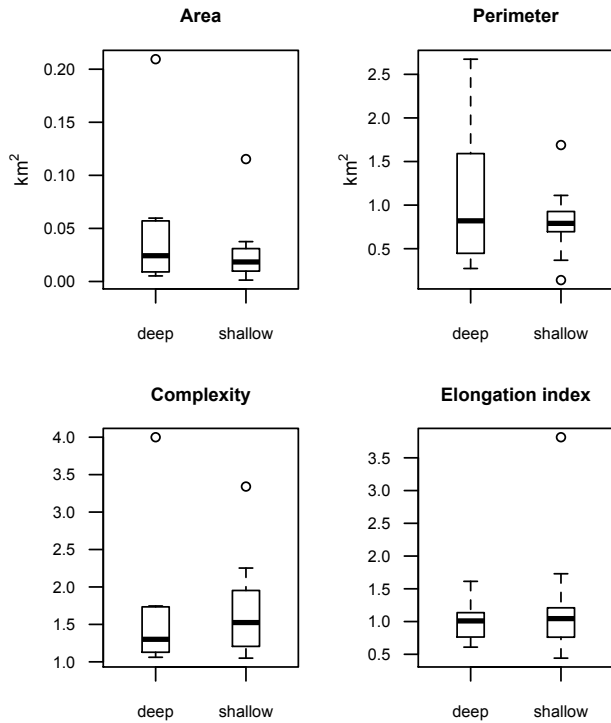
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References

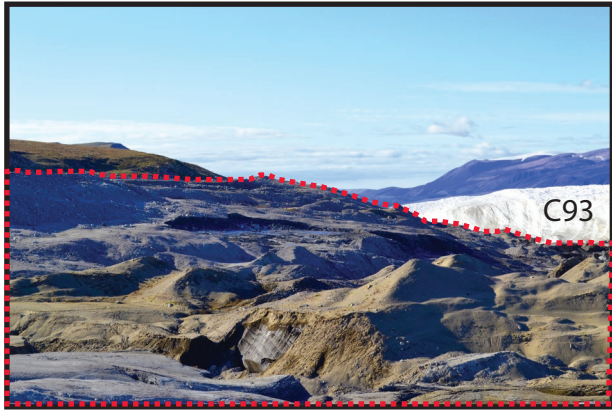
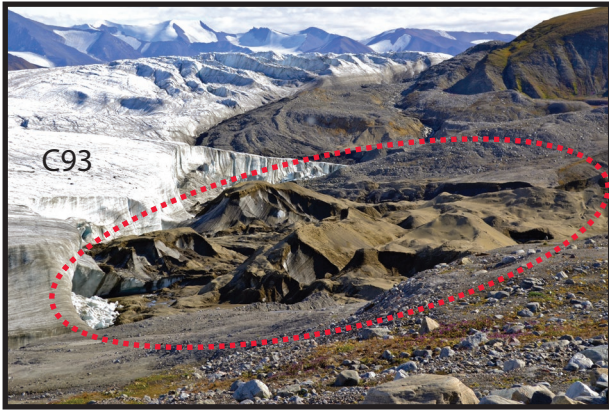
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- 15 Paquette, M., Fortier, D., Mueller, D. R., Sarrazin, D., and Vincent, W. F.: Rapid disappearance of perennial ice on Canada's most northern lake, *Geophys. Res. Lett.*, 42, 1433–1440, <https://doi.org/10.1002/2014GL062960>, 2015.



20 **Figure S2: Pair Correlation Function results for the a) valley and the b) southern plain. The function $g_{inhom}(r)$ is plotted against r , with increasing radii of analysis (in km) away from each lake. The black line shows observed $g(r)$ values, the dashed red line indicates the mean of the inhomogeneous null model and the grey area represents the simulation envelope (95% confidence interval), which is derived from the Monte-Carlo simulations of CSR. Values plotting with the grey envelope are not significantly different from a random spatial arrangement.**

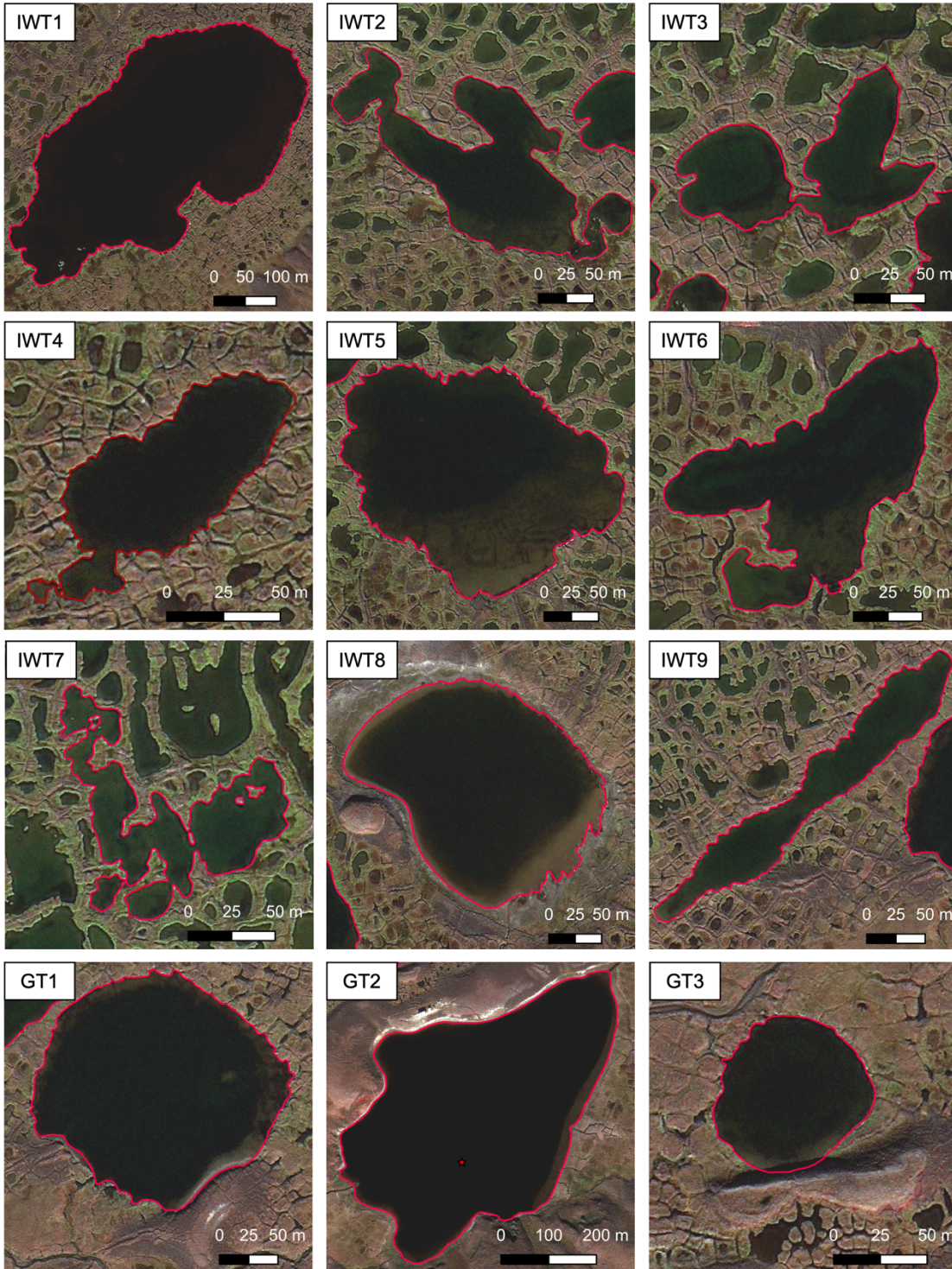


25 **Figure S3: Box plots comparing the morphometric properties (area, perimeter, complexity, elongation index) of deep (>4 m) and shallow lakes (< 4 m) obtained from all the digitised lakes. The thick line marks the median value. The bottom and the top of the box correspond to the first and third quartiles, respectively. The whiskers show the range of observed values that are not within the first and third quartile but not further away than 1.5 times the interquartile range (IQR) from the hinges, and open circles represent outliers.**



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Figure S4: Burial of glacier ice at the margin of glacier C-93.



35 Figure S5: Satellite images of lakes (background: GeoEye, 2010).

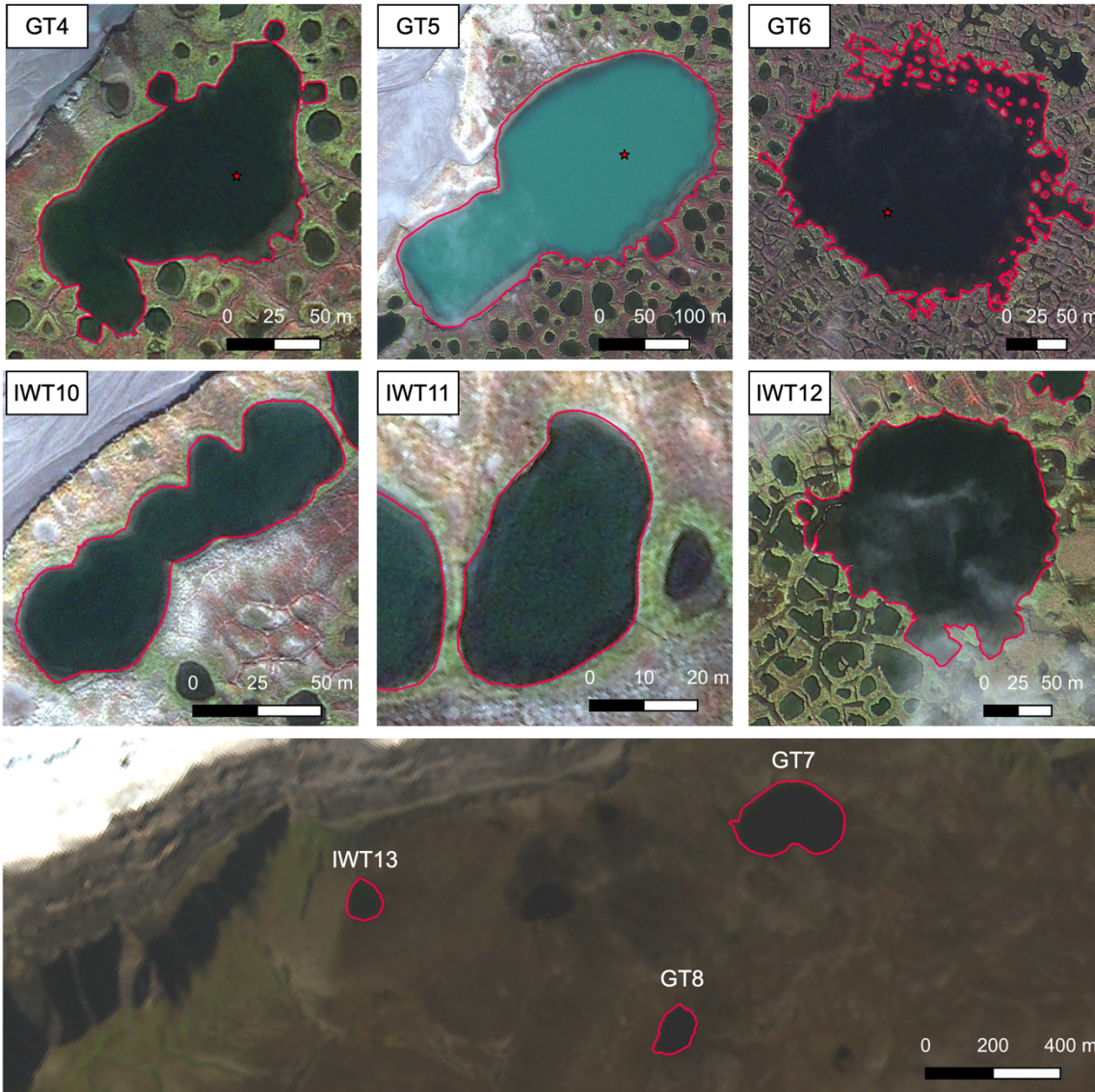


Figure S5: Continued.

S6 Computed tomography (CT) scanning of a sediment core collected in lakes K and G, Bylot Island, Nunavut, Canada

40 To further investigate the cryostratigraphic characteristics of the ice and lake sediments, all samples were observed under X-ray computed tomography (CT) scanning (Siemens SOMATOM Sensation 64) at INRS- ETE (Quebec City, Canada), as in Calmels and Allard (2004). This technique relies on the calculation of the linear attenuation coefficient that measured the density of an object passed through an X-ray beam at different angles (Boespflug et al., 1994). A CT-scan produces cross-sectional images (usually 512 by 512 pixels matrix) of an object where each pixel of the image is assigned an X-ray attenuation
45 value (μ), also called a CT number. CT numbers are then standardized using the Hounsfield scale, where the radiodensity of water (μ_{water}) is arbitrarily defined as 0 HU (Hounsfield units) according to Equation 1 (Hounsfield, 1973). Different shades of gray are assigned specific CT numbers to create the displayed image using a specific image processing software (Fiji) dedicated to DICOM (Digital Imaging and Communications in Medicine) images. In permafrost samples, unconsolidated sediments and rock (high density minerals) appear white, as the attenuation of these materials is very high. Gas inclusions and water appear
50 black and other materials, such as ice, can have various shades of grey depending on their density.

$$\text{HU value} = \frac{\mu - \mu_{\text{water}}}{\mu_{\text{water}}} \times 1000 \quad (1)$$

This tool helps refine cryostratigraphic characterization of permafrost cores as it can reveal characteristics otherwise difficult
55 or even impossible to observe with the naked eye. It allows visualization and characterization of the internal components and structures of the frozen sample, such as ice, grain-size variations, layer orientation in space, and gas inclusions. From a quantitative perspective, it has been used to segment images into regions of ice, gas and sediment in order to quantify the volumetric content of the scanned sample (Calmels et al., 2010; Dillon et al., 2008). Images were processed with FIJI to map X-ray attenuation coefficients on longitudinal images and to visualize sedimentary structures. The resulting images are
60 displayed in greyscale, with darker grey representing a lower X-ray attenuation. Greyscale values are expressed as CT numbers, which are complex units related to the mineralogy, organic matter content, grain size, and bulk density.

References

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